



Proceeding Paper

# Analysis of Optical Properties and Radiative Forcing of Different Aerosol Types in Wuhan <sup>†</sup>

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**Abstract:** The optical and radiative properties of aerosols are governed by their types. In this paper, the optical and radiative properties of different aerosol types in Wuhan, China, have been inverted and investigated using the collected  $PM_{2.5}$  samples. The results show that  $PM_{2.5}$  (average mass concentration about 31.25  $\mu g/m^3$ ) is mainly contributed by sulfate (SO<sub>4</sub>) and organic carbon (OC) (22% and 52%, respectively), while aerosol optical depth (AOD, average about 0.28) is mainly contributed by SO<sub>4</sub> and black carbon (EC) (22% and 19%, respectively). SO<sub>4</sub> and nitrate (NO<sub>3</sub>) have a negative radiative forcing at the top of the atmosphere (TOA) and a cooling effect, while OC and EC have a positive radiative forcing and a heating effect. Moreover, EC has the most significant effect on the radiative forcing in Wuhan, contributing up to 61% and 73% at the bottom of the atmosphere (BOA) and atmosphere (ATM), respectively, while contributing up to 75% to the atmospheric heating rate (about  $1\sim2$  K/day).

Keywords: aerosol optical properties; PM2.5; direct radiative forcing; atmospheric heating rate

# 1. Introduction

Aerosols are one of the most influential factors in the Earth system, as an important component of the atmosphere. Aerosols affect the Earth's energy balance via direct and indirect radiative effects, the direct radiative effect being due to the aerosol absorption and scattering of solar radiation, and the indirect radiative effect coming from the fact that aerosols act as cloud condensation nuclei changing the life cycle and physical properties of clouds [1]. Additionally, aerosols can affect human health as respirable particulate matter and reduce visibility in highly polluted regions [2]. However, these effects are usually highly uncertain because aerosols are much more varied in composition and much more heterogeneous in spatial and temporal distributions than greenhouse gases, and there is still a lack of sufficient aerosol information at larger spatial and temporal scales [3].

For decades, recognition of this fact has motivated tremendous efforts, and a great number of in situ measurements and optical remote sensing works have provided valid information on the optical and microphysical properties of aerosols, which have contributed to reducing the uncertainty of aerosols in the global and regional climate system [4,5]. Based on a large number of observation data, various aerosol classification methods have been developed to assess the global and regional distribution characteristics of aerosols in order to improve the level of understanding of aerosols. For example, using a combination of aerosol optical depth (AOD) and Angstrom exponent (AE) index to achieve aerosol classification, or combining AE with single scattering albedo (SSA), as well as using parameters such as AROD [6–8]. Moreover, numerous studies have discussed in depth the components and sources of regional aerosols. Hammer et al. have assessed the trends in global particulate matter concentrations using methods such as satellites and modeling [4]. Studies in China, India, and Europe have also revealed sources and drivers of



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particulate matter at smaller spatial scales [9–11]. Source analysis of aerosols in Beijing and Wuhan reveals the dominant influence of traffic and industrial emissions [2,12–15]. Studies have shown that different aerosol types may have different or even opposite radiative effects [16–18]. Therefore, improving the level of understanding of the optical properties and radiative effects of different aerosol types can effectively limit the aerosol direct and indirect radiative forcing and enhance the accuracy of the assessment of aerosol environmental and climate effects.

Wuhan is a major city in central China with more serious pollution and higher aerosol loading. It is mainly attributed to the rapidly growing industry and the consequent increase in energy consumption in Wuhan. The high ambient humidity and stable planetary boundary layer also limit the pollutant dispersion, which further increases the aerosol loading. Although considerable studies have investigated the composition, sources, and distribution of aerosols in Wuhan, the aerosol radiative effect depends more on the combined contribution of multiple aerosol types, and up to now, there are no relevant studies involving the radiative effects of different aerosol types and their relative contributions in Wuhan [2,13–15]. In this study, an in situ measurement experiment of atmospheric aerosols is carried out in Wuhan. The mass concentrations (MC) of different aerosol types are investigated and analyzed by the collected PM<sub>2.5</sub> samples, and the radiative forcing and atmospheric heating rates of aerosol types are subsequently simulated and discussed based on the inverted aerosol optical properties. This paper aims to reveal the radiative properties of different aerosol types in Wuhan and to limit the uncertainty of aerosol properties and radiative forcing in Wuhan.

#### 2. Materials and Methods

## 2.1. In Situ Measurement of Aerosols

The sampling site is located at Baishazhou Avenue in Wuhan, China, as shown in Figure 1. With a total population of about 17 million, developed industry, and high car ownership, the aerosols over Wuhan are mainly from various anthropogenic sources. In this study, nine sets (days) of  $PM_{2.5}$  data collected during 2021 are used, with a sampling duration of 20 h or 24 h. The equipment and operation of the experiment are described in detail in the earlier study [19], and the sampling results are shown in Figure 2. The results show that the MC of PM<sub>2.5</sub> ranges from  $26\sim37 \,\mu\text{g/m}^3$ , with mean values of about  $31.25 \,\mu\text{g/m}^3$ , and  $SO_4^{2-}$  and  $NO_3^{-}$  range from  $0.463 \sim 1.04 \,\mu\text{g/m}^3$  and  $5.015 \sim 8.19 \,\mu\text{g/m}^3$ , with mean values of about 6.94  $\mu$ g/m<sup>3</sup> and 0.67  $\mu$ g/m<sup>3</sup>, respectively. The MC of organic carbon (OC) and elemental carbon (EC) range from 9.22~16.31 μg/m<sup>3</sup> and 0.95~2.03 μg/m<sup>3</sup>, with mean values of about 11.82  $\mu$ g/m<sup>3</sup> and 1.63  $\mu$ g/m<sup>3</sup>, respectively. PM<sub>2.5</sub> is mainly contributed by SO<sub>4</sub> and OC, accounting for about 22% and 52%, respectively. Several studies show that water-soluble ions such as  $SO_4^{2-}$  and  $NO^{3-}$  in Wuhan mainly originate from vehicle exhaust emissions and organic matter combustion, while EC mainly comes from industrial production and coal-fired heating. In addition, the burning of straw in the surrounding countryside is also one of the main sources of EC, while OC comes more from the incomplete combustion of biomass and fossil fuels [14,15].

#### 2.2. Aerosol Optical Properties

In this study, the Optical Properties of Aerosols and Clouds (OPAC) model is used to implement the calculation of aerosol optical properties [20]. This package provides aerosol optical properties at 61 wavelengths between 0.3  $\mu$ m and 40  $\mu$ m and 8 relative humidities and is widely used for the calculation and inversion of the optical properties of mixed aerosols with multiple components. The core idea of the inversion is to calculate the AOD by inputting the number density of aerosols and matching it with the observation data. Specifically, the corresponding number densities are first calculated from the MC of the different aerosol types obtained using sampling. Subsequently, parameters such as number density, height profile, and relative humidity are input into the OPAC program, and the aerosol optical properties such as AOD, SSA, and asymmetry factor (ASY) are

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calculated and recorded. Finally, the AOD is compared with the observation data, and if the relative error is less than 10%, the aerosol optical properties are retained; otherwise, the parameters are adjusted and recalculated. The AOD used for comparison is the MODIS Aerosol Product from the Level 2 (MOD 04) data (Dark Target algorithm) at 500 nm.

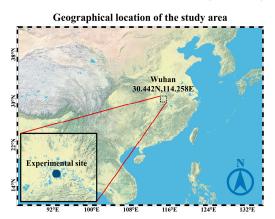


Figure 1. Geographical information and latitude and longitude of the study area.

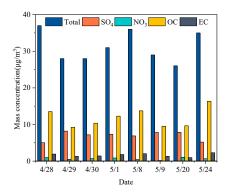


Figure 2. Mass concentrations of PM<sub>2.5</sub> and different aerosol types.

#### 2.3. Aerosol Radiation Forcing

Subsequently, the radiative forcing of different aerosol types in Wuhan is simulated using the vector radiative transfer model developed in the previous study [21], and the simulation results are compared and verified using the SBDART (Santa Barbara DISORT Atmospheric Radiative Transfer model) model. The SBDART model is developed based on the discrete coordinate method and the LOWTRAN7 model to deal with various radiative transfer problems in atmospheric energy balance studies [22]. The radiative fluxes at the top of the atmosphere (TOA) and bottom of the atmosphere (BOA) affected by different aerosol types are calculated by aerosol optical properties and radiative transfer model with short-wave radiation at wavelengths of 0.4–3  $\mu$ m, atmospheric profiles set to mid-latitude summer model, and surface albedo, water vapor concentration and ozone concentration of 0.2, 1.19 g/cm² and 0.406 ppm, respectively. Radiative forcing is defined in terms of changes in the net radiative flux due to variations in the atmospheric medium or other climatic perturbations; according to the radiative flux results from the SBDART model, the aerosol direct radiative forcing is calculated as follows,

$$\Delta F = F \downarrow -F \uparrow, \tag{1}$$

$$\Delta F_{aerosol}^{ATM} = \left(\Delta F_{aerosol}^{TOA} - \Delta F_{clear}^{TOA}\right) - \left(\Delta F_{aerosol}^{BOA} - \Delta F_{clear}^{BOA}\right),\tag{2}$$

where  $\Delta$  represents the difference in radiation flux, the subscripts aerosol and clear represent the case with and without aerosol, respectively, and ATM represents atmospheric

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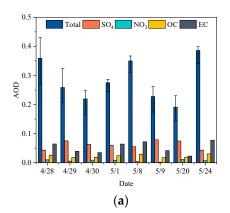
absorption. Further, according to the aerosol radiative forcing, the atmospheric heating rate can be calculated as follows,

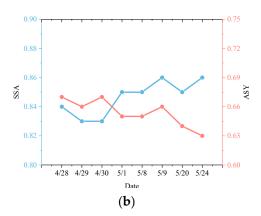
$$dT/dt = -(1/\rho c_p) \cdot (d\Delta F_{aerosol}/dz), \tag{3}$$

where  $c_p$  represents the constant pressure-specific heat capacity with the value of 1004.67 Jkg<sup>-1</sup>K<sup>-1</sup>,  $\rho$  is the air density varying with pressure, z is the altitude, and the atmospheric heating rate dT/dt is in K/day.

#### 3. Results and Discussion

The AOD inversion results of different aerosol types at a wavelength of 500 nm are shown in Figure 3a. The results show that the total aerosol AOD during the test period ranges from 0.19 to 0.39, with an average of about 0.28. The average AOD of  $SO_4$ ,  $NO_3$ , OC, and EC are 0.062, 0.008, 0.023, and 0.052, respectively, which account for about 22%, 3%, 8%, and 19% of the total AOD, with SO4 being the highest and NO3 the lowest. It can also be found that  $SO_4/NO_3$  does not differ much in MC and AOD, ranging from about 5 to 15, respectively, indicating that the optical properties and radiative effects of these two aerosols are similar. However, OC/EC in AOD (0.3–0.8) is much smaller than MC (6–10), which indicates that EC plays a greater role in the radiation effect.





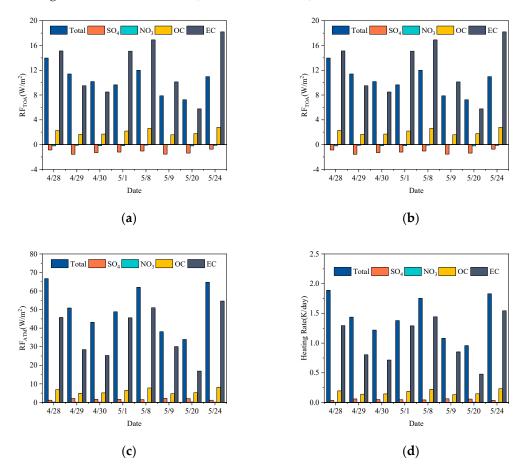
**Figure 3.** (a) Inversion results of AOD for different aerosol types; (b) Variation of aerosol optical properties with date.

The SSA and ASY of aerosols at different dates are shown in Figure 3b. The results show that the aerosol optical properties change with date, which is due to the change of different aerosol types in the mixed aerosol. From April to May, SSA shows an increasing trend, while ASY shows a decrease. Similar changes have been reported in several studies, with enhanced scattering (decreased absorption) and some decrease in the size of aerosols from spring to summer in Wuhan [23]. However, the optical properties of different aerosol types do not vary much, and the SSA (ASY) of SO4, NO3, OC, and EC are about 0.98 (0.65), 0.99 (0.65), 0.83 (0.75), and 0.42 (0.35), respectively. SO4 and NO3 have the strongest scattering properties, while EC has the strongest absorption properties, and its size is also the smallest.

The results of the radiative forcing simulations for different aerosol types, including TOA, BOA, and ATM, are given in Figure 4a–c. The results show that the aerosol radiative forcing at TOA ranges from about  $8\sim14~\rm W/m^2$  with an average of about  $10~\rm W/m^2$ , while at BOA and ATM ranges from about  $-25\sim-55~\rm W/m^2$  and  $30\sim70~\rm W/m^2$  (average of about  $-40~\rm and~50~\rm W/m^2$ ), respectively. The radiative forcing of different aerosol types differs significantly. SO<sub>4</sub> and NO<sub>3</sub> have negative radiative forcing in TOA (about  $-1\sim-2~\rm W/m^2$  and  $0\sim-1~\rm W/m^2$ ), indicating a cooling effect on the Earth system by reflecting solar radiation, while OC and EC have positive radiative forcing in TOA, showing a heating effect. Among all aerosol types, EC has the most significant contribution, not only with the largest positive radiative forcing at TOA (6~18~W/m²) but also up to 61% and 73% at

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BOA and ATM, respectively ( $10\sim35 \text{ W/m}^2$  and  $15\sim55 \text{ W/m}^2$ , respectively), reflecting an extremely strong absorption of radiation. However, the effect of OC, which also plays a heating effect, is much smaller ( $0\sim3 \text{ W/m}^2$  at TOA).



**Figure 4.** Direct radiative forcing of aerosol types at (a) TOA; (b) BOA; (c) ATM; and (d) Atmospheric heating rates of different aerosol types.

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