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Observation and Modeling of the Equilibrium Slope Response of a High-Energy Meso-Macrotidal Sandy Beach

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Abstract: Beach slope is a critical parameter to, e.g., beach safety, wave reflection at the coast and longshore transport rate. However, it is usually considered as a time-invariant and profile-average parameter. Here, we apply a state-of-the-art equilibrium model to hindcast beach slope variability from the time scales of days to years at the high-energy meso-macrotidal sandy beach of Truc Vert, southwest France. We use 9 years of bimonthly beach surveys to compute beach slope time series at different elevations. Results show that beach slope exhibits an equilibrium response with contrasting behaviors along two distinct areas of the beach profile. From 0 to 2 m above mean sea level, which is located under the berm crest, a slope response predominantly at the storm time scale is observed. The beach slope steepens under low energy waves, with the equilibrium model explaining up to 40% of the observed beach slope variability. In contrast, from 2.5 to 4 m above mean sea level, which is above the berm crest, the beach slope steepens under high-energy waves. Within this region of the beach profile, the response time scale increases upwards from seasonal (~2.5 m) to seasonal (~4 m), with the model explaining up to 65% of the observed beach slope variability. Such behaviors are found to be enforced by the berm dynamics developing from the end of the winter to early autumn, providing new perspectives to model and predict beach slope on sandy beaches.

Keywords: beach slope; equilibrium model; berm; seasonal response; storm-scale response

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1. Introduction

Sandy beaches occupy approximately one-third of the global ice-free coastline [1] and provide precious natural, structural and socio-economical resources to coastal communities [2,3]. Sandy beaches are, however, threatened by climate-change-driven sea-level rise [4,5], combined with other factors such as the natural shortage of sediment supply and human activities interfering with natural processes [1]. Monitoring, understanding, modeling and predicting sandy beach evolution is therefore of great importance.

Amongst other beach morphology parameters, beach slope is critical from many perspectives. For instance, beach slope largely impacts wave breaking type and intensity, from spilling to surging, through plunging [6–8]. Beach slope is thus important for, e.g., beach safety and lifeguarding as plunging and dumping waves at the shoreline (shorebreak waves) at steep beaches can cause severe spine injuries to those caught in the impact zone [9–11]. Beach slope controls the reflection at the coast of gravity and infragravity waves [12], and is also a critical parameter in many empirical and semi-empirical formulas used to estimate, e.g., wave runup [13] and longshore drift [14]. In such formulae, beach slope is typically assumed as a single time-invariant parameter. However, beach slope is not homogenous along the beach profile and constantly evolves in response to the storm, seasonal and interannual variations in incident wave conditions. Such beach slope time and space variability is still poorly understood.

The work of Wright and Short [15], based on earlier seminal studies (e.g., [16–18]), classified different beach types into a comprehensive single-bar beach state model. This classification goes from a reflective (steep) beach state to a dissipative (gently sloping) beach state, all the way through a range of intermediate, barred, beach states. Although this classification has been mostly used as a "static" classification for modal wave conditions [19], a fundamental aspect was the development of a framework addressing dynamic beach state behavior [18]. In brief, the transitions between the different beach types were found to be driven by a disequilibrium between the surf-zone hydrodynamics and the beach [20], thus indicating a dynamic equilibrium behavior of beach states and, presumably, beach slope.

Equilibrium-based models have been developed over the last few decades to explain the time evolution of different beach parameters. Such a computationally cheap modeling approach has been primarily used to simulate and explain shoreline change from the time scales of storms to years and decades [21–24] (see review of [25]). Although such models perform substantially better using time-varying free parameters [26,27], good results were obtained both for shoreline hindcasting [22,28,29] and forecasting [30] using time-invariant parameters. Such models have also been applied to different shoreline proxies at meso- to macro-tidal beaches showing a large variability in model skill and free parameter values along the beach profile [22,24]. It is noteworthy that equilibrium models also showed a good ability to reproduce the time variability of cross-shore sandbar behavior [31], embayed beach rotation [32], or grain size [33,34]. However, to our knowledge equilibrium models have never been tested on beach slopes.

In this paper, we investigate the beach slope equilibrium response at the high-energy meso-macrotidal beach of Truc Vert, southwest France. Using 9 years of bimonthly topographic surveys, we apply a state-of-the-art equilibrium model to explore model skill for beach slopes computed across different sections of the beach profile, from mid-tide up to the surbaerial beach. We show that two distinct sections of the profile show a clear out-of-phase equilibrium response, thus opening new perspectives to model and predict beach slope on sandy beaches.

2. Study Site and Data

2.1. Truc Vert Beach

Truc Vert is an open sandy beach located on the Gironde coast, southwest France (Figure 1a). It is a remote beach located kilometers away from the first carpark, which has never been nourished or affected by hard structures and tourism. Such settings motivated the deployment of a number of intensive field experiments [35] as well as the development of a monitoring program which started in 2003 [36].

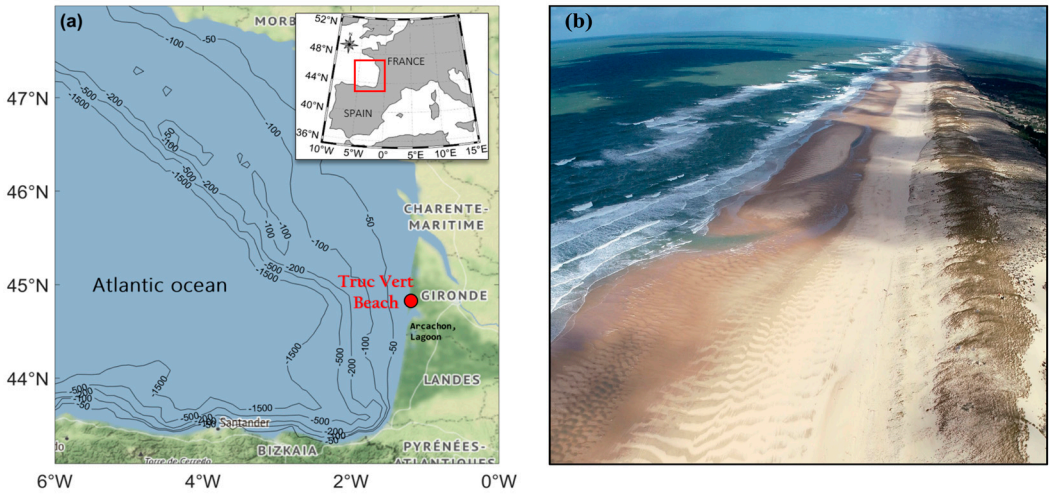

Figure 1. (**a**) Location map of Truc Vert beach in southwest France. (**b**) UAV photograph of Truc Vert (Ph. V. Marieu).

Truc Vert beach is mostly intermediate double-barred and backed by a well-developed aeolian dune (~20–25 m high and ~250 m large) separating the beach from a large area of state-owned forest (Figure 1b). The tidal range is meso-macrotidal, with a ~3.7 m annual mean spring tidal range and a maximum astronomical tide range reaching up to 5 m [37]. Incident waves are generated by extra-tropical storms tracking eastwards across the North Atlantic Ocean. The wave climate is energetic and strongly seasonally modulated with a monthly averaged significant wave height *Hs* ranging from 1.1 m in July with a dominant W–NW direction to 2.4 m in January with a dominant W direction [38]. Incident winter wave energy shows large interannual variability enforced by regional, large-scale, climate modes of atmospheric variability [37]. Truc Vert can be exposed to severe windstorms, with 10 m hourly wind speed exceeding 30 m/s, causing large coastal dune morphological changes [39]. The sediment consists of well-mixed fine to medium sand with a mean grain size of approximately 0.35–0.40 mm [40]. Over the last few decades, the shoreline position at Truc Vert has been relatively stable [37].

2.2. Beach Slope Data

This study is based on the Truc Vert beach GNSS topographic monitoring program that was implemented in 2003 [36]. The beach survey's alongshore coverage prior to 2012 was shorter than 1000 m, thus not offering optimal coverage given the strongly alongshore non-uniform beach morphology enforced by ~400 m spaced rip channel systems. Therefore, here, we used the data collected between 2012 and 2021 (1500–2000 m alongshore coverage), with a notable 2-month gap in 2020 due to the COVID-19 pandemic. Beach surveys typically extend from approximately a −1.5 m (low tide mark) to 6 m (dune foot) elevation above mean sea level (AMSL). Each topographic survey is interpolated on a regular grid in local cross-shore/longshore coordinates. The reader is referred to [36] for details on the topographic surveys and interpolation techniques.

The monitoring program from October 2010 to June 2021 resulted in 181 digital elevation models (DEM) computed in local cross-shore/longshore coordinates. Figure 2 shows an example of a topographic survey (Figure 2b) and its corresponding superimposed 20 m spaced interpolated cross-shore profiles (Figure 2a), highlighting a large alongshore variability. Beach slopes were computed on 1 m elevation sections every 0.1 m (i.e., [−1.0;0.0], [−0.9;0.1], . . . , [5.0;6.0]) of each 20 m alongshore spaced interpolated beach profile. This segmentation was based on preliminary tests showing it best emphasizes the vertical variability of beach slope response. For each date and each 1 m elevation section, beach slopes computed at each 20 m spaced cross-shore transect were then averaged alongshore in order to investigate the equilibrium response of beach slope for different parts of the beach profile.

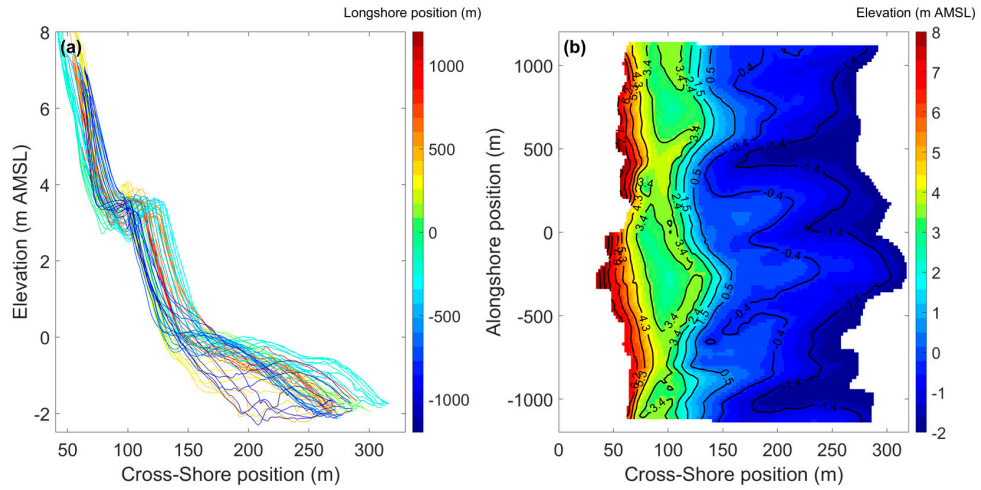

Figure 2. Truc Vert beach DEM on 11 September 2018: (**a**) superimposed 20 m spaced beach profiles and (**b**) planview DEM.

2.3. Wave Data

This study used approximately 13.5 years of continuous wave outputs gathered from a calibrated wave hindcast from January 2008 to June 2021 [28,29]. We used the time series of significant wave height (*Hs*), peak period (*Tp*) and mean direction (*Dm*) at an approximately 54 m depth offshore of Truc Vert beach [28,29]. These data were then used to force the equilibrium model with a 3-hour time step.

3. Equilibrium Profile Model

3.1. Equilibrium Model

The local beach slope time series was hindcast using the equilibrium model approach of [41], which so far has been mostly used to simulate shoreline change on cross-shore-transport-dominated beaches. The time change of local beach slope, S, is given as a function of the time-varying, wave power P and energy disequilibrium $\Delta\Omega$:

$$\frac{dS}{dt} = C^{\pm} P^{0.5} \Delta\Omega \,, \tag{1}$$

where $C^{\pm}$ is the change rate coefficient for erosion velocity (C^{-}) and accretion velocity (C^{+}), which represent efficiency rates determining the rate of beach slope change in response to the wave force. $\Delta\Omega = \Omega - \Omega_{eq}$ is the energy disequilibrium, with Ω being the Dean number given by:

$$\Omega = \frac{H_s}{w_s T_p} \,, \tag{2}$$

where w_s is the sediment fall velocity calculated using Soulsby's (1997) formula [42]. The equilibrium energy, Ω_{eq}, is computed through:

$$\Omega_{eq}(t) = \frac{\sum_{j=0}^{2\Phi/\Delta t} \Omega_j 10^{-j\Delta t/\Phi}}{\sum_{j=0}^{2\Phi/\Delta t} 10^{-j\Delta t/\Phi}} \,, \tag{3}$$

with Δt being the sampling interval and Φ being the memory decay of the system, in days. This implies that Ω_{eq} incorporates all past beach state information for the past 2Φ days and constantly evolves in time and maintains a weighted "memory" of antecedent incident wave conditions.

The wave power in Equation (1) is given by:

$$P = EC_g \,, \tag{4}$$

where E is the wave energy given by:

$$E = \frac{1}{16}\rho g H_s^2 \,, \tag{5}$$

with ρ being the water density, g the gravitational acceleration and C_g the wave group velocity computed using the wave linear theory:

$$C_g = \frac{\Omega}{k}\left[\frac{1}{2} + \frac{kh^2}{\sin h(2kh)}\right] \,, \tag{6}$$

where Ω is the angular frequency of waves, k is their wavenumber and h the water depth.

3.2. Model Parameter Optimization

The model has four free parameters, S_0 the initial slope, Φ, C^{-} and C^{+}, that were found by minimizing the root mean square error between the computed and measured beach slopes. The error minimum was found using a simulated annealing non-linear optimization algorithm [43]. Such a probabilistic approach can overcome local RMSE minima in this

4-parameter space and find a global optimum. Simulating annealing optimization already gave good results with beach equilibrium models [22,24]. Table 1 shows the search intervals used to initialize the optimization algorithm, for each free parameter. The model was initialized with broad intervals (including positive and negative erosion and accretion rates) in order to be able to find the optimal values corresponding to all the sections of the beach profile, which readily show a large range of behaviors. The simulated annealing was performed over 3,000,000 iterations.

Table 1. Initialization of the 4 free model parameters.

S_0 (m)	Φ (days)	C^- ($s^{-1}.W^{-0.5}$)	C^+ ($s^{-1}.W^{-0.5}$)
[−0.5;0.5]	[1;850]	[−0.000008;0.0000081]	[−0.000025;0.000024]

4. Results

4.1. Beach Profile Change in Time and Space

Figure 3 illustrates the seasonal beach profile shape change at a given transect over a given year. Starting from an eroded, featureless, concave shape profile in the middle of the winter (January 2018), results show the progressive building of berm throughout spring, summer and early autumn before the berm is smoothed out during the subsequent winter. The berm formed at a cross-shore distance around 100 m, and rose by over two meters before culminating at 4+ m AMSL in early autumn, making this section of the profile the most dynamic. The higher part of the profile (>5 m AMSL) is steep and barely evolves. In contrast, the lower part of the profile is systematically gently sloping owing to the overall concave shape of the beach profile. There is also the occasional presence of the inner bar (June 2018, Figure 3).

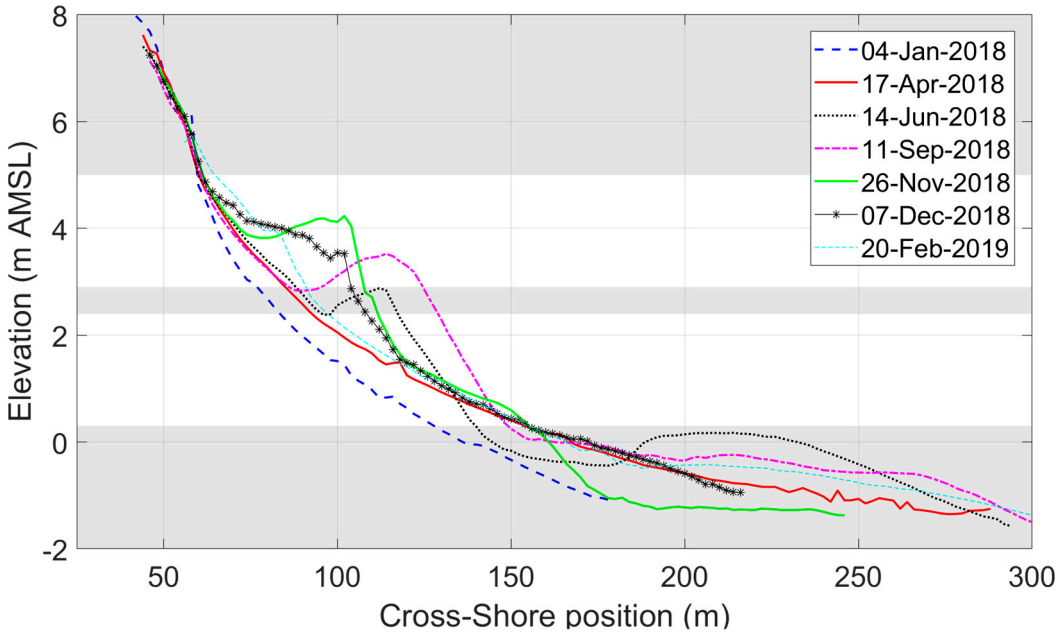

Figure 3. Superimposed beach profiles at a given transect covering over an entire year illustrating typical seasonal morphological changes at Truc Vert beach.

Figure 4 provides more insight into the beach profile variability from the time scales of weeks, to interannual, through seasonal. A clear seasonal signal can be depicted, particularly in the berm area described in Figure 3. Figure 4 also reveals a strong interannual variability, which is marked by the winter of 2013/14 which caused outstanding erosion. This is particularly striking for the upper part of the beach profile as the dune toe was

severely scarped during this winter and has been mostly, slowly, recovering since then. Non-surprisingly, the time series of the alongshore-averaged profiles (Figure 4b) shows a smoother variability than that of the single profile (Figure 4a). These morphological changes are essentially driven by temporal changes in incident wave conditions, which are used to drive the equilibrium model below.

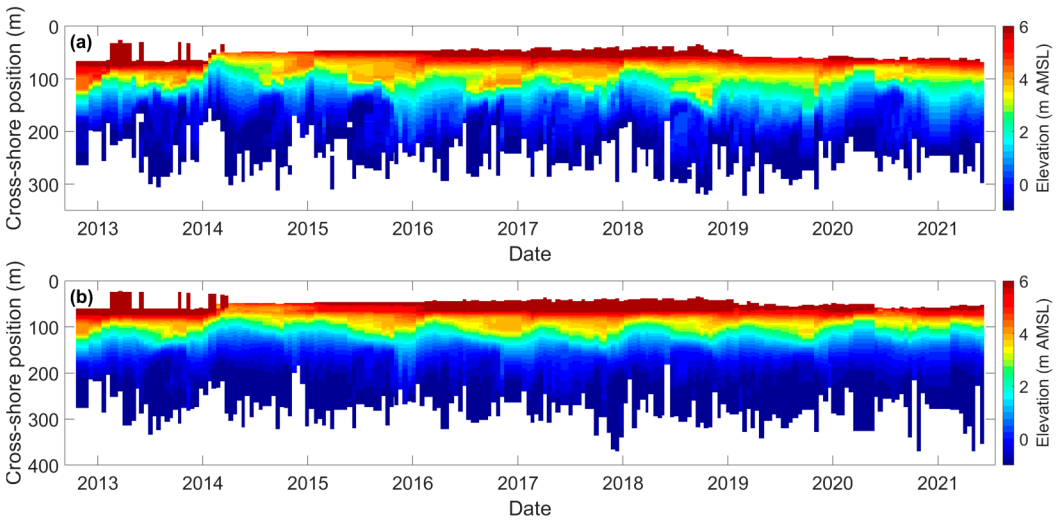

Figure 4. Cross-shore profile evolution (elevation colored) from 17 October 2012 to 10 June 2021. (**a**) Single profile located in the center of the domain and (**b**) alongshore-averaged profile.

4.2. Equilibrium Beach Slope Response

Figure 5 shows the equilibrium model skill to simulate beach slope response (Figure 5a,b) and the corresponding model free parameters (Figure 5c–e) as the function of the beach elevation where the slope is computed. There is large variability in the ability of the model to hindcast beach slope response depending on elevation. However, two distinct areas where R^2 is > 0.2 can be depicted (white areas in Figure 5), i.e., where the model explained more than 20% of the observed beach slope variability. For the area in the lower part of the profile, R^2 mostly ranges between 0.3 and 0.4, with Φ < 150 days and negative erosion and accretion rate coefficients. This means that the beach slope in this region of the profile shows fair equilibrium response characteristics, with slope changes at the scale of individual storms and with the slope steepening (flattening) during low- (high-)energy waves. This contrasts with the other area, located in the higher part of the profile (Figure 5). Overall, the equilibrium model shows better skill than in the lower area, with R^2 even locally exceeding 0.65. Beach slope also shows a slightly larger response rate, with positive values, meaning that the slope steepens (flattens) during high- (low-)energy waves, i.e., out-of-phase with the lower part. Interestingly enough, within this higher area, beach memory Φ rapidly increases upwards, eventually reaching the upper value of Φ = 850 days. This means that beach slope variability tends to predominate at seasonal scale when going upwards, rather than at the storm scale.

This is further illustrated in Figure 6 which shows the time series of observed and modeled beach slopes at five different vertical locations along the profile. In line with the greyish areas in Figure 5 where R^2 < 0.2, the model fails to simulate beach slope variability at these locations (Figure 6b–d). In contrast, the model shows fair skill at the lower part of the profile (Figure 6e), and behaves even better in the higher part of the profile (Figure 6c). In line with accretion/erosion rate coefficients of the opposite sign in Figure 5, the areas clearly show out-of-phase beach slope signals, with beach slope in the higher part being approximately twice as large as in the lower part.

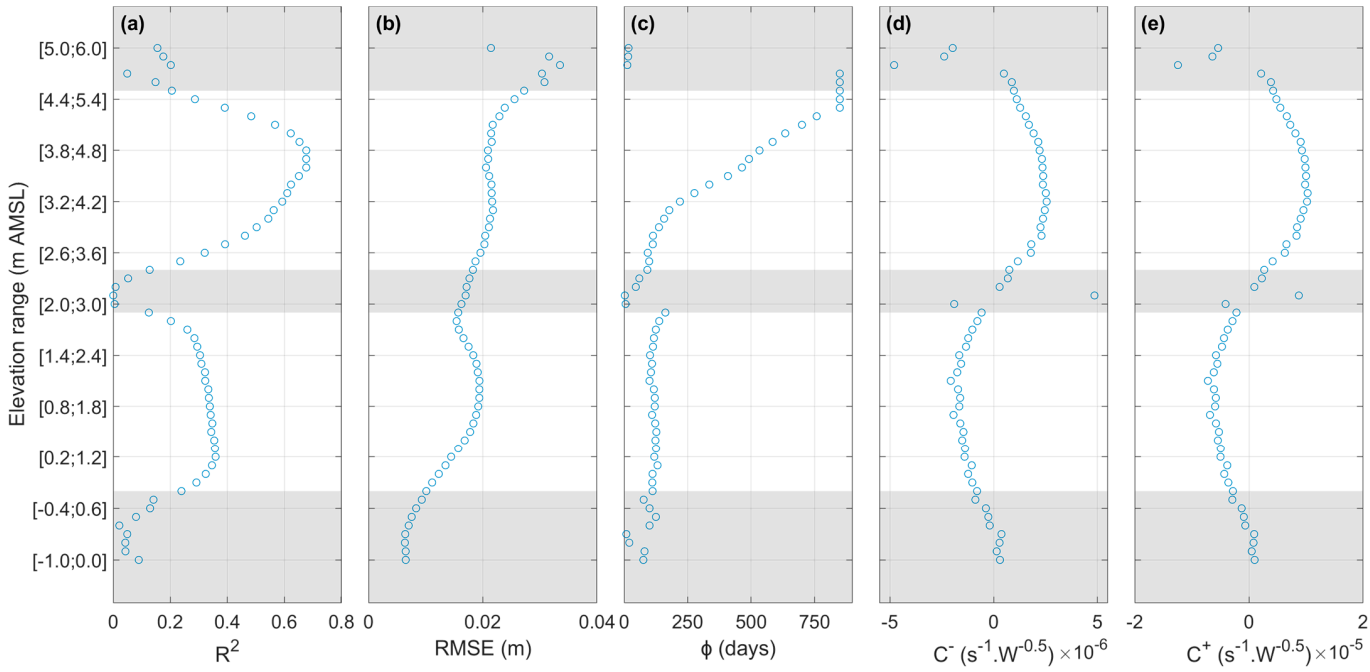

Figure 5. (**a**) R^2 and (**b**) RMSE between the observed and modeled beach slope and corresponding (**c**) Φ memory, (**d**) C^- erosion rate coefficient and (**e**) C^+ accretion rate coefficient free parameters, found as a function of the elevation where the beach slope is computed. The shaded areas are where $R^2 < 0.2$ which we considered as elevations where beach slope does not show an equilibrium response.

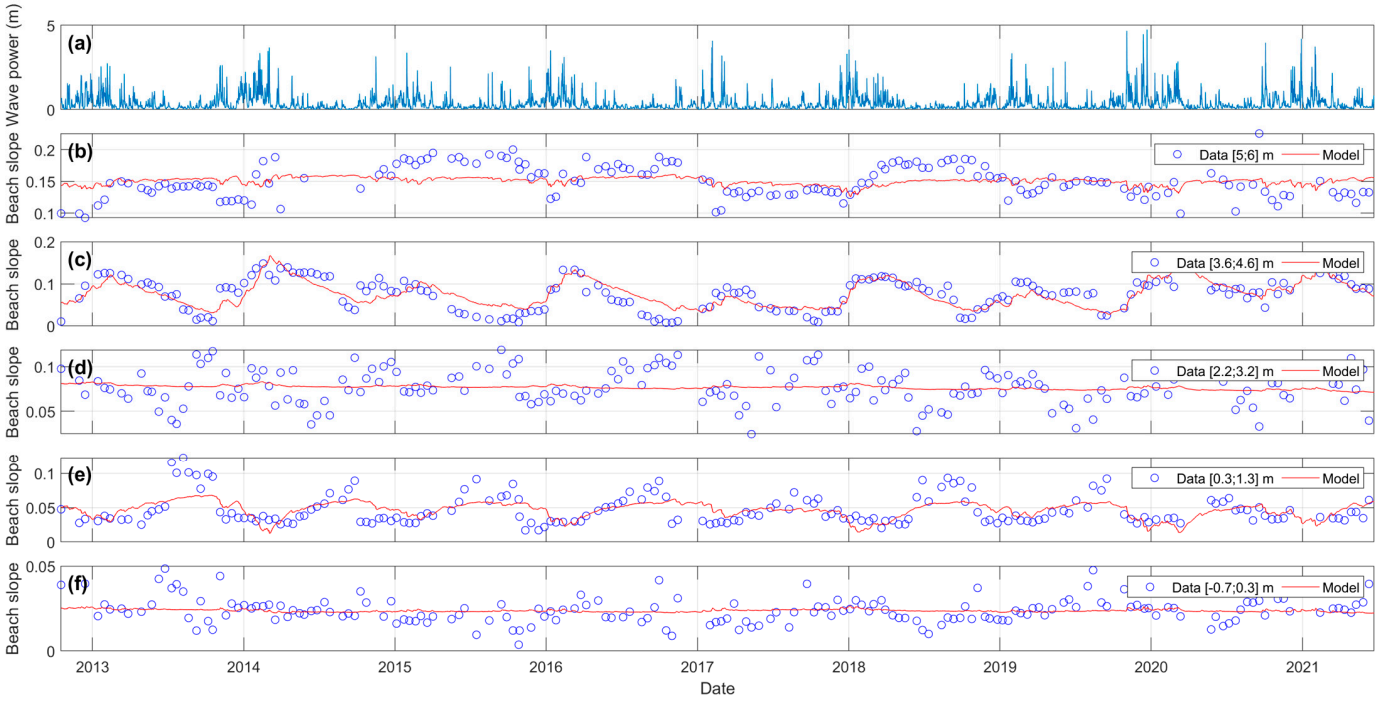

Figure 6. Time series of (**a**) wave power and (**b**–**f**) observed (circles) and modeled (solid line) beach slope between (**b**) 5–6 m; (**c**) 3.6–4.6 m; (**d**) 2.2–3.2 m (**e**) 0.3–1.3 m and (**f**) −0.7–0.3 m.

5. Discussion

Consistent with other sandy beach morphological indicators, such as shoreline, cross-shore sandbar position or grain size [31–33], beach slope also shows a clear equilibrium response to changes in incident wave conditions. On the meso-macrotidal beach of Truc Vert,

southwest France, such a response strongly depends on the beach elevation around which slope is computed, consistent with earlier work using different shoreline proxies [22,44]. The best equilibrium response is found in the higher part of the profile, around 4 m AMSL, that is, well above the highest astronomical tide level (~2.5 m AMSL). This may be seen as a surprising result, which will be discussed later, as this region of the beach profile is only rarely activated by surf zone processes and could be assumed to be decoupled from wave force. At this elevation, the equilibrium model explains over 65% of the observed beach slope variability. This is similar to the shoreline variance explained by the equilibrium model for the optimal shoreline proxy at Truc Vert of around 1.5 m AMSL. Our results therefore demonstrate that equilibrium models can be successful at simulating beach slope variability from the time scales of days to years, with similar skill as for shoreline hindcasting.

In contrast with the shoreline equilibrium response at Truc Vert beach, with a single area of consistent equilibrium shoreline response ranging approximately from 0 to 4 m AMSL, beach slope response shows two distinct opposing areas ranging approximately from 0 to 2.5 m AMSL and 3 to 4.5 m AMSL (Figure 5). The lower part of the profile responds at the storm scale, with the slope decreasing with larger waves. This is in line with existing beach state models [45,46] with more dissipative, gently sloping, beach profiles under storm conditions. In contrast, the higher part of the profile steepens with larger waves and increasingly responds by moving upwards at the seasonal time scale (increased memory Φ). Such opposing behavior can be explained by the berm dynamics. The area below 2 m AMSL is systematically under the crest of the berm and reflects the beach slope response to incident wave conditions during the daily high stage of the tide. The area ranging from 3 to 4.5 m AMSL, located higher along the profile, is systematically steeper during winter (Figure 3). This is because during winter storm wave events, surf zone and swash zone processes smooth out the berm, resulting in a featureless concave winter profile (Figure 3). However, during berm development from spring to early autumn, thus with a seasonal signature, this sector is located in the lee of the berm crest built up through swash processes, resulting in a progressively decreasing beach slope as seasonal mean wave conditions become less energetic. This explains why the best equilibrium response is found in the higher part of the profile, well above the highest astronomical tide level of around 4 m AMSL. Between these two areas forming a transition between two contrasting behaviors, beach profile does not show an equilibrium response. Similar observations can be made at the higher (>4.5 m AMSL) and lower (<0 m) parts of the profile, where a poor correlation is systemically found, for different reasons. Above 4.5 m AMSL, the beach slope is steep (>0.1) and evolves without any apparent link with incident wave conditions (Figure 6b). In contrast, below 0 m AMSL, the beach slope is low (<0.05) and also shows variations which are not always readily related to changes in incident wave energy. This is consistent with shoreline response at this elevation [22], because this part of the profile is strongly influenced by the inner-bar dynamics. Given that the cross-shore inner-bar dynamics at Truc Vert beach are influenced largely by tide-range variations [47], this is not captured by the equilibrium model. Such findings suggest that an alternative approach could be to develop an equilibrium berm response model rather than a multiple, elevation-dependent, equilibrium beach slope model, such as that developed here.

In our contribution, we only applied an equilibrium model based on that developed for the shoreline by [41]. However, different equilibrium response frameworks exist [21,22,29,48,49]. The model of [21] was also tested here, showing similar skill. However, the interpretation of the model free parameters is more complex [50], and outputs were thus not shown herein. We used time-invariant free parameters, but we hypothesize that investigating time-varying parameters, such as in [27], may further improve model skill. Another avenue for model improvement is to account for tide range, as berm development was found to depend on spring-neap tide cycles at the microtidal beach of Narrabeen, Australia, using high-frequency Lidar data [51]. We assumed a sediment grain size constant in both time and space. However, large variations (0.2–0.7 mm) in grain size over tens

of meters linked with morphological variability have been observed at Truc Vert [40]. In addition, grain size has been found elsewhere to largely vary in time in response to changes in wave conditions [34]. Including changes in grain size could improve beach slope model skill, but there is currently no framework for this. Finally, given the wealth of long-term beach monitoring programs [52–55], applying such a model to different sites to search for a parametrized version of the free parameters, similar to what [23] did for shoreline, would help to better understand the influence of the general settings (e.g., sediment size, wave climate, tide regime) on beach slope response.

6. Conclusions

Beach slope is usually considered as a time-invariant and profile-average parameter. By applying a simple equilibrium model to a 9-year time series of bimonthly beach surveys of a high-energy meso-macrotidal beach, we found that beach slope can locally exhibit a strong equilibrium response. A state-of-the-art equilibrium model and a global optimization simulating annealing algorithm were combined to reveal two distinct areas with contrasting behaviors. First, at 0–2 m AMSL, below the berm crest, equilibrium slope response was found at the storm timescales, with slope steepening under low energy waves. At 2.5–4 m AMSL, above the berm crest, an opposite behavior was observed, with the equilibrium response time scale moving increasingly, seasonally, upwards. This study provides new perspectives to better understand, model and predict beach slope and sandy beaches. It also motivates the further development of beach monitoring programs.

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Data Availability Statement: The Truc Vert beach monitoring data and wave forcing time series used herein are publicly available at [29,36], respectively.

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Conflicts of Interest: The authors declare no conflict of interest.

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