

Article

Improving the Estimation of the Occurrence Time of an Impending Major Earthquake Using the Entropy Change of Seismicity in Natural Time Analysis

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Abstract: This article is focused on a new procedure concerning a more accurate identification of the occurrence time of an impending major earthquake (EQ). Specifically, we first recapitulate that, as was recently shown [P. Varotsos et al., *Communications in Nonlinear Science and Numerical Simulation* **125** (2023) 107370], natural time analysis of seismicity supplemented with the non-additive Tsallis entropy S_q leads to a shortening of the time window of an impending major EQ. This has been shown for the Tohoku mega-EQ of magnitude M_9 that occurred in Japan on 11 March 2011, which is the largest event ever recorded in Japan. Here, we also show that such a shortening of the time window of an impending mainshock can be achieved for major, but smaller EQs, of the order of M_8 and M_7 . In particular, the following EQs are treated: the Chiapas $M_{8.2}$ EQ, which is Mexico's largest EQ for more than a century that took place on 7 September 2017 near the coast of Chiapas state in Mexico, the 19 September 2017 $M_{7.1}$ EQ that occurred within the Mexican flat slab, and the $M_{7.1}$ Ridgecrest EQ on 6 July 2019 in California.

Keywords: natural time; earthquakes; order parameter; entropy; criticality; seismic electric signals

To the memory of the Academician Seiya Uyeda who first commented on our Chiapas earthquake results during his last visit in Athens, Greece, in November 2017.

1. Introduction

A new view of time, termed natural time χ has been introduced by the first three authors in 2001, see, e.g., [1]. It has been found to be useful in various disciplines, for example it has been employed by Rundle, Turcotte, and coworkers [2–8] as a basis for a new methodology to estimate the current level of seismic risk, termed Nowcasting Earthquakes (which was recently reviewed in [9–11]). In this new view, novel dynamical features hidden behind time series in complex systems can emerge but cannot when the analysis is carried out within the frame of conventional time. Natural time analysis (NTA) enables the study of the dynamical evolution of a complex system and identifies when the system

enters a critical stage [12–14]. Earthquakes (EQs) exhibit complex correlations in time, space, and magnitude (e.g., [15–21]), and it has been widely accepted [12,22,23] that the occurrence of large earthquakes is likely associated with the proximity of crustal volumes to a critical breaking point. We clarify, however, that, as discussed by several authors, seismic activity, in general, does not require critical behavior (e.g., [24–26]) and the scaling laws of earthquakes can be derived without any reference to criticality (e.g., [27]).

The variance of natural time, i.e., $\kappa_1 = \langle \chi^2 \rangle - \langle \chi \rangle^2$, is a useful quantity in identifying the approach to a critical point and in NTA of the seismicity, κ_1 may be considered as an order parameter of seismicity [28]; see also Chapter 6 of [12,29].

NTA also revealed [30,31] that, upon studying the Olami–Feder–Christensen EQ model [32]—probably [33] the most studied non-conservative (supposedly) self-organized criticality model, see also Caruso and Kantz [34]—the entropy change of seismicity under time reversal ΔS exhibits [12,30] an evident minimum ΔS_{min} before large avalanches, which correspond to strong EQs. In general, it has been shown that the quantity ΔS_{min} plays a major role upon approaching a dynamic phase transition [35] (critical point) in which the mainshock constitutes the new phase [12].

NTA revealed the identification of the epicentral area of an impending major earthquake as well as demonstrated [36,37] that the initiation of a Seismic Electric Signals (SES) activity (i.e., series of precursory variations of the electric and of the accompanying magnetic field of the Earth [38–41]) has a lead time [12] from a few weeks to $5\frac{1}{2}$ months before a strong EQ, which almost coincides with the minimum β_{min} of the fluctuations β of the order parameter κ_1 of seismicity.

In NTA of seismicity, the occurrence of the criticality condition $\kappa_1 = 0.070$ [12,13,42] marks that a strong EQ is going to occur within a time window from a few hours to 1 week (up to 11 days or so, [43]). It is, of course, of major importance to shorten this time window before an impending mainshock. In particular, focusing on this period, we examine whether there exist additional variations of the entropy change ΔS under time reversal as well as possible changes of other entropic measures. Along these lines, very recently Varotsos et al. [14] found that before the M_9 Tohoku EQ in Japan that occurred on 11 March 2011, which is the largest EQ ever recorded in Japan, shortly after the validity of the criticality condition $\kappa_1 = 0.070$, the quantity ΔS of NTA together with the Tsallis entropy exhibited distinct simultaneous changes. Here, we extend such an investigation for three smaller EQs, i.e., the Chiapas $M_{8.2}$ EQ in Mexico—which is Mexico’s largest EQ for more than a century—the case of the 19 September 2017 $M_{7.1}$ EQ that occurred within the Mexican flat slab, and the $M_{7.1}$ Ridgecrest EQ on 6 July 2019 in California, see Figure 1.

This work is structured as follows: In Section 2, we summarize the background of NTA and then explain in Section 3 the general procedure that we apply [12] to estimate the occurrence time of an impending mainshock. The main results of our investigation on what happens *after* the validity of the criticality condition $\kappa_1 = 0.070$ are presented in Section 4. This section is divided into four subsections: Section 4.1 contains the results for the Chiapas $M_{8.2}$ EQ on 7 September 2017 and Section 4.2 the results concerning the $M_{7.1}$ EQ that hit central Mexico on 19 September 2017 killing more than 300 people in Morelo’s region as well as in Mexico City. In Section 4.3, the $M_{7.1}$ Ridgecrest EQ on 6 July 2019 in California is studied. Finally in Section 4.4 the main results for the M_9 Tohoku EQ in Japan on 11 March 2011 obtained in [14], as mentioned above, are recapitulated. A discussion follows in Section 5 and, finally, our main conclusions are summarized in Section 6.

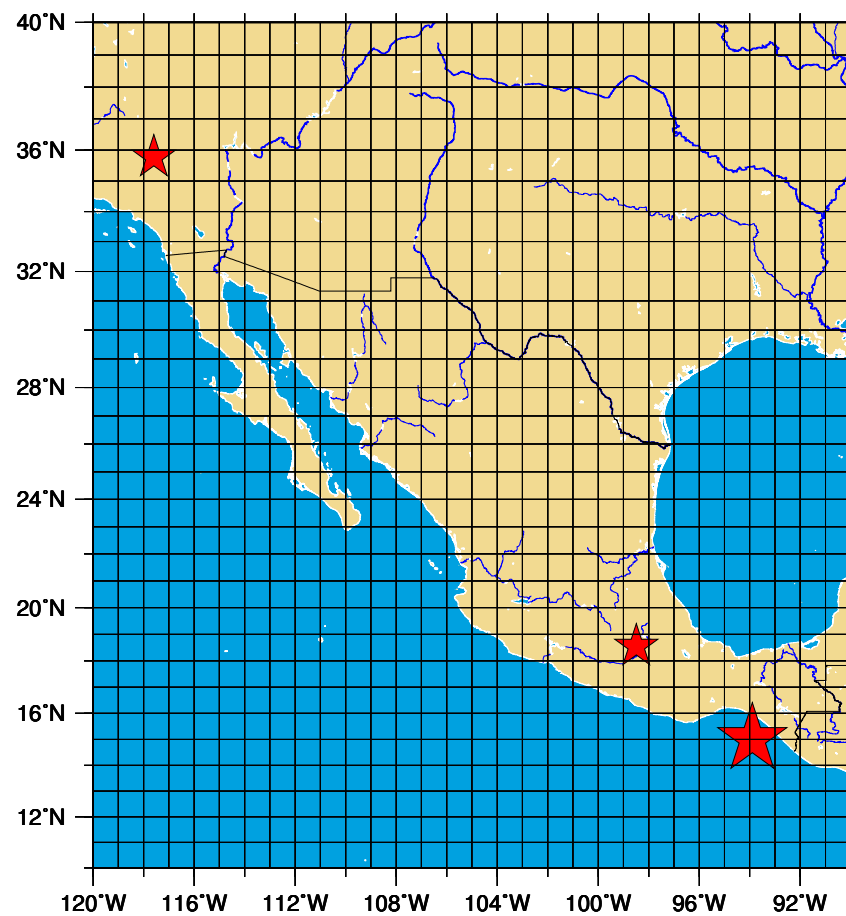


Figure 1. Map depicting the epicenters, from top to bottom, of the $M7.1$ Ridgecrest EQ on 6 July 2019, the 19 September 2017 $M7.1$ EQ that occurred within the Mexican flat slab, and the Chiapas $M8.2$ EQ in Mexico.

2. Natural Time Analysis: Background

Natural time analysis is based on a new concept of time that was put forward by Varotsos et al. [1], as already mentioned—see also Chapter 2 of [12] and the Appendix of [13]—and has been found useful to uncover important features hidden in complex systems time series from seismology to cardiology [44–49] and from atmospheric sciences [50,51] to complex networks [52]. Furthermore, various authors (e.g., Pasari and coworkers [53–56]) have applied the methodology of Nowcasting Earthquakes based on natural time to the contemporary earthquake hazards in several areas like West–Northwest Himalaya [53], Java Island [54], Sulawesi Island [55], and Sumatra in Indonesia [56].

For a series of N events, which actually is a temporal point pattern, see, e.g., [57], natural time serves as an index for the occurrence of the k -th event and is given by $\chi_k = k/N$. In NTA, the conventional interevent time is ignored, but we preserve the occurrence order and the energy Q_k of the events. For seismicity $Q_k \propto 10^{1.5M_k}$, where the moment magnitude M [58] is considered, see, e.g., [59]. In NTA, the evolution of the pair (χ_k, p_k) , where $p_k = Q_k / \sum_{k=1}^N Q_k$ is the normalized energy for the k -th event, is studied. The normalized power spectrum is defined as $\Pi(\omega) \equiv |\Phi(\omega)|^2$ where $\Phi(\omega) = \sum_{k=1}^N p_k \exp(i\omega\chi_k)$ and ω denotes the angular natural frequency. Since all the moments of the distribution of the p_k , can be deduced from $\Pi(\omega)$ in the limit $\omega \rightarrow 0$ (see p. 130 of [12]), $\Pi(\omega)$ is considered for ω close to zero. The Taylor expansion of $\Pi(\omega)$ is given by

$$\Pi(\omega) = 1 - \kappa_1\omega^2 + \kappa_2\omega^4 + \dots \quad (1)$$

where

$$\kappa_1 = \sum_{k=1}^N p_k (\chi_k)^2 - \left(\sum_{k=1}^N p_k \chi_k \right)^2. \quad (2)$$

The quantity κ_1 corresponds to the variance $\langle \chi^2 \rangle - \langle \chi \rangle^2$ of natural time. It is a key quantity when studying seismic catalogs [52,59–61], because Varotsos et al. [28] proved that it is the order parameter of seismicity through which one can identify when the system reaches the critical point. Note that EQs, as already mentioned in the Introduction, are widely considered [12,22,23] as critical phenomena (the mainshock being the new phase) exhibiting complex correlations in time, space and magnitude (e.g., [15,17–19,44,62,63]).

In NTA, the entropy, S , is given by:

$$S = \langle \chi \ln \chi \rangle - \langle \chi \rangle \ln \langle \chi \rangle \quad (3)$$

where $\langle f(x) \rangle = \sum_{k=1}^N p_k f(x_k)$. It shows [64] concavity, positivity, and Lesche stability [65,66]. S is a dynamic entropy and its value S_u for a uniform (u) distribution [12] is $S_u = 0.096$. The application of the time reversal operator $\mathcal{T} p_k = p_{N-k+1}$ to S , leads to the entropy under time reversal, S_- :

$$S_- = \sum_{k=1}^N p_{N-k+1} \left(\frac{k}{N} \right) \ln \left(\frac{k}{N} \right) - \left[\sum_{k=1}^N p_{N-k+1} \left(\frac{k}{N} \right) \right] \ln \left[\sum_{k=1}^N p_{N-k+1} \left(\frac{k}{N} \right) \right] \quad (4)$$

The difference, $\Delta S = S - S_-$, represents an important quantity with a physical meaning studied in detail in [12,30,35,67]. As already mentioned, ΔS is a key parameter, which may determine the approach to a dynamic phase transition [12,35,68].

The fluctuations of ΔS are studied as follows: Using a moving window of length i (number of successive events) sliding through the EQ catalog, the time series of S_i and $(S_-)_i$ are obtained. By means of the standard deviation $\sigma(\Delta S_i)$ of the time series of $\Delta S_i \equiv S_i - (S_-)_i$, the complexity measure Λ_i [12,45,69] is defined

$$\Lambda_i = \frac{\sigma(\Delta S_i)}{\sigma(\Delta S_{100})}, \quad (5)$$

where the denominator corresponds [69] to the standard deviation $\sigma(\Delta S_{100})$ of the time series of ΔS_i with $i = 100$ events. The quantity Λ_i quantifies how the statistics of ΔS_i time series is changed upon increasing the scale from 100 EQs to a longer scale, e.g., $i = 500$ EQs.

3. Identification of the Occurrence Time of the Impending Mainshock

The relevant procedure has been reviewed by Varotsos et al. [70]. For SES activities has been shown [1,12] that:

$$\Pi(\omega) = \frac{18}{5\omega^2} - \frac{6\cos\omega}{5\omega^2} - \frac{12\sin\omega}{5\omega^3}, \quad (6)$$

which simplifies to:

$$\Pi(\omega) \approx 1 - 0.07\omega^2 \quad (7)$$

when $\omega \rightarrow 0$. Equation (7), when considering Equation (1), results in

$$\kappa_1 = 0.070. \quad (8)$$

The above critical condition has been proven for several EQ models (see, e.g., [12,42]), as well as it has been found of usefulness in a variety of other EQ precursory phenomena [71–79].

The NTA of seismicity in the future epicentral area begins upon the starting of SES activity [28,43,80] because the latter is emitted when the focal area enters the critical stage [12,13,81]. Hence, we form EQ time series in natural time for the future epicentral

area, each time when a small EQ of magnitude M exceeds a certain threshold M_{thres} . The value of $\Pi(\omega)$ for $\omega \rightarrow 0$ and the one of κ_1 for each of the EQ time series is estimated and compared to that of Equation (6) for $\omega \in [0, \pi]$. Furthermore, the two entropies S and S_- are also calculated. The established criteria to assure a true coincidence of the EQ time series with that of critical state are the following [12,28,82,83]: (i) The “average” distance $\langle D \rangle$ between the curves of $\Pi(\omega)$ of the evolving seismicity and Equation (6) should be $\langle D \rangle < 10^{-2}$. (ii) The final approach of the evolving $\Pi(\omega)$ to that of Equation (6) must be from below as depicted by the arrow in Figure 7.1 of [12] (as well as in Figure 5 of [84]). This means that, before strong EQs, κ_1 finally approaches from above that of the critical state, i.e., $\kappa_1 = 0.070$. (iii) The entropies S and S_- should be less than S_u at the coincidence. (iv) The time of the true coincidence should not change significantly upon varying M_{thres} , since the process (critical dynamics) has to be self-similar.

4. Results

In Section 4.1, we mainly focus on the improvement of the identification of the occurrence time of the $M8.2$ Chiapas EQ on 7 September 2017 in Mexico, which has been first attempted by Sarlis et al. [60], and in Section 4.2, we present the results for the $M7.1$ EQ on 19 September 2017 in the Mexican flat slab. In Section 4.3, we treat the case of the $M7.1$ Ridgecrest EQ on 6 July 2019 in California and finally, in Section 4.4 we summarize the results for the $M9$ Tohoku EQ on 11 March 2011 that, as mentioned, have been published by Varotsos et al. [14].

4.1. Results on the $M8.2$ Chiapas EQ

The following result was deduced by the NTA of the Mexican seismicity from 1 January 1988 until the $M8.2$ Chiapas EQ on 7 September 2017 by using sliding natural time window lengths consisting of a number of events that would occur in a few months, which is approximately equal with the average value of the lead time of SES activities. Almost 1.5 months before the $M8.2$ Chiapas EQ the minimum $\beta_{\min}$ was found approximately on 27 July 2017 as the deepest minimum during the almost thirty year period investigated [60] (cf. while $\Delta S_{\min}$ has been earlier observed, i.e., on 14 June 2017, as previously mentioned).

Second, we make use of the experimental fact [37] that an SES activity initiates almost simultaneously with $\beta_{\min}$. Hence, an SES activity should have started approximately around 27 July 2017.

Third, we now plot in Figure 2 the quantity ΔS (a) and the Tsallis non-additive entropy S_q [85–91] (b) versus the conventional time in the Chiapas region (see Figure 2 of [31]) during the period almost one week before the $M8.2$ EQ. The seismic data analyzed come from the seismic catalog of the National Seismic Service (SSN) of the Universidad Nacional Autónoma de México (www.ssn.unam.mx, accessed on 23 November 2017) and to assure catalog completeness a magnitude threshold $M_{\text{thres}} = 3.5$ has been imposed as in [31]. Starting from the above date of $\beta_{\min}$ on 27 July 2017, and considering the seismicity in Chiapas area (which is the candidate area), we find that the criticality condition $\kappa_1 = 0.070$ had been fulfilled (together with the fact that the aforementioned criteria for recognizing a true coincidence were satisfied) from 21:08 CDT (=UTC-5) on 5 September 2017 until 16:06 CDT on 6 September 2017. This latter period is shaded in Figure 2. An additional inspection of this figure reveals the following: Approximately on 02:05 CDT on 7 September 2017 a simultaneous change appears both on ΔS of NTA, see Figure 2a, and on S_q , see Figure 2b, i.e., almost 10 h after the aforementioned validity of the criticality condition $\kappa_1 = 0.070$. It is challenging that, after the system entered the critical stage, both these quantities, i.e., ΔS and S_q , exhibit an almost transient change of similar shape upon the occurrence of a $M4.8$ EQ on 7 September 2017. The $M8.2$ Chiapas EQ occurred at 23:49 CDT on 7 September 2017, i.e., several hours later.

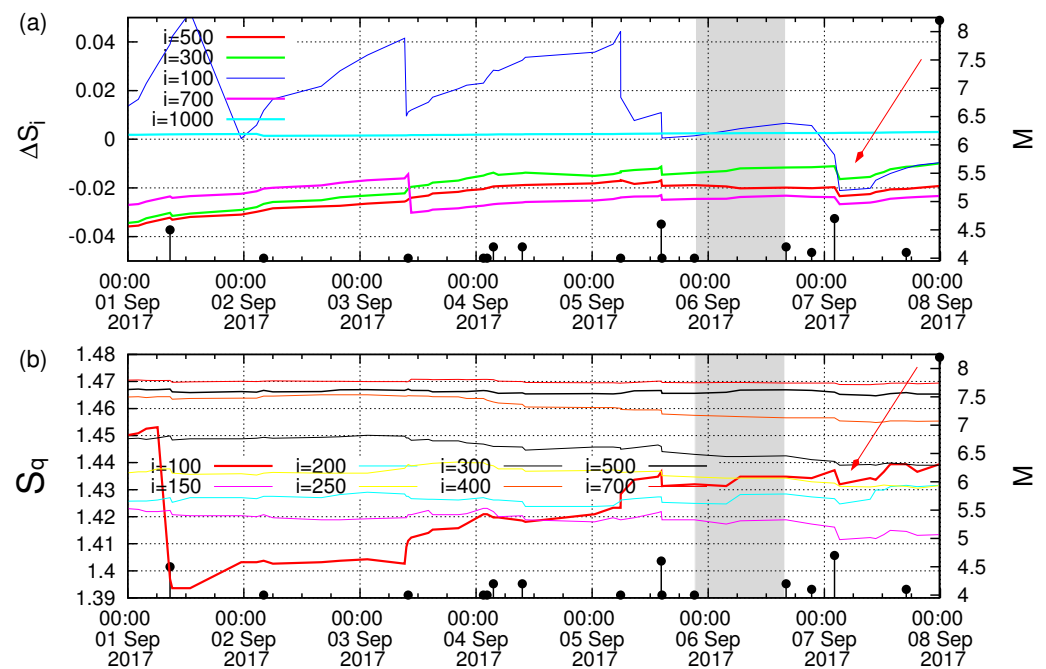


Figure 2. Plot of ΔS in NTA (a) and the Tsallis entropy S_q (b) versus the conventional time in the Chiapas region in Mexico before the occurrence of the M8.2 EQ on 7 September 2017. The red arrows indicate the simultaneous change of ΔS_i and S_q after the validity of the criticality condition $\kappa_1 = 0.070$. The shaded region corresponds to the period during which this criticality condition is obeyed [60].

Fourth, in view of the above observation (i.e., that after the system entered the critical stage, there exists the aforementioned similarity between the changes of the quantities ΔS and S_q), we are inspired to examine whether the precursory minimum ΔS_{min} on 14 June 2017 also exists in S_q . Hence, we plot in Figure 3b the S_q versus the conventional time for various scales from $i = 100$ to 700 events in the Chiapas region from 10 June 2017 until the M8.2 EQ on 7 September 2017. We find that there exists a minimum for the shorter scales which can be better seen in Figure 3c when we plot S_q versus the conventional time, but in an expanded scale. An inspection of the latter figure shows that the precursory minimum on 14 June 2017 we reported long ago in [31] on the basis of ΔS (see also Figures 4 and 5) also appears in S_q . We emphasize, however, that ΔS and S_q are *not* equivalent. In particular, the minimum observed when ΔS is used is *unique* during the whole period studied (see [31]), which is *not* the case of S_q , see Figure 3a where upon using S_q several minima appear. In other words, only when employing ΔS a unique minimum ΔS_{min} can be distinguished on 14 June 2017, while upon using S_q alone we cannot isolate—among several minima in Figure 3a—which is of truly precursory nature.

Fifth, in [69] we found that the change of the complexity measure Λ_i , i.e., $\Delta\Lambda_i(\equiv \Lambda_i(t) - \Lambda_i(t_{EQ}))$, upon the occurrence of the minimum ΔS_{min} , showed increased fluctuations of the entropy change under time reversal that obeyed the seminal work on phase transitions by Lifshitz and Slyozov [92] and independently by Wagner [93] (LSW). These works showed that the characteristic size of the minority phase droplets grows with conventional time t as $t^{1/3}$. Hence, we now turn to investigate whether it also holds for the Tsallis entropy S_q . In Figure 6, we plot the quantity $\Delta S_q = (S_q)_0 - S_q$ of the Tsallis non-additive entropy versus $t - t_0$ in the Chiapas region in Mexico before the occurrence of the M8.2 EQ on 7 September 2017 and $t_0 = 0.011$ days just after the M7.0 Guatemala EQ on 14 June 2017. $(S_q)_0$ corresponds to the value of S_q just before the latter EQ. Figure 6 reveals that ΔS_q obeys the $t^{1/3}$ time growth of the characteristic size of the minority phase droplets of the seminal LSW phase transition theory since the slope of the three straight lines drawn in this log-log plot is $1/3$.

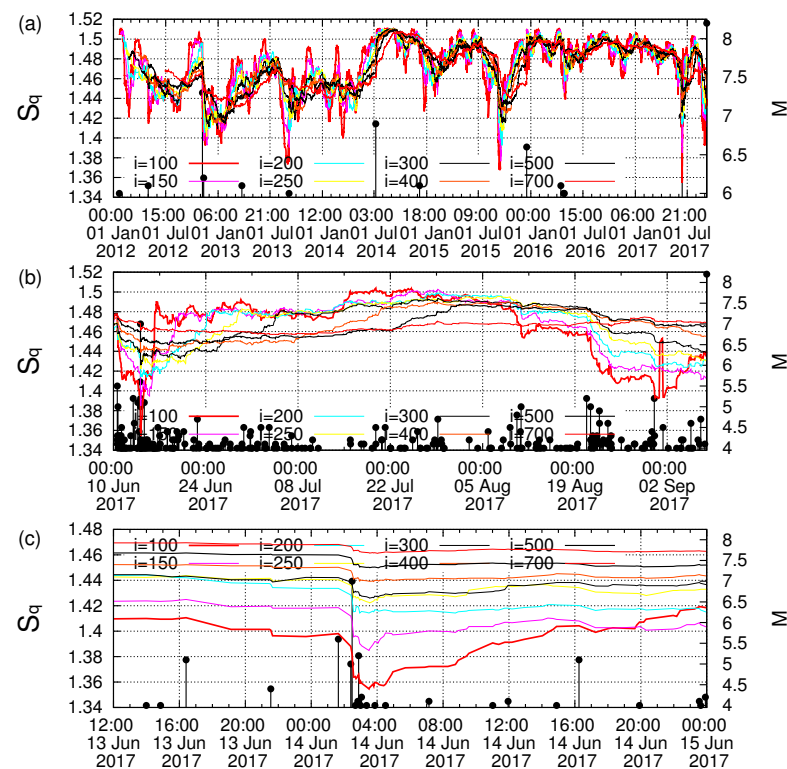


Figure 3. Plot of S_q versus the conventional time for various scales, i.e., $i = 100$ to 700 events, in Chiapas region from 1 January 2012 (a) or from 10 June 2017 (b) until the M8.2 Chiapas EQ. In (c), we plot an excerpt of (b), but in an expanded time scale during only the period in which ΔS in NTA showed a minimum on 14 June 2017 as described in [31]. See also Figures 4 and 5.

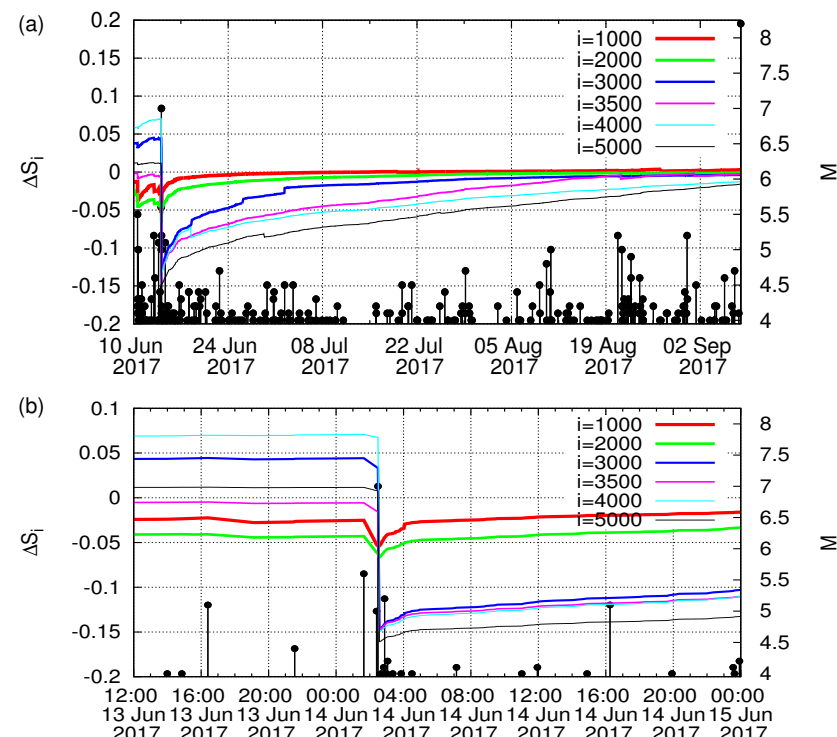


Figure 4. The same as Figure 2a, i.e., plot of ΔS_i versus the conventional time, but for the following periods: from 10 June 2017 until the M8.2 Chiapas EQ (a,b) in the expanded time scale from 12:00 CDT on 13 June until 00:00 CDT on 15 June 2017.

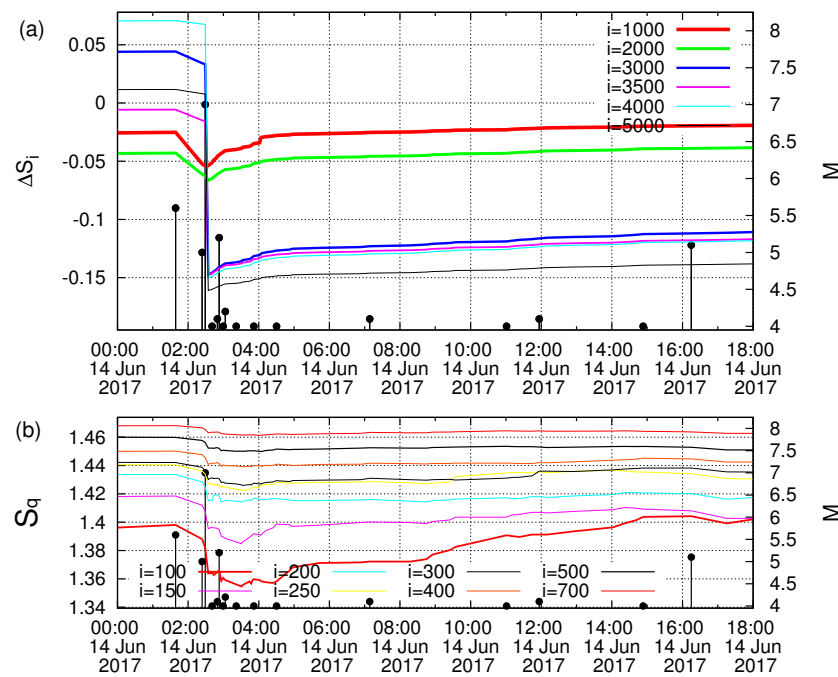


Figure 5. The same as Figure 2, but in expanded scale for the following periods: (a) ΔS_i from 00:00 CDT on 14 June 2017 until 18:00 CDT on 14 June 2017; (b) S_q from 00:00 CDT on 14 June 2017 until 18:00 CDT on 14 June 2017.

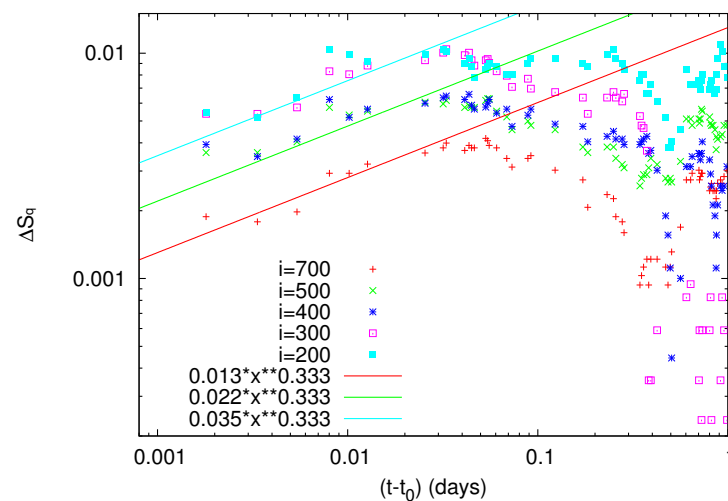


Figure 6. Plot of $\Delta S_q = (S_q)_0 - S_q$ of the Tsallis non-additive entropy versus $t - t_0$, where $t_0 = 0.011$ days, in Chiapas region in Mexico before the M8.2 EQ occurrence on 7 September 2017 and just after the M7.0 Guatemala EQ on 14 June 2017. $S_0(q)$ corresponds to the value of S_q just before the latter EQ. The slope of the straight lines of this log-log plot is $1/3$ which reveals the $t^{1/3}$ behavior of the LSW theory (see the text).

4.2. Results on the M7.1 EQ on 19 September 2017 within the Mexican Flat Slab

The first results on the NTA on this EQ have been published by Flores-Márquez et al. [84]. In short, the seismicity in the flat slab region has been analyzed in natural time from 1995 until the M7.1 EQ occurrence in 2017 by investigating the entropy change ΔS under time reversal together with the variability β of the order parameter of seismicity. The EQ catalog of SSN from 1 September 1995 until 24 September 2017, was used [84]. Considering the area of the flat slab and taking just the EQs with epicenters situated between 40 and 60 km of Moho depths, Flores-Márquez et al. [84] plot their spatial distribution in the upper panel of their Figure 2. To assure catalog completeness, a magnitude threshold

$M_{\text{thres}} = 3.5$ was imposed after studying the cumulative frequency magnitude distribution as in [84]. The quantity ΔS was minimized on 21 June 2017 which is approximately one week after the minimum ΔS_{min} that preceded the $M8.2$ Chiapas EQ on 7 September 2017. In particular, a clear minimum ΔS_{min} appeared on 21 June 2017 upon a $M4.8$ EQ occurrence, approximately 3 months before the $M7.1$ EQ occurrence on 19 September 2017. A β minimum was also observed during the period February–March 2017. In addition, it was shown that, after ΔS_{min} , the seismicity order parameter κ_1 starts diminishing by gradually approaching the critical value 0.070 around the end of August and the beginning of September 2017.

The criteria that assured the true coincidence of the EQ time series with that of critical state [12,82] have been checked during the period after 21 February 2017. This was obtained on the basis of the variability minimum which was unique during the whole period studied. A more detailed inspection uncovers that the second criterion for the true coincidence starts to be obeyed on 21 June 2017: The order parameter κ_1 after 21 June 2017 starts diminishing from values $\kappa_1 > 0.070$ and finally approaching from above the value $\kappa_1 = 0.070$ around the end of August and the beginning of September (this period is shaded in Figure 7). In other words, $\Pi(\omega)$ starts to follow the behavior indicated by the red arrow observed in Figure 5 of [84] on the date that ΔS_i showed the minimum, i.e., just after 21 June 2017.

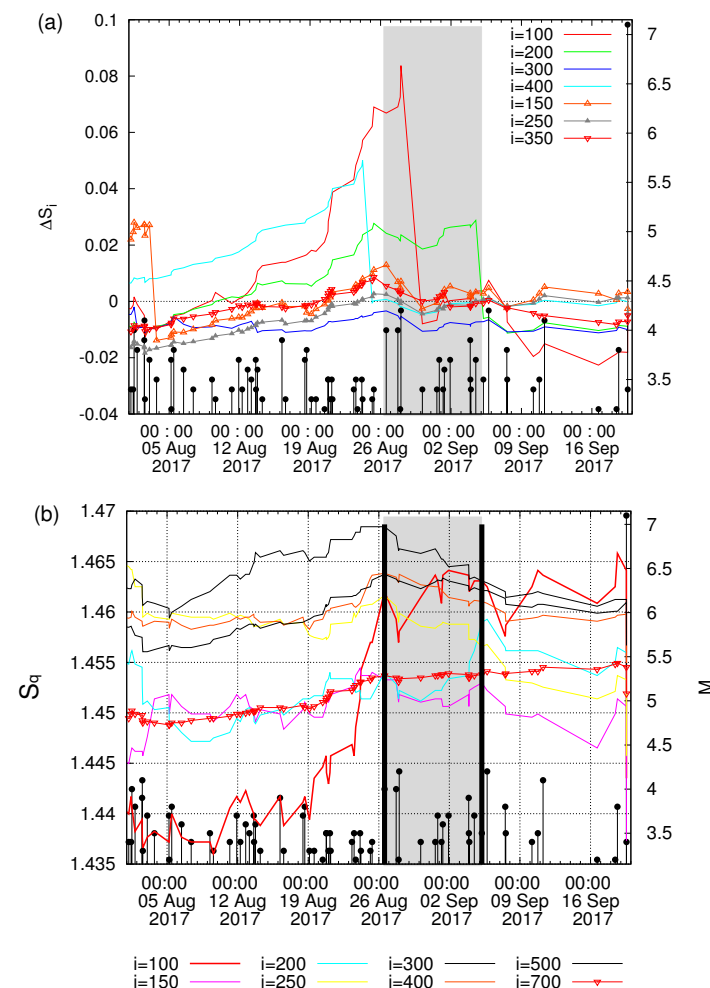


Figure 7. Plot of ΔS in NTA (a) and the Tsallis entropy S_q (b) versus the conventional time (CDT) before the occurrence of the $M7.1$ EQ on 19 September 2017 within the Mexican flat slab. The shaded region corresponds to the period during which the criticality condition $\kappa_1 = 0.070$ is obeyed [84].

The shaded area in Figure 7 indicates the period during which the condition $\kappa_1 = 0.070$ holds. After this period and in particular during the last few days before the $M7.1$ EQ occurrence,

a similarity of ΔS_i and S_q emerges in Figure 7. In this Figure, for better visualization of their changes the threshold $M_{\text{thres}} = 3.2$ has been adopted. Specifically, a number of transient changes of similar shape on both quantities appear, see for example the local minima of both ΔS_i and S_q in Figure 7a,b upon the occurrence of an event on 16 September 2017, as well as the local maxima upon the occurrence of a M3.8 EQ on 18 September 2017.

4.3. Results on the M7.1 Ridgecrest EQ on 6 July 2019 in California

Three EQs of moment magnitudes $M = 6.4, 5.4$, and 7.1 , north and northeast of the town of Ridgecrest, California, occurred in 2019. The latter, being the mainshock, was the most powerful EQ in the California state in 20 years, i.e., after the Hector Mine EQ in 1999. Specifically, at 17:33 UTC on 4 July 2019, a M6.4 foreshock occurred followed by a series of additional EQs, the strongest of which was a M5.4 at 11:07 UTC on 5 July. Finally, at 03:20 UTC on 6 July, a larger M7.1 EQ occurred in the Ridgecrest area followed by thousands of aftershocks. A map depicting all the $M \geq 2.0$ EQs reported by the Southern California Earthquake Center (SCEC) since 1 January 2004 is shown in Figure 9.7 of [13]. The EQ catalog of SCEC, available at http://service.scedc.caltech.edu/ftp/catalogs/SCEC_DC/, was used (accessed on 1 August 2019). Following previous work [94], we considered a magnitude threshold $M_{\text{thres}} = 2.0$ and focused in the analysis of seismicity within the geographic polygon shown in Figure 1 of [95] that covers the Southern California Seismic Network reporting area for local events. The β_W values have been computed for sliding natural time window lengths $W = 200, 250, 300, 350, 400, 450$, and 500 events and plotted in Figure 9.8 of [13] along with the Detrended Fluctuation Analysis [96] (DFA) exponent α_{300} —deduced from the EQ magnitude time series consisting of 300 events—versus the conventional time since 1 January 2004 until the M7.1 Ridgecrest EQ occurrence on 6 July 2019.

As can be seen in Figure 9.9a of [13], a few months before the 2019 M7.1 Ridgecrest EQ all β_W values exhibited an initial decrease after approximately 10 February 2019 and gradually approached their minimum values from 29 May to 5 June 2019. When these minima are reached, the corresponding DFA exponent α_{300} values were above $\alpha = 0.5$ (pointing to the presence of long-range correlations), as can be visualized in Figure 9.9b of [13]. In other words, the temporal correlations between EQ magnitudes deduced by means of DFA revealed that, while on 5 June 2019 the fluctuations of the order parameter of seismicity exhibited a minimum $\beta_{\min}$ showing long-range correlations ($\alpha_{300} \approx 0.6$), they have been destroyed before the M7.1 EQ since a behavior close to random ($\alpha = 0.5$) and subsequently anticorrelated ($\alpha < 0.5$) was observed almost upon of the M6.4 EQ occurrence on 4 July 2019.

Studying natural time analysis of seismicity from the minimum $\beta_{\min}$ of the variability on 5 June 2019, we find [59] that the criticality condition $\kappa_1 = 0.070$ is obeyed at 22:41 UTC on 2 July 2019 almost 3 days before the M6.4 EQ occurrence. By the same token as in Sections 4.1 and 4.2, we now depict in Figure 8 the quantities ΔS_i and S_q versus the conventional time. This figure shows that transient changes of these quantities *start* upon the M5.4 EQ occurrence on 5 July 2019 (marked by the red arrows “start”) and *end* in the beginning of 6 July 2019 (marked also with the red arrows “end”) upon a M3.46 EQ occurrence at 00:13 UTC (see Figure 3 of [59]), almost 3 h before the M7.1 Ridgecrest EQ. In other words, there exist transient changes on both quantities ΔS_i and S_q , which start as well end upon the occurrence of the events M5.4 EQ on 5 July and the M3.46 EQ on 6 July, respectively.

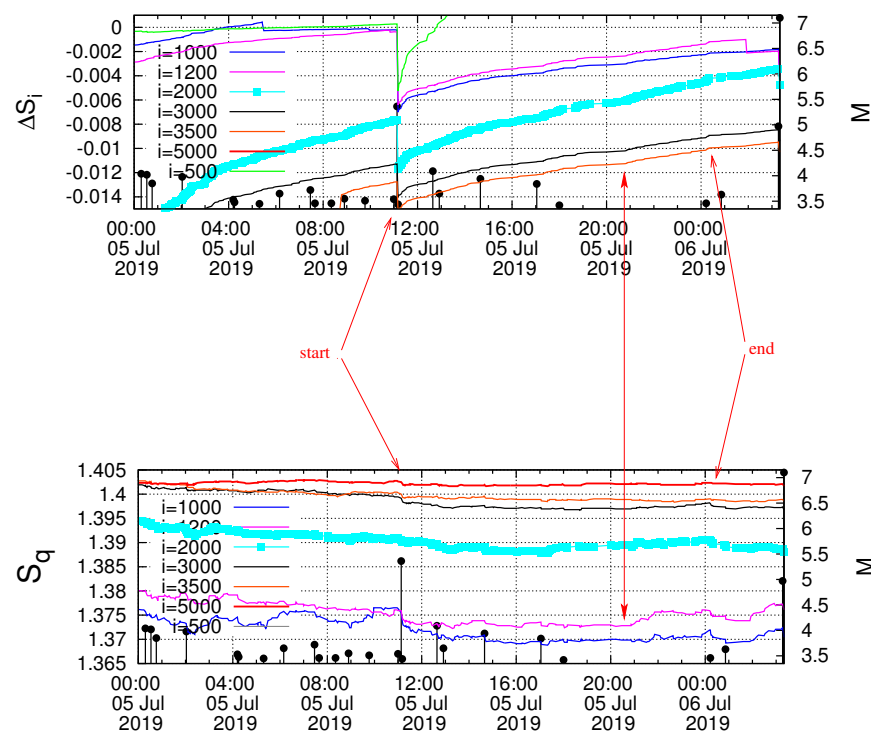


Figure 8. The quantities ΔS_i and S_q versus the conventional time before the $M7.1$ Ridgecrest EQ. Transient changes of these quantities *start* upon the $M5.4$ EQ occurrence on 5 July 2019 (marked by the red arrows “start”) and *end* in the beginning of 6 July 2019 (they are also marked with the red arrows “end”) upon a $M3.46$ EQ occurrence at 00:13 UTC, almost 3 h before the $M7.1$ Ridgecrest EQ occurrence. The criticality condition $\kappa_1 = 0.070$ was obeyed [59] at 22:41UTC on 2 July 2019 that precedes the period depicted in Figure.

4.4. Results on the $M9$ Tohoku EQ Occurrence on 11 March 2011

Approximately a day before the $M9$ Tohoku EQ occurrence on 11 March 2011 the following results have been obtained [97]: NTA revealed that the order parameter κ_1 of seismicity, and in particular from 08:36 to 13:14 LT on 10 March 2011, obeyed the critical condition $\kappa_1 = 0.070$ which signals that the main shock is going to occur within the next few days (up to around 11 days or so [43]). The data come from the Japan Meteorological Agency (JMA) seismic catalog upon setting a magnitude threshold $M_{\text{thres}} = 3.5$ to assure data completeness as in [80]. The following two important findings have been observed [97] just before this period: First, the Tsallis entropic index q exhibited distinct changes at 03:16 LT and 06:24 LT on 10 March 2011. Second, upon the $M7.3$ foreshock occurrence on 9 March 2011, a prominent increase Δq of the Tsallis entropic index q was observed that showed a scaling behavior with a characteristic exponent $1/3$ which conforms to the seminal work by LSW on phase transitions predicting that the time growth of minority phase droplets grows with time t as $t^{1/3}$. As for the prefactor A in the relation $\Delta q = A(t - t_0)^c$ of [97] of LSW theory, it increases when the scale i decreases, see Figure 3 of [97], in contrast to the complexity measure Δ_i quantifying the ΔS fluctuations for which the LSW prefactor A increases upon increasing the scale i [98].

The following additional results have just been deduced by Varotsos et al. [14]: After the above mentioned period from 08:36 at 13:14 LT on 10 March 2011 in which the system entered the critical stage, simultaneous changes appear at 18:00 and 20:00 LT on 10 March 2011 on both the quantity ΔS (see Figure 3a,b of Ref. [14]) and the Tsallis entropy S_q (see Figure 3c,d of Ref. [14]). These simultaneous changes can be seen with much difficulty when the computation is made in the entire Japanese region, but are evident when computed in the future epicentral region. A few hours later, the changes $\Delta \Delta_i$ of all the complexity measures $\Delta \Delta_{2000}$, $\Delta \Delta_{3000}$ and $\Delta \Delta_{4000}$ show a simultaneous variation almost

around 00:00 LT on 11 March 2011, i.e., almost several hours before the *M*9 EQ occurrence (see Figure 9 of Ref. [98]).

5. Discussion

Following [99], fluctuations are the dominant feature of criticality and their presence is spatially boundless while their duration is finite as they offer an ever-changing landscape. Their appearance is such that the critical state looks the same on any length scale, i.e., it is statistically scale invariant.

The fluctuations of a second order phase transition are usually studied by means of the Landau free energy in order to investigate the order parameter fluctuations and the validity of mean field theory. Close enough to the critical point, the fluctuations will become as important as the mean. The region near the critical point where fluctuations are important is known as the critical region. Close enough to the critical point, mean field theory breaks down and fluctuations are always important. The benchmark of the critical state is, as mentioned, its self similarity expressed by spatial and temporal scale invariance of correlations.

Concerning the interrelation between the order parameter fluctuations and non-extensivity, Tsallis et al. [100] showed that for asymptotically scale-invariant systems, it is S_q with $q \neq 1$, and not S_{BG} (cf. the subscript *BG* stands for Boltzmann–Gibbs entropy in which $q = 1$), the entropy which matches standard Clausius-like, prescriptions of classical thermodynamics. They also noted that the subtle case of thermodynamic critical points, where correlations at all scales exist. There, we can still refer to S_{BG} , but it exhibits singular behavior which is due to the fractal structure of the correlation clusters existing at critical points; an instructive description in non-extensive terms of such situations has been advanced in [101,102] the results of which show [102] that the departure from *BG* statistics and the applicability of q -statistics is due in part to the presence of the long range correlations in space and in time taking place at criticality. This sheds light on the point why we investigate here whether a simultaneous appearance in the changes of ΔS with those in S_q is observed after the system enters the critical stage. We recall, as mentioned in the Introduction, that ΔS plays a major role upon approaching a dynamic phase transition (critical point) in which the mainshock constitutes the new phase.

Finally, we note that since both ΔS and S_q are experimentally stable [65,66], as shown in [64,103], respectively, we do not expect the above findings to change within a plausible experimental error in the determination of the EQ magnitudes.

6. Summary and Main Conclusions

For the case of the *M*8.2 Chiapas EQ on 7 September 2017:

(a) The ΔS_{min} appears on 14 June 2017 almost three months in advance. As for the validity of the criticality condition $\kappa_1 = 0.070$, the corresponding criteria for the true coincidence hold from 21:08 CDT on 5 September 2017 until 16:06 CDT on 6 September 2017.

(b) After the latter period, ΔS exhibits a transient change several hours before the mainshock occurrence being strikingly similar in shape to an S_q change, see Figure 2.

(c) Both measures $\Delta\Lambda_i$ and ΔS_q obey the seminal LSW theory of phase transitions according to which the characteristic size of the minority phase droplets grows with the time as $t^{1/3}$. This, however, should not be misinterpreted that these measures are equivalent for the reason that will be explained at the end of this Section.

We now proceed to the results obtained here for the *M*7.1 EQ within the Mexican flat slab on 19 September 2017: After the validity of the criticality condition $\kappa_1 = 0.070$ lasting around from the end of August 2017 until the beginning of September 2017, we found transient changes of ΔS , accompanied by simultaneous changes of S_q from 11 September 2017 until 18 September 2017, i.e., almost one day before the *M*7.1 EQ occurrence within the Mexican flat slab, see Figure 7.

More or less similar phenomena appear before the *M*7.1 Ridgecrest EQ. In particular, transient changes ΔS_i accompanied by simultaneous changes of S_q start upon the *M*5.4 EQ

on 5 July 2019, i.e., several hours before the mainshock, until just a few hours before the M7.1 Ridgecrest EQ.

Finally, we recall that in [14] we have shown that shortening of the time window of the impending 2011 Tohoku M9 EQ has been achieved to several hours by combining NTA of seismicity, ΔS , and the Tsallis entropy.

After recapitulating the aforementioned results, we emphasize the following:

There exists a superiority of NTA compared to S_q as it emerges from the following facts: The ΔS minimum observed before the M8.2 Chiapas EQ on 14 June 2017, see Figure 7 of [31], is *unique* during the whole period studied. This is *not*, however, the case for the S_q , for example, see Figure 3c,d of [14] and Figure 3a where upon using S_q several minima appear, none of which can be distinguished as precursory minimum when using S_q alone. On the other hand, the minimum ΔS_{min} (see also Figure 8 of [31]) being accompanied by an evident increase in the complexity measure Λ_i can be uniquely distinguished as precursor.

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