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Vein Formation and Reopening in a Cooling Yet Intermittently Pressurized Hydrothermal System: The Single-Intrusion Tongchang Porphyry Cu Deposit

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Abstract: Porphyry deposits are the dominant sources of copper and major sources of several base and precious metals. They are commonly formed via the repeated emplacement of hydrous magmas and associated fluid exsolution. As a result, mineralized hydrothermal veins may undergo multiple deposition and reopening processes that are not fully accounted for by existing fluid models. The Tongchang porphyry Cu deposit is a rare example of being related to a single intrusion. The simplicity in intrusive history provides an ideal starting point for studying fluid processes in more complex multi-intrusion porphyry systems. Detailed scanning electron microscope (SEM) cathodoluminescence imaging (CL) revealed rich microtextures in quartz and anhydrite that point to a fluid timeline encompassing early quartz deposition followed by fluid-aided dynamic recrystallization, which was succeeded by an intermediate stage of quartz dissolution and subsequent deposition, and ended with a late stage of continuous quartz deposition, brecciation, and fracturing. Vein reopening is more common than expected. Fifteen out of seventeen examined vein samples contained quartz and/or anhydrite that was older or younger than the vein age defined by vein sequences. Thermobarometry and solubility analysis suggests that the fluid events occurred in a general cooling path (from 650 °C to 250 °C), interspersed with two episodes of fluid pressurization. The first episode occurred at high-T (>500 °C), under lithostatic conditions alongside dynamic recrystallization, whereas the second one took place at a lower temperature (~400 °C), under lithostatic to hydrostatic transition conditions. The main episode of chalcopyrite veining took place subsequent to the second overpressure episode at temperatures of 380–300 °C. The results of this study reaffirm that thermal and hydraulic conditions are the main causative factors for vein reopening and growth in porphyry deposits.

Keywords: CL; quartz; anhydrite; solubility; hydrothermal process

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1. Introduction

Porphyry deposits are the world's dominant sources of copper and molybdenum as well as important reservoirs of other base and precious metals [1]. They are one of the areas of focus of critical raw material studies [2]. The deposits are normally found along the convergent plate tectonic margins, formed via optimally-integrated magmatic and hydrothermal processes rooted in the subduction lithospheric mantle [3–7]. Despite several decades of studies, many details still await to be revealed, particularly regarding the mechanism of vein growth, reopening, and sulfide deposition [8–11].

The most widely-used porphyry fluid model classifies hydrothermal veins as early A-type quartz vein (700–550 °C), transitional B-type molybdenite quartz vein (500–400 °C), and late D-type pyrite veins (350–250 °C) based on vein morphology, mineralogy, and chronology. However, this model assumes that the veins are formed in a single event that

precipitates one generation of vein mineral. This assumption rarely holds for most porphyry deposits due to vein reopening, as revealed by several SEM-CL studies [12–18]. Erroneous interpretations on fluid source and conditions may result if they are based on bulk analyses of fluid inclusions, mineral chemistry, and isotopes of reopened veins.

Studies have shown that vein reopening in magmatic hydrothermal deposits may be linked to the number of igneous intrusions, which dictates the thermal and hydraulic properties of the system [8,19]. Ore deposits with a single intrusion may experience less reopening compared to multi-intrusion ones due to monotonic cooling, and thus less complicated hydraulic fracturing. The Tongchang Cu deposit is genetically associated with a single granodiorite intrusion [20]. This study presents a novel fluid model for the Tongchang Cu deposit by using integrated techniques of SEM-CL imaging [13,21,22], TitanQ geothermobarometry [23–27], fluid inclusion microthermometry [13,28–30], and mineral solubility analysis [31,32].

2. Regional Geology

Most porphyry deposits form in terrains overriding subduction zones, where crustal magma chambers are fed by arc magmas [4]. Repetitive magma replenishment causes magma chambers to evolve chemically and mechanically, and drives the intrusion of multiple intrusive stocks [33]. In contrast, porphyry deposits with a single intrusion are rarely reported. Their scarcity is likely linked to the unique tectonic histories of the host terrain. The Tongchang porphyry deposit is a mono-intrusion system located in the inland of the South China Block (SCB), over 500 km from contemporaneous subduction belts [34] (Figure 1a). Interestingly, the deposit is situated within a much older subduction zone that sutured the SCB in the Neoproterozoic [35–37]. Despite the multiple reworking events in the Paleozoic and Mesozoic eras, geochemical and isotopic studies of the porphyry magmas have established connections with the Proterozoic subduction event [20,25,38,39]. For the sake of relevance and clarity, we describe Neoproterozoic and Mesozoic geology below.

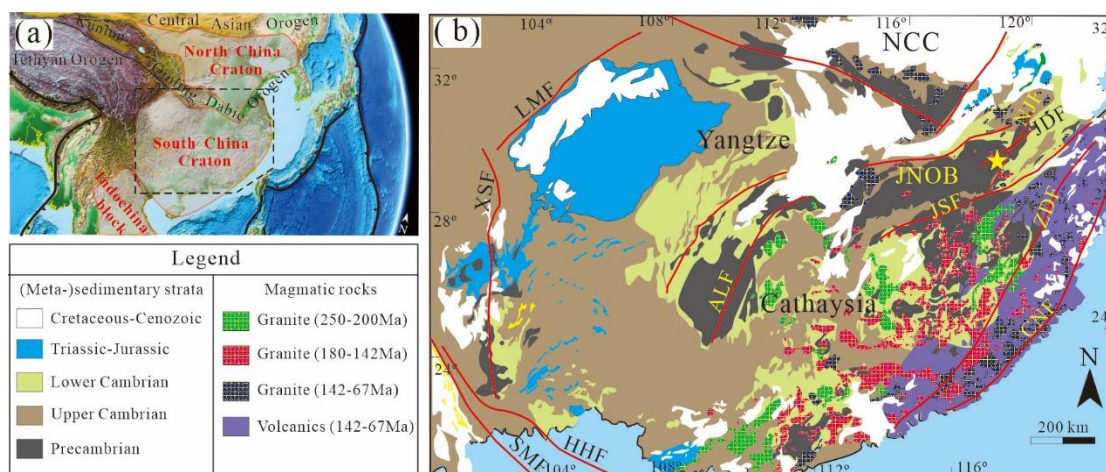


Figure 1. (a) A satellite map showing the location of the South China Craton (SCC) and surrounding geologic terrains; (b) A geological map showing major lithologies and structures in the Phanerozoic era of SCC (adapted from [34]). AHF: Ailaoshan-Honghe Fault; CNF: Changle-Nanao Fault; JDF: Jingdezhen Fault; JNOB: Jiangnan Orogenic Belt; JSF: Jiangshan-Shaoxing Fault; LMF: Longmen-shan Fault; HHF: Honghe Fault; NCC: North China Craton; NJF: Northeast Jiangxi Fault; QDOB: Qingling-Dabie Orogenic Belt; SMF: Song-Ma Fault; XMF: Xiangtan-Miluo Fault; XSF: Xianshuihe Fault; ZDF: Zhenghe-Dapu Fault. The yellow stars mark the location of the Dexing (including Tongchang) porphyry district.

2.1. Neoproterozoic Geology

The eastern SCB comprises the Jiangnan Orogen and Cathaysia block [40] (Figure 1b), separated by the Jiangshan—Shaoxing Fault [37]. The orogen consists of three tectonic domains separated by the Northeast Jiangxi [41] (Figure 1b) and Xiangtan-Miluo faults [42]. The northeast domain is a volcanic arc terrain of arc-derived granite and volcanic rocks (970–880Ma [43]), with minor turbidite, ophiolitic mélange, and blueschists [44]. The central and southwest domains expose low-grade metamorphosed volcano-sediments and intruding S-type granites [41] (860–800 Ma). The volcano-sedimentary sequences are deposited in two periods, separated by an angular unconformity [45]. The underlying sequences are dominated by flysch-like successions of sandstone, siltstone, slate, and phyllite deposited at 860–825Ma, including the Shuangqiaoshan Group in the northern Jiangxi and equivalents in other provinces [41,46]). The overlying sequence consists of slightly metamorphosed supracrustal rocks deposited in the Nanhua rift (e.g., the Banxi Group and equivalents [47]). The Cathaysia block comprises the Paleoproterozoic Badu Complex and Wuyishan granites as the basement, overlain by early Neoproterozoic volcanic successions (1000–900 Ma [48,49]). These rock associations have been interpreted as products of slab subduction [34,41,42,47,50,51].

2.2. Mesozoic Geology

The SCB is unconformably overlain by upper Neoproterozoic to lower Paleozoic, mid-Devonian to mid-Triassic, upper Triassic to lower Jurassic, and Cretaceous strata [34] (mainly clastic, volcanic and carbonate rocks, Figure 1b). The angular unconformities formed in three orogenic events in the Silurian, Triassic, and Jurassic eras [34]. Contemporaneous igneous activity was developed in large volumes. The Early Paleozoic (460–400 Ma) and Triassic (250–200 Ma) magmatic rocks are mainly S-type granites [52,53]. The Jurassic and Cretaceous igneous rocks are mainly A-type granite, syenite, and volcanic rocks, formed in two stages: i.e., 180–142 Ma and 142–67 Ma [53]. The early to middle Jurassic (190–170 Ma) is considered a period of magmatic quiescence with rock exposures only in the Dexing, Nanling and Qin-Hang regions [54]. Numerous tectonic models have been put forward for Mesozoic tectonics [34,53,55–62] (Figure 1d), all suggesting that the Dexing region had undergone an intraplate extension during this time period [20,38,63].

3. Deposit Geology

The Tongchang deposit is one of three mineralizing centers in the Dexing Cu district (Figure 2). Along with Zhushahong and Fujiawu, the three deposits appear in the intersections of regional and local faults that delineate a NW-SE alignment. The three granodiorite porphyries are identical in mineralogy and rock chemistry [20,38], and emanated from a common batholith at depths confirmed by borehole drillings. Alteration patterns are clearly separated despite the fact that the intrusions are close to one another (Figure 2). The ore types are similar except for different Au and Mo contents.

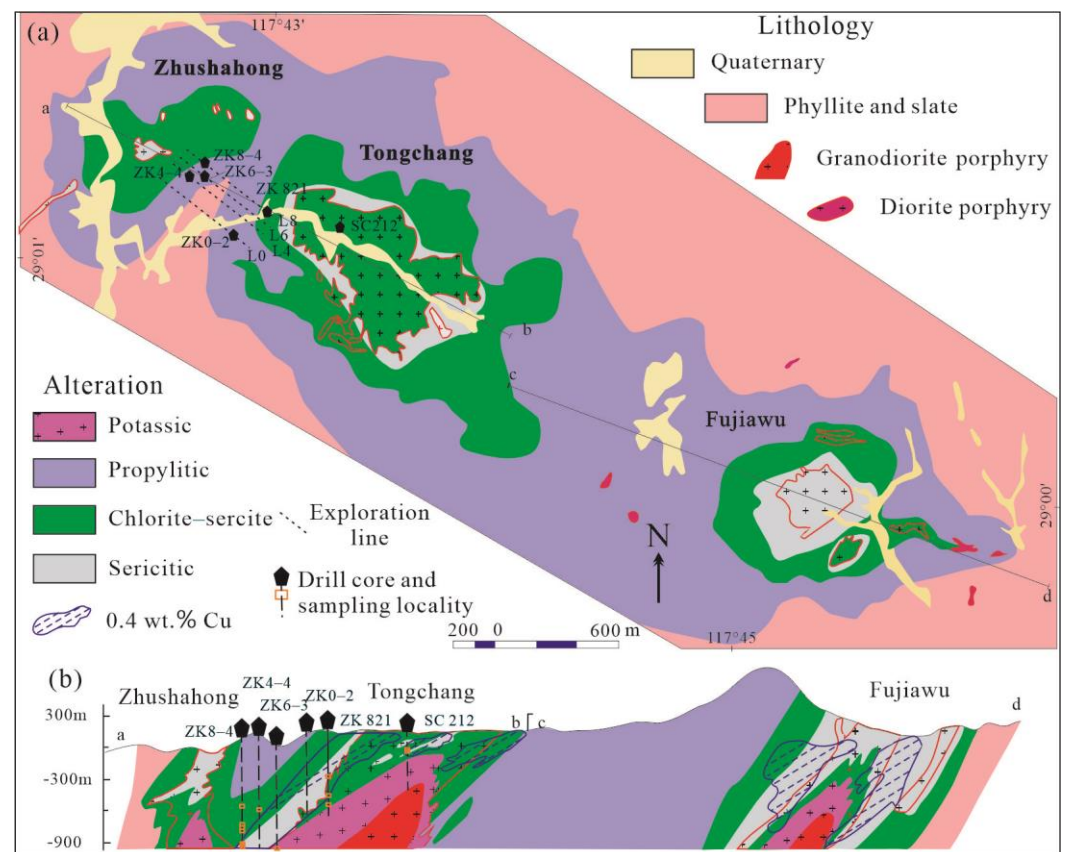


Figure 2. (a) A geological map of the Dexing porphyry Cu district; (b) Two cross sections (a–b, and c–d) of the Cu district (adapted from [19,38]). Dash lines denoted with L0–8 represents exploration lines; black pentagons represent the locations of drill hole, and the sampling localities are marked with small rectangles.

3.1. Wall Rock and Intruding Porphyries

Wall rocks of the Dexing district are phyllite and tuffaceous slate of the Shuangqiaoshan Group (Figure 3a,b). Biotite and actinolite hornfels are developed in the wall rocks surrounding the granodiorites (Figure 3c). The granodiorites are obliquely emplaced, and account for over 95 vol.% of igneous intrusions in this district. They show porphyritic texture in the peripheral and upper part of the intrusions (Figure 3d), and the center and lower portions grade into an equigranular texture (Figure 3e). Phenocrysts and groundmass have similar mineral composition to andesine, K-feldspar, quartz, hornblende, and biotite. Both the granodiorite and wall rocks are hydrothermally altered and mineralized (Figure 3f). LA-ICP-MS zircon U–Pb dating suggested that the granodiorite formed at 171 ± 3 Ma [20,38,63], contemporaneous with Cu–Mo mineralization (molybdenite Re–Os age of ~ 170 Ma [64]). The porphyry magmas are interpreted as a partial melt of Neoproterozoic subduction-modified lithosphere [20,38]. There is minor (<5 vol.%) post-ore diorite porphyry and aplite (Figure 3g–i). The diorite truncates mineralized veins (Figure 3h,i), and has been dated at ca. 154 Ma by the zircon U–Pb method [65].

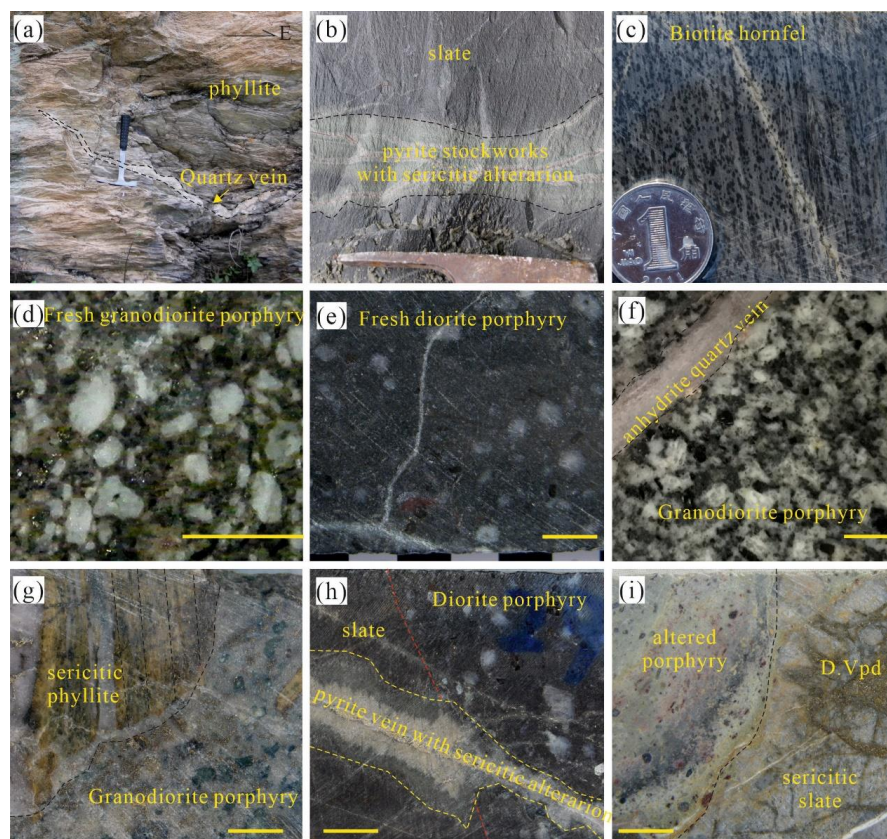


Figure 3. Major lithologies (phyllite, slate, hornfel, and porphyries) and their spatial relation at the Tongchang porphyry Cu deposit. (a) an outcrop of phyllite; (b) a slate with pyrite stockworks and sericitic alteration halo; (c) a biotite hornfel; (d) a fresh granodiorite porphyry; (e) a fresh diorite porphyry; (f) a granodiorite porphyry crosscut by an anhydrite quartz vein; (g) a granodiorite porphyry truncating phyllite, both of which are altered by sericitic alteration; (h) a dioritic porphyry truncating slate, both of which are crosscut by a pyrite vein; (i) an altered porphyry truncating D-type pyrite vein (D.Vpd) hosted in the slate. The scale in (c) is a Chinese one-dime coin.

3.2. Hydrothermal Alteration

Typical hydrothermal alteration types are developed in the Tongchang deposit including potassic, propylitic, chlorite-sericite, and sericitic assemblages [66].

The potassic alteration is preserved only in the deep parts of the porphyry stock. At shallow levels, potassic alteration is largely obscured by chlorite-sericite overprints (Figure 4a). In the potassic altered granodiorite, porphyritic textures are preserved, but primary minerals are widely replaced by secondary K-feldspar, biotite, and iron oxides. Propylitic alteration is pervasive in the peripheral rocks. Rock textures are largely unaffected with carbonation and hydration reactions.

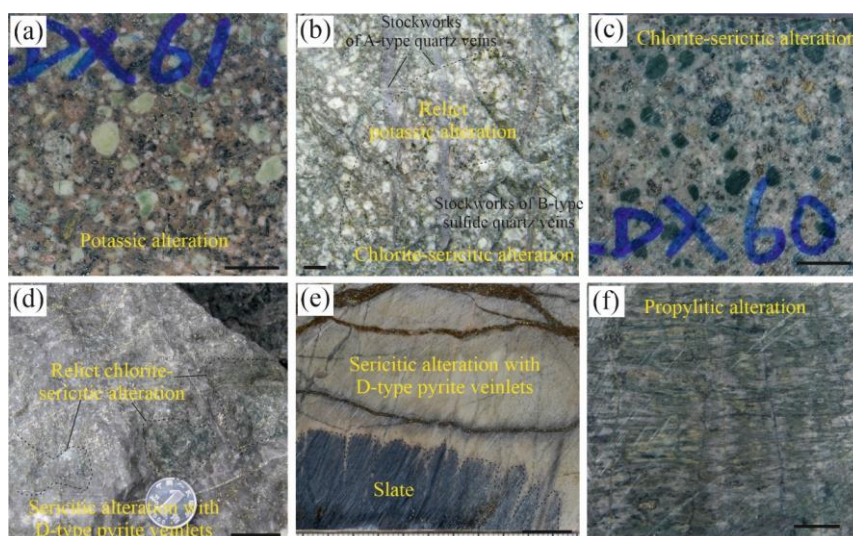


Figure 4. Major types of hydrothermal alteration at the Tongchang porphyry Cu deposit. (a) Potassic alteration in the granodiorite porphyry; (b) Chlorite-sericite overprinting on potassic alteration in the granodiorite porphyry; (c) Chlorite-sericitic alteration in the granodiorite porphyry; (d) Sericitic overprinting on chlorite-sericitic alteration in the granodiorite porphyry; (e) Sericitic alteration with pyrite-dominated veins in the slate; (f) Propylitic alteration in the phyllite. Scale bars are 1 cm. The scale in (d) is a Chinese one-dime coin.

Moving away laterally from the granodiorite-wall rock contact, chlorite–sericite alterations overprint on potassic and propylitic alterations (Figure 4b,f). While rock textures are largely preserved, most previous minerals were converted to chlorite, sericite, chalcopyrite, rutile, quartz, and anhydrite (Figure 4d).

The spatial distribution of sericitic alterations is structurally controlled. It is mainly developed in the granodiorite-wall rock contacts and surrounding faults and fractures. It obliterated earlier rock textures and converted all mafic minerals and feldspars into muscovite, rutile, pyrite, and quartz (Figure 4c,e).

3.3. Hydrothermal Veinlets and Mineralization

Stockworks of veins and veinlets are widely developed alongside the hydrothermal alterations. A previous work recognized three vein types conformable to the typical A–B–D classification, as mentioned previously. [67] further distinguished four subtypes of A-type veins, seven subtypes of B-type veins, and three subtypes of D-type veins. The early A veins are mostly wavy quartz dominant veins with variable amounts of K-feldspar, biotite, and magnetite. The transitional B veins are mainly straight-walled chlorite and quartz veins with pyrite, chalcopyrite, molybdenite, and bornite, along the vein centerline and edges. The late D pyrite-dominant veins have variable amounts of chalcopyrite and quartz. A detailed vein petrography is provided in Section 4.1.

The district has a total resource of 1870 Mt ore at an average grade of 0.46% Cu. Individually, Tongchang has 5.2 Mt Cu at 0.45%, 0.128 Mt Mo at 0.01%, 215 t Au at 0.19 g/t and 1279 t Ag; Fujiawu has 2.57 Mt Cu at 0.5% and 0.168 Mt Mo at 0.03%; and Zhushahong has 0.6 Mt Cu at 0.42% [20]. Over 85% of the ore is localized in the chlorite–sericite and sericitic alteration zones, displaying a hollow cylinder shape (Figure 1d). Copper minerals are predominantly chalcopyrite (>90%), with lesser bornite, tennantite, and chalcocite. The molybdenum is mainly in molybdenite. Gold is recovered as native gold and electrum [68]. Sulfide occurrences are different in the chlorite-sericite and sericitic zones. In the former, copper sulfides and molybdenite are primarily contained in the B-type chalcopyrite veins and altered mafic minerals with high chalcopyrite to pyrite ratios (2–0.5). In the sericitic zones, copper sulfides occur primarily as inclusions in pyrite and occasionally as tennantite aggregates [64] with low chalcopyrite to pyrite ratios (<0.1). Molybdenite is

contained in D-type pyrite veins and altered wall rocks. Native Au, electrum and Au-rich pyrite are primarily hosted in the sericitic zone.

4. Samples and Methods

Sixteen representative rock samples were selected to represent different alteration and vein types. The sampling localities are marked in Figure 2 and Table 1, which include mining platforms (20 and 230 m) and drill cores of the Tongchang open pit (SC212), as well as drill cores (ZK0-2, ZK4-4, ZK6-3, ZK8-4, ZK821). Detailed descriptions of sample petrography are given in the Appendix A.

Table 1. A summary of sampling locality and petrography of the analyzed rock samples of the Tongchang porphyry Cu deposit.

Sample NO.	Sampling Location	Vein type	Host Rock Alteration
12LDX37	Drill hole SC212 171m	A.V _{hq}	The rock is significantly affected by chlorite-sericite alteration. Primary mafic minerals are replaced by chunks of chlorite, sericite, hematite, muscovite, and chalcopyrite. Feldspar phenocrysts are replaced by sericite. Quartz phenocrysts remain largely unaffected. Groundmass is replaced by fine-grained hematite and chalcopyrite. The rock is intensively crosscut by hematite quartz veins (V _{hq}), which are subsequently cut by chalcopyrite pyrite quartz veins and veinlets (V _{cpq}).
12LDX40	Drill hole SC212 175m		Feldspar phenocrysts are completely replaced by sericite. Primary mafic minerals are replaced by chlorite, rutile, and muscovite. Groundmass consists of muscovite, quartz, sericite, and chalcopyrite.
10DX172	Drill hole ZK8-4 933m	A.V _{kq}	Plagioclase phenocrysts are partly (less than 50% of the surface) altered by sericite along cleavages and cracks. K-feldspar phenocrysts are completely replaced by sericite. Unidentified mineral pseudomorphs of inner sericite and outer chlorite, hematite and magnetite are present. Primary mafic minerals are replaced by chunks of small biotite, chlorite, hematite, and magnetite. Groundmass consists of sericite, quartz, biotite, hematite, and magnetite.
10DX197	Drill hole ZK8-4 777m		K-feldspar phenocrysts are partly altered by muscovite and chlorite, and plagioclase are pseudomorphed by sericite and chlorite. Quartz grains are overgrown and enlarged. Primary biotites are pseudomorphed by muscovite, chlorite, rutile, and quartz. Numerous chunks of chlorite, sericite, muscovite, rutile, chalcopyrite, epidote, and calcite are present.
12DXF05	Open pit 230m platform		Feldspar phenocrysts are overgrown by K-feldspar, which are partly altered by sericite. Primary biotites are replaced by quartz, sericite, chalcopyrite, pyrite, and rutile. Quartz phenocrysts remain largely unaffected. Numerous chunks of chlorite are present. Groundmass consists of fine-grained quartz.
10DX168	Drill hole ZK8-4 943m	A.V _{aq}	K-feldspar phenocrysts are pseudomorphed by sericite, and some anhedral K-feldspar grains are partly altered by sericite, chlorite, and epidote. Clusters of chlorites, sericite, anhydrite, epidote and molybdenite are present. Some quartz phenocrysts are overgrown, and others remain unaffected.
12DX500	Drill hole ZK821 620m		K-feldspar phenocrysts are largely replaced by sericite, and plagioclase is slightly altered by sericite. Primary biotite remains largely unaffected, and some biotites are related by small, unoriented biotite, which are replaced by chlorite, sericite, quartz, hematite and magnetite. Groundmass consists of quartz, sericite, and chlorite.
10DX22	Drill hole ZK8-4 587m	B.V _{cpq}	The host rock is a diorite porphyry. Feldspars are completely replaced by sericite in the center and chlorite in the rim. Groundmass mainly consists of small biotite, sericite, and minor amounts of quartz.
10DX261	Open pit 20m platform		Feldspar phenocrysts are completely replaced by sericite and muscovite. Primary biotites are pseudomorphed by chlorite, rutile, sericite and chalcopyrite. The rock is intensively cut by chlorite-muscovite-chalcopyrite-pyrite veins.

12DX328	Drill hole ZK821 499m	B.V _{macpq}	The center of feldspar phenocryst is replaced by chlorite while rim is replaced by sericite. Primary mafic minerals are replaced by chlorite, rutile, chalcopyrite, hematite, and magnetite. Aggregations of chlorite-rutile- magnetite, chlorite-chalcopyrite-rutile-anhydrite are present. Groundmass consists of quartz, anhydrite, rutile, and muscovite. Molybdenite occurs in both alteration and vein.
10DX164	Drill hole ZK8-4 952m	D.V _{mqp}	Feldspar phenocrysts are pseudomorphed by sericite. Mafic mineral pseudomorphs are not common. Where observed, they consist of chlorite and pyrite. The rock is intensely silicified. Quartz in the groundmass is medium-sized, intergrowing with abundant muscovite and chlorite. Sulfides are dominated by pyrite and occur as aggregations and bands.
10DX220	Drill hole ZK8-4 601m		The immediate host rock consists of strong Sericitic alteration halos of the molybdenite quartz pyrite vein. In the halo, primary minerals except quartz are completely replaced by large muscovite, rutile, quartz, and pyrite. No mineral pseudomorphs are eliminated. Groundmass consists of fine-grained quartz and large muscovite.
10DX140	Drill hole ZK0-2 337m	D.V _{pd}	The rock is a sericitic altered phyllite, consisting of oriented sericite, quartz, pyrite, and rutile.
10DX145	Drill hole ZK4-4 410m		It is a sericitic altered phyllite, consisting of oriented sericite, quartz, pyrite, and rutile.
10DX158	Drill hole ZK6-3 1067m		It is a slate with Propylitic alteration, consisting of randomly distributed chlorite, sericite, and quartz.
10DX201	Drill hole ZK8-4 738m		It is a breccia consisting of pyrite, quartz, and carbonate.

4.1. Vein Petrography

Three subtypes of A-type veins have been selected: hematite-quartz (A.V_{hq}), K-feldspar-quartz (A.V_{kq}), and anhydrite-quartz (A.V_{aq}). In the abbreviation, A denotes A-type, and the subscripts denote vein minerals in increasing mounts from left to right. The same abbreviation rule is used for all vein types.

A.V_{hq} veins have irregular vein boundaries without discernible alteration halos. They consist of about 95% quartz and 5% hematite (Figure 5a). They are commonly crosscut and reopened by chlorite-sericite-chalcopyrite-magnetite veins. A.V_{kq} veins consist of about 80–90% quartz and 10–20% K-feldspar, and are commonly enveloped by potassic halos (Figure 5b). Similarly, they are commonly crosscut and reopened by veinlets of chalcopyrite, muscovite, pyrite, and carbonate. A.V_{aq} veins consist of about 70% quartz, 20–30% anhydrite and variable amounts of pyrite, chalcopyrite, chlorite, carbonate, rutile, molybdenite, apatite and epidote (Figure 5c,d). They are commonly sandwiched by weak potassic halos. No crosscutting relationships have been observed between these three vein types.

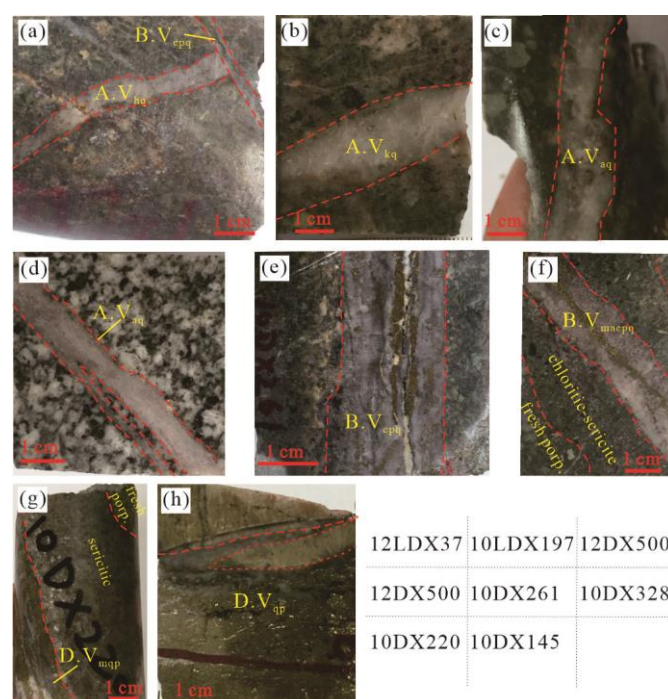


Figure 5. Hydrothermal vein types recognized in the Tongchang porphyry Cu deposit. (a) An A-type hematite quartz vein in the granodiorite porphyry (A.V_{hq}), crosscut by a B-type chalcopyrite pyrite quartz vein (B.V_{cpq}); (b) An A-type K-feldspar quartz vein in the granodiorite porphyry (A.V_{kq}); (c,d) A-type anhydrite quartz veins in the granodiorite porphyry (A.V_{aq}); (e) A B-type chalcopyrite pyrite quartz vein in the granodiorite porphyry (B.V_{cpq}); (f) A B-type molybdenite anhydrite chalcopyrite pyrite quartz vein in the granodiorite porphyry (B.V_{macpq}) with chlorite-sericite halos; (g) A D-type quartz pyrite vein in the granodiorite porphyry with sericitic halo (D.V_{qp}); (h) A D-type pyrite-dominated veins (D.V_{pd}) in phyllite.

Two B subtype veins have been selected, including chalcopyrite-pyrite-quartz (B.V_{cpq}) and molybdenite-anhydrite-chalcopyrite-pyrite quartz vein (B.V_{macpq}). They are characterized by centerlines of sulfides and straight vein boundaries without alteration halos. B.V_{cpq} veins consist of about 70–90% quartz and 10–30% sulfides, with variable pyrite to chalcopyrite ratios (Figure 5e). B.V_{macpq} veins differ from B.V_{cpq} veins with additional molybdenite and anhydrite in the vein and chloritic alteration halos. They commonly consist of about 70% quartz, 10% pyrite, 5–10% chalcopyrite, 5% anhydrite and minor amounts of molybdenite (Figure 5f). B.V_{cpq} and B.V_{macpq} crosscut B.V_{hq} veins but no other crosscutting relations have been observed between B subtypes.

Two D subtype veins have been selected including molybdenite-quartz-pyrite (D.V_{mqp}) and pyrite-dominant (D.V_{pd}). D.V_{mqp} veins containing about 70–80% pyrite and 15–20% quartz and 5% molybdenite, bordered by centimeter-wide sericitic alteration halos (Figure 5g). D.V_{mqp} veins are crosscut by D.V_{pd} veins, but their relationships with other vein types is unclear. D.V_{pd} veins are dominated by pyrite (>95%) with minor amounts of quartz (<5%), and are commonly bordered by sericitic alteration halos (Figure 5h). D.V_{pd} veins usually crosscut A.V_{hq}, B.V_{cpq}, and D.V_{mqp} veins.

4.2. Analytical Methods

300 µm-thick sections were made by conventional cutting, grinding, and polishing. Mineral textures and parageneses were studied with petrographic microscopes and scanning electron microscopes (SEM) equipped with secondary electron (SE), backscattered electron (BSE), cathodoluminescence (CL), and energy dispersive spectroscopy (EDS) detectors. A total of 774 grey-scale CL, 579 color CL, 782 BSE high-resolution images, and 42 X-ray maps were taken. Trace element concentrations in different generations of quartz were measured in situ by LA-ICP-MS. Overall, 656 analyses for quartz from 16 samples were made. After all analyses, the actual ablation spots were confirmed by new CL imagery.

4.2.1. SEM-CL Microscopy

Greyscale CL imagery for anhydrite and quartz were firstly obtained on a Tescan Vega3 SEM equipped with a Centaurus CL detector and a photomultiplier at the Western Washington University. The accelerating voltage was 15 kV, and the probe current was about 10 nA for both quartz and anhydrite. The photomultiplier had a spectral range from ~300 nm to ~600 nm. A later CL study was carried out at the Institute of Geology and Geophysics, Chinese Academy of Sciences (IGGCAS), with a Nova NanoSEM 450 SEM equipped with a Gatan MonoCL4 panchromatic detector. An acceleration voltage of 15 kV and beam current of 9–12 nA were used.

Color CL images of quartz and anhydrite were collected using a Tescan Vega3 SEM equipped with a Gatan ChromaCL2 multichromatic detector at the GeoRessources Lab, CNRS (UMR7359), Université de Lorraine. An acceleration voltage of 15 kV and beam current of ~10 nA was used. The CL detector records spectra in ranges of red, green, and blue, and produces an RGB image. A dwell time of 300 µs was set for each pixel. During a separate session aiming to test the lifespan of luminescence, images were taken after the samples were exposed to electron beams for a certain amount of time.

4.2.2. In-Situ LA-ICP-MS Trace Elements

Trace elements in quartz were first measured with an Agilent 7500 ce ICP-MS coupled with a NewWave 213 nm Nd:YAG solid state laser ablation system at WWU. A laser fluence of 15 to 20 J/cm², a repetition rate of 10 Hz, and a spot size of 65 µm were used. Helium was used as a carrier gas and mixed with make-up argon gas before entering the ICP torch. The Octupole Reaction System (ORS) was operated in hydrogen reaction mode (H₂ flow rate of 2 mL/min) to increase sensitivity. Oxide production was monitored with ¹³²Th¹⁶O/¹³²Th ratios, which were always less than 0.2%. Twenty-two masses were monitored: ⁷Li, ⁹Be, ¹¹B, ²³Na, ²⁷Al, ²⁹Si, ³⁰Si, ³¹P, ³⁹K, ⁴³Ca, ⁴⁹Ti, ⁵⁵Mn, ⁵⁶Fe, ⁶³Cu, ⁷¹Ga, ⁷⁴Ge, ⁷⁵As, ⁸⁸Sr, ¹¹⁸Sn, ¹²¹Sb, ¹³⁷Ba, and ²⁰⁸Pb. The dwelling time for ⁴⁹Ti was 0.1 s, and it was 0.01–0.03 s for the other masses. About 30 s of gas blank was collected as the baseline, which was followed by 60 s of data acquisition with laser ablation. Instrument performance was checked by a rough on-the-fly data reduction with the GLITTER program. Data were reduced with care later on using GLITTER. ³⁰Si was used as the internal standard using stoichiometric values. NIST 610 glass was used as an external standard, with reference values from [69]. A natural quartz standard (QZ7 [70]) and a synthetic quartz with a known chemical composition was used as quality controls. Results for QZ7 were within the range of

recommended values for all elements except lower Al values. A correction factor was obtained by dividing the recommended value by measured values of Qz7 (ranging from 0.51 to 0.76, Table 4), which was applied to unknowns of the same analytical session. Additional analyses were made with an Agilent 7900 ICP-MS equipped with a MICRO/Las Pro 193 nm ArF excimer laser ablation system at the Institute of Geochemistry of the Chinese Academy of Sciences (IGCAS). A repetition rate of 10 Hz and an energy density of 10 J/cm² were used. Spot size of the laser beam was 44 µm. NIST 610 glass was used as an external standard and analyzed twice every 10 analyses. An internal standard-independent calibration strategy, implemented in the ICPMSDataCal program, was used for quantification [71]. NIST 612, GSE-1G, and Qz7 were analyzed as quality controls. Analytical uncertainties of most elements for both setups were around 5–10% (1σ).

Although every effort was taken to avoid mineral and fluid inclusions, the influence of sub-microscopic inclusions was inevitable. During data reduction, abnormal spikes and bumps were filtered and excluded from integration. Even so, values of Na, K, Ca, Al, Mn, Fe, Cu, Zn, Sr, As, and Pb were too high to be real. Co-variations in Na-K-Al, Ca-Sr, Fe-Cu, Ca-Pb, K-Ca-Al, Cu-As, and Fe-Mn was interpreted to be the result of feldspar, carbonate, chalcopyrite/pyrite, tennantite, mica, and fluid inclusions, and were used as the basis to discard those values.

4.2.3. Fluid Inclusion Microthermometry

Fluid inclusions were classified as primary, secondary and pseudosecondary, following the principles of [72]. Phase transition temperatures were measured on a Linkam THMS 600 cooling/heating stage at the GeoResources Lab of the Université de Lorraine. The stage was calibrated against the melting point of solid CO₂ (−56.6 °C) and ice (0 °C), and a homogenization temperature (165 °C) of well-characterized natural fluid inclusions. The heating rate varied between 0.2 and 10 °C/min. When fluid inclusions approached the phase transition, the rate was lowered to 0.2–0.5 °C/min. Fluid salinity, density and isochores were calculated using the HokieFlincks_H₂O-NaCl program [73].

5. Results

5.1. SEM-CL Imaging

Quartz growth sequencing was based on CL microtextures and vein crosscutting relations. The earliest quartz generation (Qz1, Figure 6a), most common in the A.V_{hq}, B.V_{cpq}, and B.V_{macpq} veins, was characterized by variable grain sizes and straight grain boundaries. The grains commonly show undulose extinction under polarized light. They exhibit mottled CL with faint zoning, and abundant microfractures that are filled with CL-dark “splatters” and “cobwebs” (Figure 6a).

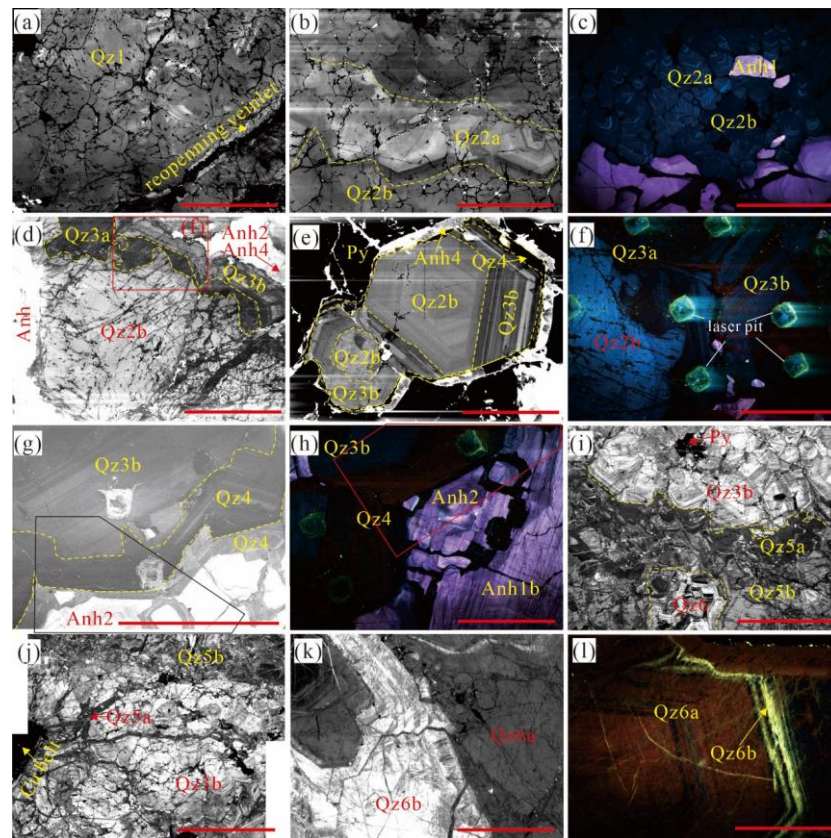


Figure 6. CL microtextures of quartz at the Tongchang porphyry Cu deposit. (a) The earliest CL-homogeneous quartz (Qz1) reopened by later veinlet; (b,c) The second generation (Qz2a and Qz2b) that shows dark blue CL with faint zoning. Also shown are three generations of anhydrite (purple); (d) The third generation of quartz (Qz3a and Qz3b) resorbing and overgrowing on early Qz2b. Note that the Qz3a has lower CL intensity than Qz3b does; (e) Qz3b directly overgrown on Qz2b, showing conformable contact with the surrounding pyrite; (f) Qz2 and Qz3 displayed blue CL with different intensities; (g,h) Coprecipitation of CL-dark Qz4 (overgrown blue-CL Qz3b) and Anh4 (cementing purple-CL Anh2 breccia); (i) Qz5a and Qz5b overgrew on Qz3b. (j) Qz5a fill in fractures in Qz3a. Note that Qz5a has lower CL than Qz5b; (k) CL-homogeneous Qz6a and CL-oscillatory Qz6b. Note that Qz6b has bright CL. (l) Qz6a shows red CL, whereas Qz6b shows bright blue C. Scale bars are 300 μm .

The second quartz generation of (Qz2), common in A.V_{kq}, A.V_{aq}, B.V_{cpq}, B.V_{macpq}, D.V_{pq}, and D.V_{pd} comprises anhedral grains with contrast sizes and irregular boundaries filled with CL-dark quartz. Qz2 grains are further divided into larger Qz2a grains that have concentric CL zoning and smaller Qz2b grains that show mottled CL (Figure 6b). The contacts between Qz2a and Qz2b can be both transitional and crosscutting. They form discrete layers in most samples. They also contain “splatters” and “cobwebs”. Under color CL, both grains show long-lived blue, with lower intensity for Qz2b (Figure 6c). Qz2 grains truncate Q1 locally, suggesting a later deposition.

The third generation of quartz (Qz3), common in the A.V_{kq}, B.V_{macpq}, D.V_{mpq}, D.V_{pq}, and D.V_{pd} veins, comprises a CL-dark core (Qz3a) and an overgrowth with oscillatory zoning (Qz3b) (Figure 6d). Some Qz3 grains contain a relict Qz1/Qz2 core overgrown either by Qz3a or Qz3b (Figure 6e). They also show long-lived blue CL (Figure 6f).

The fourth generation of quartz (Qz4), mainly observed in the A.V_{aq}, B.V_{cpq}, and B.V_{macpq} veins, is low in abundance. The Qz4 grains show dark, red CL and occur as thin rims or aggregates coating earlier quartz (Figure 6e,g,h). Qz4 crosscuts Qz3b grains locally.

The fifth generation of quartz (Qz5) commonly consists of a CL-dark core (Qz5a) and oscillatory overgrowth (Qz5b) (Figure 6i). They are observed in the D.V_{mpq}, D.V_{pq}, and

D.V_{pd} veins, and are intimately associated with pyrite. Locally, they occur as dark bands crosscutting or filling the interstitial space of earlier quartz (Figure 6j).

The latest generation of quartz (Qz6) comprises a CL-dark to grey interior (Qz6a) and a very bright-CL exterior (Qz6b) (Figure 6k). They were observed in the A.V_{aq}, D.V_{pq}, and D.V_{pd}. Qz6a exhibited an oscillation of red and blue CL, whereas Qz6b displayed a bright yellowish blue (Figure 6l).

Four generations of anhydrite have been identified. The first generation of anhydrite (Anh1) exhibited bright, homogeneous to patchy CL (Figure 7a). They are subdivided to Anh1a and Anh1b according to CL intensity. Anh1a grains have higher CL intensity and irregular boundaries (Figure 7a) relative to Anh1b grains that are commonly overgrown, recrystallized, or crosscut by later anhydrites.

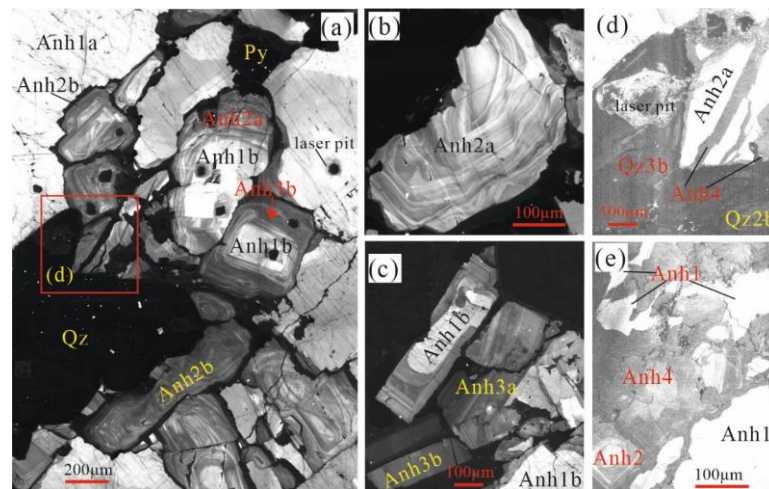


Figure 7. Cathodoluminescence (CL) microtextures of vein anhydrite at the Tongchang porphyry Cu deposit. (a) The first generation of anhydrite (Anh1), rimmed and surrounded by CL-oscillatory Anh2, coexisting with pyrite and quartz; (b) Anh2a showing “wavy” oscillatory zoning; (c) A cluster of euhedral Anh3a and Anh3b grains, with Anh3a showing oscillatory zoning, whereas Anh3b showed homogeneous CL; (d,e) CL-dark Anh4 cemented fragments of earlier anhydrite grains including Anh1 and Anh2.

The second generation (Anh2) shows bright to grey CL with “wavy” oscillatory boundaries (Figure 7a,b), with Anh2a being brighter and Anh2b darker.

The third generation of anhydrite (Anh3) occurred as euhedral crystals overgrowing Anh1 or infilled interstitial spaces of Anh1 and Anh2 (Figure 7a,c). They are subdivided to Anh3a and Anh3b according to CL intensity and zoning patterns (Figure 7c). Anh3a is brighter in CL and shows evident oscillatory zoning in contrast to Anh3b.

The last generation of anhydrite (Anh4) is anhedral and CL-dark (Figure 7d). Anh4 may cement Anh1 and Anh2 fragments (Figure 7e). All anhydrite generations except Anh3 have been found in both A.V_{aq} and B.V_{macpq}.

Mineralogical and textural interpretations suggest a sequential co-precipitation of Qz1 and hematite, Qz2a and Anh1a, Qz2b and Anh1b, Qz3 and Anh2, along with pyrite and rutile, Qz4, and Anh4 along with chalcopyrite, pyrite, rutile, molybdenite, and Qz5 along with pyrite, apatite, rutile, and gypsum (Table 2).

Table 2. A summary of generations, CL textures, mineral assemblages and Ti contents of quartz and anhydrite of the Tongchang porphyry Cu deposit.

Qz	Anh	Other Vein Minerals	CL	Ti (ppm)	Occurrence in Sample
Qz1		hematite	Qz: granular, homogeneous to faint zoning, grey.	27-43	10DX261, 12LDX37, 12LDX40
Qz2a	Anh1a	K-feldspar, hematite, magnetite	Qz: subhedral, oscillatory zoning, bright. Anh: homogeneous, bright.	80-220	10DX22, 10DX172, 10DX328, 12DXF05, 12DX500
Qz2b	Anh1b	hematite, magnetite	Qz: homogeneous to oscillatory zoning, bright. Anh: homogeneous, bright	50-70	10DX22, 10DX140, 10DX168, 10DX172, 10DX197, 10DX328, 12DXF05, 12DX500
Qz3a	Anh2a	rutile, hematite, magnetite	Qz: homogeneous, dark. Anh: euhedral oscillatory zoning	7-16	10DX140, 10DX168, 10DX197, 10DX220, 10DX261, 10DX328
Qz3b	Anh2b	epidote, rutile, pyrite, hematite, magnetite	Qz: euhedral oscillatory zoning. Anh: euhedral oscillatory zoning	2-35	
Qz4	Anh4	chlorite, sericite, rutile, bornite, chalcopyrite, pyrite, magnetite, monazite, apatite, epidote	Qz: homogeneous, dark. Anh: subhedral to euhedral homogeneous, dark.	3-17	10DX168, 10DX261, 10DX328, 12LDX37
Qz5a		sericite, muscovite, pyrite,	Qz: homogeneous, dark.	8-10	10DX164
Qz5b		molybdenite, rutile, apatite, gypsum	Qz: euhedral oscillatory zoning.	1-15	10DX140, 10DX145, 10DX164, 10DX201, 10DX220
Qz6a		Not observed	Qz: euhedral oscillatory zoning, dark to grey.	3-8	10DX140, 10DX158, 10DX201
Qz6b		Not observed	Qz: euhedral oscillatory zoning, very bright.	0.3-2	

5.2. Fluid Inclusions

Fluid inclusions at Tongchang were named using the nomenclature of [74] based on the relative volume fraction of bubbles and solid species (e.g., B15HS denotes 15% vapor phase, a halite daughter, and an opaque solid). Fluid inclusion assemblages (FIA) and their origin were determined with the aid of CL microtextures. Microthermometric temperatures and estimated salinity, density, and pressure are listed in Table 3.

Table 3. Fluid inclusion microthermometry of the Tongchang porphyry Cu deposit.

Type	Host Qz	T _{m,ice} (°C)	T _{m,halite} (°C)	Thv (°C)	Salinity (wt.%)	Density (g/cm ³)	dP/dT (bar/°C)	Homo. Mode
B15HS	primary in	\	458.4	369.7	54.2	1.2	19.1	halite dissolution
	QZ2a in	\	469.2	382.8	55.4	1.2	18.7	
	12DX500	\	512.1	415.3	60.4	1.2	20.3	
	primary in	−1.1	387.9	279.7	46.7	1.2	23.7	
	Qz2b in	\	396.1	269.4	47.4	1.2	26.2	
B80	10DX168	\	403.5	403.5	\	\	\	critical
		−4.2	344.6	344.6	6.7	0.7	8.5	
		\	356.8	356.8	\	\	\	
		\	316.5	316.5	\	\	\	
		\	320.7	320.7	\	\	\	
B50	primary in	\	300.6	300.6	\	\	\	
	Qz3b in	\	359.4	359.4	\	\	\	
	10DX168	\	359.4	359.4	\	\	\	
		−3.6	343.4	343.4	5.9	0.7	8.3	
		−2.7	299.6	299.6	4.5	0.8	10.4	
B30		−5.6	284.5	284.5	8.7	0.8	11.9	to liquid
		−3.2	287.5	287.5	5.3	0.8	11.2	
		−3.3	275.6	275.6	5.4	0.8	11.9	
	secondary	\	272.6	272.6	\	\	\	
	in Qz2b	\	255.5	255.5	\	\	\	
	and Qz3b	\	226.2	226.2	\	\	\	
	primary in	\	262	262	\	\	\	
	10DX168	−4.8	248.9	248.9	7.6	0.9	13.6	
		\	257.5	257.5	\	\	\	
		\	278.3	278.3	\	\	\	
B10		\	248.9	248.9	\	\	\	
		\	198.5	198.5	\	\	\	
		\	212.3	212.3	\	\	\	
	secondary	\	213.7	213.7	\	\	\	
	in Qz2b in	\	210.8	210.8	\	\	\	
	10DX168	−2.6	207.9	207.9	4.3	0.9	15.3	
		−3.4	186.6	186.6	5.6	0.9	16.6	

Due to intense recrystallization, Qz1 grains are devoid of primary FIAs, but flooded with secondary FIAs in healed fractures. Qz2a and Qz2b quartz host primary FIAs of B15HS and B80 inclusions (Figure 8a,b). B15HS inclusions in Qz2a and Qz2b showed different temperatures of halite dissolution and vapor disappearance. In Qz2a grains, B15HS showed halite dissolution at 458–512 °C and vapor disappearance at 370–415 °C, whereas those in Qz2b grains showed lower halite-dissolution temperatures (388–396 °C) and vapor disappearance temperatures (270–280 °C). One B80 inclusion in Qz2b was homogenized at 404 °C via critical mode.

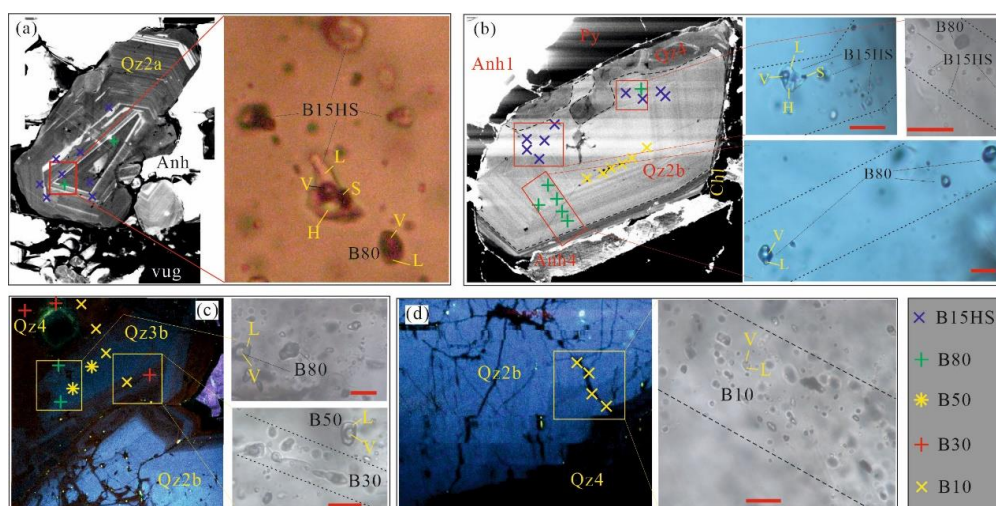


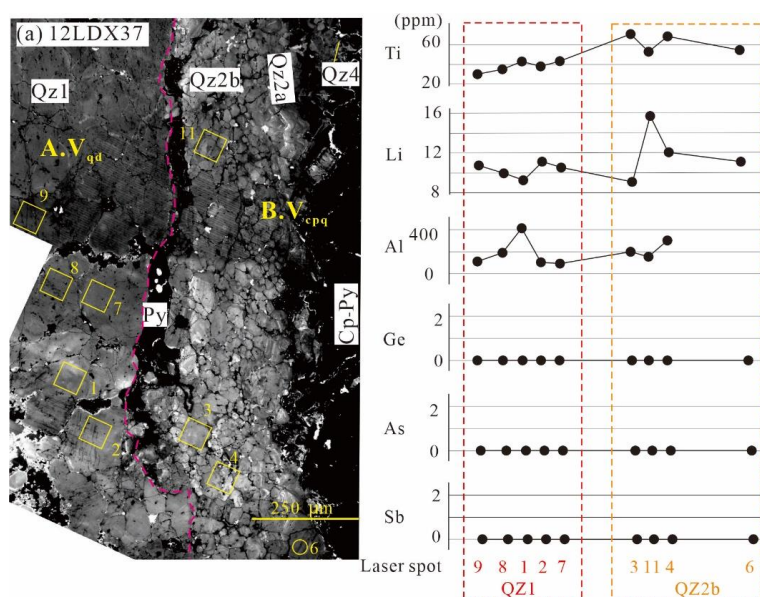
Figure 8. Petrography of fluid inclusions hosted in different generations of quartz in the Tongchang porphyry Cu deposit. The coexistence of brine (B15HS) and vapor-rich (B80) inclusions in Qz2a (a) and Qz2b (b), which distributed along CL growth zones indicative of primary origin; primary B50 inclusions in Qz3b with secondary trails of B80 and B30 inclusions (c); secondary trail of B10 inclusion in Qz2b (d). “B10” denotes the vapor phase and its volume proportion relative to the total inclusion volume; “L” denotes aqueous liquid, “V” denotes vapor, “H” denotes a halite daughter phase, and “S” denotes unidentified solids. CL growth zones are marked with dashed curves in the inclusion images. Scale bars are 10 μm .

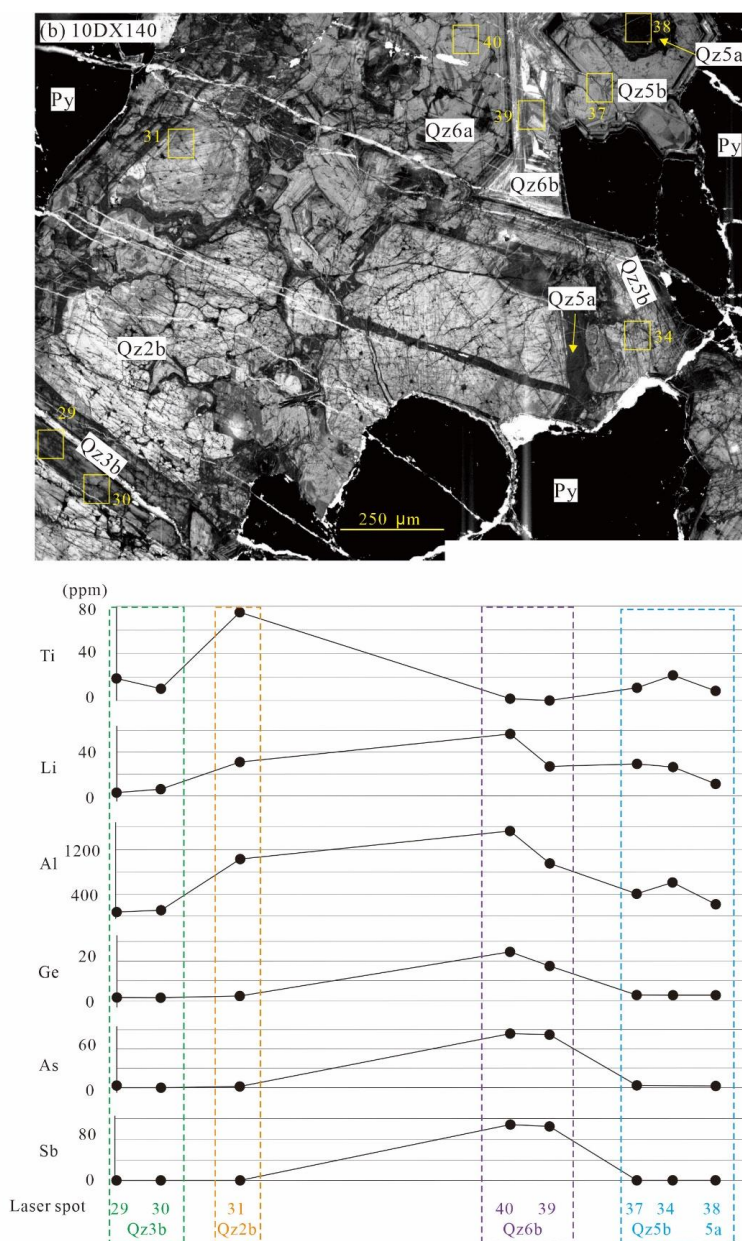
Qz3b contains primary B50 inclusions (Figure 8c). They had ice melting temperatures of -2.7 – -4.2 $^{\circ}\text{C}$ and were homogenized at 300 – 350 $^{\circ}\text{C}$ via critical mode or vapor expansion.

Qz4 commonly contains only B30 inclusions, which are considered primary. They exhibited ice melting at -3.2 – -5.6 $^{\circ}\text{C}$ and were homogenized at 250 – 300 $^{\circ}\text{C}$. They also occurred as secondary inclusions in Qz2 and Qz3 (Figure 8c). B10 occurred mainly as secondary inclusions in Qz2 and Qz3. It had ice melting temperatures at -2.6 – -3.4 $^{\circ}\text{C}$ and was homogenized at 187 – 208 $^{\circ}\text{C}$.

5.3. Trace Elements

Results are shown and depicted in Figure 9, Table 4, and the Appendix A. In all quartz generations, only Ti, Li, Na, Al, K, and Ge occurred in appreciable concentrations. Arsenic and Sb becomes significant in Qz5 and Qz6.





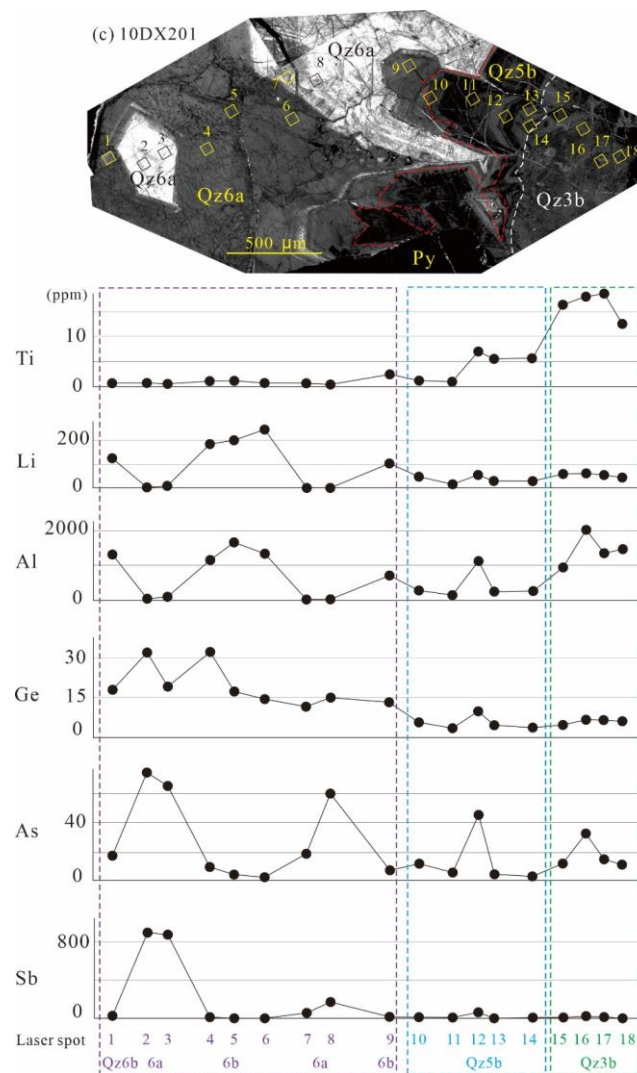


Figure 9. Sketches showing quartz CL textures and ablation spots for trace elements. (a) B.V_{cpq} vein crosscutting the A.V_{qd} vein (12LDX37). Qz1 had intermediate Ti contents (20–30 ppm); Qz2 had higher Ti contents (50–70 ppm). Note that Cu-Fe sulfides coprecipitated with Qz4, which reopened Qz2 grains. Qz2 grains, in turn, crosscut Qz1 grains; (b) In D.V_{pd} pyrite veins (10DX140), Qz3 had lower Ti contents (10–20 ppm) compared to earlier quartz; (c) A profile of trace element analyses encompassing Qz3, Qz5, and Qz6 (10DX201). Qz5 had variable yet lower Ti contents (1–15 ppm). Qz6a and Qz6b had elevated contents of Ge, As and Sb.

Titanium contents show characteristic patterns among quartz generations. It increases from 26.8 ± 10.3 ppm (1σ , $N = 59$) in Qz1 to 127 ± 37 ppm (1σ , $N = 42$) in Qz2a and then decreases to 53.9 ± 16.1 ppm (1σ , $N = 163$) in Qz2b, 10.7 ± 4.0 ppm (1σ , $N = 21$) in Qz3a, 15 ± 7.9 ppm (1σ , $N = 131$) in Qz3b, 6.74 ± 4.22 ppm (1σ , $N = 24$) in Qz4, 8.40 ± 0.88 ppm (1σ , $N = 4$) in Qz5a, 7.66 ± 4.06 ppm (1σ , $N = 66$) in Qz5b, 2.67 ± 2.19 ppm (1σ , $N = 19$) in Qz6a, and 0.699 ± 0.454 ppm in Qz6b.

Lithium increases slightly from 11 ppm in Qz1 to av. 19 ppm in Qz3b, then decreases to 8 ppm in Qz4, and increases to 70 ppm in Qz6a which decreases to 29 ppm in Qz6b. Al increases from 172 ppm in Qz1 to 578 ppm in Qz3b, and then decreases to 300 ppm in Qz4, and then increases to 923 ppm in Qz6a which decreases to 718 ppm in Qz6b. Ge increases from 1 ppm in Qz1 to 14 ppm in Qz6b. As increases from 1 ppm in early quartz to 31 ppm in Qz6b, and Sb increases from 0 ppm (below detection limits) in early quartz to 20 ppm in Qz6a and 207 ppm in Qz6b.

Table 4. LA-ICP-MS trace elements in quartz of the Tongchang porphyry Cu deposit.

Spot No.	CL	Li	Na	Al	K	Ca	Ti	Mn	Fe	Cu	Zn	Ga	Ge	As	Sr	Sn	Sb	Pb	Al Factor	Al Corr	P min (kbar)	P max (kbar)	Tmin (°C)	Tmax (°C)
12LDX37-1	Qz1	9.26	54.6	290	229	<86.01	43.2	1.72	v.d.	1.00	0.78	0.205	1.10	0.490	1.03	0.13	1.59	1.35	0.7	414	0.5	1.5	461	538
12LDX37-2	Qz1	11.2	187	74.3	49.2	<39.88	38.0	v.d.	v.d.	v.d.	v.d.	0.0207	1.28	0.304	0.31	0.119	0.792	v.d.	0.7	106	0.5	1.5	453	528
12LDX37-7	Qz1	10.6	v.d.	65.5	v.d.	v.d.	43.4	v.d.	v.d.	v.d.	v.d.	v.d.	1.36	v.d.	v.d.	0.100	0.773	v.d.	0.7	93.6	0.5	1.5	462	538
12LDX37-8	Qz1	10.0	120	134	104	<88.69	35.0	v.d.	v.d.	v.d.	v.d.	0.0440	^{0.83} ₁	0.350	0.39	^{<0.06} ₂	0.587	2.57	0.7	191	0.5	1.5	448	522
12LDX37-9	Qz1	10.8	689	74.3	v.d.	v.d.	29.7	v.d.	v.d.	v.d.	v.d.	v.d.	v.d.	0.399	v.d.	0.142	0.467	v.d.	0.7	106	0.5	1.5	438	510
12LDX37-14	Qz1	6.59	v.d.	118	56.9	240	30.5	v.d.	v.d.	v.d.	v.d.	0.0270	^{0.87} ₀	<0.43	1.16	0.207	v.d.	v.d.	0.7	168	0.5	1.5	439	512
12LDX37-15	Qz1	7.32	v.d.	v.d.	v.d.	109	35.9	v.d.	v.d.	v.d.	v.d.	0.0310	1.1	0.670	v.d.	0.098	0.561	v.d.	0.7	v.d.	0.5	1.5	450	524
12LDX37-36	Qz1	13.5	18.1	¹⁷⁶ ₂	599	445	36.7	v.d.	v.d.	v.d.	<1.01	0.346	1.41	0.420	2.56	^{0.054} ₀	0.378	0.116	0.7	2518	0.5	1.5	451	525
12LDX37-34	Qz1	7.30	^{<2.7} ₁	135	28.17	234	28.8	0.711	v.d.	1.01	<0.94	0.223	1.31	0.680	0.171	^{<0.07} ₉	^{<0.08} ₀	0.227	0.7	192	0.5	1.5	436	507
12LDX37-40	Qz1	10.3	^{<3.7} ₈	87.1	<6.91	^{<140.4} ₅	43.9	0.144	2.26	0.730	1.65	0.107	^{0.86} ₀	0.580	0.221	^{<0.08} ₈	^{<0.06} ₇	0.09	0.7	124	0.5	1.5	463	539
12LDX37-42	Qz1	10.2	141	70.6	60.8	v.d.	32.3	v.d.	v.d.	v.d.	v.d.	0.216	1.58	0.360	v.d.	0.240	0.199	1.85	0.7	101	0.5	1.5	443	516
12LDX37-43	Qz1	7.84	18.9	58.8	5.13	<88.42	23.2	v.d.	v.d.	v.d.	0.950	0.247	1.13	<0.33	0.587	^{<0.08} ₃	0.317	1.02	0.7	84.0	0.5	1.5	422	492
12LDX37-44	Qz1	10.5	v.d.	61.3	v.d.	457	24.5	v.d.	v.d.	v.d.	v.d.	0.079	1.06	0.270	v.d.	0.162	0.098	v.d.	0.7	87.6	0.5	1.5	426	496
12LDX37-45	Qz1	11.3	^{<2.9} ₀	88.8	<5.24	^{<103.2} ₆	24.9	0.143	4.14	<0.31	1.50	0.143	^{0.84} ₀	<0.28	0.054	^{<0.08} ₁	^{<0.08} ₀	0.096	0.7	127	0.5	1.5	427	497
12LDX37-46	Qz1	10.8	6.73	71.6	8.02	<58.04	32.8	0.819	v.d.	0.418	<0.67	^{<0.013} ₃	^{0.97} ₅	0.239	1.05	0.055	^{<0.04} ₂	0.233	0.7	102	0.5	1.5	444	517
12LDX37-47	Qz1	10.3	8.08	76.1	<4.52	99.3	24.0	0.237	4.18	0.290	<0.53	0.130	1.19	<0.33	0.773	0.190	0.206	0.103	0.7	109	0.5	1.5	424	494
12LDX37-48	Qz1	11.2	v.d.	64.4	69.6	v.d.	23.4	v.d.	v.d.	v.d.	v.d.	0.0620	^{0.85} ₀	<0.35	v.d.	0.179	0.113	v.d.	0.7	91.9	0.5	1.5	423	492
12LDX37-49	Qz1	10.3	18.8	99.5	<5.31	591	36.6	0.105	2.38	<0.34	1.63	0.348	1.50	0.730	1.29	0.095	0.679	0.259	0.7	142	0.5	1.5	451	525
12LDX37-50	Qz1	12.6	v.d.	v.d.	v.d.	v.d.	30.5	v.d.	v.d.	v.d.	2.77	0.134	1.20	1.59	v.d.	0.177	1.68	v.d.	0.7	v.d.	0.5	1.5	439	512
12LDX40-1	Qz1	12.0	83.0	67.6	28.2	614	37.8	v.d.	v.d.	v.d.	v.d.	0.112	1.64	0.640	v.d.	^{0.098} ₀	0.516	v.d.	0.6	113	0.5	1.5	453	528
12LDX40-2	Qz1	12.0	37.7	66.7	6.78	<97.70	15.4	1.01	v.d.	0.610	<1.05	0.285	1.49	1.24	v.d.	0.387	1.202	v.d.	0.6	111	0.5	1.5	398	463
12LDX40-3	Qz1	10.1	39.0	65.0	30.0	202	13.7	v.d.	v.d.	v.d.	v.d.	0.168	0.98	1.01	0.74	0.226	0.847	v.d.	0.6	108	0.5	1.5	391	455
12LDX40-4	Qz1	11.8	29.4	73.5	<3.95	<75.82	16.5	1.241	v.d.	1.26	0.900	0.191	1.14	0.52	0.214	0.255	0.288	1.40	0.6	122	0.5	1.5	402	468

12LDX40-5	Qz1	8.14	47.5	64.4	8.45	<87.48	9.7	v.d.	v.d.	v.d.	v.d.	0.252	1.37	<0.26	0.788	0.109	0.568	0.85	0.6	107	0.5	1.5	372	432
12LDX40-6	Qz1	14.1	21.6	54.5	<2.64	<52.57	13.8	0.998	v.d.	0.403	0.910	0.160	1.15	0.260	1.39	0.160	0.732	0.887	0.6	90.9	0.5	1.5	392	456
12LDX40-7	Qz1	7.17	117	53.3	<3.48	214	15.9	v.d.	v.d.	v.d.	v.d.	0.0450	$\frac{0.89}{1}$	0.330	0.564	0.156	0.293	3.00	0.6	88.8	0.5	1.5	400	465
12LDX40-8	Qz1	7.94	108	v.d.	v.d.	67.2	20.0	v.d.	v.d.	v.d.	v.d.	0.369	$\frac{0.84}{7}$	0.340	0.473	0.033	$\frac{<0.02}{8}$	v.d.	0.6	v.d.	0.5	1.5	413	481
12LDX40-9	Qz1	13.8	9.40	61.9	<2.82	88.7	21.4	0.0530	v.d.	0.340	<0.52	0.298	0.96	1.71	$\frac{0.084}{0}$	0.142	0.683	0.464	0.6	103	0.5	1.5	417	486
12LDX40-10	Qz1	12.9	31.8	53.7	25.3	<64.58	21.4	v.d.	v.d.	v.d.	v.d.	0.415	1.66	0.28	v.d.	0.559	0.839	v.d.	0.6	89.5	0.5	1.5	418	486
12LDX40-11	Qz1	12.1	56.4	73.8	31.5	$\frac{<135.2}{6}$	13.2	v.d.	v.d.	v.d.	v.d.	0.061	1.00	<0.43	0.137	0.285	0.614	v.d.	0.6	123	0.5	1.5	389	453
12LDX40-12	Qz1	13.3	64.0	55.1	32.4	<58.82	17.6	v.d.	v.d.	v.d.	v.d.	0.321	$\frac{0.96}{2}$	1.12	0.036	0.103	0.681	v.d.	0.6	91.9	0.5	1.5	406	472
12LDX40-13	Qz1	12.6	34.8	56.9	22.1	<80.39	15.4	0.838	4.08	1.23	v.d.	0.245	1.49	0.520	$\frac{0.093}{7}$	0.783	0.26	v.d.	0.6	94.8	0.5	1.5	398	463
12LDX40-14	Qz1	14.6	29.1	73.6	26.8	138	10.6	1.55	v.d.	<0.23	8.82	0.226	1.31	0.530	0.746	$\frac{<0.06}{2}$	0.296	v.d.	0.6	123	0.5	1.5	377	438
12LDX40-15	Qz1	12.8	v.d.	90.1	56.1	<55.79	18.7	v.d.	v.d.	v.d.	v.d.	0.135	1.25	1.78	0.601	0.514	0.79	v.d.	0.6	150	0.5	1.5	409	476
12LDX40-16	Qz1	15.6	18.4	94.5	8.78	$\frac{<128.3}{2}$	27.0	0.152	3.88	1.56	<1.10	1.39	1.11	2.51	0.459	0.231	0.467	v.d.	0.6	157	0.5	1.5	432	503
12LDX40-17	Qz1	16.3	10.5	71.5	9.96	<98.31	28.6	0.312	5.13	1.83	2.75	0.500	$\frac{0.94}{0}$	2.38	v.d.	0.446	1.73	v.d.	0.6	119	0.5	1.5	435	507
12LDX40-18	Qz1	12.3	104	66.6	36.6	201	24.5	v.d.	v.d.	v.d.	v.d.	0.238	1.46	0.780	v.d.	1.06	1.47	v.d.	0.6	111	0.5	1.5	426	496
12LDX40-19	Qz1	17.4	50.4	78.9	22.9	162	42.6	v.d.	v.d.	v.d.	v.d.	0.098	1.07	1.59	1.44	0.537	1.51	v.d.	0.6	131	0.5	1.5	461	537
12LDX40-20	Qz1	13.8	$\frac{<2.3}{0}$	61.7	5.52	v.d.	15.4	<0.037	2.57	0.48	<0.42	0.437	0.91	0.510	1.12	0.147	0.146	1.7	0.6	103	0.5	1.5	398	463
12LDX40-21	Qz1	14.9	18.6	64.6	11.0	<65.02	13.1	1.26	v.d.	1.07	1.88	0.160	1.02	1.7	v.d.	0.257	0.233	1.49	0.6	108	0.5	1.5	389	452
12LDX40-22	Qz1	11.2	59.2	52.8	19.3	<92.12	16.0	v.d.	v.d.	v.d.	v.d.	0.510	1.00	0.720	0.645	0.191	0.130	v.d.	0.6	88	0.5	1.5	400	466
12LDX40-23	Qz1	11.8	18.5	77.5	<4.66	<79.06	14.2	0.138	3.77	0.270	<0.48	0.253	1.38	0.580	0.163	0.095	0.283	0.307	0.6	129	0.5	1.5	393	457
12LDX40-24	Qz1	13.6	40.5	67.6	13.1	<74.70	24.7	v.d.	v.d.	v.d.	v.d.	0.221	1.11	0.560	0.058	0.118	0.15	v.d.	0.6	113	0.5	1.5	426	496
12LDX40-25	Qz1	8.41	56.5	v.d.	137	105	16.7	v.d.	v.d.	v.d.	v.d.	0.203	1.30	0.670	v.d.	0.093	$\frac{<0.06}{3}$	0.734	0.6	v.d.	0.5	1.5	403	469
10DX328-13	Qz1	7.24	41.5	57.2	4.37	151	33.1	v.d.	v.d.	v.d.	<0.77	0.424	1.40	0.300	v.d.	0.128	0.151	0.911	0.76	75.3	0.5	1.5	444	517
10DX328-14	Qz1	8.40	97.7	58.3	<4.11	v.d.	37.8	v.d.	v.d.	v.d.	<0.86	0.102	1.98	0.880	v.d.	0.151	0.116	0.603	0.76	76.6	0.5	1.5	453	528
10DX328-16	Qz1	13.4	117	73.0	13.6	<79.63	31.4	v.d.	v.d.	2.93	1.31	0.099	1.78	0.360	v.d.	0.144	0.058	v.d.	0.76	96.1	0.5	1.5	441	514
10DX328-35	Qz1	13.2	118	71.1	32.0	<68.97	28.4	v.d.	v.d.	v.d.	3.92	0.098	1.33	0.175	0.194	0.096	0.346	v.d.	0.76	93.5	0.5	1.5	435	506
10DX328-36	Qz1	11.5	133	60.5	18.4	v.d.	32.7	v.d.	v.d.	v.d.	3.00	0.524	2.15	1.35	v.d.	0.451	0.304	v.d.	0.76	79.5	0.5	1.5	444	517

10DX328-37	Qz1	7.15	25.0	83.1	8.2	<74.08	31.1	v.d.	v.d.	v.d.	<0.77	0.229	1.24	0.310	v.d.	0.101	0.263	0.651	0.76	109	0.5	1.5	441	513
10DX261-12	Qz1	12.5	11.0	124	29.2	<51.06	43.0	0.313	v.d.	v.d.	<0.39	0.116	$\frac{0.90}{9}$	0.367	v.d.	1.81	0.257	0.097	0.62	163	0.5	1.5	461	537
10DX261-13	Qz1	16.7	5.42	v.d.	62.5	<67.40	57.6	0.336	v.d.	v.d.	v.d.	0.102	1.45	<0.22	2.23	0.116	0.210	v.d.	0.62	v.d.	0.5	1.5	480	560
10DX261-18	Qz1	15.3	167	v.d.	v.d.	<43.16	32.0	v.d.	v.d.	v.d.	v.d.	0.456	1.40	0.310	0.500	0.077	0.159	v.d.	0.62	v.d.	0.5	1.5	442	515
10DX261-19	Qz1	16.7	35.8	498	292	<55.98	36.7	v.d.	v.d.	0.492	v.d.	0.936	$\frac{0.90}{9}$	0.770	0.987	0.135	0.404	0.612	0.62	656	0.5	1.5	451	525
10DX261-20	Qz1	12.6	5.48	67.4	<3.49	<69.43	15.9	0.705	6.44	v.d.	v.d.	0.419	$\frac{0.73}{1}$	0.740	0.084	$\frac{0.04}{9}$	0.166	v.d.	0.62	88.6	0.5	1.5	400	465
10DX261-26	Qz1	12.7	28.3	85.7	15.7	229	19.1	0.300	6.27	0.620	v.d.	0.795	1.69	v.d.	v.d.	v.d.	v.d.	v.d.	0.62	113	0.5	1.5	411	478
10DX261-29	Qz1	11.5	96.5	84.2	25.8	<74.43	34.3	v.d.	v.d.	v.d.	v.d.	v.d.	1.12	0.166	10.4	$\frac{0.070}{0}$	0.292	v.d.	0.62	111	0.5	1.5	447	520
10DX261-28	Qz1	10.2	51.2	104	36.8	<89.38	22.3	v.d.	v.d.	v.d.	v.d.	0.793	1.40	0.430	0.844	0.188	0.914	0.731	0.62	137	0.5	1.5	420	489
10DX261-34	Qz1	12.7	$\frac{2.5}{8}$	120	23.4	89.0	21.7	0.226	v.d.	v.d.	1.22	0.348	1.29	0.390	0.180	$\frac{0.087}{0}$	0.189	0.36	0.62	158	0.5	1.5	418	487
12DXF05-5	Qz2 _a	13.9	16.5	75.7	<6.89	204	152	1.43	7.93	0.730	<1.68	0.155	1.09	0.390	0.832	$\frac{0.09}{8}$	$\frac{0.09}{0}$	0.296	0.52	146	0.5	1.5	543	635
12DXF05-6	Qz2 _a	14.8	25.5	59.2	<3.92	<75.23	56.9	1.32	7.42	0.202	0.550	0.171	1.10	$\frac{0.18}{8}$	1.41	$\frac{0.057}{0}$	$\frac{0.05}{7}$	0.491	0.52	114	0.5	1.5	479	559
12DXF05-7	Qz2 _a	15.9	v.d.	207	108	215	193	v.d.	v.d.	v.d.	v.d.	0.179	$\frac{0.64}{4}$	0.920	1.46	$\frac{0.07}{2}$	0.304	v.d.	0.52	399	0.5	1.5	558	653
12DXF05-8	Qz2 _a	15.4	v.d.	50.3	29.4	<67.91	165	v.d.	v.d.	v.d.	v.d.	0.045	0.81	0.240	v.d.	$\frac{0.05}{4}$	0.09	v.d.	0.52	96.8	0.5	1.5	548	641
12DXF05-9	Qz2 _a	13.9	$\frac{3.8}{3}$	56.7	<6.52	$\frac{127.1}{5}$	167	<0.070	2.39	<0.25	<0.99	0.133	0.81	<0.35	0.136	0.219	$\frac{0.07}{5}$	0.979	0.52	109	0.5	1.5	549	643
12DXF05-10	Qz2 _a	12.8	v.d.	53.5	66.0	101	90.6	v.d.	v.d.	v.d.	v.d.	0.114	$\frac{0.75}{7}$	0.233	v.d.	$\frac{0.04}{6}$	$\frac{0.03}{6}$	v.d.	0.52	103	0.5	1.5	510	595
12DXF05-11	Qz2 _a	13.1	83.8	57.9	7.30	<72.94	174	v.d.	v.d.	v.d.	v.d.	0.195	1.04	$\frac{0.17}{1}$	1.48	$\frac{0.05}{8}$	$\frac{0.05}{2}$	0.939	0.52	111	0.5	1.5	552	645
12DXF05-12	Qz2 _a	23.3	71.7	225	110	105.9	97.6	v.d.	v.d.	v.d.	v.d.	0.114	$\frac{0.87}{1}$	<0.20	0.349	0.228	0.042	1.29	0.52	433	0.5	1.5	515	601
12DXF05-13	Qz2 _a	27.2	9.92	313	43.5	<70.04	164	0.657	v.d.	$\frac{0.13}{0}$	1.25	0.0620	$\frac{0.91}{1}$	$\frac{0.14}{7}$	0.369	0.087	$\frac{0.04}{5}$	0.161	0.52	602	0.5	1.5	548	641
12DXF05-14	Qz2 _a	17.4	v.d.	v.d.	133	<84.66	130	v.d.	v.d.	v.d.	v.d.	0.0520	$\frac{0.65}{1}$	0.306	v.d.	$\frac{0.06}{9}$	0.057	v.d.	0.52	v.d.	0.5	1.5	533	623
12DXF05-15	Qz2 _a	27.5	49.8	218	62.5	<50.53	157	v.d.	v.d.	v.d.	v.d.	0.0490	$\frac{0.71}{9}$	$\frac{0.09}{3}$	0.233	0.072	$\frac{0.02}{4}$	0.753	0.52	420	0.5	1.5	545	638

12DXF05-16	Qz2 a	14.0	124	56.8	46.8	<61.76	110	v.d.	v.d.	v.d.	v.d.	0.0810	1.08	0.276	0.154	0.095	<0.03 7	v.d.	0.52	109	0.5	1.5	522	610
12DXF05-19	Qz2 a	14.1	v.d.	v.d.	73.5	<59.93	142	v.d.	v.d.	v.d.	v.d.	0.0440	1.01	0.406	v.d.	0.210	0.091	v.d.	0.52	v.d.	0.5	1.5	539	630
12DXF05-26	Qz2 a	22.3	55.5	224	96.3	95.32	180	v.d.	v.d.	v.d.	v.d.	0.0850	1.03	<0.19 4	v.d.	0.048 0	<0.04 0	0.809	0.52	430	0.5	1.5	554	648
12DXF05-27	Qz2 a	11.2	v.d.	54.0	28.3	<57.03	109	v.d.	v.d.	v.d.	v.d.	0.0188	0.89 4	1.06	0.513	<0.04 8	0.085	v.d.	0.52	104	0.5	1.5	522	610
12DXF05-28	Qz2 a	13.5	80.2	67.0	13.2	<67.38	159	v.d.	v.d.	v.d.	v.d.	<0.010 4	0.80 6	<0.20 7	0.297	0.065 0	0.084	v.d.	0.52	129	0.5	1.5	546	639
12DXF05-29	Qz2 a	11.4	45.7	70.0	14.2	<79.92	152	v.d.	v.d.	v.d.	v.d.	0.100	1.21 4	0.270	0.926	0.078 0	<0.04 6	v.d.	0.52	135	0.5	1.5	543	635
10DX140-50	Qz2 a	11.7	30.2	106	16.3	114	101	0.441	v.d.	0.990	0.560	0.0810	1.32	0.360	1.14	0.139	0.625	0.472	0.56	189	0.5	1.5	517	604
12DX500-1	Qz2 a	25.8	10.8	127	18.5	83.0	101	0.568	v.d.	0.305	<0.33	0.167	0.72 0	0.450	0.169	0.058 0	0.046	0.262	0.56	226	0.5	1.5	517	604
12DX500-2	Qz2 a	31.2	13.5	126	14.7	v.d.	117	0.433	v.d.	<0.12 9	<0.37	0.0117	1.18	<0.14 1	v.d.	0.053 0	0.187	0.654	0.56	226	0.5	1.5	526	615
12DX500-3	Qz2 a	21.5	46.6	108	29.0	329	101	0.800	v.d.	0.252	<0.30	0.114	0.80 6	0.194	1.10	0.119	0.152	0.465	0.56	193	0.5	1.5	517	604
12DX500-4	Qz2 a	27.0	18.0	106	13.5	<89.59	201	0.410	v.d.	0.400	<1.28	0.0530	0.97 0	0.810	4.10	0.118	0.179	0.745	0.56	190	0.5	1.5	560	656
12DX500-5	Qz2 a	26.0	13.6	97.2	<4.73	270	121	0.178	3.99	<0.18 6	<0.00	<0.016 2	0.84 0	<0.23	0.569	0.141	<0.05 3	0.119	0.56	174	0.5	1.5	529	618
12DX500-6	Qz2 a	29.7	37.0	146	42.6	v.d.	123	v.d.	v.d.	1.45	2.33	0.236	0.79 5	<0.24	v.d.	0.191	0.119	0.81	0.56	260	0.5	1.5	530	619
12DX500-7	Qz2 a	30.9	49.6	v.d.	v.d.	v.d.	227	v.d.	v.d.	0.448	1.64	0.206	0.92 5	0.690	v.d.	0.042	0.079	0.415	0.56	v.d.	0.5	1.5	567	664
12DX500-8	Qz2 a	28.0	13.2	86.7	10.7	47.0	147	0.251	v.d.	0.392	0.370	0.121	0.96 4	<0.13 1	0.165	0.079	0.089	0.096	0.56	155	0.5	1.5	541	633
10DX172-12	Qz2 a	9.73	v.d.	v.d.	v.d.	v.d.	112	10.4	v.d.	31.6	3.52	0.454	0.90 8	1.96	v.d.	0.187	0.383	0.744	0.51	v.d.	0.5	1.5	524	612
10DX172-20	Qz2 a	14.3	75.3	74.5	19.9	120	108	0.389	2.55	0.82	1	0.0830	0.69 4	0.280	0.241	0.068	0.093	0.243	0.51	146	0.5	1.5	521	609
10DX172-21	Qz2 a	15.6	19.5	64.4	<4.82	<93.21	91.7	0.297	4.09	0.38	<1.08	0.0740	0.76 0	0.420	v.d.	<0.06 7	<0.06 4	0.113	0.51	126	0.5	1.5	511	596

10DX172-22	Qz2 _a	15.7	50.2	78.8	10.8	<67.38	118	0.127	2.24	0.215	<1.06	0.0241	$\frac{0.79}{0}$	0.580	0.483	0.046	$\frac{<0.04}{5}$	<0.026	0.51	154	0.5	1.5	527	616
10DX328-1	Qz2 _a	17.7	v.d.	116	v.d.	v.d.	85.3	3.58	v.d.	1.30	<0.88	0.036	$\frac{0.54}{0}$	<0.26	v.d.	$\frac{<0.04}{5}$	0.076	0.946	0.76	153	0.5	1.5	506	591
10DX328-2	Qz2 _a	17.2	6.92	121	v.d.	<88.55	100	0.423	7.76	0.800	<1.18	$\frac{<0.020}{9}$	1.21	0.320	0.569	$\frac{<0.06}{3}$	$\frac{<0.06}{4}$	0.049	0.76	159	0.5	1.5	517	603
10DX328-3	Qz2 _a	21.5	v.d.	159	v.d.	79.83	87.6	0.133	4.90	$\frac{<0.23}{8}$	<0.64	0.0519	$\frac{0.85}{9}$	<0.22	0.211	0.074	0.143	0.073	0.76	209	0.5	1.5	508	593
10DX328-5	Qz2 _a	13.5	62.4	87.7	8.34	<82.29	68.2	v.d.	v.d.	v.d.	0.99	0.305	1.39	0.340	0.139	0.187	0.078	0.839	0.76	115	0.5	1.5	491	573
10DX328-6	Qz2 _a	17.1	v.d.	90.2	17.0	78.9	96.8	v.d.	v.d.	v.d.	1.66	$\frac{<0.008}{4}$	1.59	<0.22	$\frac{0.099}{0}$	0.088	$\frac{<0.03}{0}$	v.d.	0.76	119	0.5	1.5	514	601
10DX328-7	Qz2 _a	15.2	70.4	97.3	16.8	<49.84	90.1	v.d.	v.d.	v.d.	1.56	<0.016	1.19	0.270	0.193	0.112	$\frac{0.046}{0}$	v.d.	0.76	128	0.5	1.5	510	595
10DX328-10	Qz2 _a	15.3	v.d.	133	43.1	85.8	79.0	v.d.	v.d.	v.d.	4.63	0.0950	1.23	0.730	0.551	0.155	$\frac{0.093}{0}$	v.d.	0.76	176	0.5	1.5	501	585
10DX328-11	Qz2 _a	16.5	9.55	v.d.	v.d.	v.d.	96.8	v.d.	v.d.	v.d.	v.d.	0.0430	1.1	0.580	v.d.	0.09	$\frac{<0.08}{5}$	v.d.	0.76	v.d.	0.5	1.5	514	601
10DX22-4	Qz2 _a	15.3	10.8	78.9	<3.48	<65.83	130	0.314	4.10	0.927	0.390	0.107	$\frac{0.66}{1}$	$\frac{<0.16}{0}$	v.d.	$\frac{<0.06}{5}$	0.108	0.105	0.55	143	0.5	1.5	533	624
10DX22-5	Qz2 _a	15.9	35.2	v.d.	185	210	160	1.31	v.d.	0.214	1.06	0.125	$\frac{0.85}{4}$	0.306	0.582	0.186	0.161	0.253	0.55	v.d.	0.5	1.5	546	639
10DX22-7	Qz2 _a	15.8	27.8	v.d.	273	75.6	143	0.947	v.d.	0.291	<0.46	0.133	$\frac{0.93}{3}$	$\frac{<0.18}{9}$	0.565	0.108	0.177	0.17	0.55	v.d.	0.5	1.5	539	630
10DX22-20	Qz2 _a	17.6	3.34	76.4	4.1	<49.37	129	0.156	4.67	1.47	<0.58	0.160	1.06	0.224	0.982	0.093	$\frac{0.048}{0}$	0.184	0.55	139	0.5	1.5	533	623
10DX168-2	Qz2 _b	13.1	$\frac{<3.1}{7}$	55.6	<5.46	100	57.2	<0.064	1.41	$\frac{<0.19}{2}$	<0.83	0.229	0.45	0.710	$\frac{0.011}{8}$	0.209	$\frac{<0.05}{4}$	0.142	0.58	95.9	0.5	1.5	480	559
10DX168-4	Qz2 _b	9.34	10.8	50.5	<7.86	$\frac{<148.2}{9}$	47.9	0.509	3.93	0.900	<1.14	0.028	0.76	<0.48	0.755	$\frac{<0.09}{9}$	$\frac{<0.10}{5}$	0.069	0.58	87.1	0.5	1.5	468	546
10DX168-5	Qz2 _b	12.0	$\frac{<3.6}{4}$	47.9	<6.26	$\frac{<113.2}{0}$	62.6	<0.069	1.3	0.720	1.17	0.310	1.25	<0.34	$\frac{0.007}{3}$	$\frac{<0.08}{1}$	$\frac{<0.08}{7}$	0.082	0.58	82.6	0.5	1.5	486	566
10DX168-7	Qz2 _b	7.93	34.7	62.4	<5.49	$\frac{<102.7}{8}$	53.6	v.d.	v.d.	v.d.	<0.75	0.0267	1.01	$\frac{<0.15}{7}$	$\frac{0.029}{5}$	0.072	0.239	0.385	0.58	108	0.5	1.5	476	554
10DX168-8	Qz2 _b	11.7	18.0	56.2	3.87	<52.12	62.2	0.279	2.35	0.446	<0.37	0.0310	0.97	0.165	$\frac{0.026}{9}$	0.134	0.152	0.234	0.58	96.9	0.5	1.5	485	566

10DX168-10	Qz2 b	9.88	26.9	56.5	<3.90	150	64.7	0.494	1.47	0.920	0.44	0.0310	^{0.93} ₇	0.470	^{0.038} ₉	0.136	0.427	0.179	0.58	97.3	0.5	1.5	488	569
10DX168-14	Qz2 b	10.9	47.4	61.2	7.44	<93.37	61.4	0.461	2.81	0.640	<0.84	0.0370	0.9	^{<0.20} ₄	0.298	0.110	0.140	0.534	0.58	106	0.5	1.5	484	565
10DX168-16	Qz2 b	7.67	48.0	45.0	13.1	<74.34	39.7	0.125	v.d.	v.d.	<0.37	0.0590	1.22	0.362	0.227	0.058	^{0.091} ₀	0.156	0.58	77.6	0.5	1.5	456	531
10DX168-26	Qz2 b	8.81	49.7	46.5	14.0	357	16.2	v.d.	v.d.	v.d.	<0.43	0.0270	1.54	0.540	v.d.	0.063	0.215	0.343	0.58	80.2	0.5	1.5	401	467
10DX168-29	Qz2 b	6.70	62.8	58.6	5.43	v.d.	49.2	1.32	v.d.	v.d.	<0.37	0.0940	1.14	0.950	v.d.	0.047	0.257	0.253	0.58	101	0.5	1.5	470	548
10DX168-32	Qz2 b	10.8	68.9	87.0	20.3	82.2	61.0	0.146	5.20	v.d.	<0.00	0.0211	^{0.83} ₁	v.d.	v.d.	0.032	0.130	0.089	0.58	150	0.5	1.5	484	564
10DX168-35	Qz2 b	8.42	57.7	v.d.	70.3	^{<104.0} ₈	54.7	v.d.	v.d.	v.d.	v.d.	0.0550	^{0.88} ₀	1.14	0.143	0.234	0.788	v.d.	0.58	v.d.	0.5	1.5	477	556
10DX168-36	Qz2 b	18.5	19.7	81.7	9.18	106	65.3	0.0980	2.35	v.d.	<0.48	0.125	1.09	0.640	^{0.096} ₀	^{<0.07} ₂	0.224	0.751	0.58	141	0.5	1.5	488	570
10DX168-37	Qz2 b	15.1	57.2	95.5	9.5	<62.55	72.9	0.120	4.80	v.d.	<0.59	0.0910	^{0.86} ₈	0.231	0.122	0.18	0.305	0.182	0.58	165	0.5	1.5	496	578
10DX168-38	Qz2 b	25.3	^{<3.1} ₃	77.1	<5.46	124	68.7	<0.051	1.60	<0.21	<0.00	0.843	1.41	0.590	^{0.024} ₉	^{<0.06} ₈	0.391	0.309	0.58	133	0.5	1.5	492	574
10DX168-39	Qz2 b	13.3	25.1	53.7	<3.86	<68.75	31.8	<0.036	1.415	v.d.	<0.52	0.222	1.38	1.18	^{0.090} ₀	0.066	0.176	0.252	0.58	92.5	0.5	1.5	442	515
10DX168-40	Qz2 b	7.35	30.2	45.7	5.56	<77.67	45.5	<0.047	0.870	v.d.	<0.89	0.0880	^{0.71} ₉	<0.30	^{0.069} ₈	^{<0.06} ₁	0.317	0.066	0.58	78.8	0.5	1.5	465	542
DX168GY-01	Qz2 b	8.96	^{0.34} ₉	69.7	0.00	20.4	36.2	0.291	v.d.	0.000	0.261	0.010	^{0.94} ₉	0.267	0.002	0.228	0.016	0.005	v.d.	v.d.	0.5	1.5	450	524
DX168GY-02	Qz2 b	8.59	v.d.	194	398	82.3	59.7	v.d.	v.d.	v.d.	14.5	0.127	^{0.78} ₇	0.636	v.d.	0.454	0.064	v.d.	v.d.	v.d.	0.5	1.5	483	563
DX168GY-03	Qz2 b	11.9	1.03	101	0.00	29.2	51.5	0.411	0.864	0.099	0.000	0.003	^{0.70} ₁	0.179	0.002	0.320	0.001	0.000	v.d.	v.d.	0.5	1.5	473	551
DX168GY-05	Qz2 b	19.9	19.3	242	20.4	0.00	50.8	1.25	7.195	0.315	0.263	0.000	^{0.83} ₅	0.393	0.111	0.257	0.038	0.062	v.d.	v.d.	0.5	1.5	472	550
DX168GY-06	Qz2 b	11.9	23.6	126	7.05	0.00	50.5	0.000	1.449	0.030	0.000	0.018	^{0.85} ₂	0.780	0.484	0.282	0.172	0.000	v.d.	v.d.	0.5	1.5	472	550
DX168GY-07	Qz2 b	15.1	64.7	164	18.7	75.1	50.1	0.355	4.855	1.523	0.111	0.015	^{0.68} ₉	0.721	0.078	0.207	0.134	0.050	v.d.	v.d.	0.5	1.5	471	549

DX168GY-09	Qz2 b	14.6	17.9	147	6.99	v.d.	51.5	0.729	v.d.	0.179	0.055	0.000	0.53 1	0.104	v.d.	0.248	0.041	0.045	v.d.	v.d.	0.5	1.5	473	551
DX168GY-10	Qz2 b	15.9	3.47	152	3.36	69.9	59.7	0.417	1.602	0.054	0.145	0.000	0.63 7	0.278	0.063	0.168	0.012	0.003	v.d.	v.d.	0.5	1.5	483	563
DX168GY-13	Qz2 b	14.5	34.2	153	12.7	34.4	37.1	0.000	0.245	0.081	0.546	0.004	0.82 8	0.238	0.051	0.120	0.029	0.014	v.d.	v.d.	0.5	1.5	452	526
DX168GY-14	Qz2 b	12.4	7.32	110	3.67	70.6	41.4	0.000	1.981	0.000	0.000	0.009	0.79 6	0.221	0.235	0.278	0.029	0.003	v.d.	v.d.	0.5	1.5	459	534
DX168GY-18	Qz2 b	12.9	1.66	110	1.33	45.4	41.4	0.000	2.381	0.305	0.005	0.014	0.65 9	0.257	0.011	0.168	0.008	0.000	v.d.	v.d.	0.5	1.5	459	535
DX168GY-19	Qz2 b	18.3	2.64	189	10.4	32.9	59.6	0.581	2.914	0.000	0.288	0.007	0.74 9	0.128	0.009	0.146	0.018	0.001	v.d.	v.d.	0.5	1.5	482	563
DX168GY-20	Qz2 b	15.5	7.09	159	8.50	26.3	59.1	0.000	1.435	0.000	0.101	0.008	0.70 1	0.140	0.034	0.293	0.014	0.006	v.d.	v.d.	0.5	1.5	482	562
DX168GY-21	Qz2 b	11.6	8.61	98.1	2.59	5.86	48.5	0.199	0.845	0.118	0.155	0.003	0.87 1	0.197	0.060	0.125	0.010	0.007	v.d.	v.d.	0.5	1.5	469	547
DX168GY-22	Qz2 b	8.40	31.8	76.1	15.1	56.8	42.0	0.967	3.474	0.399	0.553	0.014	1.13 7	0.155	0.021	0.197	0.007	0.111	v.d.	v.d.	0.5	1.5	460	536
DX168GY-25	Qz2 b	13.4	2.07	108	1.87	0.000	55.7	0.380	0.835	0.000	0.000	0.006	0.72 4	0.253	0.005	0.180	0.000	0.001	v.d.	v.d.	0.5	1.5	478	557
DX168GY-30	Qz2 b	12.7	8.04	108	2.58	4.58	53.9	0.060	1.552	0.474	0.004	0.012	0.85 8	0.242	0.121	0.185	0.007	0.002	v.d.	v.d.	0.5	1.5	476	555
DX168GY-31	Qz2 b	14.4	26.3	132	5.55	71.6	53.0	0.872	0.000	0.229	0.000	0.000	0.58 5	0.522	0.202	0.125	0.163	0.014	v.d.	v.d.	0.5	1.5	475	553
DX168GY-33	Qz2 b	4.30	154	81.2	18.1	28.9	38.4	0.000	4.785	2.778	0.816	0.003	0.85 0	0.328	0.657	0.003	0.103	0.266	v.d.	v.d.	0.5	1.5	454	529
DX168GY-38	Qz2 b	14.9	4.05	135	2.80	41.2	56.5	0.000	1.672	0.165	0.000	0.008	0.68 0	0.321	0.094	0.180	0.004	0.002	v.d.	v.d.	0.5	1.5	479	558
DX168GY-Q39	Qz2 b	14.8	0.99 2	123	3.91	24.7	49.9	0.000	1.374	0.000	0.229	0.002	0.80 5	0.200	0.026	0.199	0.005	0.004	v.d.	v.d.	0.5	1.5	471	549
DX168GY-43	Qz2 b	16.1	0.12	130	1.43	69.0	57.1	0.000	2.982	0.000	0.000	0.010	0.75 7	0.192	0.005	0.159	0.088	0.002	v.d.	v.d.	0.5	1.5	480	559
DX168GY-44	Qz2 b	12.9	v.d.	103	0.35	41.1	48.3	0.000	0.334	0.015	0.200	0.002	0.60 0	0.242	0.002	0.170	0.040	0.003	v.d.	v.d.	0.5	1.5	469	546
DX168GY-45	Qz2 b	16.6	2.54	142	1.55	53.2	54.8	0.187	0.979	0.100	0.000	0.009	0.74 7	0.159	0.011	0.276	0.100	0.000	v.d.	v.d.	0.5	1.5	477	556

DX168GY-47	Qz2 b	7.89	10.3	72	3.04	12.8	34.5	0.000	0.636	1.968	0.000	0.011	0.76 4	0.254	0.122	0.191	0.050	0.000	v.d.	v.d.	0.5	1.5	447	521
DX168GY-48	Qz2 b	11.4	31.6	111	6.64	43.5	51.5	0.000	2.735	11.94	0.000	0.009	0.91 0	0.310	0.043	0.091	0.057	0.024	v.d.	v.d.	0.5	1.5	473	551
DX168GY-51	Qz2 b	16.9	121	176	16.2	62.6	53.6	0.000	1.997	0.607	0.435	0.003	0.61 5	0.516	0.066	0.144	0.097	0.054	v.d.	v.d.	0.5	1.5	475	554
12DX500-9	Qz2 b	24.3	<2.4 5	82.9	<4.24	v.d.	82.0	0.206	2.840	0.327	<0.59	0.160	0.87 7	0.380	v.d.	0.190	<0.06 7	0.118	0.56	148	0.5	1.5	503	588
12DX500-10	Qz2 b	24.7	5.62	87.0	<3.83	v.d.	86.7	0.166	2.810	<0.14 9	1.800	0.076	1.18 0	<0.19	0.063	0.113	<0.04 4	0.264	0.56	155	0.5	1.5	507	592
12DX500-11	Qz2 b	33.5	34.9	141	101	v.d.	62.1	3.06	v.d.	1.31	0.700	0.046	0.68 0	<0.14	0.436	0.209	0.374	1.015	0.56	251	0.5	1.5	485	566
12DX500-12	Qz2 b	22.1	v.d.	83.2	166	v.d.	56.9	v.d.	v.d.	v.d.	v.d.	0.126	1.30	1.00	1.302	0.382	0.371	v.d.	0.56	149	0.5	1.5	479	559
10DX172-3	Qz2 b	14.2	6.37	74.0	<2.83	<54.07	83.1	0.256	3.07	0.138	0.760	0.170	0.82 4	<0.16 2	0.152	0.050 0	0.062	0.147	0.51	145	0.5	1.5	504	589
10DX172-4	Qz2 b	17.0	9.11	74.4	4.14	<56.49	65.4	0.163	3.21	1.11	0.280	0.0480	0.64 2	0.500	0.104	0.120	<0.03 6	0.236	0.51	146	0.5	1.5	489	570
10DX172-5	Qz2 b	13.9	29.7	62.1	<2.63	<46.96	50.7	0.0930	1.70	0.264	0.51	0.0407	0.85 3	0.316	0.355	0.088 0	<0.03 3	0.032	0.51	122	0.5	1.5	472	550
10DX172-6	Qz2 b	12.6	v.d.	61.0	36.1	223	47.1	v.d.	v.d.	1.136	3.72	0.121	0.63 7	0.410	0.715	0.089 0	0.062	1.14	0.51	120	0.5	1.5	467	544
10DX172-7	Qz2 b	13.0	52.3	65.6	12.8	93.6	53.0	v.d.	v.d.	v.d.	<0.40	0.157	0.72 1	0.580	0.234	<0.03 9	<0.03 4	0.39	0.51	129	0.5	1.5	475	553
10DX172-8	Qz2 b	11.1	78.0	86.6	9.75	v.d.	61.8	v.d.	v.d.	v.d.	v.d.	0.150	0.71 0	1.52	v.d.	0.095 0	0.104	1.33	0.51	170	0.5	1.5	485	565
10DX172-9	Qz2 b	11.6	18.4	79.1	<6.55	<115.8 8	48.4	0.0760	5.58	<0.26	0.850	<0.034	1.01	0.710	0.314	<0.08 3	<0.06 7	<0.042	0.51	155	0.5	1.5	469	546
10DX172-10	Qz2 b	15.4	148	64.0	70.5	<79.24	37.1	v.d.	v.d.	v.d.	v.d.	0.210	0.78 1	0.620	0.604	<0.06 3	0.115	v.d.	0.51	125	0.5	1.5	452	526
10DX172-13	Qz2 b	15.6	15.3	75.5	25.2	<60.79	56.6	0.129	2.66	0.185	<0.00	0.146	0.80 6	0.310	0.302	0.065 0	<0.04 6	0.121	0.51	148	0.5	1.5	479	559
10DX172-14	Qz2 b	14.0	63.3	v.d.	588	<70.88	51.1	v.d.	v.d.	v.d.	0.940	0.0450	0.52 9	0.253	v.d.	0.068 0	<0.05 2	0.542	0.51	-	0.5	1.5	472	551
10DX172-15	Qz2 b	16.1	34.4	235	138	99.9	57.4	1.415	v.d.	0.396	<0.42	0.110	0.69 0	0.152	0.799	0.122	<0.03 5	0.21	0.51	460	0.5	1.5	480	560

10DX172-16	Qz2 b	16.8	20.3	59.9	<4.43	<82.32	63.1	0.416	3.67	0.210	<0.41	0.142	$\frac{0.88}{0}$	1.02	0.852	0.161	$\frac{<0.04}{9}$	0.164	0.51	117	0.5	1.5	486	567
10DX172-17	Qz2 b	14.9	73.1	97.8	33.8	<75.05	56.8	v.d.	v.d.	v.d.	<0.61	<0.024	$\frac{0.85}{2}$	<0.20	v.d.	$\frac{0.090}{0}$	$\frac{<0.04}{7}$	0.421	0.51	192	0.5	1.5	479	559
10DX172-18	Qz2 b	13.9	93.3	117	46.5	127	47.2	v.d.	v.d.	v.d.	<0.88	0.112	$\frac{0.65}{2}$	0.190	0.424	0.116	0.088	0.737	0.51	229	0.5	1.5	467	544
10DX172-19	Qz2 b	13.1	110	73.0	15.0	110	80.2	2.22	v.d.	3.54	<0.27	0.0224	$\frac{0.78}{2}$	0.490	0.621	$\frac{0.048}{0}$	0.179	0.413	0.51	143	0.5	1.5	502	586
10DX197-8	Qz2 b	20.9	41.3	468	11.2	188	63.5	0.079	2.51	$\frac{<0.18}{4}$	<0.00	0.110	2.9	2.82	0.184	$\frac{<0.06}{0}$	0.734	1.02	0.58	806	0.5	1.5	487	568
10DX197-17	Qz2 b	13.8	23.0	92.5	<6.66	$\frac{<118.5}{0}$	37.3	0.41	3.94	0.580	<0.88	0.113	1.21	<0.26	0.149	0.096	$\frac{<0.06}{6}$	0.181	0.58	159	0.5	1.5	452	527
10DX197-20	Qz2 b	21.5	7.8	88.5	7.48	<52.51	59.1	0.262	3.27	0.119	0.530	0.0214	1.41	0.260	$\frac{0.031}{1}$	$\frac{<0.03}{2}$	0.087	0.123	0.58	153	0.5	1.5	482	562
10DX197-21	Qz2 b	8.8	26.3	72.7	<5.30	<91.44	40.5	0.594	v.d.	v.d.	v.d.	1.14	1.25	v.d.	v.d.	0.290	0.349	v.d.	0.58	125	0.5	1.5	457	533
10DX197-22	Qz2 b	16.0	35.0	86.0	12.9	102	53.2	0.508	5.36	$\frac{<0.14}{9}$	<0.75	0.211	1.03	0.980	0.087	0.128	0.101	0.136	0.58	148	0.5	1.5	475	554
10DX197-24	Qz2 b	17.8	14.8	82.6	10.1	103	41.8	0.251	3.22	0.890	0.370	$\frac{<0.015}{4}$	1.1	$\frac{<0.14}{2}$	$\frac{0.059}{9}$	0.179	0.069	0.071	0.58	142	0.5	1.5	459	535
10DX197-25	Qz2 b	17.0	22.6	77.5	<3.95	<69.28	22.5	0.376	3.14	0.353	<0.71	0.102	1.36	0.840	0.250	0.071	0.19	0.037	0.58	134	0.5	1.5	421	489
10DX197-32	Qz2 b	16.0	50.0	76.0	20.3	<59.14	42.2	v.d.	v.d.	v.d.	<0.68	<0.00	1.06	0.370	$\frac{0.060}{9}$	$\frac{<0.04}{3}$	$\frac{<0.04}{9}$	0.666	0.58	131	0.5	1.5	460	536
10DX197-33	Qz2 b	16.2	11.7	83.0	<4.09	<72.67	45.4	0.411	3.55	0.300	<0.68	$\frac{<0.012}{5}$	1.14	$\frac{<0.21}{6}$	1.02	0.102	0.056	0.048	0.58	143	0.5	1.5	465	541
10DX197-34	Qz2 b	18.5	v.d.	118	94.9	71.9	48.7	v.d.	v.d.	v.d.	v.d.	0.0440	1.36	0.960	0.187	0.148	0.08	v.d.	0.58	204	0.5	1.5	469	547
10DX197-36	Qz2 b	13.2	$\frac{10.0}{7}$	65.3	<4.41	<77.81	33.5	0.291	4.09	$\frac{<0.17}{8}$	0.780	0.295	1.21	0.650	$\frac{0.017}{9}$	$\frac{<0.05}{8}$	$\frac{<0.04}{5}$	0.052	0.58	113	0.5	1.5	445	518
10DX197-37	Qz2 b	16.1	$\frac{29.7}{4}$	109	13.7	<56.56	58.2	0.194	4.71	0.281	0.560	0.0180	$\frac{0.94}{9}$	1.60	0.111	$\frac{0.076}{0}$	0.138	0.061	0.58	188	0.5	1.5	481	561
10DX197-40	Qz2 b	11.6	$\frac{73.5}{4}$	v.d.	286	<51.82	36.6	v.d.	v.d.	v.d.	v.d.	0.184	$\frac{1.34}{1}$	v.d.	0.290	$\frac{0.074}{0}$	0.184	0.414	0.58	v.d.	0.5	1.5	451	525
12DXF05-1	Qz2 b	14.0	$\frac{40.6}{4}$	54.6	13.9	<55.57	69.2	v.d.	v.d.	v.d.	<0.66	0.0330	1.11	$\frac{<0.20}{0}$	0.950	$\frac{0.064}{0}$	0.069	0.711	0.52	105	0.5	1.5	492	574

12DXF05-3	Qz2 b	12.6	97.7 2	43.7	33.0	120	80.1	v.d.	v.d.	v.d.	v.d.	0.0440	1.06	<0.13 2	0.093 0	0.074 0	0.143	1.69	0.52	84	0.5	1.5	502	586
12DXF05-20	Qz2 b	14.9	v.d.	111	83.6	<89.02	81.8	v.d.	v.d.	v.d.	v.d.	0.256	0.97 7	<0.17 6	v.d.	0.176	0.125	v.d.	0.52	214	0.5	1.5	503	587
12DXF05-21	Qz2 b	13.4	41.6 6	69.8	42.5	<104.9 1	78.8	v.d.	v.d.	0.290	2.51	0.155	0.76	<0.30	0.128	<0.07 1	<0.07 6	v.d.	0.52	134	0.5	1.5	501	585
12DXF05-23	Qz2 b	15.8	v.d.	66.7	72.4	<71.69	72.7	v.d.	v.d.	v.d.	v.d.	0.126	1.13	<0.20	v.d.	0.106	<0.04 2	v.d.	0.52	128	0.5	1.5	496	578
12DXF05-24	Qz2 b	15.3	v.d.	v.d.	29.6	<104.5 2	69.2	v.d.	v.d.	v.d.	v.d.	0.517	1.16	1.03	0.925	0.076 0	0.112	v.d.	0.52	v.d.	0.5	1.5	492	574
12DXF05-30	Qz2 b	12.2	105	65.2	26.0	<45.88	75.2	v.d.	v.d.	v.d.	v.d.	0.0082	0.92 5	0.245	0.550	0.081 0	<0.03 1	v.d.	0.52	125	0.5	1.5	498	581
12LDX37-3	Qz2 b	9.08	80.4 2	138	42.4	v.d.	70.8	v.d.	v.d.	v.d.	2.39	0.306	2.56	0.280	v.d.	<0.05 6	1.50	v.d.	0.7	197	0.5	1.5	494	576
12LDX37-4	Qz2 b	12.1	v.d.	213	64.7	v.d.	68.6	v.d.	v.d.	v.d.	v.d.	0.390	1.75	1.09	v.d.	v.d.	v.d.	v.d.	0.7	304	0.5	1.5	492	574
12LDX37-6	Qz2 b	11.1	159	v.d.	v.d.	61.0	54.7	v.d.	v.d.	v.d.	0.980	0.492	1.47	0.390	v.d.	0.231	0.474	v.d.	0.7	v.d.	0.5	1.5	477	556
12LDX37-11	Qz2 b	15.7	v.d.	108	v.d.	v.d.	52.4	v.d.	v.d.	v.d.	v.d.	0.062	1.51	v.d.	v.d.	0.220	0.737	v.d.	0.7	155	0.5	1.5	474	553
12LDX37-16	Qz2 b	17.3	20.5	223	54.4	77.7	47.1	0.976	v.d.	v.d.	1.2	0.220	1.82	<0.24	0.540	0.114	1.02	0.374	0.7	319	0.5	1.5	467	544
12LDX37-17	Qz2 b	13.4	17.3	176	27.6	80.8	42.1	0.770	v.d.	v.d.	1.51	0.248	1.16	v.d.	0.253	0.167	0.807	v.d.	0.7	252	0.5	1.5	460	536
12LDX37-23	Qz2 b	13.4	75.3	125	21.5	<59.13	54.6	2.28	v.d.	4.21	1.81	0.039	1.24	0.146	0.582	0.074 0	1.13	v.d.	0.7	178	0.5	1.5	477	556
12LDX37-24	Qz2 b	10.7	33.8	142	39.1	267	74.2	v.d.	v.d.	v.d.	<0.00	0.033	1.73	0.190	v.d.	0.118	0.564	1.69	0.7	203	0.5	1.5	497	580
12LDX37-25	Qz2 b	12.7	80.0	430	252	<82.48	41.5	v.d.	v.d.	v.d.	1.14	0.236	1.30	1.10	1.70	0.127	0.271	v.d.	0.7	614	0.5	1.5	459	535
12LDX37-26	Qz2 b	13.2	54.6	172	46.3	369	51.5	v.d.	v.d.	v.d.	<0.93	0.433	2.89	0.730	1.36	0.133	0.233	1.59	0.7	245	0.5	1.5	473	551
12LDX37-27	Qz2 b	10.8	49.7	171	29.9	<94.15	55.4	v.d.	v.d.	v.d.	<0.00	0.201	1.51	1.46	v.d.	0.193	1.85	v.d.	0.7	245	0.5	1.5	478	557
12LDX37-29	Qz2 b	10.2	40.6	257	97.0	<55.01	51.4	0.831	v.d.	1.58	1.65	0.114	1.95	0.520	v.d.	<0.03 7	0.692	v.d.	0.7	367	0.5	1.5	473	551

10DX22-1	Qz2 b	15.5	29.3	292	88.0	<86.45	79.6	v.d.	v.d.	0.680	1.83	0.0220	1.27	<0.20	0.916	0.162	0.212	0.236	0.55	531	0.5	1.5	501	585
10DX22-2	Qz2 b	12.6	6.67	83.3	6.00	<74.98	54.7	0.325	4.70	2.93	0.530	0.0640	$\frac{0.56}{9}$	$\frac{<0.20}{5}$	0.054	0.071	0.125	0.079	0.55	151	0.5	1.5	477	556
10DX22-3	Qz2 b	14.8	9.05	77.8	12.1	<60.13	64.1	0.734	5.76	0.405	<0.66	0.140	$\frac{0.74}{1}$	0.266	1.54	$\frac{<0.04}{8}$	0.179	0.377	0.55	141	0.5	1.5	487	568
10DX22-6	Qz2 b	16.5	33.4	267	118	<53.61	55.0	1.24	v.d.	3.78	<0.68	0.0820	$\frac{0.90}{6}$	0.182	0.503	0.096	0.148	0.35	0.55	486	0.5	1.5	477	556
10DX22-8	Qz2 b	14.4	$\frac{<1.8}{1}$	81.0	7.33	<57.46	38.0	0.123	4.47	0.332	0.570	0.0390	0.92	0.370	v.d.	$\frac{0.094}{0}$	0.119	0.152	0.55	147	0.5	1.5	453	528
10DX22-9	Qz2 b	16.5	3.46	74.1	<2.62	<46.77	47.1	0.150	5.43	0.100	0.590	0.0290	$\frac{0.50}{8}$	$\frac{0.067}{0}$	0.541	0.069	0.246	0.128	0.55	135	0.5	1.5	467	544
10DX22-10	Qz2 b	16.2	11.1	378	131	<97.09	79.6	v.d.	v.d.	<0.20	v.d.	0.158	0.92	$\frac{<0.15}{0}$	0.722	$\frac{<0.07}{6}$	0.187	0.426	0.55	687	0.5	1.5	501	585
10DX22-11	Qz2 b	15.6	11.2	73.0	3.98	149	48.6	0.187	4.55	0.519	<0.37	0.108	$\frac{0.67}{9}$	0.206	1.52	$\frac{0.063}{0}$	0.122	0.25	0.55	133	0.5	1.5	469	547
10DX22-12	Qz2 b	17.0	107	163	78.5	77.5	60.7	8.25	v.d.	v.d.	v.d.	0.114	$\frac{0.49}{5}$	$\frac{<0.14}{6}$	0.225	$\frac{0.083}{0}$	0.150	1.26	0.55	297	0.5	1.5	484	564
10DX22-13	Qz2 b	18.2	5.63	112	21.4	145	30.1	0.306	5.72	$\frac{<0.11}{8}$	0.760	0.219	$\frac{1.07}{5}$	0.328	8.50	0.127	0.194	0.323	0.55	203	0.5	1.5	439	511
10DX22-14	Qz2 b	17.1	$\frac{<1.6}{2}$	66.1	<2.69	<47.42	49.1	0.143	3.37	0.162	<0.23	0.121	$\frac{0.74}{5}$	0.116	$\frac{0.027}{4}$	0.142	$\frac{0.096}{0}$	0.109	0.55	120	0.5	1.5	470	548
10DX22-15	Qz2 b	14.2	$\frac{<2.0}{5}$	68.4	4.91	<63.07	35.3	0.121	3.12	$\frac{<0.12}{6}$	<0.33	0.191	$\frac{0.75}{9}$	0.750	0.631	$\frac{0.067}{0}$	0.589	0.135	0.55	124	0.5	1.5	449	522
10DX22-16	Qz2 b	17.1	$\frac{<2.3}{4}$	79.4	5.94	77.9	56.0	0.202	4.33	$\frac{<0.15}{7}$	<0.39	0.115	$\frac{0.89}{6}$	0.708	0.338	$\frac{0.096}{0}$	0.145	0.162	0.55	144	0.5	1.5	478	558
10DX22-17	Qz2 b	16.5	2.05	85.2	<2.87	<52.04	25.4	0.110	3.91	$\frac{<0.10}{7}$	<0.49	0.0398	$\frac{0.96}{5}$	0.215	$\frac{0.038}{9}$	$\frac{0.074}{0}$	$\frac{0.061}{0}$	0.09	0.55	155	0.5	1.5	428	498
10DX22-18	Qz2 b	17.1	$\frac{<1.8}{4}$	84.4	<3.06	<59.32	73.3	0.199	4.17	0.234	<0.31	$\frac{<0.014}{7}$	$\frac{0.64}{2}$	0.331	$\frac{0.034}{3}$	$\frac{<0.05}{3}$	$\frac{<0.03}{6}$	<0.027	0.55	154	0.5	1.5	496	579
10DX22-19	Qz2 b	16.0	4.56	67.4	6.96	<43.90	60.2	0.208	4.01	$\frac{<0.10}{8}$	<0.37	0.217	$\frac{0.94}{8}$	0.570	0.107	$\frac{0.045}{0}$	0.162	0.143	0.55	123	0.5	1.5	483	563
10DX22-21	Qz2 b	19.3	30.7	270	113	96.7	89.2	0.930	11.5	0.126	<0.53	0.127	$\frac{0.89}{6}$	0.122	0.547	$\frac{<0.02}{9}$	0.110	0.136	0.55	491	0.5	1.5	509	594
10DX22-22	Qz2 b	18.1	21.9	107	25.6	<48.70	32.2	0.416	6.55	0.130	0.460	0.0950	$\frac{0.75}{4}$	0.243	0.130	$\frac{0.092}{0}$	$\frac{0.042}{0}$	0.183	0.55	195	0.5	1.5	443	516

10DX22-23	Qz2 b	13.4	13.4	104	21.0	<72.42	22.1	0.719	v.d.	0.255	<0.55	0.571	0.75 9	0.520	0.168	0.077 0	<0.06 2	0.272	0.55	189	0.5	1.5	420	488
10DX22-24	Qz2 b	16.7	22.8	89.8	17.5	225	38.1	0.593	v.d.	0.544	<0.36	0.0830	1.11	0.445	1.55	0.096 0	0.076 0	0.213	0.55	163	0.5	1.5	453	528
10DX22-25	Qz2 b	17.5	8.37	154	26.2	<61.85	42.6	0.497	v.d.	0.229	1.37	0.310	0.84 7	0.317	0.313	0.071 0	0.058 0	0.18	0.55	281	0.5	1.5	461	537
10DX22-26	Qz2 b	15.1	4.04	81.1	<3.19	<59.39	42.2	0.196	3.72	<0.12 5	<0.30	0.0136	0.83 5	0.161	0.094	0.123	0.055 0	0.028	0.55	148	0.5	1.5	460	536
10DX22-27	Qz2 b	16.2	4.08	85.4	6.06	<57.88	28.2	v.d.	v.d.	v.d.	<0.54	0.0393	0.77	<0.15 9	0.068 9	0.069	0.087 0	0.076	0.55	155	0.5	1.5	434	506
10DX22-28	Qz2 b	13.8	4.77	77.5	<4.01	<78.13	22.1	0.272	3.79	0.386	<0.54	0.0910	0.61 5	0.174	0.162	0.104	0.120 0	0.184	0.55	141	0.5	1.5	419	488
10DX22-29	Qz2 b	15.2	39.2	69.5	12.62	<42.37	18.3	0.831	7.57	0.094 0	<0.49	0.0470	0.79 4	<0.10 1	0.097 7	0.121	0.045 0	0.257	0.55	126	0.5	1.5	408	475
10DX22-31	Qz2 b	16.8	v.d.	77.1	154	58.0	38.4	v.d.	v.d.	v.d.	v.d.	0.100	0.71 9	0.339	2.17	0.087	0.051 0	v.d.	0.55	140	0.5	1.5	454	529
10DX22-32	Qz2 b	17.2	8.67	152	40.1	<52.38	48.7	0.327	7.01	0.206	0.630	0.0800	0.88 9	<0.13 7	0.241	0.163	0.130	0.079	0.55	277	0.5	1.5	469	547
10DX22-33	Qz2 b	14.6	<1.7 5	78.3	3.27	<54.53	21.3	0.116	3.68	<0.08 8	<0.62	0.111	0.57 2	0.174	0.133	0.116	0.078	<0.022	0.55	142	0.5	1.5	417	485
10DX22-34	Qz2 b	14.7	<2.0 0	79.6	<3.37	<61.83	40.8	0.147	3.68	0.214	<0.54	0.0730	0.99 3	0.199	0.015 6	<0.04 8	<0.04 0	<0.021 6	0.55	145	0.5	1.5	458	533
10DX22-35	Qz2 b	15.8	<1.5 8	73.9	<2.67	<47.37	40.0	0.091	v.d.	<0.08 1	<0.41	<0.011 3	0.79 5	0.112	0.062 3	0.269	0.099	0.074	0.55	134	0.5	1.5	457	532
10DX22-36	Qz2 b	15.1	8.39	88.2	9.04	<66.10	35.1	0.544	5.8	<0.10 7	1.14	0.0330	1.23	0.17	0.202	0.066	<0.04 1	0.151	0.55	160	0.5	1.5	448	522
10DX22-37	Qz2 b	17.8	7.59	154	36.2	76.6	45.3	0.332	6.86	<0.11 1	<0.39	0.0262	0.65 5	<0.12 8	0.132	0.086	<0.03 3	0.081	0.55	280	0.5	1.5	465	541
10DX22-38	Qz2 b	16.3	<1.7 1	76.7	<2.89	50.9	65.9	0.152	3.53	<0.10 4	<0.46	0.0820	0.73 5	0.141	0.023 0	0.120	<0.04 1	0.104	0.55	139	0.5	1.5	489	570
10DX22-39	Qz2 b	16.6	<1.7 2	85.0	<2.92	<54.19	65.6	0.162	3.62	0.123	0.470	0.0730	0.65 6	0.292	0.023 0	<0.04 0	0.075	0.03	0.55	154	0.5	1.5	489	570
10DX22-40	Qz2 b	17.1	<1.7 0	100	5.85	<52.00	63.3	0.261	5.6	<0.09 9	0.750	0.0254	0.79 3	<0.11 2	1.31	0.145	0.078	0.135	0.55	182	0.5	1.5	486	567
10DX22-41	Qz2 b	16.8	7.47	101	14.7	<54.98	33.2	0.266	5.55	<0.10 4	<0.39	0.0840	0.75 3	0.083	0.093 9	0.061 0	0.198	0.123	0.55	184	0.5	1.5	445	518

10DX22-42	Qz2 b	16.7	11.2	83.0	13.6	<45.62	38.2	0.323	5.94	0.373	<0.61	0.0540	$\frac{0.76}{2}$	0.224	4.46	$\frac{0.066}{0}$	0.148	0.15	0.55	151	0.5	1.5	454	528
10DX328-9	Qz2 b	15.2	6.46	77.4	<2.77	59.6	60.2	0.669	v.d.	0.350	1.70	0.0210	$\frac{0.96}{0}$	0.22	0.6	0.141	$\frac{<0.03}{9}$	0.218	0.55	102	0.5	1.5	483	563
10DX328-18	Qz2 b	19.9	50.4	247	11.9	v.d.	59.2	1.93	v.d.	3.06	2.21	0.0380	1.40	<0.24	1.32	0.247	0.131	0.606	0.76	325	0.5	1.5	482	562
10DX328-28	Qz2 b	24.0	v.d.	v.d.	74.2	<80.38	64.2	v.d.	v.d.	v.d.	5.60	0.0540	1.73	0.520	0.828	0.233	0.061	v.d.	0.76	v.d.	0.5	1.5	487	568
10DX328-29	Qz2 b	16.1	v.d.	v.d.	51.0	<59.11	53.2	v.d.	v.d.	v.d.	4.25	0.0400	2.44	1.63	0.428	$\frac{<0.03}{9}$	0.197	v.d.	0.76	v.d.	0.5	1.5	475	554
10DX328-30	Qz2 b	21.3	v.d.	v.d.	45.5	94.8	62.6	v.d.	v.d.	v.d.	v.d.	0.0660	2.02	0.820	v.d.	$\frac{0.075}{0}$	0.171	v.d.	0.76	v.d.	0.5	1.5	486	566
10DX328-31	Qz2 b	27.7	v.d.	333	64.9	<53.44	71.5	v.d.	v.d.	v.d.	v.d.	0.0340	1.70	0.800	0.368	$\frac{0.056}{0}$	$\frac{<0.03}{9}$	v.d.	0.76	438.2	0.5	1.5	494	577
10DX328-33	Qz2 b	13.9	v.d.	96	30.8	205	41.9	v.d.	v.d.	v.d.	v.d.	0.0460	1.87	0.177	v.d.	$\frac{0.084}{0}$	0.165	v.d.	0.76	125.9	0.5	1.5	460	535
10DX328-34	Qz2 b	7.6	v.d.	76	34.2	82.2	47.8	v.d.	v.d.	v.d.	v.d.	$\frac{<0.016}{9}$	1.40	0.600	v.d.	$\frac{0.075}{0}$	0.1	v.d.	0.76	99.7	0.5	1.5	468	546
10DX164-35	Qz2 b	16.6	52.8	158	29.5	<67.45	55.4	0.137	1.06	$\frac{<0.15}{1}$	<0.39	0.0690	2.45	0.730	0.240	0.114	0.473	0.142	0.66	239.7	0.5	1.5	478	557
10DX145-20	Qz2 b	7.9	86.2	v.d.	11.4	v.d.	25.6	v.d.	v.d.	1.25	2.26	v.d.	2.68	2.56	v.d.	0.259	0.411	v.d.	0.62	v.d.	0.5	1.5	428	499
10DX140-2	Qz2 b	13.8	76.4	158	44.2	95.1	58.3	v.d.	v.d.	v.d.	v.d.	0.211	1.24	1.14	0.228	0.127	0.687	v.d.	0.56	282.6	0.5	1.5	481	561
10DX140-7	Qz2 b	13.0	26.6	186	16.2	$\frac{<198.9}{7}$	41.9	1.01	v.d.	<0.65	<2.24	0.0870	2.38	2.34	0.642	$\frac{<0.13}{5}$	3.43	0.206	0.56	332.4	0.5	1.5	460	535
10DX140-17	Qz2 b	6.86	123	101	30.6	168	52.6	v.d.	v.d.	v.d.	v.d.	0.0184	1.63	1.14	v.d.	$\frac{<0.02}{9}$	1.10	0.869	0.56	179.9	0.5	1.5	474	553
10DX140-19	Qz2 b	8.49	v.d.	v.d.	85.2	$\frac{<144.7}{7}$	46.0	v.d.	v.d.	v.d.	v.d.	0.106	2.54	v.d.	1.14	$\frac{<0.15}{2}$	0.737	v.d.	0.56		0.5	1.5	466	543
10DX140-20	Qz2 b	8.15	v.d.	117	53.3	63.7	24.8	v.d.	v.d.	v.d.	v.d.	0.0245	1.96	1.76	1.08	$\frac{0.064}{0}$	v.d.	v.d.	0.56	208.2	0.5	1.5	426	496
10DX140-18	Qz2 b	5.62	188	96	23.3	<50.47	77.2	v.d.	v.d.	v.d.	v.d.	$\frac{<0.010}{7}$	1.51	19.8	0.794	$\frac{0.048}{0}$	v.d.	v.d.	0.56	171.9	0.5	1.5	499	583
10DX140-23	Qz2 b	7.27	108	128	38.4	<53.75	81.1	v.d.	v.d.	v.d.	v.d.	0.0257	1.78	1.46	1.03	0.123	2.13	v.d.	0.56	229.4	0.5	1.5	503	587

10DX140-24	Qz2 b	8.78	v.d.	138	50.0	<41.37	84.0	v.d.	v.d.	v.d.	v.d.	<0.014 6	1.59	1.01	1.50	0.104	0.842	v.d.	0.56	246.7	0.5	1.5	505	589
10DX140-25	Qz2 b	13.5	96.7	147	36.3	<44.07	73.8	v.d.	v.d.	1.61	1.49	0.112	1.81	1.25	1.48	0.075 0	1.02	0.77	0.56	263.2	0.5	1.5	497	579
10DX140-27	Qz2 b	9.61	v.d.	171	125	<48.33	84.3	v.d.	v.d.	v.d.	v.d.	0.0950	1.81	v.d.	1.80	0.132	14.7	v.d.	0.56	306.1	0.5	1.5	505	590
10DX140-31	Qz2 b	31.2	165	102 5	312	84.5	75.5	v.d.	v.d.	v.d.	<0.39	0.0460	2.5	2.09	1.20	0.072	2.74	0.953	0.56	1830	0.5	1.5	498	581
10DX140-58	Qz2 b	10.0	122	143	44.4	<62.66	94.1	3.35	v.d.	2.21	2.48	0.0257	1.76	1.69	0.277	0.161	3.13	1.67	0.56	256.0	0.5	1.5	512	598
10DX140-59	Qz2 b	11.4	279	262	66.8	88.6	47.6	v.d.	v.d.	v.d.	6.43	0.0770	2.52	2.52	0.425	0.109	1.85	v.d.	0.56	467.3	0.5	1.5	468	545
10DX140-68	Qz2 b	7.35	190	85	43.8	172	99.7	v.d.	v.d.	v.d.	v.d.	0.0210	1.69	0.400	0.735	0.144	0.663	v.d.	0.56	151.7	0.5	1.5	516	603
10DX140-79	Qz2 b	10.8	7.07	99	8.26	<58.06	93.8	0.258	5.54	0.178	0.250	0.0280	1.13	<0.09 0	1.047	0.111	<0.03 7	0.086	0.56	176.7	0.5	1.5	512	598
10DX140-51	Qz2 b	5.85	110	91	37.2	<98.97	55.3	v.d.	v.d.	v.d.	v.d.	0.190	1.61	0.680	0.995	1.15	3.52	v.d.	0.56	163.1	0.5	1.5	478	557
10DX140-55	Qz2 b	19.7	44.7	190	7.12	<103.8 6	41.3	0.565	7.26	0.98	1.08	0.0330	2.16	1.65	6.21	3.45	1.52	0.669	0.56	339.4	0.5	1.5	459	534
10DX140-57	Qz2 b	5.06	v.d.	73	34.7	<49.62	54.2	v.d.	v.d.	v.d.	v.d.	<0.00	1.39	1.20	0.60	0.195	1.08	v.d.	0.56	130.0	0.5	1.5	476	555
10DX140-60	Qz2 b	8.20	188	114	24.7	87.3	60.2	v.d.	v.d.	0.98	2.99	0.0750	1.89	1.58	1.99	0.117	1.44	v.d.	0.56	203.3	0.5	1.5	483	563
10DX140-64	Qz2 b	6.47	38.9	62	15.2	v.d.	40.6	0.46	4.41	<0.15 0	<0.83	<0.012 9	1.61	<0.16 6	v.d.	0.096	<0.05 4	0.411	0.56	109.9	0.5	1.5	458	533
10DX140-74	Qz2 b	10.4	v.d.	215	120	189	45.9	v.d.	v.d.	v.d.	v.d.	0.0680	2.02	6.64	0.685	0.173	6.51	v.d.	0.56	383.7	0.5	1.5	465	542
10DX140-26	Qz3 a	35.1	95.9	625	107	114	16.3	1.89	v.d.	2.79	1.07	v.d.	v.d.	v.d.	v.d.	0.082	1.38	0.936	0.56	1115.6	0.5	1.5	401	420
10DX140-28	Qz3 a	8.71	55.9	245	40.4	303	6.81	0.469	v.d.	v.d.	<0.32	0.0560	2.39	2.30	1.24	0.093	1.81	0.749	0.56	437.7	0.5	1.5	353	371
10DX220-33	Qz3 a	24.7	27.2	542	13.2	<74.17	9.81	0.095	1.302	<0.17 4	<0.56	<0.023 9	2.85	3.08	1.13	0.121	0.863	0.477	0.66	821.4	0.5	1.5	373	391
10DX220-34	Qz3 a	17.9	3.83	140	<4.32	<79.28	4.65	<0.056	<0.13 6	0.250	<0.45	0.0620	2.3	0.820	5.5	0.368	0.531	0.598	0.66	212.7	0.5	1.5	334	351

10DX145-40	Qz3 _a	3.16	63.7	305	74.2	<63.20	14.6	0.954	64.5	0.830	<0.37	0.110	1.81	0.820	1.12	0.061	1.12	1.31	0.62	492	0.5	1.5	395	414
10DX145-41	Qz3 _a	2.15	316	227	45.8	v.d.	13.4	1.16	v.d.	0.500	<0.68	0.0620	1.87	1.66	v.d.	0.087	0.891	0.85	0.62	366	0.5	1.5	390	409
10DX145-43	Qz3 _a	4.25	v.d.	548	115	v.d.	21.3	1.58	v.d.	1.45	1.00	0.303	1.53	1.33	v.d.	0.239	0.503	0.928	0.62	884	0.5	1.5	417	437
10DX145-46	Qz3 _a	1.38	182	234	30.7	v.d.	14.5	0.572	v.d.	2.68	<0.49	0.0850	1.93	1.22	v.d.	0.618	1.77	0.878	0.62	378	0.5	1.5	395	413
10DX145-47	Qz3 _a	1.91	v.d.	331	76.6	v.d.	12.3	1.33	v.d.	1.81	<0.87	0.116	2.51	1.55	v.d.	0.265	2.16	0.579	0.62	534	0.5	1.5	385	404
10DX145-44	Qz3 _a	4.11	v.d.	271	54.8	v.d.	16.2	2.7	v.d.	3.46	0.52	0.0330	1.96	0.240	1.19	0.272	1.06	1.55	0.62	438	0.5	1.5	401	420
10DX145-48	Qz3 _a	3.65	v.d.	262	36.4	<42.52	12.9	v.d.	v.d.	v.d.	v.d.	0.0210	2.2	0.780	0.545	0.338	1.05	v.d.	0.62	422	0.5	1.5	388	407
10DX145-49	Qz3 _a	1.62	148	79	15.5	123	9.42	2.34	v.d.	1.09	1.09	<0.010 ₉	1.81	0.330	v.d.	0.136	0.756	1.49	0.62	128	0.5	1.5	370	389
10DX158-15	Qz3 _a	4.00	63.2	52	14.5	122	9.02	1.14	3.29	1.92	<0.65	0.228	1.60	0.760	0.307	0.052	0.539	0.327	0.52	100	0.5	1.5	368	386
10DX158-16	Qz3 _a	3.60	88.3	50	22.8	<84.13	7.18	v.d.	v.d.	0.304	1.14	<0.027	1.40	0.740	0.313	0.084	0.518	0.361	0.52	95.7	0.5	1.5	356	374
10DX158-18	Qz3 _a	5.18	41.6	89	27.4	<73.60	9.91	v.d.	v.d.	v.d.	<0.67	0.0430	1.99	1.33	v.d.	0.101	0.938	v.d.	0.52	170	0.5	1.5	373	391
10DX158-20	Qz3 _a	2.15	83.4	43	15.7	87.75	7.95	v.d.	v.d.	v.d.	<0.70	0.166	1.44	1.69	v.d.	0.103	0.163	0.201	0.52	83.1	0.5	1.5	361	379
10DX158-26	Qz3 _a	2.80	71.9	70	9.65	169	7.36	1.33	2.25	0.940	<0.98	0.169	1.59	0.720	1.20	0.105	0.303	0.153	0.52	134	0.5	1.5	357	375
10DX158-27	Qz3 _a	3.17	159	110	21.4	134	7.9	1.29	4.15	2.45	1.00	0.0183	1.7	1.22	0.415	0.148	0.352	0.249	0.52	211	0.5	1.5	361	379
10DX158-28	Qz3 _a	3.36	116	86	27.3	<64.67	7.66	1.85	8.53	1.50	<0.58	<0.00	1.96	1.68	0.908	0.125	0.713	0.396	0.52	165	0.5	1.5	359	377
10DX158-1	Qz3 _a	0.87 ₀	31.7	62	<4.40	<80.15	8.07	1.54	6.32	1.05	0.680	0.0580	1.35	1.25	1.99	<0.05 ₉	0.843	0.94	0.52	118	0.5	1.5	362	380
10DX158-2	Qz3 _a	0.70 ₄	89.8	31	<3.28	<61.14	7.34	0.648	2.74	1.20	<0.54	0.0450	1.51	1.61	0.504	0.052 ₀	0.805	0.243	0.52	58.8	0.5	1.5	357	375
10DX140-10	Qz3 _b	43.9	49.6	468	41.3	131.6	19.4	0.233	3.25	0.403	0.280	0.0420	4.92	4.66	0.769	0.095 ₀	2.31	0.306	0.56	836	0.5	1.5	412	431

10DX140-11	Qz3 b	23.7	124	986	91.7	<65.06	21.7	v.d.	v.d.	v.d.	<0.72	0.234	11.6	v.d.	0.270	0.205	4.52	0.445	0.56	1761	0.5	1.5	418	438
10DX140-12	Qz3 b	23.5	205	¹¹⁴ ₇	58.7	155	28.6	1.36	v.d.	2.16	1.24	0.347	17.5	23.0	3.16	0.125	11.8	0.72	0.56	2048	0.5	1.5	435	455
10DX140-29	Qz3 b	2.75	96.7	78	23.6	270	19.4	0.499	6.12	2.86	0.530	0.068	1.74	1.56	0.333	0.117	1.10	0.829	0.56	140	0.5	1.5	411	431
10DX140-30	Qz3 b	6.53	105	112	28.9	<46.63	10.5	v.d.	v.d.	v.d.	0.310	0.132	1.79	0.700	0.791	0.088	0.521	0.362	0.56	200	0.5	1.5	377	395
10DX140-41	Qz3 b	53.2	43.5	560	23.5	v.d.	15.5	0.112	7.24	0.240	<0.74	0.035	2.09	1.81	v.d.	0.182	0.762	v.d.	0.56	1000	0.5	1.5	398	417
10DX140-42	Qz3 b	8.79	162	101	18.0	249	7.92	v.d.	v.d.	1.87	v.d.	0.0200	2.93	0.930	v.d.	0.052	0.272	v.d.	0.56	180	0.5	1.5	361	379
10DX140-49	Qz3 b	5.02	74.2	100	28.6	<46.86	20.0	4.75	v.d.	0.464	v.d.	0.0094	1.99	0.630	v.d.	0.097	0.762	1.03	0.56	178	0.5	1.5	413	433
10DX140-52	Qz3 b	2.33	41.3	60	25.5	v.d.	16.4	1.15	v.d.	0.78	<1.35	0.376	1.74	1.12	0.479	0.270	1.33	v.d.	0.56	107	0.5	1.5	402	421
10DX140-53	Qz3 b	43.9	74.3	621	43.8	97.9	26.3	0.133	4.30	1.12	0.760	0.201	2.91	4.79	v.d.	0.623	2.73	v.d.	0.56	1108	0.5	1.5	430	450
10DX140-54	Qz3 b	30.8	75.8	529	30.6	<86.92	12.6	0.337	7.85	0.42	0.530	0.519	4.55	v.d.	0.389	^{<0.06} ₈	1.39	1.66	0.56	944	0.5	1.5	387	405
10DX140-56	Qz3 b	26.5	v.d.	380	88.4	v.d.	11.1	2.52	3.70	3.63	1.24	4.64	5.01	3.79	v.d.	3.86	2.39	11.5	0.56	678	0.5	1.5	380	398
10DX140-61	Qz3 b	13.7	93.0	240	42.7	v.d.	5.86	0.621	v.d.	3.01	<0.39	0.0230	1.92	2.54	v.d.	0.185	1.81	0.295	0.56	429	0.5	1.5	346	363
10DX140-62	Qz3 b	51.5	50.5	495	26.6	128	17.8	0.197	4.22	0.69	^{<0.22} ₄	0.0590	3.00	0.960	0.254	0.183	0.371	0.164	0.56	884	0.5	1.5	406	426
10DX140-63	Qz3 b	22.4	41.0	215	19.4	v.d.	6.24	0.067	1.17	0.836	<0.29	0.202	2.44	2.05	v.d.	0.147	1.33	0.336	0.56	384	0.5	1.5	349	366
10DX140-65	Qz3 b	32.4	73.5	395	40.6	111	16.7	0.855	v.d.	1.37	<0.22	0.0111	5.25	1.30	0.379	0.072	1.45	0.259	0.56	705	0.5	1.5	403	422
10DX140-66	Qz3 b	22.5	24.2	192	9.09	178	6.48	1.11	1.44	^{<0.14} ₀	<0.57	0.0730	1.87	0.680	0.348	^{<0.03} ₇	0.224	0.123	0.56	343	0.5	1.5	351	368
10DX140-67	Qz3 b	35.1	34.2	297	20.7	<41.59	7.79	0.118	2.62	0.427	<0.43	0.0960	2.54	1.87	0.129	^{0.057} ₀	0.989	0.182	0.56	530	0.5	1.5	360	378
10DX140-75	Qz3 b	5.5	87.4	133	36.0	134	9.97	v.d.	v.d.	v.d.	v.d.	^{<0.022} ₀	2.94	1.27	v.d.	^{0.053} ₀	0.606	v.d.	0.56	238	0.5	1.5	374	392

10DX140-76	Qz3 b	13.2	34.6	209	81.6	268	11.1	0.576	v.d.	0.530	<0.90	$\frac{<0.019}{7}$	2.29	0.66	0.656	0.149	0.719	0.28	0.56	372	0.5	1.5	380	398
10DX140-77	Qz3 b	2.84	36.1	43	7.58	<69.44	6.24	0.327	2.37	0.177	0.910	$\frac{<0.012}{7}$	1.98	0.42	3.33	$\frac{<0.04}{2}$	0.181	0.166	0.56	76.4	0.5	1.5	349	366
10DX140-78	Qz3 b	2.60	135	47	11.2	361	7.57	v.d.	v.d.	v.d.	1.44	$\frac{<0.012}{7}$	2.09	1.72	3.70	0.214	0.6	0.731	0.56	83.8	0.5	1.5	359	377
10DX158-29	Qz3 b	1.24	103	41	11.8	<53.36	11.3	0.477	1.88	0.870	1.36	0.0198	1.28	1.04	0.393	$\frac{0.051}{0}$	1.73	0.192	0.52	79.3	0.5	1.5	380	399
10DX158-3	Qz3 b	1.11	60.0	54	8.26	<71.84	12.5	0.425	2.14	0.482	<0.44	0.0750	1.05	2.23	0.793	$\frac{0.055}{0}$	1.72	0.243	0.52	103	0.5	1.5	386	405
10DX158-4	Qz3 b	8.33	35.2	63	15.3	<47.02	14.4	0.173	1.72	0.930	<0.55	0.0350	1.23	0.630	$\frac{0.094}{3}$	$\frac{0.082}{0}$	0.431	0.69	0.52	122	0.5	1.5	394	413
10DX158-13	Qz3 b	$\frac{0.58}{4}$	59.2	51	7.37	<82.15	15.9	0.495	2.79	1.26	1.00	<0.028	1.53	0.620	4.75	0.171	0.524	0.183	0.52	98.0	0.5	1.5	400	419
10DX158-14	Qz3 b	4.25	106	61	15.0	112	19.2	1.29	6.68	0.960	<0.94	0.0430	1.40	0.740	0.450	0.201	0.645	0.281	0.52	117	0.5	1.5	411	430
10DX158-19	Qz3 b	10.3	55.4	90	16.9	86.8	13.3	0.577	5.20	0.363	0.890	0.160	2.11	1.60	v.d.	$\frac{0.065}{0}$	0.419	0.298	0.52	172	0.5	1.5	389	408
10DX145-1	Qz3 b	19.0	54.0	662	184	183	4.49	0.098	v.d.	1.26	1.72	0.315	0.44	0.890	0.121	$\frac{0.071}{0}$	1.3	1.7	0.62	1068	0.5	1.5	332	349
10DX145-2	Qz3 b	17.2	111	326	36.5	<86.48	19.3	0.372	v.d.	0.950	1.44	0.364	1.75	2.06	0.191	0.126	1.52	0.23	0.62	526	0.5	1.5	411	430
10DX145-7	Qz3 b	14.8	63.7	699	98.8	<83.92	2.49	1.09	6.36	2.22	2.68	0.132	2.21	7.72	0.258	$\frac{0.096}{0}$	4.71	0.2	0.62	1128	0.5	1.5	305	321
10DX145-8	Qz3 b	21.9	86.2	611	49.7	<92.88	20.1	v.d.	v.d.	0.940	2.13	0.068	2.22	2.07	v.d.	0.308	2.52	0.191	0.62	985	0.5	1.5	414	433
10DX145-10	Qz3 b	35.0	76.2	592	112	<93.09	6.15	1.06	v.d.	2.22	<0.00	0.292	1.85	1.82	0.164	0.282	2.77	0.202	0.62	955	0.5	1.5	348	365
10DX145-11	Qz3 b	46.1	65.6	655	6.45	<75.81	4.76	0.252	4.72	0.580	1.00	0.204	1.92	1.55	0.952	$\frac{0.089}{0}$	1.29	0.141	0.62	1057	0.5	1.5	335	352
10DX145-13	Qz3 b	26.2	13.3	496	59.9	<92.88	2.31	0.29	5.78	0.870	<0.54	0.215	3.24	1.63	0.110	$\frac{<0.08}{5}$	1.35	0.164	0.62	800	0.5	1.5	302	318
10DX145-16	Qz3 b	6.07	19.2	109	<5.73	$\frac{<109.6}{5}$	5.33	0.757	v.d.	1.12	<1.03	1.36	1.29	0.56	v.d.	$\frac{<0.07}{7}$	0.379	0.218	0.62	176	0.5	1.5	341	358
10DX145-23	Qz3 b	4.71	v.d.	174	17.1	$\frac{<104.9}{3}$	14.7	v.d.	v.d.	v.d.	v.d.	0.0540	2.25	1.34	0.498	0.137	0.149	v.d.	0.62	281	0.5	1.5	395	414

10DX145-24	Qz3 b	2.92	252	176	40.2	<72.64	13.7	v.d.	v.d.	v.d.	v.d.	0.132	2.42	3.04	v.d.	0.087	0.441	v.d.	0.62	283	0.5	1.5	391	410
10DX145-30	Qz3 b	5.17	115	296	36.0	120	17.8	v.d.	v.d.	3.53	1.90	0.0320	2.13	<0.28	0.474	0.112	0.096	1.50	0.62	477	0.5	1.5	406	426
10DX145-33	Qz3 b	47.3	12.8	959	12.8	232	4.41	<0.058	0.58	0.510	<1.00	0.615	4.98	1.62	0.549	0.184	0.950	0.2	0.62	1546	0.5	1.5	331	348
10DX145-3	Qz3 b	65.2	159	$\frac{240}{3}$	12.0	113	15.9	0.148	5.46	1.23	<0.67	0.208	3.37	4	0.076	0.085	v.d.	0.457	0.62	3876	0.5	1.5	400	419
10DX145-21	Qz3 b	13.7	128	$\frac{139}{7}$	231	<97.73	57.2	3.61	v.d.	2.41	0.610	0.238	3.56	1.54	0.539	0.234	0.567	0.756	0.62	2253	0.5	1.5	480	501
10DX145-25	Qz3 b	5.01	155	269	105	v.d.	26.5	v.d.	v.d.	0.350	v.d.	0.107	2.49	0.410	v.d.	0.221	0.095	v.d.	0.62	434	0.5	1.5	430	450
10DX145-27	Qz3 b	7.30	169	224	28.0	113	22.7	v.d.	v.d.	v.d.	v.d.	0.0180	3.05	1.05	0.429	0.069	0.200	v.d.	0.62	361	0.5	1.5	421	441
10DX145-28	Qz3 b	8.14	2.45	129	<3.88	82.4	15.0	0.065	2.76	$\frac{<0.15}{6}$	1.00	0.257	2.34	0.209	0.522	0.059	0.168	0.126	0.62	208	0.5	1.5	396	415
10DX164-20	Qz3 b	19.0	95.9	280	73.82	<48.85	35.1	0.656	5.18	1.27	$\frac{<0.25}{1}$	0.0290	2.68	2.76	1.95	0.136	6.82	0.304	0.66	425	0.5	1.5	448	469
10DX164-34	Qz3 b	11.4	$\frac{153}{5}$	128	38.7	v.d.	24.8	0.708	v.d.	0.930	<0.53	0.0440	1.92	1.66	v.d.	v.d.	v.d.	v.d.	0.66	194	0.5	1.5	426	446
10DX164-36	Qz3 b	18.7	74.8	159	24.6	<47.23	26.8	1.61	v.d.	1.05	1.23	<0.023	2.02	0.870	0.825	0.318	0.279	0.558	0.66	241	0.5	1.5	431	451
10DX164-39	Qz3 b	17.5	57.7	118	13.6	<43.00	23.9	0.124	1.88	1.89	<0.34	$\frac{<0.013}{1}$	2.08	0.730	0.155	$\frac{<0.03}{3}$	0.301	0.51	0.66	179	0.5	1.5	424	444
10DX164-50	Qz3 b	7.81	94.5	129	34.8	408	19.0	v.d.	v.d.	v.d.	0.840	0.366	1.93	0.750	v.d.	1.25	0.921	0.919	0.66	196	0.5	1.5	410	430
10DX168-1	Qz3 b	9.18	16.1	28.2	<3.09	<54.87	16.0	<0.031	0.809	$\frac{<0.12}{3}$	$\frac{<0.29}{9}$	0.217	1.25	0.440	v.d.	0.122	0.065	0.068	0.58	48.7	0.5	1.5	400	419
10DX168-3	Qz3 b	9.91	6.7	38.0	<5.20	<97.84	20.5	0.0930	1.19	0.460	0.890	0.354	1.44	0.710	4.10	0.119	0.108	0.192	0.58	65.6	0.5	1.5	415	434
10DX168-13	Qz3 b	6.68	98.9	42.7	<4.87	313	13.0	0.103	5.56	v.d.	<0.47	0.0820	1.64	<0.20	1.56	0.225	0.097	1.10	0.58	73.6	0.5	1.5	388	407
10DX168-18	Qz3 b	13.1	5.79	61.7	<5.21	<91.06	25.5	<0.057	0.442	<0.19	0.460	0.225	1.43	0.220	$\frac{0.022}{9}$	$\frac{<0.07}{6}$	0.121	0.24	0.58	106	0.5	1.5	428	448
10DX168-20	Qz3 b	6.97	23.6	28.1	<4.10	v.d.	14.9	<0.039	1.35	0.320	<0.66	0.161	1.46	$\frac{<0.17}{9}$	v.d.	0.095	0.109	0.908	0.58	48.4	0.5	1.5	396	415

10DX168-23	Qz3 b	9.91	<3.4 2	52.0	<5.98	<108.8 6	23.3	<0.039	0.321	1.35	0.360	0.0630	1.66	0.340	0.014 2	<0.07 4	0.857	0.3	0.58	89.6	0.5	1.5	423	442
10DX168-25	Qz3 b	12.4	<2.3 8	41.0	<4.18	<72.94	20.4	0.059	0.738	<0.14 7	0.890	0.157	1.46	<0.20 8	0.238	<0.06 4	0.212	0.959	0.58	70.7	0.5	1.5	415	434
10DX168-30	Qz3 b	6.87	14.5	30.9	<4.02	775	16.4	<0.035	0.953	v.d.	0.620	0.424	1.58	0.480	v.d.	<0.05 8	0.122	0.435	0.58	53.3	0.5	1.5	402	421
10DX168-31	Qz3 b	5.16	30.0	26.6	<4.98	v.d.	11.8	0.0720	3.78	0.490	<0.45	0.144	1.48	0.790	v.d.	<0.06 0	0.121	14.8	0.58	45.8	0.5	1.5	383	401
10DX168-33	Qz3 b	13.5	13.8	52.9	7.82	<56.26	19.4	0.0580	2.72	0.249	<0.43	0.120	1.86	0.260	v.d.	<0.04 9	0.202	0.121	0.58	91.2	0.5	1.5	412	431
10DX168GY-Q04	Qz3 b	5.68	v.d.	44.9	0.383	37.4	17.0	0.965	1.90	0.044	0.000	0.009	1.06 4	0.215	0.000	0.095	0.008	0.024	v.d.	v.d.	0.5	1.5	404	423
10DX168GY-Q52	Qz3 b	8.80	46.4	116	9.36	121	17.5	0.218	2.25	1.94	0.000	0.003	1.33 8	0.236	0.521	0.244	0.045	0.082	v.d.	v.d.	0.5	1.5	406	425
10DX197-1	Qz3 b	11.3	104. 1	55.5	12.7	110	3.90	v.d.	v.d.	v.d.	1.01	<0.01	1.92	v.d.	0.967	<0.04 1	0.260	0.556	0.58	96	0.5	1.5	326	342
10DX197-2	Qz3 b	22.5	62.3	126	6.87	260	5.38	0.258	1.62	v.d.	0.68	0.0136	2.78	v.d.	0.089 0	0.084	0.376	0.228	0.58	217	0.5	1.5	341	359
10DX197-3	Qz3 b	2.76	72.2	44.1	7.14	118	2.42	0.892	3.13	1.62	<0.56	0.0770	1.82	10.6	1.71	<0.04 1	0.533	0.089	0.58	76	0.5	1.5	304	320
10DX197-4	Qz3 b	2.30	46.2	23.3	<4.26	91.7	3.47	<0.038	<0.11 8	<0.13 9	1.43	0.262	2.24	0.600	v.d.	0.146	0.055	0.089	0.58	40	0.5	1.5	320	337
10DX197-5	Qz3 b	4.02	36.3	23.1	<4.80	v.d.	15.3	0.053	0.454	<0.18 4	<0.64	0.376	1.22	v.d.	v.d.	0.090 0	0.198	0.192	0.58	40	0.5	1.5	398	417
10DX197-6	Qz3 b	27.6	35.2	126	7.03	76.1	13.5	0.128	2.05	0.410	<0.36	0.0680	2.24	1.60	v.d.	<0.04 9	<0.05 0	0.155	0.58	218	0.5	1.5	390	409
10DX197-7	Qz3 b	35.2	94.0	221	38.2	239	10.8	0.225	1.87	<0.17 7	<0.83	0.170	1.54	1.71	0.648	<0.06 8	0.156	0.131	0.58	380	0.5	1.5	378	396
10DX197-9	Qz3 b	11.8	17.6	79.9	<5.80	193	8.03	0.268	1.03	<0.23	<1.38	0.0590	1.57	1.36	1.48	0.135	<0.06 7	0.156	0.58	138	0.5	1.5	362	380
10DX197-10	Qz3 b	16.8	28.1	95.5	9.43	96.2	19.4	0.072	0.87	0.221	<0.26 2	0.0690	2.16	2.33	0.151	0.062	0.152	0.049	0.58	165	0.5	1.5	412	431
10DX197-11	Qz3 b	9.65	33.4	62.7	<2.80	<50.51	9.17	v.d.	v.d.	0.568	0.43	<0.010 8	1.50	<0.13 2	0.619	0.121	0.122	0.076	0.58	108	0.5	1.5	369	387
10DX197-12	Qz3 b	23.6	180	171	40.0	<53.95	8.42	v.d.	v.d.	v.d.	<0.36	0.116	1.67	0.670	0.149	0.138	0.115	1.24	0.58	294	0.5	1.5	364	382

10DX197-13	Qz3 b	10.3	98.1	65.4	4.90	93.4	16.7	v.d.	v.d.	0.551	<0.51	0.0590	2.47	0.660	0.099	0.064	0.248	0.644	0.58	113	0.5	1.5	403	422
10DX197-14	Qz3 b	6.88	17.8	47.3	<3.21	<55.06	17.1	0.106	0.232	<0.11 4	0.53	<0.008 9	1.77	0.650	0.558	0.081	0.227	0.032	0.58	82	0.5	1.5	404	423
10DX197-15	Qz3 b	22.7	19.9	153	9.92	<69.01	4.20	0.042	0.74	0.235	<0.63	0.0680	3.37	0.710	0.040 5	0.115	0.370	0.57	0.58	264	0.5	1.5	329	346
10DX197-16	Qz3 b	1.34	20.0	20.4	<7.39	<131.7 8	9.23	0.102	0.624	0.370	<1.41	0.265	1.17	0.880	0.041	0.170	<0.08 6	0.194	0.58	35	0.5	1.5	369	387
10DX197-19	Qz3 b	15.3	73.0	66.0	7.18	<85.37	8.13	0.440	1.42	<0.19	1.2	<0.014 2	1.07	<0.23	0.069	0.130	0.106	0.415	0.58	114	0.5	1.5	363	380
10DX197-26	Qz3 b	26.7	<2.8 8	107	<5.05	<88.73	9.28	<0.044	0.398	<0.17 9	<0.61	0.281	1.74	0.250	1.89	<0.05 2	<0.05 8	0.149	0.58	184	0.5	1.5	370	388
10DX197-27	Qz3 b	22.3	<1.7 7	131	<3.09	<56.89	26.0	<0.031	1.14	<0.12 0	<0.65	0.0790	1.98	1.46	0.003 8	0.076 0	0.224	<0.022	0.58	226	0.5	1.5	429	449
10DX197-28	Qz3 b	4.14	37.0	37.5	8.66	<48.32	9.95	0.210	2.18	v.d.	0.51	<0.007 8	2.01	0.500	0.047 3	0.122	<0.03 2	0.156	0.58	65	0.5	1.5	373	392
10DX197-29	Qz3 b	3.26	38.2	26.4	<3.59	<65.55	8.92	0.238	5.59	0.360	<0.78	0.803	1.41	0.450	1.10	0.072 0	<0.03 5	0.234	0.58	45	0.5	1.5	367	386
10DX197-30	Qz3 b	26.0	31.6	155	11.3	<53.62	28.4	v.d.	v.d.	0.438	<0.28	0.0780	2.38	0.620	0.705	0.075 0	0.305	0.149	0.58	267	0.5	1.5	435	455
10DX197-31	Qz3 b	33.7	11.0	157	10.8	<62.36	9.53	0.109	2.7	0.233	<0.59	0.0510	0.84 5	0.470	0.027 4	0.064 0	0.080 0	0.125	0.58	271	0.5	1.5	371	389
10DX197-35	Qz3 b	31.0	45.5	339	23.1	<51.87	31.8	v.d.	v.d.	0.380	1.07	0.0590	2.69	v.d.	v.d.	0.095 0	0.321	1.92	0.58	585	0.5	1.5	442	462
10DX197-38	Qz3 b	32.0	10.3	176	4.98	<53.99	22.0	0.119	2.1	0.756	<0.45	<0.00	2.07	0.670	1.68	0.059 0	0.077	0.032	0.58	304	0.5	1.5	419	439
10DX197-39	Qz3 b	14.8	55.7	94.7	<3.46	<60.32	15.2	v.d.	v.d.	<0.15 1	<0.49	0.249	1.9	1.55	1.12	0.102	0.157	0.500	0.58	163	0.5	1.5	397	416
10DX201-15	Qz3 b	58.4	139	928	79.7	<70.20	16.3	0.176	v.d.	0.670	<0.61	0.0560	4.6	11.7	0.94	<0.04 2	10.7	0.110	0.61	1522	0.5	1.5	401	420
10DX201-16	Qz3 b	61.1	218	201 8	401	<69.26	17.9	0.573	2.74	1.45	<0.82	0.108	6.68	32.8	0.831	0.112	29.6	0.093	0.61	3307	0.5	1.5	407	426
10DX201-17	Qz3 b	54.4	186	135 0	207	59.94	18.5	0.613	4.26	1.47	<0.29	0.0482	6.58	14.7	1.24	<0.04 0	12.4	0.168	0.61	2213	0.5	1.5	409	428
10DX201-18	Qz3 b	43.8	153	146 2	473	<72.32	12.4	2.56	v.d.	4.36	1.31	0.0630	6.06	11.2	v.d.	0.184	13.2	0.566	0.61	2397	0.5	1.5	386	404

10DX201-28	Qz3 b	49.3	68.3	465	66.4	75.72	11.1	0.922	v.d.	0.226	1.20	0.167	4.09	2.87	1.99	0.131	3.07	0.18	0.61	763	0.5	1.5	380	398
10DX201-29	Qz3 b	52.2	134	136 8	164	<74.85	15.7	0.643	4.95	1.158	<0.41	0.102	5.21	13.4	2.28	<0.04 4	9.56	0.226	0.61	2243	0.5	1.5	399	418
10DX201-31	Qz3 b	55.7	131	121 8	217	<52.89	14.9	0.984	v.d.	0.708	0.350	0.0700	5.07	11.8	2.23	0.050 0	8.34	0.300	0.61	1997	0.5	1.5	396	415
10DX201-32	Qz3 b	44.2	176	128 7	196	330.7	18.7	0.595	1.33	1.44	<0.37	0.0800	6.76	19.9	1.80	0.151	17.5	0.113	0.61	2109	0.5	1.5	410	429
10DX201-33	Qz3 b	46.9	103	118 1	216	<81.39	17.6	1.489	v.d.	1.06	<0.70	0.0630	6.54	15.4	4.74	<0.06 0	9.03	0.410	0.61	1936	0.5	1.5	406	425
10DX201-34	Qz3 b	58.5	130	886	131	63.89	12.0	0.418	1.35	0.359	<0.21 0	0.0132	7.45	10.6	0.928	0.037 0	7.63	0.139	0.61	1453	0.5	1.5	384	402
10DX201-35	Qz3 b	48.0	156	864	96.1	<58.67	9.99	1.54	v.d.	0.793	0.660	0.0550	7.75	12.2	0.267	0.098 0	7.27	0.182	0.61	1416	0.5	1.5	374	392
10DX201-36	Qz3 b	48.2	153	115 9	163	316	18.4	1.612	v.d.	1.45	<0.00	0.184	7.32	16.5	1.89	0.065 0	8.89	0.891	0.61	1899	0.5	1.5	408	428
10DX220-1	Qz3 b	16.4	26.4	218	6.35	171	13.0	0.598	v.d.	v.d.	1.48	0.111	1.39	<0.26	0.279	0.094 0	0.471	0.150	0.66	330.1	0.5	1.5	388	407
10DX220-2	Qz3 b	6.94	40.9	108	4.79	<55.50	10.5	0.338	4.72	1.62	0.550	<0.013 3	1.76	<0.17	0.315	0.130	0.400	0.111	0.66	164.3	0.5	1.5	376	395
10DX220-3	Qz3 b	7.79	41.0	70.2	<3.16	<58.32	8.72	1.395	v.d.	2.37	0.920	0.0280	1.19	0.077	0.073	0.103	0.190	0.344	0.66	106.3	0.5	1.5	366	384
10DX220-4	Qz3 b	10.4	42.9	318	43.0	<137.7 8	13.5	2.02	v.d.	3.1	0.640	0.0190	2.22	1	1.21	<0.10 2	0.474	0.295	0.66	482.0	0.5	1.5	391	409
10DX220-5	Qz3 b	7.18	96.8	207	38.0	<100.9 4	11.1	2.10	v.d.	3.07	<0.76	0.0670	1.65	0.790	0.845	0.169	0.468	0.421	0.66	313.3	0.5	1.5	379	398
10DX220-6	Qz3 b	6.13	81.1	83.3	12.8	<101.5 1	9.94	3.53	v.d.	3.86	0.67	0.277	1.28	0.690	0.395	0.106	0.128	0.684	0.66	126.2	0.5	1.5	373	392
10DX220-7	Qz3 b	8.44	121	209	30.3	124.2	13.9	2.08	v.d.	v.d.	1.55	0.161	2.05	1.39	0.504	0.140	0.345	0.42	0.66	316.7	0.5	1.5	392	411
10DX220-8	Qz3 b	6.55	60.7	97.8	12.7	<96.30	9.15	1.699	v.d.	v.d.	0.990	0.080	1.31	0.570	0.105	<0.06 7	<0.08 4	0.417	0.66	148.1	0.5	1.5	369	387
10DX220-9	Qz3 b	2.16	187	100	25.1	<58.73	9.85	v.d.	v.d.	v.d.	v.d.	0.340	1.32	0.960	1.05	0.047	0.173	v.d.	0.66	151.7	0.5	1.5	373	391
10DX220-10	Qz3 b	3.52	51.5	184	26.4	149	6.49	v.d.	v.d.	v.d.	<1.14	0.223	1.89	0.800	v.d.	<0.10 1	0.476	0.591	0.66	279.3	0.5	1.5	351	368

10DX220-13	Qz3 b	5.98	101	903	321	144	15.8	v.d.	v.d.	v.d.	0.520	0.447	3.87	2.41	v.d.	0.114	0.976	0.500	0.66	1368.6	0.5	1.5	399	418
10DX220-14	Qz3 b	15.7	53.2	173	18.0	<104.0 5	11.6	2.74	v.d.	1.00	1.21	0.0360	2.78	0.330	0.225	<0.07 2	0.216	0.542	0.66	261.6	0.5	1.5	382	400
10DX220-15	Qz3 b	3.50	103	48.3	20.4	73.8	9.36	v.d.	v.d.	v.d.	v.d.	0.0286	1.18	0.540	0.188	0.084 0	0.072	0.721	0.66	73.1	0.5	1.5	370	388
10DX220-16	Qz3 b	4.17	135	v.d.	78.3	<86.63	17.3	v.d.	v.d.	v.d.	<1.03	0.0550	2.70	1.14	0.282	0.081 0	0.641	0.227	0.66	v.d.	0.5	1.5	405	424
10DX220-18	Qz3 b	2.93	170. 8	109	9.95	<108.6 0	4.20	v.d.	v.d.	v.d.	1.33	0.0400	1.68	2.19	0.595	0.184	0.106	0.433	0.66	165.0	0.5	1.5	329	346
10DX220-22	Qz3 b	12.6	30.9	116	5.64	95.0	10.1	1.166	v.d.	2.35	<0.53	<0.00	1.39	<0.24	0.126	0.079 0	0.308	0.448	0.66	175.3	0.5	1.5	374	393
10DX220-23	Qz3 b	14.3	91.8	633	144	<50.76	16.9	v.d.	v.d.	1.45	1.39	0.165	2.90	2.00	1.44	0.121	0.426	0.628	0.66	959.7	0.5	1.5	403	422
10DX220-24	Qz3 b	20.1	98.1	143	44.2	<56.27	8.91	v.d.	v.d.	0.544	<0.55	<0.00	2.01	1.15	0.198	0.282	0.218	1.33	0.66	216.2	0.5	1.5	367	385
10DX220-25	Qz3 b	14.0	331	140	104	76.1	7.01	v.d.	v.d.	v.d.	v.d.	<0.00	1.59	1.26	0.347	0.189	0.513	v.d.	0.66	211.6	0.5	1.5	355	372
10DX220-26	Qz3 b	12.8	79.3	65.1	36.9	259	8.50	v.d.	v.d.	0.836	1.77	0.164	1.73	0.560	1.15	0.413	0.430	v.d.	0.66	98.6	0.5	1.5	365	383
10DX220-27	Qz3 b	8.4	71.4	56.6	17.8	<66.10	10.2	3.05	v.d.	1.02	<0.41	0.189	1.17	0.340	1.05	0.247	0.189	1.49	0.66	85.8	0.5	1.5	375	393
10DX220-28	Qz3 b	22.4	45.2	144	29.4	<62.52	10.3	1.45	v.d.	0.87	0.960	0.0470	1.83	0.550	2.90	0.069	0.145	0.930	0.66	217.6	0.5	1.5	375	393
10DX220-31	Qz3 b	22.0	4.85	92.2	<3.01	<55.67	24.5	0.183	2.19	0.44	<0.33	<0.00	1.97	0.260	0.024 8	0.171	0.446	1.04	0.66	139.6	0.5	1.5	426	445
10DX220-32	Qz3 b	10.2	159	93.1	38.3	185	16.8	v.d.	v.d.	1.21	1.18	0.177	1.47	0.260	v.d.	0.119	0.409	v.d.	0.66	141.0	0.5	1.5	403	422
10DX220-36	Qz3 b	18.0	104	194 1	89.4	<89.06	28.7	<0.060	0.675	0.366	0.310	0.0550	2.81	8.95	0.036 9	<0.07 4	1.37	0.455	0.66	2941.5	0.5	1.5	435	455
10DX220-37	Qz3 b	23.0	78.1	170 3	66.1	137	20.4	0.580	1.40	<0.24	<0.00	0.0700	2.98	6.48	0.174	<0.07 7	0.687	0.454	0.66	2581.0	0.5	1.5	415	434
10DX220-38	Qz3 b	28.3	26.8	731	22.2	<102.3 6	16.9	0.063	2.15	0.200	<1.00	0.0400	3.09	1.90	0.850	<0.05 0	0.515	0.278	0.66	1107.8	0.5	1.5	404	423
10DX328-17	Qz3 b	19.0	23.6	110	9.23	81.4	24.8	v.d.	v.d.	0.629	<0.50	0.0460	1.43	<0.21 3	0.065 1	0.111	0.139	0.292	0.76	145.0	0.5	1.5	426	446

10DX328-19	Qz3 _b	16.1	103	83.2	14.4	v.d.	22.0	v.d.	v.d.	3.81	3.26	<0.011 3	1.10 2	<0.22	v.d.	0.112	<0.04 4	0.857	0.76	109.5	0.5	1.5	419	439
10DX328-20	Qz3 _b	19.7	13.0	111	5.790	<59.76	24.5	0.380	v.d.	<0.17 2	<0.61	0.048	1.21 0	<0.17 2	0.164	0.157	<0.04 0	0.088	0.760	146.4	0.500	1.5	426	445
10DX328-21	Qz3 _b	9.86	10.8	67.8	<4.56	<96.43	26.2	0.924	v.d.	2.840	1.590	0.136	0.81 0	<0.31	0.134	0.079	0.130	0.223	0.760	89.1	0.500	1.5	430	450
10DX328-22	Qz3 _b	13.8	49.0	139	18.3	650	32.3	v.d.	v.d.	v.d.	<1.06	0.045	1.72 0	0.420	1.056	0.148	<0.06 7	0.572	0.760	182.6	0.500	1.5	443	463
10DX168GY-Q49	Qz4	4.66	1.33	45.8	1.462	0.000	7.84 2	v.d.	0.167	v.d.	0.185	0.021	1.61 6	0.207	0.015	0.236	0.062	0.000	v.d.	v.d.	0.100	1.5	353	378
10DX168GY-Q29	Qz4	12.7	1.46	107	3.294	51.7	8.18 8	0.070	0.494	0.049	0.067	0.000	1.54 2	0.208	0.009	0.193	0.059	0.000	v.d.	v.d.	0.100	1.5	356	381
10DX168GY-Q35	Qz4	10.2	44.4	120	10.4	75.2	7.33 5	0.238	1.694	0.678	0.486	0.019	2.51 6	0.789	0.394	0.123	0.326	0.042	v.d.	v.d.	0.100	1.5	350	375
10DX168-17	Qz4	25.6	<3.3 1	156	<5.79	<105.5 1	6.43 0	0.071	0.862	0.252	<0.56	0.058	2.29 0	0.700	0.275	0.100	0.189	0.242	0.580	268.6	0.100	1.5	343	368
10DX168-19	Qz4	4.74	17.9	22.4	<4.62	312	7.69 0	0.063	0.773	0.327	0.610	0.062	1.54 0	0.460	0.429	<0.06 6	0.200	0.692	0.580	38.6	0.100	1.5	352	377
10DX168-22	Qz4	v.d.	123	v.d.	5.810	<79.49	9.84 0	0.120	2.120	0.588	<0.58	0.649	2.12 0	0.760	0.153	<0.06 2	0.202	0.346	0.580	v.d.	0.100	1.5	366	391
DX168GY-11	Qz4	2.21	14.1	30.8	1.61	46.0	6.57	v.d.	v.d.	v.d.	v.d.	0.036	1.52	1.00	0.076	0.155	0.087	0.023	v.d.	v.d.	0.1	1.5	344	369
DX168GY-12	Qz4	3.14	30.1	51.8	7.04	41.3	14.7	0.0766	1.85	0.817	0.507	0.000	0.57 9	1.07	0.735	0.260	0.262	0.078	v.d.	v.d.	0.1	1.5	388	414
DX168GY-15	Qz4	2.64	48.2	51.4	6.29	72.0	7.65	0.71	0.838	0.096	0.000	0.017	1.57	1.01	1.21	0.193	0.209	0.014	v.d.	v.d.	0.1	1.5	352	377
DX168GY-16	Qz4	4.68	54.2	59.9	6.03	23.5	7.57	0.000	1.10	0.790	0.102	0.014	1.38	1.21	0.748	0.199	0.378	0.002	v.d.	v.d.	0.1	1.5	352	377
DX168GY-17	Qz4	3.73	65.1	104	32.6	v.d.	1.08	v.d.	v.d.	1.44	1.73	0.052	0.40 6	0.683	v.d.	0.070	0.443	0.891	v.d.	v.d.	0.1	1.5	264	285
DX168GY-23	Qz4	1.46	16.1	26.1	2.13	37.7	2.60	0.342	2.51	1.24	0.276	0.000	1.70	0.363	0.219	0.261	0.033	0.061	v.d.	v.d.	0.1	1.5	300	323
DX168GY-24	Qz4	2.67	3.83	27.7	0.580	29.4	4.23	v.d.	v.d.	0.26	0.000	0.003	1.73	0.325	0.012	0.149	0.047	0.014	v.d.	v.d.	0.1	1.5	323	346
DX168GY-26	Qz4	2.86	0.31 2	32.7	0.497	38.4	4.23	0.000	2.57	0.000	0.000	0.001	1.80	0.239	0.011	0.089	0.042	0.000	v.d.	v.d.	0.1	1.5	323	346
DX168GY-28	Qz4	3.26	v.d.	24.0	0.856	54.7	2.93	0.297	0.000	0.091	0.000	0.008	1.46	0.287	0.004	0.146	0.030	0.004	v.d.	v.d.	0.1	1.5	306	329
DX168GY-36	Qz4	1.67	3.92	30.1	0.336	85.4	2.88	0.000	1.64	0.015	0.000	0.000	1.88	0.101	0.177	0.201	0.017	0.006	v.d.	v.d.	0.1	1.5	305	328
DX168GY-37	Qz4	10.0	36.6	167	18.2	v.d.	2.30	0.000	v.d.	0.317	0.000	1.00	1.36	4.25	v.d.	0.360	0.345	0.145	v.d.	v.d.	0.1	1.5	295	317
DX168GY-40	Qz4	3.07	v.d.	26.7	2.12	26.3	2.69	0.000	0.000	0.133	0.141	0.020	2.27	0.198	0.171	0.210	0.031	0.041	v.d.	v.d.	0.1	1.5	302	325
DX168GY-41	Qz4	1.59	0.24	19.8	0.324	59.0	2.70	0.370	0.000	0.115	0.549	0.000	1.69	0.029	0.002	0.085	0.057	0.001	v.d.	v.d.	0.1	1.5	302	325
DX168GY-46	Qz4	1.04	1.51	21.7	0.398	72.1	5.99	0.000	0.000	0.058	0.000	0.002	1.67	0.257	0.002	0.181	0.019	0.001	v.d.	v.d.	0.1	1.5	340	364

DX168GY-50	Qz4	2.47	14.2	29.4	2.49	73.9	3.48	0.008	1.80	0.080	1.53	0.014	1.38	0.137	0.027	0.018	0.080	0.030	v.d.	v.d.	0.1	1.5	313	337
10DX261-23	Qz4	14.1	v.d.	81.5	185	<85.49	14.7	v.d.	v.d.	v.d.	v.d.	0.182	1.63	0.770	v.d.	0.132	0.147	v.d.	0.62	131	0.1	1.5	388	414
10DX261-24	Qz4	12.3	21.2	86.5	28.3	<90.66	17.4	v.d.	v.d.	v.d.	<0.66	0.245	1.39	<0.24	0.841	$\frac{0.073}{0}$	0.681	0.655	0.62	140	0.1	1.5	397	424
10DX261-31	Qz4	60.9	54.5	569	18.7	119	10.9	1.01	v.d.	0.880	1.67	0.159	3.93	2.03	0.976	$\frac{<0.07}{8}$	9.47	0.19	0.62	919	0.1	1.5	371	397
10DX164-37	Qz5 _a	10.6	50.5	168	56.0	v.d.	7.65	0.550	v.d.	1.05	<0.97	0.430	3.47	1.15	v.d.	0.499	0.669	v.d.	0.66	255	0.1	0.5	352	359
10DX164-38	Qz5 _a	10.3	52.5	115	10.2	$\frac{<100.1}{8}$	7.59	<0.071	2.86	3.10	<0.83	0.145	3.19	0.790	1.23	0.204	0.827	0.505	0.66	174	0.1	0.5	352	359
10DX164-47	Qz5 _a	13.1	59.0	171	29.9	v.d.	9.74	<0.100	v.d.	0.670	<1.16	0.290	4.41	1.73	1.48	$\frac{<0.10}{4}$	4.02	1.75	0.66	260	0.1	0.5	365	372
10DX140-38	Qz5 _a	10.9	63.9	226	31.4	<77.06	8.63	0.436	v.d.	0.860	<0.95	0.0570	3.07	2.12	v.d.	$\frac{<0.05}{1}$	1.70	v.d.	0.56	403	0.1	0.5	358	366
10DX140-1	Qz5 _b	25.5	134	306	61.8	79.8	4.69	0.149	4.68	3.33	3.22	0.150	4.99	7.85	0.641	0.963	6.09	v.d.	0.56	546	0.1	0.5	328	334
10DX140-3	Qz5 _b	24.8	57.4	503	182	<81.74	7.28	0.225	v.d.	<0.28	0.580	0.879	3.76	3.65	0.687	0.538	1.19	v.d.	0.56	898	0.1	0.5	350	357
10DX140-4	Qz5 _b	22.8	81.8	607	41.5	622	1.52	0.740	4.38	<1.07	<3.88	0.280	7.03	16.0	0.694	0.230	13.6	0.92	0.56	1083	0.1	0.5	277	284
10DX140-5	Qz5 _b	15.3	v.d.	v.d.	169	$\frac{<139.9}{8}$	9.12	v.d.	v.d.	v.d.	v.d.	0.283	5.87	v.d.	0.526	$\frac{<0.08}{7}$	v.d.	v.d.	0.56	v.d.	0.1	0.5	361	369
10DX140-6	Qz5 _b	7.30	90.2	638	17.8	$\frac{<238.7}{4}$	7.55	v.d.	v.d.	v.d.	v.d.	v.d.	5.77	v.d.	1.02	$\frac{<0.17}{1}$	8.27	v.d.	0.56	1139	0.1	0.5	351	359
10DX140-8	Qz5 _b	19.0	60.7	221	13.9	<93.17	9.19	0.204	v.d.	0.740	1.80	0.264	1.97	2.66	v.d.	0.121	2.33	0.635	0.56	394	0.1	0.5	362	369
10DX140-9	Qz5 _b	17.0	11.1	218	<3.95	<66.17	4.35	0.0960	2.84	0.500	0.990	0.0220	1.89	0.880	0.285	0.155	0.681	0.227	0.56	388	0.1	0.5	324	331
10DX140-13	Qz5 _b	9.46	44.1	127	19.1	296	3.71	0.634	4.48	0.770	<0.38	0.048	2.36	1.84	v.d.	$\frac{0.090}{0}$	2.89	0.656	0.56	227	0.1	0.5	316	323
10DX140-14	Qz5 _b	27.2	119	$\frac{100}{2}$	143	90.5	13.6	1.05	v.d.	1.05	<0.00	0.215	8.45	v.d.	0.977	0.320	12.5	0.552	0.56	1789	0.1	0.5	383	391
10DX140-21	Qz5 _b	20.3	107	325	57.2	<43.63	10.4	v.d.	v.d.	0.530	v.d.	0.0200	3.53	1.32	0.981	0.112	3.50	0.897	0.56	580	0.1	0.5	368	376
10DX140-22	Qz5 _b	10.4	140	180	62.2	114	19.8	v.d.	v.d.	v.d.	v.d.	0.214	2.30	1.08	1.37	0.111	0.918	1.68	0.56	322	0.1	0.5	405	413

10DX140-34	Qz5 _b	26.1	131	608	86.7	v.d.	21.8	v.d.	v.d.	0.435	<0.00	0.137	2.92	v.d.	v.d.	0.266	2.30	v.d.	0.56	1086	0.1	0.5	411	419
10DX140-37	Qz5 _b	29.4	52.5	375	42.8	317	11.2	0.757	v.d.	0.830	<0.95	0.202	2.94	1.64	1.56	^{<0.06} ₃	1.97	0.589	0.56	670	0.1	0.5	372	380
10DX140-43	Qz5 _b	30.6	29.5	v.d.	v.d.	256	11.2	1.62	v.d.	0.720	<0.53	0.0840	2.14	1.48	v.d.	0.369	1.18	1.22	0.56	v.d.	0.1	0.5	243	249
10DX140-44	Qz5 _b	7.24	64.1	98.2	25.2	<54.86	8.12	v.d.	v.d.	v.d.	3.10	^{<0.013} ₁	2.31	1.28	0.62	0.154	0.761	0.995	0.56	175	0.1	0.5	264	270
10DX140-46	Qz5 _b	46.8	73.5	410	29.7	v.d.	14.9	0.0480	v.d.	^{<0.15} ₅	<0.00	0.068	2.62	0.350	0.359	^{0.050} ₀	0.637	0.752	0.56	731	0.1	0.5	373	380
10DX140-47	Qz5 _b	23.4	48.1	334	17.2	<67.59	9.10	<0.031	v.d.	0.343	1.14	0.254	6.24	2.17	v.d.	0.156	0.900	0.571	0.56	597	0.1	0.5	355	362
10DX140-48	Qz5 _b	23.7	70.0	364	17.8	<67.44	12.0	0.0510	5.10	^{<0.14} ₆	<0.55	<0.00	3.46	3.34	1.75	0.165	0.950	0.865	0.56	651	0.1	0.5	389	396
10DX145-4	Qz5 _b	40.3	99.4	¹¹⁹ ₀	31.2	220	3.21	0.0600	1.19	0.850	<0.64	1.17	4.65	5.40	0.275	0.084	v.d.	0.139	0.62	1919	0.1	0.5	361	369
10DX145-5	Qz5 _b	20.2	31.2	527	38.9	^{<147.6} ₆	2.61	<0.101	2.11	0.870	2.01	0.118	3.06	1.10	0.277	^{<0.09} ₈	2.52	<0.097	0.62	851	0.1	0.5	376	384
10DX145-12	Qz5 _b	28.8	41.3	868	241	<48.18	2.87	0.304	6.85	0.780	0.410	0.229	3.57	2.19	0.073	0.427	2.02	0.082	0.62	1399	0.1	0.5	310	316
10DX145-14	Qz5 _b	36.3	93.2	v.d.	v.d.	108	8.58	0.822	v.d.	1.75	1.04	0.159	4.01	2.40	0.431	0.176	2.00	<0.055	0.62	v.d.	0.1	0.5	300	307
10DX145-15	Qz5 _b	49.5	69.3	¹¹² ₈	17.4	<99.11	8.63	0.109	3.36	<0.22	0.890	0.563	2.50	2.57	^{0.048} ₁	0.152	2.98	0.178	0.62	1819	0.1	0.5	305	311
10DX145-17	Qz5 _b	45.6	95.4	937	22.0	<64.58	5.99	1.86	7.14	0.880	<0.44	0.214	4.44	4.9	0.195	0.077	2.39	0.071	0.62	1512	0.1	0.5	358	365
10DX145-22	Qz5 _b	28.4	12.4	457	7.88	75.2	2.84	0.330	6.18	<0.15	<0.75	0.348	4.91	4.13	0.055	0.072	3.13	0.385	0.62	736	0.1	0.5	358	366
10DX145-26	Qz5 _b	26.1	52.9	392	11.3	<72.40	4.30	2.02	v.d.	1.33	<0.50	0.374	4.27	2.33	0.832	^{<0.04} ₇	2.56	0.336	0.62	632	0.1	0.5	340	347
10DX145-31	Qz5 _b	30.2	^{<3.7} ₇	634	<6.78	^{<123.3} ₀	3.68	<0.091	1.42	<0.30	<0.75	0.407	2.52	2.53	^{<0.03} ₄	0.175	0.731	0.076	0.62	1023	0.1	0.5	304	311
10DX145-32	Qz5 _b	26.1	^{<3.8} ₈	522	<7.31	^{<137.5} ₁	2.39	<0.060	<0.23	0.690	0.580	0.400	2.60	0.730	1.30	0.145	0.457	<0.085	0.62	841	0.1	0.5	323	330
10DX145-34	Qz5 _b	13.1	49.0	217	13.4	143	6.67	0.114	2.45	0.917	<0.38	0.0402	1.85	0.410	0.335	0.078	0.257	0.042	0.62	350	0.1	0.5	316	323

10DX145-35	Qz5 b	28.7	9.82	691	140	<89.15	2.75	0.209	7.74	0.230	<0.68	0.144	3.17	1.93	0.161	0.747	1.31	0.098	0.62	1114	0.1	0.5	297	303
10DX145-36	Qz5 b	44.9	25.8	¹⁵⁵ ₉	328	<68.26	5.70	0.609	v.d.	0.470	<0.00	0.491	v.d.	v.d.	v.d.	0.169	v.d.	0.237	0.62	2514	0.1	0.5	345	352
10DX145-37	Qz5 b	51.6	68.5	¹⁴³ ₅	42.6	<57.96	10.9	0.199	2.21	0.910	^{<0.29} ₄	^{<0.009} ₅	3.64	1.28	0.417	^{0.094} ₀	3.68	0.09	0.62	2314	0.1	0.5	303	309
10DX145-38	Qz5 b	29.6	11.1	600	12.0	<80.65	3.70	<0.047	0.705	<0.17	1.35	0.0280	1.81	0.140	1.71	^{0.078} ₀	2.05	0.166	0.62	968	0.1	0.5	337	344
10DX145-39	Qz5 b	39.1	93.0	¹²⁰ ₂	251	106	4.38	0.129	v.d.	0.520	<0.85	0.726	3.31	0.190	v.d.	0.133	0.91	<0.068	0.62	1939	0.1	0.5	371	378
10DX145-45	Qz5 b	8.17	v.d.	415	120	v.d.	7.27	1.702	v.d.	1.58	<0.52	0.0520	2.87	1.09	v.d.	0.414	1.94	1.78	0.62	669	0.1	0.5	316	323
10DX158-5	Qz5 b	8.64	99.0	167	37.2	168	3.67	0.309	0.893	0.952	^{<0.31} ₄	0.0920	2.82	4.07	0.859	0.086	3.71	0.176	0.52	321	0.1	0.5	324	331
10DX158-7	Qz5 b	0.93	102	47.2	13.8	<60.44	3.15	0.120	0.368	0.599	<0.38	0.0562	2.78	3.01	v.d.	0.056	2.87	0.048	0.52	90.7	0.1	0.5	350	357
10DX158-21	Qz5 b	5.21	76.1	123	36.9	68.2	5.69	v.d.	v.d.	v.d.	<0.39	1.10	3.37	5.51	v.d.	0.242	7.29	0.766	0.52	236	0.1	0.5	316	323
10DX158-22	Qz5 b	6.19	78.3	100	32.8	120	3.07	1.06	2.35	0.986	<0.35	0.0182	3.03	2.11	1.35	^{<0.04} ₀	2.83	0.12	0.52	192	0.1	0.5	309	316
10DX164-2	Qz5 b	v.d.	42.3	v.d.	<4.18	<83.61	6.07	1.27	3.83	1.41	0.890	0.0200	2.50	1.02	0.338	^{<0.07} ₀	1.02	0.116	0.66	v.d.	0.1	0.5	337	344
10DX164-4	Qz5 b	23.0	37.7	173	19.2	118	6.51	0.0870	1.88	0.207	<0.38	0.118	2.50	1.55	0.377	0.084	2.02	0.112	0.66	263	0.1	0.5	308	314
10DX164-10	Qz5 b	17.5	11.5	101	v.d.	v.d.	14.3	0.2900	4.59	0.353	<1.27	<0.028	3.06	0.790	1.83	0.113	0.503	0.331	0.66	152	0.1	0.5	340	347
10DX164-11	Qz5 b	19.5	31.8	145	v.d.	v.d.	8.83	0.655	3.90	3.32	1.76	0.0490	2.51	1.38	2.41	0.182	0.323	0.346	0.66	219	0.1	0.5	344	351
10DX164-12	Qz5 b	29.0	v.d.	191	v.d.	107	13.7	1.41	4.32	1.29	<0.28	0.0490	3.46	2.37	1.8	^{0.087} ₀	2.18	0.296	0.66	290	0.1	0.5	386	394
10DX164-14	Qz5 b	10.7	126	158	31.5	v.d.	8.15	1.11	v.d.	1.30	v.d.	0.0220	2.88	4.81	v.d.	0.190	4.73	1.52	0.66	239	0.1	0.5	360	367
10DX164-15	Qz5 b	27.8	78.7	211	6.48	<85.47	7.51	<0.059	1.56	0.424	0.850	0.0730	4.37	3.07	v.d.	0.112	3.41	<0.056	0.66	319	0.1	0.5	384	391
10DX164-17	Qz5 b	12.4	62.6	115	29.8	v.d.	7.27	0.902	4.09	0.62	1.25	0.0200	1.90	<0.23	v.d.	^{<0.06} ₈	0.475	0.117	0.66	175	0.1	0.5	355	363

10DX164-21	Qz5 _b	11.3	18.4	108	11.5	101	5.42	<0.048	0.268	0.249	0.540	0.0084	1.47	0.57	v.d.	0.080 ₀	0.259	0.089	0.66	163	0.1	0.5	351	358
10DX164-24	Qz5 _b	14.1	27.1	126	26.9	<86.99	10.7	0.577	4.09	2.73	<1.01	0.183	2.02	<0.24 ₉	0.055	<0.07 ₄	0.272	0.160	0.66	190	0.1	0.5	350	357
10DX164-25	Qz5 _b	14.1	52.8	397	136	220	11.4	0.299	6.99	1.87	<1.31	0.260	3.13	3.49	1.25	<0.08 ₅	16.3	<0.077	0.66	601	0.1	0.5	335	342
10DX164-42	Qz5 _b	10.0	44.4	181	33.4	<145.6 ₉	6.95	0.221	6.91	1.36	0.550	0.0860	2.99	0.720	0.597	0.110	1.75	0.210	0.66	275	0.1	0.5	370	377
10DX164-43	Qz5 _b	24.5	145	v.d.	v.d.	v.d.	9.55	0.402	v.d.	v.d.	<0.00	1.43	3.91	3.91	v.d.	0.739	5.04	0.784	0.66	v.d.	0.1	0.5	373	381
10DX164-44	Qz5 _b	18.5	77.9	254	94.2	<113.2 ₂	6.42	0.972	v.d.	1.50	<1.15	0.194	2.52	1.49	0.284	0.095	1.20	0.379	0.66	385	0.1	0.5	347	354
10DX164-45	Qz5 _b	11.8	99.5	187	49.8	295	5.17	0.480	3.45	2.73	<0.00	0.106	3.07	0.350	0.512	<0.09 ₁	0.659	0.141	0.66	283	0.1	0.5	364	371
10DX164-46	Qz5 _b	13.8	141	166	68.7	v.d.	9.84	0.372	v.d.	v.d.	<0.37	0.500	3.39	2.96	v.d.	0.662	2.48	v.d.	0.66	251	0.1	0.5	343	350
10DX164-48	Qz5 _b	11.0	53.3	108	28.7	v.d.	4.75	0.157	4.26	0.436	<0.61	0.109	2.57	1.47	v.d.	0.109	1.70	1.90	0.66	163	0.1	0.5	332	339
10DX164-48A	Qz5 _b	13.9	32.6	151	53.5	<259.2 ₉	9.14	<0.173	0.790	0.700	1.66	<0.00	2.01	<0.79	0.097	<0.20	0.940	0.200	0.66	228	0.1	0.5	366	373
10DX164-49	Qz5 _b	14.0	85.9	91.9	6.96	<67.44	11.6	0.352	5.76	0.75	0.840	0.0970	3.85	1.92	0.755	0.226	1.11	0.115	0.66	139	0.1	0.5	328	335
10DX164-55	Qz5 _b	8.13	165	41.6	12.2	v.d.	3.60	0.575	v.d.	4.40	0.970	0.0500	2.05	1.58	v.d.	0.067 ₀	0.401	0.190	0.66	63.0	0.1	0.5	362	369
10DX164-56	Qz5 _b	16.0	76.6	97.2	23.4	<50.69	9.74	1.19	v.d.	5.76	<0.48	0.0490	2.96	2.09	0.234	0.099 ₀	3.56	0.411	0.66	147	0.1	0.5	375	382
10DX164-57	Qz5 _b	29.5	19.8	173	10.0	<58.72	13.8	0.237	1.52	<0.11 ₉	<0.51	0.0530	2.7	2.16	0.041 ₅	0.134	2.25	0.074	0.66	263	0.1	0.5	315	322
10DX220-29	Qz5 _b	21.9	33.2	139	9.55	v.d.	4.77	0.255	7.67	<0.27	<1.46	1.14	2.59	2.08	v.d.	1.57	1.03	v.d.	0.66	211	0.1	0.5	365	372
10DX220-30	Qz5 _b	30.1	40.9	421	<14.1 ₉	254	10.9	<0.172	2.69	0.980	2.66	1.00	3.41	2.39	1.64	0.450	0.590	v.d.	0.66	638	0.1	0.5	384	392
10DX201-12	Qz5 _b	54.4	160	112 ₈	581	<33.50	6.91	0.661	7.70	1.93	<0.30 ₉	2.34	9.99	45.6	0.691	0.152	64.0	0.168	0.61	1849	0.1	0.5	328	335
10DX201-13	Qz5 _b	29.5	106	255	57.2	78	5.42	0.159	2.17	1.27	0.91	0.153	4.71	4.45	1.15	0.130	5.72	0.294	0.61	418	0.1	0.5	371	378

10DX201-14	Qz5 _b	28.5	75.5	251	30.2	<73.36	5.61	0.0580	0.585	0.391	0.35	0.211	3.72	2.97	1.12	0.054 ₀	7.42	0.074	0.61	412	0.1	0.5	347	354
10DX140-39	Qz6 _a	26.6	164	943	206	151	0.63 ₀	0.964	v.d.	0.770	<0.55	0.262	17.8	83.3	v.d.	<0.04 ₃	106	0.46	0.56	1684	0.1	0.5	335	342
10DX140-40	Qz6 _a	56.4	276	152 ₃	375	158	1.09	0.404	5.17	v.d.	<0.46	0.0580	24.6	84.9	v.d.	0.130	109	0.913	0.56	2721	0.1	0.5	336	343
10DX140-71	Qz6 _a	0.38 ₂	16.8	81.2	33.9	<45.85	0.65 ₀	0.0400	0.903	<0.09 ₀	<0.40	1.167	12.3	98.0	0.057 ₅	0.139	107	0.046	0.56	145	0.01	0.5	209	239
10DX164-22	Qz6 _a	3.94	73.8	112	23.4	<95.32	3.50	v.d.	v.d.	v.d.	1.04	<0.00	1.77	1.48	v.d.	0.267	0.419	0.467	0.66	169	0.01	0.5	243	277
10DX164-1	Qz6 _a	4.11	41.7	68.9	16.8	528.3	3.64	<0.045	0.747	1.21	<0.59	0.112	2.48	0.290	0.381	0.144	0.749	0.111	0.66	104	0.01	0.5	312	355
10DX164-61	Qz6 _a	1.54	62.9	28.8	17.19	<163.7 ₈	5.75	0.712	4.75	3.68	<1.27	<0.040	1.73	<0.46	0.125	<0.12 ₀	0.590	<0.106	0.66	43.6	0.01	0.5	314	357
10DX164-62	Qz6 _a	1.40	104. ₄	40.5	<22.9 ₉	<428.2 ₅	4.82	1.70	v.d.	v.d.	2.63	<0.106	1.81	1.12	0.092	<0.29	<0.26	<0.24	0.66	61.3	0.01	0.5	336	382
10DX201-1	Qz6 _a	124	156. ₅	131 ₉	309	60.5	0.84 ₀	0.887	v.d.	v.d.	<0.00	0.0350	18.0	17.88	v.d.	0.067 ₀	27.8	0.202	0.61	2163	0.01	0.5	327	372
10DX201-4	Qz6 _a	184	147. ₇	115 ₃	155	<60.33	1.07	0.180	0.952	0.225	<0.33	0.0460	32.5	9.36	1.61	0.063 ₀	13.6	0.083	0.61	1890	0.01	0.5	253	288
10DX201-5	Qz6 _a	201	132. ₆	166 ₀	147	<105.5 ₈	1.15	0.424	1.89	0.390	0.89	<0.022	17.3	4.13	1.17	0.208	3.43	<0.038	0.61	2721	0.01	0.5	262	298
10DX201-6	Qz6 _a	245	121. ₇	133 ₆	106	172	0.65 ₀	0.386	v.d.	0.516	<0.49	1.249	14.5	2.61	v.d.	0.069	2.03	0.331	0.61	2191	0.01	0.5	265	301
10DX201-9	Qz6 _a	102	77.6	709	44.3	<91.39	2.42	0.0700	0.456	3.15	0.48	0.0460	13.3	7.15	0.557	0.125	18.1	0.117	0.61	1162	0.01	0.5	243	277
10DX201-10	Qz6 _a	46.9	78.4	287	87.3	v.d.	1.20	0.424	v.d.	2.51	<0.27	0.484	5.65	11.6	6.32	0.146	11.4	0.878	0.61	470	0.01	0.5	296	336
10DX201-11	Qz6 _a	15.0	54.2	155	27.6	<82.02	0.89 ₀	0.0860	0.371	0.710	<0.67	0.257	3.39	5.61	0.627	<0.06 ₄	7.95	0.17	0.61	254	0.01	0.5	266	303
10DX201-23	Qz6 _a	149	82.6	908	39.9	<74.56	2.05	<0.042	1.07	0.183	<0.43	0.0650	11.5	8.23	0.445	<0.05 ₇	28.6	0.072	0.61	1489	0.01	0.5	255	290
10DX201-24	Qz6 _a	108	114	957	165	<66.57	2.61	0.929	v.d.	0.663	<0.70	0.0760	12.6	10.6	1.63	<0.04 ₇	20.3	<0.020 ₁	0.61	1568	0.01	0.5	288	328
10DX201-26	Qz6 _a	17.5	41.7	125	<4.40	123	7.17	0.0860	1.35	3.59	<0.40	0.0700	2.34	1.53	0.537	0.122	2.39	0.172	0.61	204	0.01	0.5	299	340

10DX201-27	Qz6 _a	45.2	120	933	235	<66.20	7.80	0.688	4.72	0.578	<0.00	1.01	8.26	36.0	0.842	$\frac{<0.04}{6}$	50.0	0.223	0.61	1530	0.01	0.5	347	395
10DX201-30	Qz6 _a	10.6	48.0	119	13.5	87.63	2.78	0.158	1.08	1.26	<0.64	0.0150	3.08	2.1	0.758	0.066	2.96	0.138	0.61	195	0.01	0.5	352	400
10DX140-70	Qz6 _b	4.69	21.1	139	30.0	142	$\frac{0.24}{0}$	0.0970	0.883	0.218	<0.61	0.474	6.47	44.1	0.155	$\frac{0.096}{0}$	34.5	0.062	0.56	248	0.01	0.5	302	343
10DX158-6	Qz6 _b	14.2	31.9	229	16.0	<87.68	$\frac{0.34}{0}$	0.808	v.d.	0.890	1.02	<0.023	4.45	2.73	1.52	$\frac{<0.07}{1}$	4.50	0.126	0.52	441	0.01	0.5	221	252
10DX158-8	Qz6 _b	2.85	48.7	114	11.1	<73.62	$\frac{0.40}{0}$	0.226	0.57	0.399	<0.63	0.0750	2.62	2.24	0.983	0.073	3.14	0.036	0.52	219	0.01	0.5	226	258
10DX158-9	Qz6 _b	16.8	121	530	131	<81.87	$\frac{0.46}{0}$	0.463	1.27	0.570	0.700	0.0340	9.35	16.3	1.37	$\frac{0.098}{0}$	25.0	<0.023	0.52	1020	0.01	0.5	231	264
10DX158-10	Qz6 _b	42.1	107	566	124	<46.76	$\frac{0.73}{0}$	0.331	5.83	1.02	<0.44	$\frac{<0.007}{7}$	10.5	13.7	0.705	$\frac{0.070}{0}$	20.4	0.095	0.52	1089	0.01	0.5	247	282
10DX158-11	Qz6 _b	37.6	136	713	220	104	$\frac{0.63}{0}$	0.970	3.56	1.02	<0.47	0.0510	10.6	10.3	2.49	$\frac{<0.05}{3}$	12.5	0.116	0.52	1371	0.01	0.5	242	276
10DX158-17	Qz6 _b	31.3	101	361	81.2	127	2.18	0.476	6.42	0.610	0.960	0.392	4.72	6.09	0.483	$\frac{<0.07}{2}$	13.2	0.875	0.52	693	0.01	0.5	291	331
10DX158-23	Qz6 _b	1.67	63.7	55.1	15.8	87.7	$\frac{0.52}{0}$	0.350	4.22	0.512	<0.63	0.0760	2.27	1.66	1.29	0.100	2.23	0.03	0.52	106	0.01	0.5	235	268
10DX158-24	Qz6 _b	8.20	40.2	79.4	7.05	<64.56	1.27	0.054	0.294	$\frac{<0.13}{8}$	0.540	0.0399	2.85	1.21	0.227	$\frac{0.098}{0}$	2.54	0.08	0.52	153	0.01	0.5	269	306
10DX158-25	Qz6 _b	56.3	84.6	701	149	<42.59	$\frac{0.99}{0}$	0.126	0.637	0.131	$\frac{<0.27}{9}$	0.0590	12.4	6.94	v.d.	$\frac{0.092}{0}$	7.36	0.043	0.52	1349	0.01	0.5	259	295
10DX164-40	Qz6 _b	80.3	104	$\frac{122}{6}$	187	$\frac{<111.0}{1}$	1.47	0.215	2.96	1.30	1.92	<0.00	5.06	11.1	v.d.	0.188	22.2	0.720	0.66	1857	0.01	0.5	275	313
10DX164-41	Qz6 _b	132	126	$\frac{113}{9}$	278	141.7	1.03	0.0900	2.29	1.75	0.460	<0.00	21.6	13.6	v.d.	0.176	27.7	0.392	0.66	1726	0.01	0.5	260	297
10DX201-2	Qz6 _b	1.67	18.8	33.9	<3.50	<63.53	$\frac{0.60}{0}$	<0.033	$\frac{<0.11}{2}$	0.336	<0.68	<0.022	32.3	74.9	0.776	0.134	897	0.033	0.61	55.6	0.01	0.5	240	274
10DX201-3	Qz6 _b	7.39	47.7	105	19.7	<46.03	$\frac{0.46}{0}$	0.0630	0.303	$\frac{<0.11}{7}$	<0.36	0.096	19.2	66.0	0.887	$\frac{0.070}{0}$	875	0.031	0.61	173	0.01	0.5	231	264
10DX201-7	Qz6 _b	1.17	20.3	26.6	<2.62	<46.57	$\frac{0.53}{0}$	<0.024	0.339	0.362	<0.36	0.0560	11.5	18.8	0.157	$\frac{<0.03}{0}$	58.3	0.031	0.61	43.7	0.01	0.5	236	269
10DX201-8	Qz6 _b	1.10	14.9	26.9	<4.62	<79.68	$\frac{0.35}{0}$	<0.046	$\frac{<0.13}{7}$	$\frac{<0.15}{6}$	<0.68	0.0520	16.1	60.9	v.d.	$\frac{<0.07}{1}$	172	0.029	0.61	44.1	0.01	0.5	222	253

10DX201-19	Qz6 _b	3.06	25.9	102	18.5	<51.18	$\frac{0.43}{0}$	<0.027	0.279	$\frac{<0.09}{8}$	0.390	0.289	13.4	47.4	0.275	0.043	232	$\frac{<0.018}{9}$	0.61	167	0.01	0.5	229	261
10DX201-20	Qz6 _b	8.32	17.9	92.5	12.4	96.72	$\frac{0.37}{0}$	0.024	0.172	0.189	<0.31	0.133	13.7	50.3	0.127	0.076	285	0.02	0.61	152	0.01	0.5	223	255
10DX201-21	Qz6 _b	1.47	24.4	23.7	<2.79	<50.10	$\frac{0.46}{0}$	$\frac{<0.021}{9}$	$\frac{<0.08}{6}$	0.109	<0.29	0.205	22.0	79.4	$\frac{0.040}{9}$	0.099	279	$\frac{<0.022}{9}$	0.61	38.8	0.01	0.5	231	264
10DX201-22	Qz6 _b	3.19	57.2	143	13.2	<60.40	$\frac{0.49}{0}$	0.234	5.02	0.160	<0.47	0.350	30.9	81.4	0.374	0.066	877	0.052	0.61	234	0.01	0.5	233	266
10DX201-25	Qz6 _b	126	196	$\frac{209}{6}$	656	<44.32	$\frac{0.73}{0}$	v.d.	v.d.	1.47	0.840	0.0990	29.5	53.8	v.d.	0.122	333	0.878	0.61	3436	0.01	0.5	247	282

v.d.: value discarded; values below detection limit are denoted with a "<" sign.

5.4. Temperature Estimation

At a chemical equilibrium, the incorporation of Ti in quartz is dependent on pressure and temperature. The dependency has been devised as a geothermometer known as TitaniQ by various experiments [23,26,27,75,76]. Different calibration equations have been derived [77], but one developed by [26], as shown below, was selected in this study based on the arguments presented in [76]. An overview of the arguments is out of the scope of this study; interested readers are referred to the publications mentioned above.

$$RT\ln X_{\text{TiO}_2}^{\text{quartz}} = -60952(\pm 3122) + 1.52(\pm 0.04) \times T - 1741(\pm 63) \times P + RT\ln a_{\text{TiO}_2}$$

At a given Ti activity (a_{TiO_2}), the mole fraction of Ti in quartz ($X_{\text{TiO}_2}^{\text{quartz}}$) increases with T but decreases with P. In a P-T space, equal $X_{\text{TiO}_2}^{\text{quartz}}$ (isopleth) defines vertical lines that have increasing slopes (Figure 10). For the Tongchang deposit, $a_{\text{TiO}_2} = 1$ can be assumed for Qz2-, Qz3- and Qz4- and Qz5-depositing fluids since the rutile was coprecipitated. For Qz1 and Qz6, $a_{\text{TiO}_2} = 0.5$ and 1, are used to calculate two isopleths for each $X_{\text{TiO}_2}^{\text{quartz}}$.

Exact pressures and temperatures can be obtained by combining TitaniQ and other geothermobarometers. Fluid inclusion isochores are an option but, unfortunately, they are only available for Qz3 and Qz4. B15HS inclusions in Qz1 and Qz2 are not used since their origin remains unclear [19] and their isochores intersect with TitaniQ isopleths at unreasonably high pressures (Figure 10). For the remaining quartz generations, a range of geologically reasonable pressures are assumed, i.e., 1500–500 bars for Qz1, Qz2, and Qz3a, 500–100 bars for Qz5, and 500–10 bars for Qz6. The assumption of decreasing pressures is supported by the increasing Al and Li contents consistent with a recent study by [78]. At a constant pressure, temperature is negatively related to a_{TiO_2} . The minimum temperatures for Qz1 and Qz6 are calculated with $a_{\text{TiO}_2} = 1$ and minimum pressure, and maximum temperatures are calculated with $a_{\text{TiO}_2} = 0.5$ and maximum pressure. Accordingly, the following averaged temperature ranges are obtained for different quartz generations (Figure 10; Table 4): Qz1 425–495 °C; Qz2a 529–618 °C; Qz2b 473–551 °C; Qz3a 374–420; Qz3b 350–450 °C; Qz4 310–390 °C; Qz5a 356–385 °C; Qz5b 344–351 °C; Qz6a 287–327 °C; Qz6b 242–276 °C.

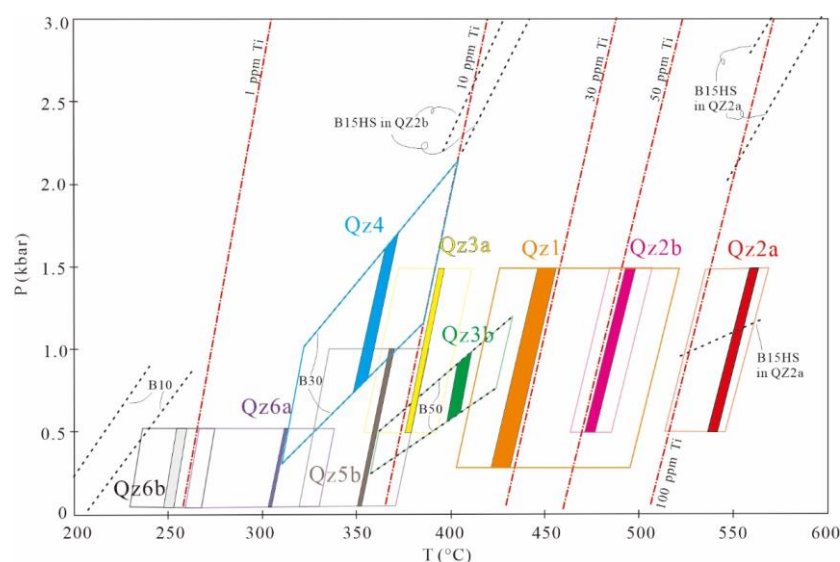


Figure 10. Pressure and temperature estimates by combining TitaniQ isopleths (dash-dotted lines in red) and fluid inclusion isochores (dash lines in black). Ti isopleths of 1 ppm, 10 ppm, 30 ppm, 50 ppm, and 100 ppm are shown; isochores of B15HS inclusions in both Qz2a and Qz2b, B50 inclusions in Qz3b, B30 inclusions in Qz4, and B10 inclusions are shown. The minimum and maximum range (solid open diamonds in colors) as well as an averaged range (solid color-filled) of Ti contents for all quartz generations are shown. Orange indicates Qz1, red indicates Qz2a, pink indicates Qz2b, yellow indicates Qz3a, green indicates Qz3b, blue indicates Qz4, dark brown indicates Qz5b, purple

indicates Qz6a, and grey indicates Qz6b. Note that only Qz3a and Qz4 had reliable fluid inclusion isochore data that allowed for the use of the interception method. For other quartz generations, a pressure range was assumed to produce a reliable temperature estimate.

6. Discussion

6.1. An Improved Timeline Based on Quartz CL Sequencing

Vein crosscut relations are often used to sort fluid events in the porphyry deposits [8]. However, this technique could be imparted by the existence of reopened veins. In those cases, a detailed look into mineral microtextures could help construct a reliable timeline of superimposed fluid events.

CL microtextures in hydrothermal quartz falls in two general categories: primary crystallization textures and secondary modification textures attributable to quartz deposition, dissolution, and deformation processes [79,80]. Concentric and euhedral oscillatory zoning in the Qz2a, Qz3, Qz5, and Qz6 grains belong to the primary crystallization textures, indicating a direct deposition from hydrothermal fluids. Mottled CL in the Qz1 and Qz2b grains is indicative of fluid-aided dynamic recrystallization. Similar Ti contents in both quartz generations might indicate a common event of recrystallization. The seemingly contradictory coexistence of un-recrystallized Qz2a and recrystallized Qz2b is likely linked to a difference in crystal sizes and shapes and resultant preferential deformation, as proposed for the Butte porphyry Cu-Mo deposit [22]. Qz2b resorbing Qz2a, “splatter” and “cobweb” in the Qz1 and Qz2 grains as well as overgrowth of Qz3 on earlier quartz grains indicate quartz dissolution. Qz3 and Qz4 fragments and fracture-filling Qz5 grains are indicative of brittle fracturing.

Based on these pieces of textural evidence, a refined timeline of fluid evolution is proposed for the Tongchang deposit. The early stage involved the deposition of the Qz1 precursor, Qz2a, and the Qz2b precursor followed by dynamic recrystallization and a first-time dissolution. The subsequent intermediate stage started with dissolution of earlier quartz grains followed by progressive deposition of Qz3a, Qz3b, and Qz4. The late stage commenced with the fracturing and brecciation of earlier quartz grains followed by the progressive deposition of Qz5 and Qz6.

Petrographic observations suggest that the early Qz1- and Qz2-depositing stage also precipitated hematite and anhydrite (Anh1) in the veins, and bornite, chalcopryrite, and potassic silicates in the alteration halos. The early intermediate stage precipitated pyrite, chalcopryrite, anhydrite (Anh3), chlorite, and sericite, along with Qz3 quartz, as well as chlorite-sericite alteration halos. The late intermediate stage precipitated abundant chalcopryrite, pyrite, anhydrite (Anh4), chlorite, sericite, and calcite in the veins. The late stages generated pyrite and sericite with Qz5 and Qz6 grains.

6.2. Fluid Cooling and Compression during Quartz Deposition, Recrystallization, and Dissolution

6.2.1. Quartz Deposition

TitaniQ thermobarometry has been widely utilized to estimate the P-T conditions of quartz deposition in magmatic [81,82] and hydrothermal environments [22,74,83]. Quartz solubility in various fluid compositions at varied P-T conditions can be used to investigate the fluid processes responsible for quartz deposition and dissolution, and, by inference, vein formation processes [31,32,84].

According to TitaniQ temperatures, early quartz deposition (Qz2a) occurred at around 529–618 °C; early intermediate quartz deposition (Qz3a and Qz3b) occurred at around 350–450 °C; late intermediate quartz deposition (Qz4) occurred at around 310–400 °C; and late quartz deposition (Qz5 and Qz6) occurred at around 240–400 °C. On a P-T phase diagram for H₂O-NaCl overlapped with quartz solubility curves (Figure 11), the quartz deposition temperature ranges are located in the single liquid phase region where quartz solubility decreases dramatically with temperature, and much less with pressure.

It is inferred that the quartz deposition is mainly the result of fluid cooling. Anhydrite has a similar prograde solubility behavior at temperatures above 350 °C in H₂O-10wt.%NaCl solutions, but shows retrograde solubility below that temperature (the grey region in the Figure 11). Increasing salinity tends to decrease the deflection point of the temperature, thus shifting the retrograde region to lower temperatures. H₂O-NaCl fluids containing silica and calcium sulfate would precipitate quartz and anhydrite upon cooling to 350 °C. Below 350 °C, quartz would continue to precipitate, whereas anhydrite deposition would cease. This inference is consistent with the coprecipitation of quartz and anhydrite in the early to intermediate stage and the disappearance of anhydrite in the late stage in the Tongchang deposit.

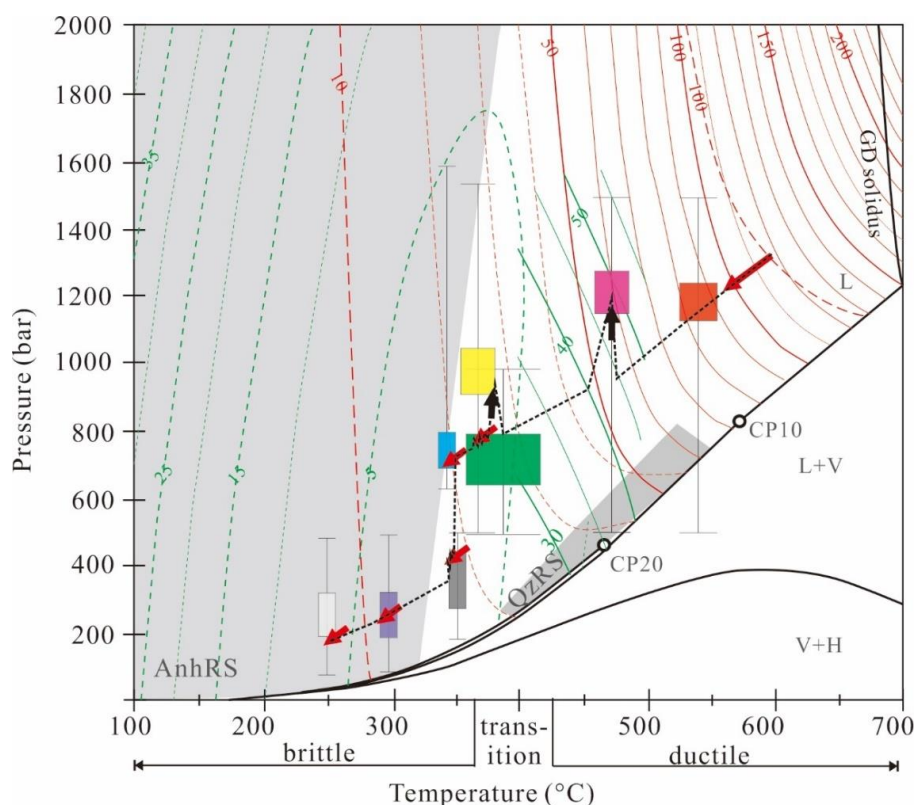


Figure 11. P-T phase diagrams for H₂O-10 wt.% NaCl and H₂O-20 wt.% NaCl aqueous systems, contoured with quartz and anhydrite solubilities (red: quartz; green: anhydrite; solid curve: H₂O-20 wt.% NaCl; dashed curve: H₂O-10 wt.% NaCl. Anhydrite data are from [85]; and quartz data are from [31]. The two areas in grey color are the retrograde solubility regions of anhydrite (AnhRS) and quartz (QzRS). The color-filled rectangles accompanied with error bars represent the deposition conditions of quartz and anhydrite (colors are the same as the in the Figure 10). The thick arrows (decompression cooling in red; quasi-isothermal compression in black) and connecting dash lines represent the possible P-T paths of the hydrothermal ore fluids. Note that the quasi-isothermal fluid overpressure occurred prior to precipitation of Qz3/Anh2, which was followed by pressure fluctuations. Pressure fluctuations also happened during the formation of Qz5. Solubility values (in mm/kgH₂O) are marked on the curves. GD solidus: H₂O-saturated granodioritic melt solidus; CP10/20: critical point for H₂O-10 wt.% NaCl or H₂O-20 wt.% NaCl. L: liquid; V: vapor; H: halite.

6.2.2. Quartz Recrystallization

Like many other minerals, quartz undergoes dynamic recrystallization during deformation to minimize free strain energy [86]. Three distinctive recrystallization mechanisms have been recognized [87], i.e., bulging recrystallization (BLG), subgrain rotation (SGR), and grain boundary migration (GBM). These mechanisms occur at different temperature regimes and result in characteristic microstructures [22]. BLG occurs at temperatures between 250 °C and 400 °C, resulting in small, recrystallized grains surrounding large, un-

recrystallized grains. SGR occurs at temperatures between 400 °C and 500 °C, generating straight and commonly polygonal grain boundaries. GBM occurs at temperatures between 500 °C and 600 °C, producing lobate grain boundaries. TitaniQ applied in a metamorphic environment [88] and experiments [76,89,90] suggests that SGR and GBM could re-equilibrate Ti in quartz, and thus that TitaniQ would record the temperature of recrystallization.

The grain boundaries shown in Qz1 and Qz2b grains are similar to those of SGR (Figure 6a,b). TitaniQ temperatures of 400–550 °C suggest SGR/GBM recrystallization consistent with the microstructures. A noteworthy feature is that less-recrystallized Qz2a grains are commonly rimmed by CL-dark edges. A similar feature in B-type veins of the Butte porphyry deposit was proposed to be a result of mismatch in lattice alignment of adjacent crystals during competitive grain growth [22]. However, we noted that the CL-dark rims in the Qz2a grains are comparable to those quartz experimentally deformed with the presence of fluid [76]. We infer that the CL-dark rims in Qz2 (Figure 6b) may have been a product of fluid-mediated dissolution-precipitation during dynamic recrystallization. If this inference is correct, the quartz dissolution event requires an increase in silica solubility, which cannot be explained by a drop in fluid pressure in a deflating magma chamber as proposed by [22]. Conversely, we consider it more likely to reflect an increase in fluid pressure and/or temperature based on a solubility analysis (Figure 11). Fluid compression is a plausible process in that overpressure in the overlying carapace that could be caused by fluid exsolution and accumulation, which is common for the deposition of early quartz veins [91]. In addition, silica solubility can be enhanced by new pulses of mafic recharge, which may introduce hotter fluids. A possible magmatic recharge was suggested by a geochemical study of the mafic enclaves of the Tongchang deposit [20].

6.2.3. Quartz Dissolution and Fracturing

Immediately after the deposition of Qz2b, a quartz- and anhydrite-dissolution took place prior to the deposition of Qz3 and Anh2 at lower temperatures (350–450 °C) and fluctuating pressures. The temperature constraints indicate a transition from lithostatic to hydrostatic regimes (Figure 11). The quartz dissolution was previously attributed to quasi-isobaric cooling through the retrograde quartz solubility region [17,31,92]. However, this explanation is not likely for Tongchang because quasi-isobaric cooling would cause lower Ti contents in later deposited Qz3b, which is inconsistent with higher Ti contents in Qz3b (around 15 ppm) than in Qz3a (around 10 ppm). Another observation not in favor of quasi-isobaric cooling is that anhydrite displays prograde solubility in the quartz retrograde solubility field, which cannot account for the coeval dissolution of anhydrite and quartz (Figure 11). Therefore, it is inferred that fluid pressurization is seemingly the only viable way to elevate the solubility for both minerals.

Quartz and anhydrite show a similar solubility behavior at temperatures between 350 °C and 400 °C and pressures below 2000 bars, corresponding to a hydrostatic regime [92]. Under hydrostatic conditions, quartz and anhydrite deforms in a brittle fashion, which is consistent with early anhydrite breccia cemented by Anh4 (Figure 7h). The quartz micro-breccia enclosed in Qz5 suggests the occurrence of another episode of fluid overpressure. With the space created by brecciation, pressure fluctuated in response to fluid exsolution and mineral precipitation, resulting in the zoning patterns in Qz5b (Figure 11).

Preexisting veins can be reopened during a later fluid event by dissolution, mechanical breaking, brecciation, and recrystallization [17]. At Tongchang, these phenomena are abundant and occurred at different stages. Dissolution and recrystallization occurred in the early to transitional stage when host rock behaved in ductile fashion, whereas breaking and brecciation occurred in the late stage when host rocks behaved in a brittle fashion.

6.3. Vein Growth Via Reopening

Hydrothermal veins in porphyry Cu deposits are commonly grouped into early quartz (A-type), transitional molybdenite quartz (B-type), and late pyrite (D-type) veins

[93]. This model was later modified with the addition of several subordinate vein types, including early magnetite (M-type) [94], early biotitic (EB-type) [95], late chalcopyrite-sericite (C-type) [95,96] (Dilles et al., 1992, and epithermal (E-type) veins [97]. Although widely used, the model faces challenges with regard to vein overprints [12].

The Tongchang porphyry deposit presents a good example of such challenges. According to their deposition temperatures, Qz1-2, Qz3 and Qz5 can be considered as indigenous quartz of A-, B-, and D- type veins [31]. Three fourths of the examined veins contain quartz that precipitated earlier or later than the vein formation defined by vein classification (Figure 12). Visual estimates suggested that the early A veins, except the A.V_{kq} vein in 10DX197, contain younger quartz (Qz3 and Qz4) in volumes <50%. The B veins particularly involve the addition of younger Qz4 quartz to older quartz. For instance, the B.V_{macpq} vein in 12DX328 comprises a central lining of pyrite, chalcopyrite, Qz4, Anh4 and carbonates, which is sandwiched by Qz1 and Qz2 quartz. Overprints in late veins are dramatic, and are characterized by the presence of earlier relict quartz (Qz1-3) and minor younger Qz6 quartz. Qz4 quartz are typically absent. D.V_{mqp} is only different from D.V_{pd} veins in that it has a larger amount of relict quartz.

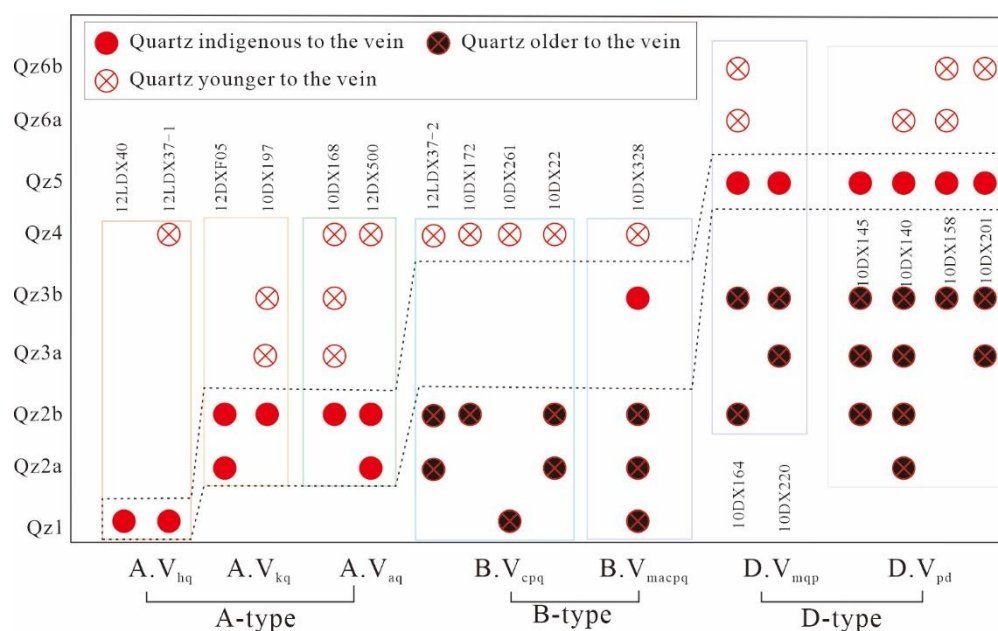


Figure 12. A qualitative estimation diagram of quartz types contained in A-type, B-type, and D-type veins examined in this study. All examined veins (17 in total) contained quartz younger or older than the vein formation, except two A-type veins. All vein contained early quartz (Qz1, Qz2a, and Qz2b), regardless of vein types. The presence of younger quartz in A-type veins indicates low-degree vein reopening, whereas the presence of older quartz in B-type and D-type veins indicates high-degree vein reopening. Note that there were two veins in the sample 12LXD37. The A.V_{hq} vein is named as 12LXD37-1, and the B.V_{cpq} vein is named 12LXD37-2.

A striking feature of the studied veins is that they all contain early quartz (Qz1-2), regardless of vein types. If these samples are representative of veining processes at Tongchang, it is plausible to infer that all the vein types are reopening early quartz-dominated veins. The present early veins are the product of slight overprints. The chalcopyrite veins (B.V_{cpq} and B.V_{macpq}) may be produced by the reopening of the earlier quartz veins by Qz4-depositing fluids, which is consistent with the conclusion of [31]. The late pyritic veins (D.V_{mqp} and D.V_{pd}) are products of overprints by Qz5-depositing fluids on earlier quartz veins.

6.4. Implications for Cu Sulfide Deposition

In the early fluid stage, hematite and anhydrite with little sulfide were deposited in the A.V_{hq} veins (Figure 13), attesting to oxidized sulfur species (SO_2 , SO_4^{2-}) at high temperatures [98,99]. In the potassic altered rocks, however, the low-sulfidation assemblage of chalcopyrite, bornite, magnetite and hematite developed [100]. These Fe-Cu sulfides were formed coevally with Qz1 and Qz2 at high temperatures (650 °C to 450 °C) (Figure 13), and the required H_2S was probably produced through the chemical reduction of sulfate in the fluids by mafic minerals such as biotite and hornblende [101]. It is worth mentioning that potassic ores of Tongchang are much less significant than in usual porphyry deposits [74], likely due to sulfide remobilization [1].

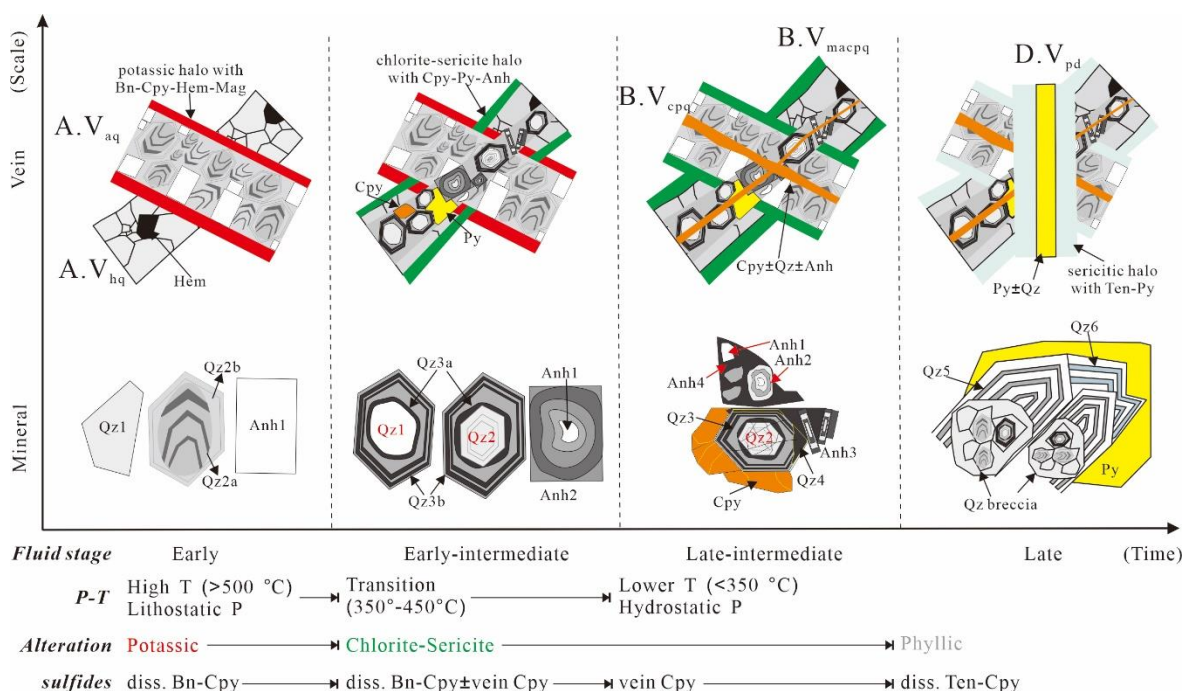


Figure 13. A mineral to vein scale model illustrating hydrothermal alterations, veins, and vein reopening over four fluid stages encompassed in the Tongchang porphyry Cu deposit. It also summarizes the temporal evolution of P-T conditions and sulfide assemblages. Hem: hematite; Bn: bornite; Cpy: chalcopyrite; Mag: magnetite; Qz: quartz; Anh: anhydrite; Py: pyrite; Ten: tennantite; diss: disseminated.

The first overprint is characterized by the deposition of pyrite ± chalcopyrite in the Qz3-Anh2 veins and chlorite-sericite alterations (Figure 13). The disappearance of ferric iron oxides and the appearance of pyrite indicate an increase in the sulfidation state with continuous cooling [100]. The H_2S required for sulfide deposition is likely derived from the SO_2 disproportionation favored by fluid cooling (350 °C to 450 °C) [98]. The remobilization of early-stage sulfides might not be a significant source of H_2S because chalcopyrite solubility is significantly impeded during this process [102]. In the alteration halos, the precipitation of anhydrite (Anh2) occurred coevally with sulfide deposition and the hydrolysis of feldspars, indicating a potential chemisorption reaction between SO_2 -bearing fluids and andesine [10].

Cu sulfides of the Tongchang deposit were mainly deposited in the vein assemblage of pyrite—chalcopyrite—Qz4—Anh4 at 300 °C–380 °C (Figure 13). The increased availability of Cu may be a result of the dissociation of CuCl_2^- complexes caused by fluid dilution in response to the entrainment of groundwaters and metamorphic fluids [103]. Another contributing factor might be the remobilization of existing sulfides associated with potassic alteration. Inflow of air saturated groundwater may significantly elevate the oxidation state and acidity of the magmatic fluids, which are favorable for chalcopyrite dissolution [102].

Although a large proportion of Cu sulfides are hosted in sericitic alteration zones, petrographic evidence of chalcopyrite inclusions and high-sulfidation assemblage of pyrite and tennantite indicates that they are remobilized from earlier ores (Figure 13). The mineral assemblage in the phyllic altered rocks reflects the stability of muscovite, quartz and pyrite, and the instability of chlorite, feldspars, mafic minerals, and chalcopyrite. This is similar to the Butte porphyry deposit, where the early ores (chalcopyrite-pyrite-magnetite) with potassic–chloritic alterations were remobilized by circulating oxidized groundwaters and re-deposited as chalcocite and covellite [104].

7. Conclusions

The Tongchang porphyry deposit represents a rare case of a porphyry copper deposit resulting from a single magmatic intrusion. It was expected that it would have undergone less fluid overprint compared to multi-intrusion counterparts. However, CL imaging of vein quartz and anhydrite revealed vein overprints as complicated (if not more complicated) as those in multi-intrusion deposits. This observation suggests that vein overprinting is likely an intrinsic property of porphyry deposits, regardless of intrusive history and other geological factors. The failure to recognize such vein overprints may result in erroneous fluid chronology and the incorrect interpretation of geochemical data (e.g., fluid inclusion and isotopes). Therefore, one should be careful when conducting geochemical analysis based on macroscopic features and a conventional vein model. Whenever possible, detailed CL imaging is an imperative.

As expected, the P–T history of Tongchang is characterized by an overall cooling and decompression trajectory. However, in contrast to monotonic cooling model, the P–T analysis of Tongchang porphyry deposit uncovered two significant episodes of fluid compression, likely corresponding to two periods of magmatic fluid accumulation under lithostatic and lithostatic to hydrostatic regimes. There might have been an increase in temperature during the second episode of compression. These irregularities in pressure should be considered for fluid models of porphyry deposits.

The Cu deposition process is better elucidated with the aid of CL petrography and integrated geothermometry. It was clearly shown that the high oxidation state of early high temperature fluids precludes Cu sulfide deposition in veins. However, Cu sulfide could be deposited via a chemical reaction with mafic minerals. The Cu sulfide deposition commences at temperatures lower than 450 °C when SO₂ disproportionation starts. Most chalcopyrite was precipitated at lower temperatures of 350 °C to 300 °C.

Author Contributions: X.L. and B.G.R. initiated the idea and work plan; X.L. performed the experiments; X.L., A.R., J.P. and B.G.R. interpreted the data; X.L. prepared the manuscript; A.R., J.P. and B.G.R. edited the manuscript. X.L. and J.P. acquired research funds. All authors have read and agreed to the published version of the manuscript.

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Conflicts of Interest: The authors declare no conflict of interest.

Appendix A

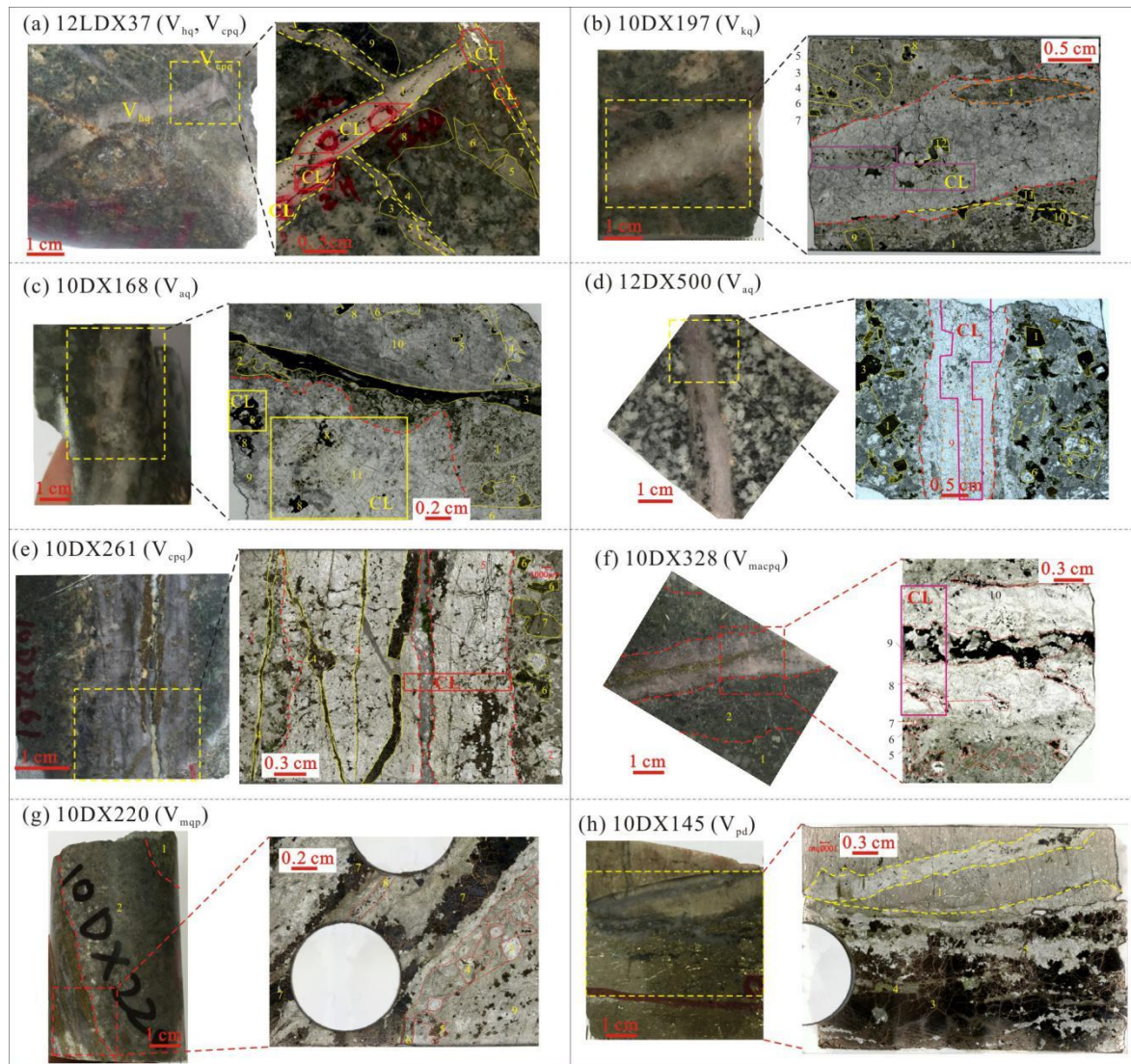


Figure A1. Petrography of studied samples. Selected vein samples and photos of the thick sections. (a) 1- hematite quartz vein; 2- muscovite chlorite chalcopyrite pyrite quartz vein; 3- chunk of chlorite, hematite, sericite and muscovite; 4- large chunk of hematite; 5- feldspar replaced by sericite; 6- residue quartz; 7- groundmass containing disseminated chalcopyrite; 8- groundmass containing abundant hematite and chalcopyrite; 9- green aggregation consisting of chlorite and sericite; (b) 1- K-feldspar partly altered by muscovite and chlorite; 2- sericite - chlorite pseudomorph after plagioclase; 3- chunk of sericite, chlorite, muscovite, chalcopyrite and rutile; 4- chunk of chlorite, chalcopyrite and rutile; 5- chalcopyrite with pyrite inclusions; 6- chunk of rutile, chalcopyrite and muscovite; 7- muscovite - chlorite - rutile - quartz after biotite; 8- chunk of muscovite, chalcopyrite, pyrite and rutile; 9- chlorite (center) - sericite (rim) after feldspar; 10- chunk of chlorite and rutile; 11- chunk of pyrite, chlorite, rutile and epidote; 12- chunk of chlorite, rutile and calcite; (c) 1- altered porphyry by chlorite, sericite, rutile and pyrite; 2- anhydrite, epidote and chlorite; 3- carbonate vein; 4- anhydrite epidote veinlet; 5- chunk of molybdenite; 6- K-feldspar partly altered by sericite, chlorite and epidote; 7- sericite pseudomorph after K-feldspar; 8- pyrite crosscut by chalcopyrite veinlets; 9- dirty quartz with abundant fluid inclusions; 10- clear quartz; 11- anhydrite; (d) 1-fresh biotite; 2- biotite slightly altered by chlorite; 3- biotite slightly altered by quartz, sericite, chlorite, hematite and magnetite; 4- chunks of secondary biotite; 5- plagioclase slightly altered by sericite; 6- residue quartz with overgrowth; 7- sericite pseudomorph after feldspar; 8- groundmass of small quartz, sericite and chlorite; 9- vein quartz and anhydrite; (e) 1- quartz vein; 2- altered porphyry; 3- chlorite hematite calcite vein; 4- chlorite muscovite chalcopyrite pyrite vein; 5- chalcopyrite pyrite chlorite quartz veinlets; 6- biotite replaced by chlorite sericite chalcopyrite and rutile; 7- feldspar completely

replaced by sericite and muscovite; (f) 1- slightly altered porphyry; 2- illitic chlorite - sericite altered porphyry; 3- sericite (rim)-chlorite (center) pseudomorph after feldspar; 4- chlorite-chalcopryrite pseudomorph after biotite; 5- sericite-rutile pseudomorph after biotite; 6- chunk of chalcopryrite, muscovite and anhydrite; 7- residue anhydrite; 8- anhydrite chalcopryrite veinlet; 9- anhydrite chalcopryrite pyrite vein; 10- chalcopryrite veinlet; (g) 1- slightly altered porphyry; 2- sericitic altered porphyry; 3- residue quartz; 4- chunk of muscovite and pyrite; 5- pyrite, muscovite and rutile pseudomorph after biotite; 6- muscovite, rutile, quartz and pyrite pseudomorph after biotite; 7- vein pyrite crosscut by carbonate veinlets; 8- quartz, muscovite, molybdenite, rutile layer; 9- fine-grained quartz and large muscovite in equal amount; (h) 1- silicified phyllite with oriented sericite, quartz and rutile; 2- pyrite quartz vein; 3- pyrite dominated vein; 4- aggregation of chlorite; 5- aggregation of quartz. Areas mapped with SEM-CL are also shown.

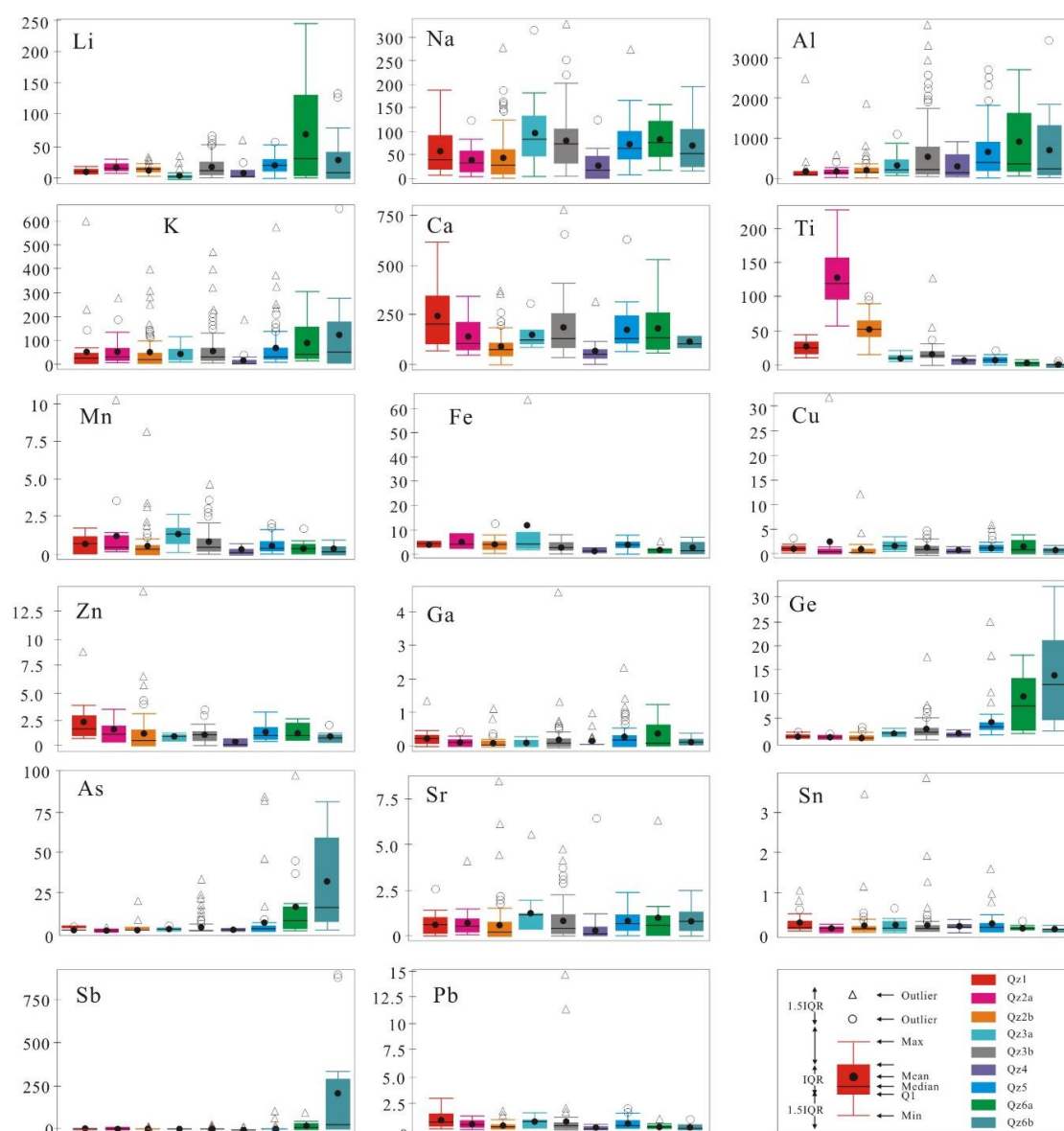


Figure A2. Tukey Boxplot of trace elements in vein quartz of the Tongchang porphyry Cu deposit.

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