

Article

Geochronology and Petrogenesis of the Early Paleozoic Jilongjie Granites in the Central South China Block: Implication for Post-Kinematic Lithospheric Delamination

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Abstract: Controversy over the geodynamic interpretation of the early Paleozoic granites in the South China Block constrains understanding of tectonic–magmatic evolution. In this paper, we present zircon U–Pb age, Hf isotope, and major and trace element data of the early Paleozoic granites in the Jilongjie region, south-central Hunan Province. A sample that yielded a weighted average $^{206}\text{Pb}/^{238}\text{U}$ age of 425 ± 3 Ma falls into the post-collisional granite field in the classification discriminant of magmatic rocks. Geochemical features indicate that the Jilongjie pluton is a shoshonitic metaluminous rock. The Jilongjie pluton's chondrite-normalized rare earth element patterns exhibit a slight enrichment of light rare earth elements (LREEs) relative to heavy rare earth elements (HREEs) with $(\text{La}/\text{Yb})_{\text{N}}$ ratios of 15.1–23.7 and weak Eu anomalies ($\text{Eu}/\text{Eu}^* = 0.68\text{--}0.78$). Zircon Hf isotope results show $\epsilon_{\text{Hf}}(t)$ ranging from -9.94 to -0.69 . Jilongjie granite's parent magma originated from a mixing of crust-derived felsic and mantle-derived mafic magmas, which then underwent fractional crystallization during its ascent. Jilongjie granite was generated through a post-collisional extensional setting associated with delamination of the thickened lithosphere.

Keywords: geochronology; petrogenesis; early Paleozoic; Jielongjie pluton; South China Block



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1. Introduction

The South China Block (SCB) was formed by the amalgamation of the Cathaysia Block in the southeast and the Yangtze Block in the northwest during the Neoproterozoic [1–5]. There is a significant amount of granite in the South China Block, which is generally considered one of the largest granite provinces in the world [6,7]. These granites are considered to have responded to the Caledonian, Indosinian, and Yanshanian tectonic events in the South China Block [7–20]. Caledonian massive granitic intrusions outcrop to the east of the Anhua–Luocheng Fault zone as batholiths and laccoliths, and are important components of the Phanerozoic granites in the eastern South China Block (Figure 1).

In recent years, many studies have been carried out on Indosinian and Yanshanian granites in the South China Block; Yanshanian granites are closely related to large-scale mineralization [21–25]. However, the petrogenesis and tectonic setting of Caledonian granites are still controversial (Table 1), with different viewpoints on tectonic settings, such as continental margin arc, oceanic–continental subduction, continental collision, and intracontinental orogeny.

In this paper, we present zircon U–Pb dating, Hf isotope results, and whole-rock geochemical data for the early Paleozoic granites in the south-central Hunan, central

South China Block, which have previously received much attention. In-depth study of the petrogenesis and tectonic setting provides information to further understanding of the SCB's early Paleozoic tectonic–magmatic evolution.

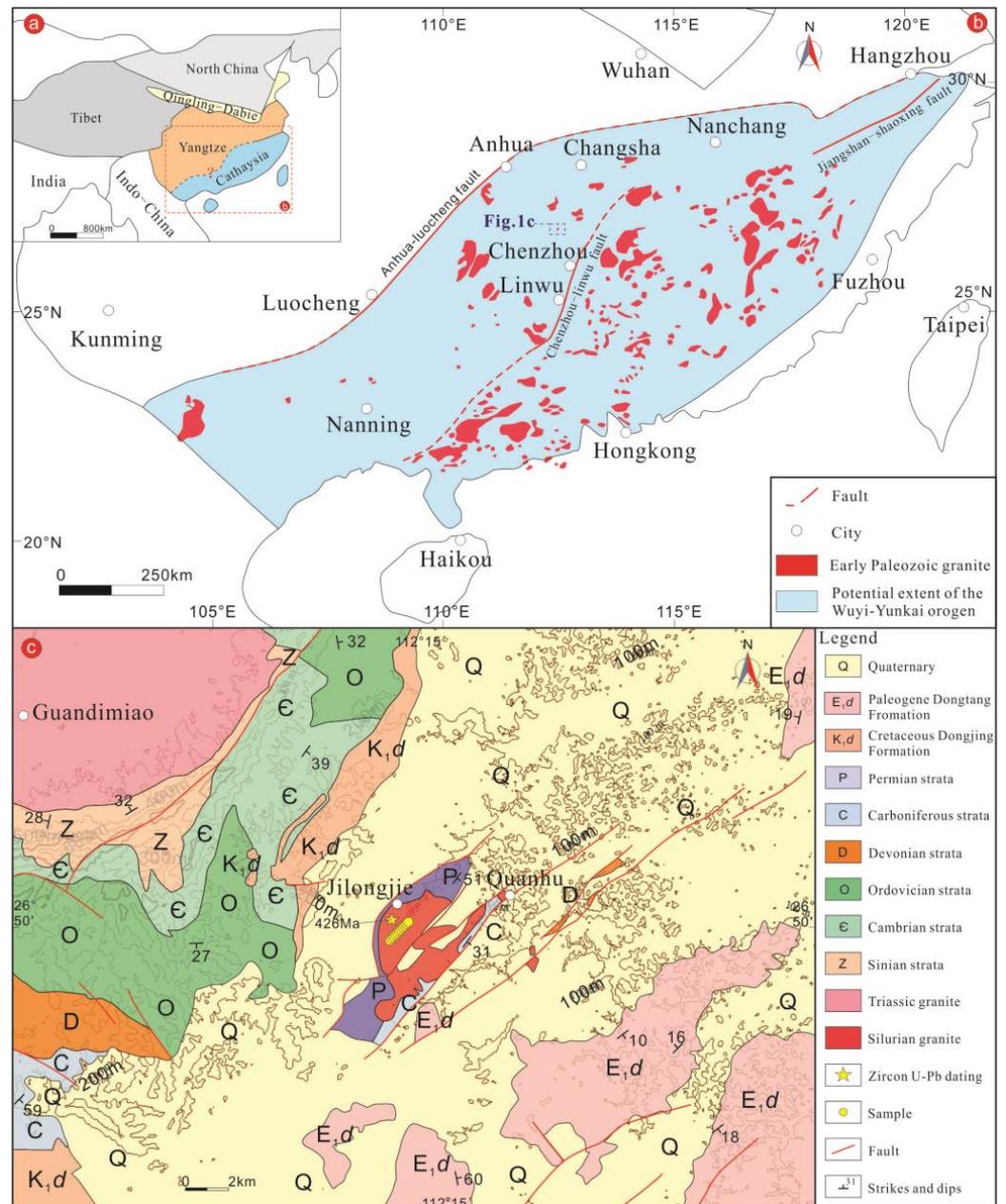


Figure 1. (a) Geotectonic map of China and its adjacent areas; (b) simplified geological map showing the distribution of the Early Middle Paleozoic granites in the eastern SBC (modified from Li, et al. [26] and Zhao, et al. [27]); (c) geological map of the Jilongjie area. Part of the data sources are shown in Table 1.

Table 1. Summary ages of the Early Middle Paleozoic rocks in SCB.

| Number | Locality | Lithology | Dating Method | Age (Ma) | Literature |
|--------|------------------------------|--|--------------------|---|------------|
| 1 | Tanghu Wangyangshan | Granite | Zircon U-Pb dating | 433 ± 2 | [28] |
| | | Granite | Zircon U-Pb dating | 434 | |
| 2 | Guidong Zhaiqian | Granodiorite | Zircon U-Pb dating | 425.5 ± 1.7 | [29] |
| | | Granite | Zircon U-Pb dating | 430.7 ± 1.9 | |
| 3 | Napeng | Granite | Zircon U-Pb dating | 418 ± 12 | [30] |
| 4 | Song Chay | Granite | Zircon U-Pb dating | 428 ± 5 | [31] |
| 5 | Zhuguang Songwang Dagu | Granite | Zircon U-Pb dating | 446.7 ± 6.3; 424.6 ± 3.7 | [32] |
| | | Foliated granite | | 440.7 ± 5.6 | |
| | | Granitic gneiss | | 421.9 ± 9.8 | |
| 6 | Yuntan Chidong Hebapu | Biotite orthogneiss | Zircon U-Pb dating | 427.1 ± 4.2 | [7] |
| | | Biotite paragneiss | | 423.0 ± 7.0 | |
| | | Granitic gneiss | | 429.6 ± 5.2 | |
| 7 | Weipu | Granite | Zircon U-Pb dating | 427.4 ± 4.0 | [33] |
| 8 | Northwestern Fujian | Granite | Zircon U-Pb dating | 437 ± 5; 437 ± 4; 440 ± 5; 441 ± 4 | [34] |
| 9 | Tianjingping | Granodiorite | Zircon U-Pb dating | 447 ± 2 | [35] |
| 10 | Sibao Weipu | Granite | Shrimp U-Pb zircon | 432 433 | [26] |
| 11 | Yunkai | Granite; gneissic granite | Zircon U-Pb dating | 430 ± 10; 443 ± 4; 437 ± 5 | [36] |
| 12 | Wugong domain | Gneissoid granite; orthogneiss; migmatite | | 455 ± 8; 455 ± 9; 456 ± 5; 443 ± 5; 424 ± 6; 452 ± 4 | [37] |
| | Northern Wuyi domain | Genissoid granite; orthogneiss; migmatite | Zircon U-Pb dating | 410 ± 10; 427 ± 15; 430 ± 9; 457 ± 6 | |
| | Southern Wuyi domain | Genissoid granite; orthogneiss | | 430 ± 6; 438 ± 3; 432 ± 6; 427 ± 4; 426 ± 6; 426 ± 8; 437 ± 3; 430 ± 6 | |
| 13 | Yunkai domain | Orthogneiss; gneissoid granite; leucosome in migmatite; paragneiss | | 450 ± 8; 449 ± 5; 443 ± 7; 415 ± 7; 435 ± 8; 452 ± 6 | [38] |
| | Wuping | Gneissic granite | | 496 ± 4; 494 ± 6 | |
| | Le'an | Biotite monzogranite | | 429 ± 2 | |
| | Zhangjiafang | Biotite monzogranite | | 440 ± 2 | |
| | Shuangzhuang | Monzogranite; biotite monzogranite | | 424 ± 4; 441 ± 3 | |
| | Pengongmiao | Monzogranite | | 405 ± 3 | |
| 14 | Wanyangshan | Monzogranite | Zircon U-Pb dating | 433 ± 4 | [39] |
| | Tanghu | Monzogranite | | 454 ± 2 | |
| | Banshanpu | Monzogranite | | 418 ± 2 | |
| | Hongxiaqiao | Granodiorite | | 432 ± 6 | |
| | Miaoershan | Monzogranite | | 400 ± 4; 415 ± 2 | |
| | Haiyangshan | Biotite monzogranite | | 429 ± 11 | |
| | Fengdingshan | Granodiorite | | 402 ± 2 | |

Table 1. Cont.

| Number | Locality | Lithology | Dating Method | Age (Ma) | Literature |
|--------|---------------------------|--|--|--|------------|
| 15 | Xuefengshan Belt | Granite | Zircon U-Pb dating | 428 ± 4; 438 ± 3; 437 ± 4; 411 ± 4; 412 ± 4; 424 ± 3 | [40] |
| 16 | Taishan | Granite | Zircon U-Pb dating | 436 ± 3; 436 ± 3; 436 ± 4; 436 ± 6; | [41] |
| 17 | Yuechenling Miaoershan | Granite | Zircon U-Pb dating | 435 ± 4; 427 ± 3; 417 ± 6 404 ± 4 | [27] |
| 18 | Northern Guangdong | Basalt; andesite; dacite, ignimbrite | Zircon U-Pb dating; Shrimp U-Pb dating | 435 ± 6; 435 ± 6 | [42] |
| 19 | Yunkai | Charnockite | Zircon U-Pb dating | 439 ± 2; 439 ± 4 | [43] |
| 20 | Doulong | Granite | Zircon U-Pb dating | 429 ± 3; 430 ± 3; 430 ± 2; 430 ± 2 | [44] |
| 21 | Shangmushui | Granodiorite | Zircon U-Pb dating | 444 ± 4 | [45] |
| 22 | Wanyangshan | Tonalite; granodiorite; monzonitic granite | Zircon U-Pb dating | 438 ± 3; 426 ± 3 | [46] |
| 23 | Daning | Granite | SHRIMP | 419.1 ± 6.4 | [47] |

2. Geological Setting and Petrography

The SCB consists of the Cathaysia Block in the southeast and the Yangtze Block in the northwest. The northeasterly trending Jiangshan–Shaoxing Fault is the present boundary between the Cathaysia Block and the Yangtze Block [25,39,48]. However, the southwestern extension is uncertain due to intensive younger tectonic modification and poor exposure.

The Cathaysia basement is considered to be predominantly composed of gneiss, schist, migmatite, amphibolite, and pyroclastic rocks from the Mesoproterozoic and Paleoproterozoic ages [14,49–51]. The Precambrian Cathaysia Block basement can be divided into the Nanling–Yunkai terrane in the southwest and the Wuyishan terrane in the northeast [52]. The oldest basement rocks in the Cathaysia Block are amphibolites (~1.80 Ga) distributed on the Wuyishan terrane [53]. Moreover, minor Mesoproterozoic granite (~1.43 Ga) was identified on Hainan Island in the south [54]. Recently, some Neoproterozoic mafic rocks were identified in the Cathaysian block [55]. The basement of the Yangtze Block is mainly composed of Proterozoic rocks, with minor Archean rocks, such as Kongling complex, dating to ca. 3.2 Ga [56–59]. Moreover, Neoproterozoic volcanic rocks appear around the Yangtze Block.

Samples for this study were collected in Jilongjie Town, 30 km southwest of Hengyang City, Hunan Province (N 26°50′17.81″; E 112°15′10.21″). The plutonic rock area is approximately 25 km²; host rocks are quartz sandstone and siltstone of the Permian Yangping and Leping Formation, as well as Carboniferous dolomite and dolomitic limestone. Plutonic rocks are in nonconformity contact with host rocks (Figure 2a). Exposed rock units in the study area also include Sinian slate, dolomite and limestone, Cambrian slate and phyllite, Ordovician phyllite and limestone, Devonian limestone, dolomite and mudstone, Carboniferous sandstone and siltstone, Cretaceous Dongjing Formation sandstone and conglomerate, Paleogene Dongtang Formation sandstone and sandy mudstone, and Quaternary sandy clay and sandy soil (Figure 1c).

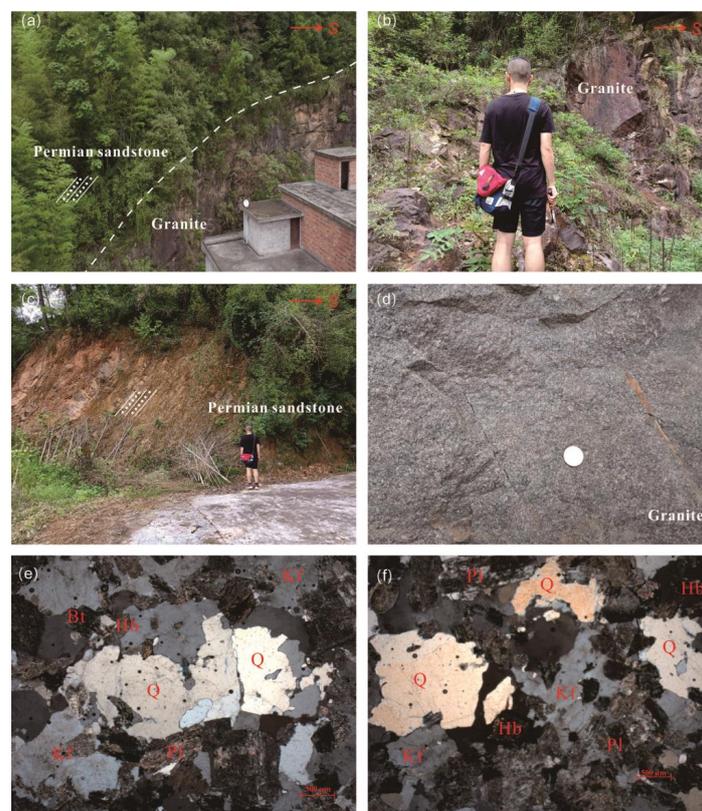


Figure 2. (a) Field contact relationship between rocks; (b) field photo of Jilongjie granite; (c) field photo of Permian sandstone; (d) hand specimen photo of Jilongjie granite; (e,f) photomicrographs of Jilongjie granite. Q—quartz; Kf—potassium feldspar; Pl—plagioclase; Bt—biotite; Hb—hornblende.

Jilongjie plutonic rocks are uniform in structure; their main minerals are K-feldspar (35%~40%), plagioclase (25%~30%), quartz (20%~30%), biotite (6%~10%), and amphibole (2%~4%); their accessory minerals include titanite, magnetite, and zircon (Figure 2e,f). Plagioclase occurs mainly as euhedral plate-like crystals, with albite twinning; K-feldspar and quartz are anhedral (Figure 2e,f).

3. Analytical Methods

Zircon U-Pb dating, major and trace element analyses of whole rock, and zircon Hf isotopic analyses were conducted at the Wuhan SampleSolution Analytical Technology Co., Ltd., Wuhan, China. Zircon U-Pb dating analysis was conducted using laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS). Detailed operating conditions for the LA system and ICP-MS instrument and data reduction were the same as described in [60]. Ion-signal intensities were acquired using an Agilent 7700e (Agilent Technology, Tokyo, Japan). In this study, the spot size was set to 32 μm and the laser frequency was set to 5 Hz. U-Pb dating and trace element calibration used zircon 91500 and glass NIST610 (Wuhan SampleSolution Analytical Technology Co., Ltd., Wuhan, China) as external standards, respectively. Off-line selection and integration of background and analyzed signals, time-drift correction, and quantitative calibration for trace element analysis and U-Pb dating were completed using ICPMSDataCal 12.2, an Excel-based software [61]. Concordia diagrams and weighted mean calculations of zircon samples were conducted using the Isoplot/Ex (version 3.0) program [62].

Major and trace element analyses of whole rock were performed using X-ray fluorescence (XRF, Rigaku, Japan) and ICP-MS (Agilent 7700e). Analytical uncertainties for major elements were generally <1 wt.%. Analytical results for AGV-2, BHVO-2, BCR-2, and RGM-2 international standards indicate that accuracies were better than 5% for most elements. The analytical procedure details were as described by Liu, et al. [63].

Hafnium isotope ratio analysis experiments were conducted in situ using a Neptune Plus multicollector (MC) ICP-MS (Thermo Fisher Scientific, Germany) in combination with a Geolas HD excimer ArF laser ablation system (Coherent, Göttingen, Germany). This laser ablation system includes a “wire” signal smoothing device that produces smooth signals even at very low laser repetition rates (as low as 1 Hz) [64]. Helium was the carrier gas within the ablation cell, and was merged with argon (makeup gas) after the ablation cell. Small amounts of nitrogen were added to the argon makeup gas flow to improve its sensitivity to Hf isotopes [65]. Compared to the standard arrangement, the addition of nitrogen in combination with the newly designed X skimmer cone and Jet sample cone in Neptune Plus improved the signal intensities of Hf, Yb, and Lu by factors of 5.3, 4.0, and 2.4, respectively. Detailed operating conditions for the laser ablation system, the MC-ICP-MS instrument, and the analytical method were the same as those described by Hu, et al. [66].

4. Results

4.1. Zircon U-Pb Geochronological Results

In this study, seventeen zircon grains in a sample (JL-U) from the Jilongjie pluton were selected for LA-ICP-MS dating. Zircon U-Pb data are listed in Table 2, and the concordia diagram is shown in Figure 3a. These zircon grains were euhedral, gray-white, or colorless. The crystals were elongated with lengths ranging from 120 to 280 μm and aspect ratios from 2:1 to 5:1. In cathodoluminescence, zircon crystals showed a clear and dense ring structure, indicating magmatic origin (Figure 3a). Zircon trace element data showed high Th/U values, which also indicate a magmatic crystallization origin [67]. Zircon trace element contents are listed in Table 3, and the chondrite-normalized REE diagram is shown in Figure 3b. Zircon crystals had similar characteristics to typical magmatic zircons [67], such as positive Ce and negative Eu anomalies, and HREE enrichment relative to LREE (Figure 3b).

Table 2. LA-ICP-MS results of zircon from the Jilongjie pluton.

| Samples and Anal. NO. | Th (ppm) | U (ppm) | Th/U Ratio | Isotopic Ratios ($\pm 1\sigma$) | | | | | | Ages ($\pm 1\sigma$ Ma) | | | Concordance | | | |
|-----------------------|----------|---------|------------|-----------------------------------|----------------------------------|----------------------------------|-----------------------------------|----------------------------------|----------------------------------|--------------------------|----|-----|-------------|-----|---|-----|
| | | | | $^{207}\text{Pb}/^{206}\text{Pb}$ | $^{207}\text{Pb}/^{235}\text{U}$ | $^{206}\text{Pb}/^{238}\text{U}$ | $^{207}\text{Pb}/^{206}\text{Pb}$ | $^{207}\text{Pb}/^{235}\text{U}$ | $^{206}\text{Pb}/^{238}\text{U}$ | | | | | | | |
| 1. JL-U-01 | 838 | 1047 | 0.80 | 0.0599 | 0.0013 | 0.5609 | 0.0110 | 0.0679 | 0.0006 | 611 | 51 | 452 | 7 | 423 | 4 | 93% |
| 2. JL-U-02 | 535 | 745 | 0.72 | 0.0602 | 0.0013 | 0.5762 | 0.0128 | 0.0692 | 0.0006 | 613 | 46 | 462 | 8 | 431 | 4 | 93% |
| 3. JL-U-03 | 963 | 1035 | 0.93 | 0.0572 | 0.0013 | 0.5332 | 0.0112 | 0.0676 | 0.0007 | 498 | 45 | 434 | 7 | 421 | 4 | 97% |
| 4. JL-U-04 | 1019 | 1029 | 0.99 | 0.0615 | 0.0012 | 0.5906 | 0.0119 | 0.0693 | 0.0007 | 657 | 44 | 471 | 8 | 432 | 4 | 91% |
| 5. JL-U-05 | 628 | 867 | 0.72 | 0.0612 | 0.0013 | 0.5828 | 0.0119 | 0.0687 | 0.0005 | 656 | 38 | 466 | 8 | 428 | 3 | 91% |
| 6. JL-U-06 | 575 | 772 | 0.75 | 0.0593 | 0.0012 | 0.5589 | 0.0107 | 0.0680 | 0.0005 | 589 | 10 | 451 | 7 | 424 | 3 | 93% |
| 7. JL-U-07 | 705 | 908 | 0.78 | 0.0520 | 0.0011 | 0.4865 | 0.0100 | 0.0679 | 0.0007 | 283 | 48 | 403 | 7 | 423 | 4 | 94% |
| 8. JL-U-08 | 824 | 973 | 0.85 | 0.0576 | 0.0012 | 0.5487 | 0.0115 | 0.0688 | 0.0006 | 522 | 44 | 444 | 8 | 429 | 3 | 96% |
| 9. JL-U-09 | 648 | 926 | 0.70 | 0.0566 | 0.0012 | 0.5334 | 0.0110 | 0.0681 | 0.0005 | 476 | 46 | 434 | 7 | 425 | 3 | 97% |
| 10. JL-U-10 | 1465 | 1032 | 1.42 | 0.0565 | 0.0013 | 0.5360 | 0.0120 | 0.0688 | 0.0006 | 478 | 52 | 436 | 8 | 429 | 4 | 98% |
| 11. JL-U-11 | 584 | 1001 | 0.58 | 0.0578 | 0.0012 | 0.5395 | 0.0108 | 0.0673 | 0.0005 | 524 | 44 | 438 | 7 | 420 | 3 | 95% |
| 12. JL-U-12 | 541 | 883 | 0.61 | 0.0570 | 0.0012 | 0.5332 | 0.0109 | 0.0676 | 0.0005 | 494 | 42 | 434 | 7 | 422 | 3 | 97% |
| 13. JL-U-13 | 710 | 951 | 0.75 | 0.0620 | 0.0012 | 0.5737 | 0.0106 | 0.0668 | 0.0005 | 676 | 43 | 460 | 7 | 417 | 3 | 90% |
| 14. JL-U-14 | 666 | 968 | 0.69 | 0.0579 | 0.0012 | 0.5563 | 0.0113 | 0.0693 | 0.0005 | 528 | 46 | 449 | 7 | 432 | 3 | 96% |
| 15. JL-U-15 | 610 | 859 | 0.71 | 0.0567 | 0.0013 | 0.5344 | 0.0115 | 0.0679 | 0.0005 | 480 | 48 | 435 | 8 | 424 | 3 | 97% |
| 16. JL-U-16 | 691 | 939 | 0.74 | 0.0565 | 0.0012 | 0.5409 | 0.0110 | 0.0691 | 0.0005 | 472 | 44 | 439 | 7 | 431 | 3 | 98% |
| 17. JL-U-17 | 1048 | 960 | 1.09 | 0.0605 | 0.0013 | 0.5807 | 0.0124 | 0.0693 | 0.0006 | 620 | 46 | 465 | 8 | 432 | 3 | 92% |

Table 3. Trace element (ppm) data of zircon from the Jilongjie pluton.

| Samples | La | Ce | Pr | Nd | Sm | Eu | Gd | Tb | Dy | Ho | Er | Tm | Yb | Lu | Hf |
|---------|------|------|-------|------|------|------|------|------|------|------|------|------|-----|------|--------|
| JL-U-01 | 14.4 | 69.1 | 3.93 | 16.7 | 6.05 | 2.02 | 17.8 | 5.62 | 64.7 | 25.2 | 123 | 27.6 | 269 | 57.4 | 11,975 |
| JL-U-02 | 4.49 | 43.6 | 2.15 | 11.3 | 5.94 | 1.92 | 14.2 | 4.13 | 44.9 | 16.3 | 74.6 | 16.2 | 159 | 35.1 | 11,808 |
| JL-U-03 | 113 | 296 | 29.2 | 117 | 19.8 | 4.28 | 23.8 | 6.17 | 65.7 | 24.0 | 113 | 24.8 | 251 | 55.4 | 10,905 |
| JL-U-04 | 0.14 | 36.0 | 0.26 | 3.04 | 3.76 | 1.18 | 14.1 | 4.64 | 52.0 | 19.5 | 93.2 | 20.3 | 201 | 43.7 | 12,041 |
| JL-U-05 | 54.2 | 156 | 16.6 | 73.8 | 12.8 | 1.82 | 15.8 | 3.66 | 42.0 | 15.8 | 74.9 | 17.6 | 182 | 41.0 | 12,484 |
| JL-U-06 | 16.0 | 63.7 | 5.14 | 23.6 | 5.41 | 1.31 | 14.0 | 4.11 | 47.6 | 17.5 | 83.7 | 18.8 | 188 | 42.1 | 11,647 |
| JL-U-07 | 5.48 | 42.2 | 1.99 | 9.89 | 4.10 | 0.72 | 11.9 | 3.50 | 43.5 | 17.0 | 84.4 | 18.7 | 192 | 43.8 | 11,190 |
| JL-U-08 | 26.1 | 88.4 | 7.05 | 29.0 | 6.72 | 1.51 | 12.3 | 3.77 | 42.7 | 16.8 | 81.0 | 18.6 | 189 | 41.9 | 11,752 |
| JL-U-09 | 7.3 | 26.6 | 0.042 | 0.97 | 2.05 | 0.60 | 9.93 | 3.25 | 39.5 | 16.1 | 78.9 | 18.3 | 189 | 43.4 | 12,680 |
| JL-U-10 | 1.79 | 54.3 | 2.52 | 16.6 | 11.5 | 3.81 | 25.3 | 6.83 | 73.2 | 25.4 | 113 | 24.5 | 233 | 50.1 | 10,947 |
| JL-U-11 | 31.8 | 113 | 10.3 | 45.7 | 9.2 | 1.55 | 17.2 | 4.77 | 58.1 | 24.2 | 126 | 29.8 | 313 | 71.2 | 12,653 |
| JL-U-12 | 2.14 | 32.9 | 0.85 | 4.54 | 2.77 | 0.61 | 11.1 | 3.64 | 47.0 | 19.4 | 98 | 23.2 | 243 | 56.0 | 12,367 |
| JL-U-13 | 22.7 | 82.2 | 6.88 | 29.6 | 6.70 | 1.36 | 13.0 | 3.75 | 41.5 | 16.1 | 78.7 | 18.2 | 186 | 41.8 | 12,362 |
| JL-U-14 | 2.84 | 36.6 | 1.03 | 5.79 | 3.13 | 0.89 | 11.4 | 3.91 | 48.0 | 20.1 | 98.8 | 22.9 | 240 | 54.4 | 12,462 |
| JL-U-15 | 16.6 | 73.5 | 5.98 | 28.6 | 6.43 | 1.39 | 13.6 | 4.14 | 47.3 | 18.9 | 90.3 | 21.1 | 216 | 49.8 | 11,889 |
| JL-U-16 | 23.3 | 91.3 | 9.15 | 44.7 | 9.49 | 1.51 | 16.3 | 4.37 | 46.4 | 17.5 | 84.3 | 19.1 | 191 | 43.1 | 12,040 |
| JL-U-17 | 7.15 | 56.7 | 3.10 | 15.8 | 6.88 | 2.70 | 19.0 | 5.70 | 61.0 | 21.3 | 96.2 | 20.8 | 203 | 43.5 | 10,285 |

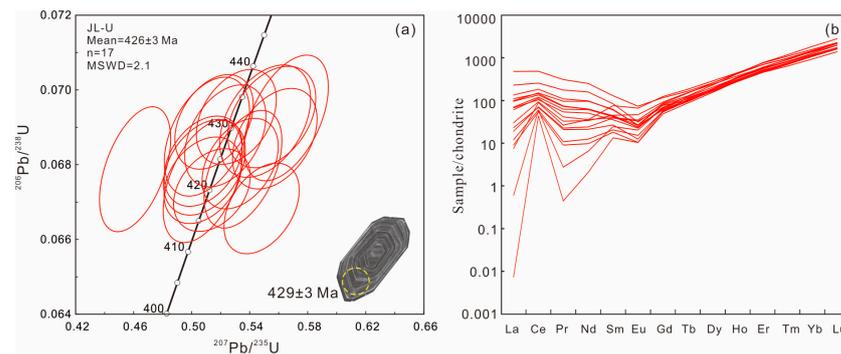


Figure 3. Cathodoluminescence (CL) and U-Pb concordia diagrams (a) and chondrite normalized REE pattern (b) of zircon crystals from the Jilongjie pluton.

Seventeen zircon grains had consistent or near-uniform $^{206}\text{Pb}/^{238}\text{U}$ ages ranging from 417 ± 3 Ma to 432 ± 3 Ma, with a weighted mean age of 426 ± 3 Ma (MSWD = 2.1, n = 17) (Figure 3a). The average age of Jilongjie plutonic rocks represents the crystallization age of the magmatic rocks, indicating that they were emplaced in the early Paleozoic.

4.2. Whole-Rock Geochemistry

Eleven representative samples were selected for whole-rock major element and trace element analyses (Table 3). Their loss on ignition (L.O.I.) range (1.83–3.04) suggests that Jilongjie plutonic rocks were insignificantly affected by alteration. Hence, we normalized the major element contents to 100% on a volatile-free basis. Jilongjie plutonic rocks had an SiO_2 content of 64.27–66.34 wt.%, K_2O content of 4.38–4.94 wt.%, Na_2O content of 2.50–2.93 wt.%, Al_2O_3 content of 12.95–14.31 wt.%, and total alkali ($\text{K}_2\text{O} + \text{Na}_2\text{O}$) content of 7.43–8.20. According to the total alkali–silica (TAS) diagram of igneous rocks proposed by Le Maitre [68], all Jilongjie plutons fall into the granite area (Figure 4a). In the $\text{K}_2\text{O} - \text{SiO}_2$ diagram, all samples fall into the shoshonite field [69]. The samples’ A/CNK ($\text{Al}_2\text{O}_3/(\text{CaO} + \text{Na}_2\text{O} + \text{K}_2\text{O})$) range was 0.80–1.04, and their A/NK ($\text{Al}_2\text{O}_3/(\text{Na}_2\text{O} + \text{K}_2\text{O})$) range was 1.28–1.53. Therefore, in the A/CNK – A/NK diagram, Jilongjie granites fall into the meta-aluminous to weakly peraluminous fields [70]. In conclusion, Jilongjie pluton is shoshonitic granite.

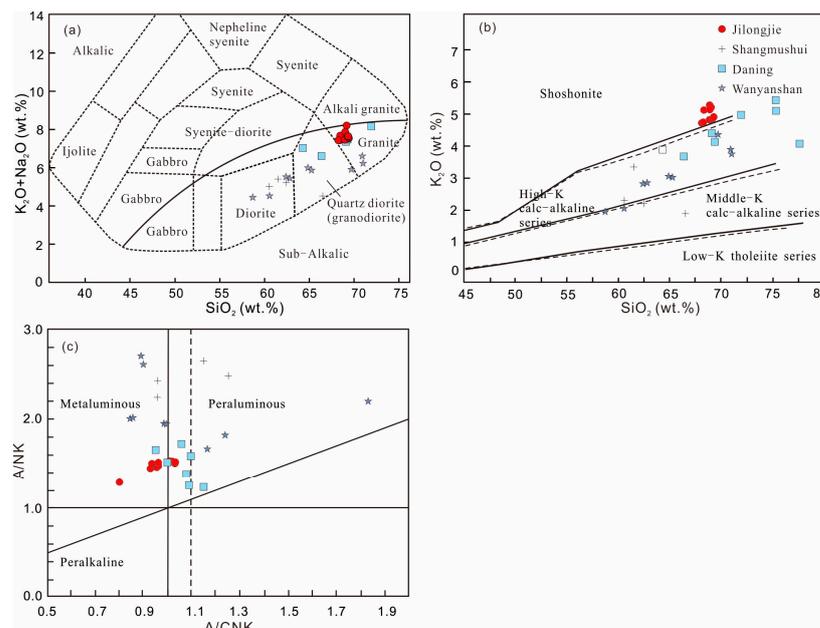


Figure 4. Geochemical characteristics of the Jilongjie pluton. (a) $\text{SiO}_2 - (\text{Na}_2\text{O} + \text{K}_2\text{O})$ diagram [68]; (b) $\text{SiO}_2 - \text{K}_2\text{O}$ diagram [69]; (c) A/CNK–A/NK diagram [70]. Data sources: Nanmushui [45]; Daning [47]; and Wanyangshan [46].

Jilongjie granite contains 155.1–239.7 ppm of rare earth elements (Σ REE). The chondrite-normalized rare earth element pattern suggests that Jilongjie granite has obvious enrichment of light rare earth elements relative to heavy rare earth elements ($(La/Yb)_N = 15.1\text{--}23.7$) (Figure 5a). The samples showed a weak negative Eu anomaly, with Eu/Eu^* in the 0.68–0.78 range (Figure 5a, Table 4). In the primitive-mantle-normalized trace element diagram, the samples clearly showed enrichment of large ion lithophile elements (such as Sr, Rb, and Ba) and depletion of high-field-strength elements (such as Nb, Ta, and Ti) (Figure 5b).

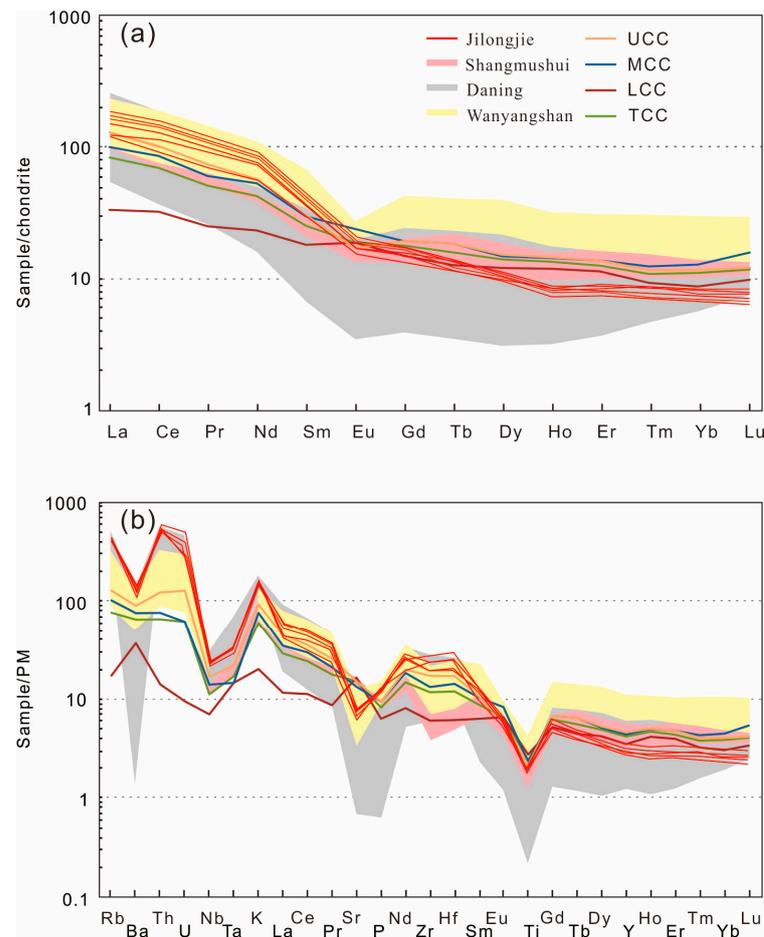


Figure 5. (a) Chondrite-normalized REE patterns; (b) Primitive-mantle-normalized trace element patterns. UCC—upper continental crust; MCC—middle continental crust; LCC—lower continental crust; TCC—total continental crust [71].

4.3. Zircon Hf Isotopic Results

Hafnium isotopic results and related parameters of 17 zircon grains of the Jilongjie granite samples are listed in Table 5. $^{176}\text{Lu}/^{177}\text{Hf}$ values were 0.001295–0.002283, $^{176}\text{Yb}/^{177}\text{Hf}$ values were 0.30414–0.56216, and $^{176}\text{Hf}/^{177}\text{Hf}$ values were 0.282502–0.282238 (Table 5). The Hf isotopic composition showed a wide variation; $\epsilon\text{Hf}(t)$ values were -9.94 to -0.69 ; model age values T_{DM} were 1.08–1.46 Ga, and two-stage model age values T_{DM2} were 1.46–2.04 Ga (Table 5).

Table 4. Major (wt.%) and trace element (ppm) concentrations of the Jilongjie granite.

| Sample | JL-1H | JL-2H | JL-3H | JL-4H | JL-5H | JL-6H | JL-7H | JL-8H | JL-9H | JL-10H | JL-11H |
|---|----------------|--------|-------|-------|-------|--------|-------|-------|-------|--------|--------|
| Rock type | Granite | | | | | | | | | | |
| SiO ₂ | 65.48 | 64.95 | 64.70 | 66.34 | 64.27 | 65.47 | 65.81 | 65.22 | 65.22 | 65.97 | 65.70 |
| TiO ₂ | 0.53 | 0.55 | 0.52 | 0.55 | 0.55 | 0.57 | 0.57 | 0.56 | 0.54 | 0.52 | 0.52 |
| Al ₂ O ₃ | 14.31 | 13.79 | 14.18 | 13.94 | 13.61 | 12.95 | 14.07 | 13.83 | 14.01 | 13.87 | 13.72 |
| Fe ₂ O ₃ ^T | 3.14 | 3.42 | 3.46 | 3.23 | 3.31 | 3.06 | 3.30 | 2.72 | 3.16 | 2.89 | 2.88 |
| MnO | 0.04 | 0.05 | 0.05 | 0.05 | 0.06 | 0.05 | 0.05 | 0.06 | 0.05 | 0.05 | 0.05 |
| MgO | 1.67 | 2.40 | 1.86 | 2.06 | 2.50 | 1.56 | 2.09 | 1.89 | 2.02 | 2.05 | 2.03 |
| CaO | 2.41 | 2.84 | 2.62 | 2.33 | 3.00 | 3.33 | 2.51 | 2.90 | 2.31 | 2.74 | 2.72 |
| Na ₂ O | 2.50 | 2.65 | 2.51 | 2.60 | 2.64 | 2.93 | 2.61 | 2.67 | 2.63 | 2.72 | 2.72 |
| K ₂ O | 4.94 | 4.44 | 4.77 | 4.62 | 4.38 | 4.87 | 4.54 | 4.79 | 4.62 | 4.56 | 4.55 |
| P ₂ O ₅ | 0.20 | 0.20 | 0.19 | 0.20 | 0.20 | 0.21 | 0.21 | 0.20 | 0.20 | 0.19 | 0.19 |
| L.O.I | 2.17 | 2.74 | 2.49 | 1.83 | 2.04 | 3.04 | 1.99 | 2.68 | 1.87 | 2.46 | 2.38 |
| Total | 99.38 | 100.01 | 99.34 | 99.73 | 99.57 | 100.03 | 99.73 | 99.52 | 99.63 | 100.01 | 99.46 |
| A/CNK | 1.03 | 0.96 | 1.01 | 1.03 | 0.94 | 0.80 | 1.02 | 0.93 | 1.04 | 0.96 | 0.96 |
| A/NK | 1.51 | 1.51 | 1.53 | 1.50 | 1.50 | 1.28 | 1.53 | 1.44 | 1.50 | 1.47 | 1.46 |
| Mg# | 51.26 | 58.13 | 51.51 | 55.79 | 59.92 | 50.22 | 55.57 | 57.85 | 55.93 | 58.41 | 58.24 |
| Trace element (ppm) | | | | | | | | | | | |
| Sc | 10.8 | 10.0 | 11.0 | 10.9 | 10.9 | 8.49 | 11.0 | 11.0 | 10.9 | 9.49 | 9.31 |
| V | 67.0 | 69.0 | 67.3 | 69.8 | 72.4 | 56.2 | 70.5 | 70.6 | 69.8 | 64.1 | 63.3 |
| Cr | 150 | 153 | 149 | 160 | 161 | 156 | 150 | 159 | 158 | 143 | 142 |
| Co | 10.8 | 11.4 | 12.8 | 10.5 | 11.9 | 6.87 | 10.4 | 7.82 | 9.89 | 9.23 | 8.93 |
| Ni | 66.3 | 66.1 | 69.9 | 61.2 | 65.3 | 37.9 | 62.1 | 46.2 | 60.5 | 53.1 | 52.0 |
| Ga | 16.5 | 16.9 | 16.7 | 16.6 | 16.7 | 15.4 | 16.4 | 16.1 | 16.7 | 16.1 | 15.6 |
| Rb | 287 | 249 | 266 | 252 | 243 | 271 | 249 | 264 | 255 | 247 | 246 |
| Sr | 144 | 154 | 139 | 155 | 144 | 126 | 161 | 154 | 165 | 155 | 155 |
| Y | 15.4 | 14.0 | 16.3 | 14.0 | 15.3 | 13.6 | 14.2 | 14.5 | 13.2 | 13.0 | 12.8 |
| Zr | 287 | 211 | 211 | 284 | 310 | 265 | 274 | 308 | 267 | 226 | 217 |
| Nb | 17.1 | 16.7 | 15.8 | 16.7 | 17.1 | 16.5 | 16.8 | 16.8 | 16.5 | 15.4 | 15.2 |
| Ba | 787 | 990 | 742 | 942 | 1012 | 918 | 836 | 935 | 885 | 923 | 931 |
| La | 41.8 | 41.9 | 37.9 | 40.2 | 36.6 | 28.6 | 44.1 | 29.5 | 41.1 | 38.8 | 35.8 |
| Ce | 88.9 | 89.3 | 81.8 | 88.0 | 81.9 | 56.8 | 97.6 | 70.3 | 89.4 | 87.6 | 79.1 |
| Pr | 10.0 | 10.4 | 9.53 | 10.4 | 9.73 | 6.61 | 11.4 | 8.72 | 10.4 | 10.1 | 9.30 |
| Nd | 38.5 | 39.7 | 36.4 | 39.2 | 37.2 | 26.4 | 42.9 | 34.6 | 39.6 | 38.3 | 35.7 |
| Sm | 5.78 | 5.97 | 5.93 | 6.02 | 5.87 | 4.39 | 6.55 | 5.80 | 5.97 | 5.79 | 5.28 |

Table 4. Cont.

| Sample | JL-1H | JL-2H | JL-3H | JL-4H | JL-5H | JL-6H | JL-7H | JL-8H | JL-9H | JL-10H | JL-11H |
|----------------------------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|
| Eu | 1.05 | 1.07 | 1.11 | 1.09 | 1.13 | 0.90 | 1.19 | 0.99 | 1.11 | 1.06 | 1.00 |
| Gd | 3.54 | 3.64 | 3.83 | 3.59 | 3.72 | 2.80 | 3.50 | 3.39 | 3.43 | 3.42 | 3.13 |
| Tb | 0.51 | 0.49 | 0.53 | 0.49 | 0.50 | 0.43 | 0.51 | 0.50 | 0.48 | 0.45 | 0.44 |
| Dy | 2.91 | 2.76 | 3.05 | 2.74 | 2.97 | 2.57 | 2.83 | 2.83 | 2.57 | 2.58 | 2.44 |
| Ho | 0.51 | 0.48 | 0.54 | 0.48 | 0.51 | 0.45 | 0.48 | 0.50 | 0.46 | 0.45 | 0.41 |
| Er | 1.52 | 1.35 | 1.62 | 1.40 | 1.52 | 1.30 | 1.50 | 1.44 | 1.39 | 1.31 | 1.24 |
| Tm | 0.22 | 0.20 | 0.24 | 0.22 | 0.22 | 0.20 | 0.22 | 0.22 | 0.22 | 0.18 | 0.18 |
| Yb | 1.53 | 1.27 | 1.54 | 1.37 | 1.45 | 1.29 | 1.41 | 1.40 | 1.30 | 1.18 | 1.13 |
| Lu | 0.21 | 0.17 | 0.22 | 0.20 | 0.21 | 0.18 | 0.21 | 0.20 | 0.19 | 0.17 | 0.16 |
| Hf | 8.04 | 6.12 | 6.06 | 8.21 | 8.79 | 7.81 | 8.04 | 9.15 | 7.98 | 6.49 | 6.37 |
| Ta | 1.40 | 1.32 | 1.43 | 1.33 | 1.35 | 1.38 | 1.36 | 1.38 | 1.35 | 1.24 | 1.19 |
| Pb | 34.1 | 35.5 | 35.4 | 37.9 | 33.6 | 39.2 | 39.4 | 40.1 | 38.8 | 36.4 | 37.3 |
| Th | 45.0 | 49.8 | 44.6 | 50.3 | 44.9 | 51.5 | 47.8 | 47.3 | 44.9 | 52.0 | 48.1 |
| U | 6.47 | 5.34 | 7.63 | 6.41 | 5.56 | 10.4 | 6.23 | 7.05 | 8.33 | 6.20 | 5.76 |
| ∑REE | 223.16 | 222.64 | 211.59 | 220.30 | 209.81 | 155.07 | 239.67 | 185.96 | 221.71 | 213.82 | 197.45 |
| (La/Yb)_N | 19.58 | 23.72 | 17.69 | 21.04 | 18.16 | 15.87 | 22.40 | 15.12 | 22.62 | 23.58 | 22.82 |
| Eu/Eu* | 0.71 | 0.71 | 0.71 | 0.72 | 0.74 | 0.78 | 0.76 | 0.68 | 0.75 | 0.73 | 0.75 |

$$\text{Eu}/\text{Eu}^* = \frac{Eu_N}{\sqrt{(Sm_N) \times (Gd_N)}}$$

Table 5. Zircon Hf isotopic in situ analysis results and related parameters.

| Sample No. | Age/Ma | $^{176}\text{Yb}/^{177}\text{Hf}$ | 2σ | $^{176}\text{Lu}/^{177}\text{Hf}$ | 2σ | $^{176}\text{Hf}/^{177}\text{Hf}$ | 2σ | $\epsilon\text{Hf}(0)$ | $\epsilon\text{Hf}(t)$ | $T_{\text{DM}}(\text{Ga})$ | $T_{\text{DM}}^{\text{C}}(\text{Ga})$ | $f_{\text{Lu/Hf}}$ |
|------------|--------|-----------------------------------|-----------|-----------------------------------|-----------|-----------------------------------|-----------|------------------------|------------------------|----------------------------|---------------------------------------|--------------------|
| JL-U-01 | 423 | 0.0356 | 0.0010 | 0.001495 | 0.000038 | 0.282353 | 0.000024 | −14.82 | −5.94 | 1.29 | 1.78 | −0.95 |
| JL-U-02 | 431 | 0.0534 | 0.0006 | 0.002153 | 0.000027 | 0.282502 | 0.000014 | −9.56 | −0.69 | 1.09 | 1.46 | −0.94 |
| JL-U-03 | 421 | 0.0353 | 0.0003 | 0.001411 | 0.000021 | 0.282422 | 0.000019 | −12.37 | −3.50 | 1.19 | 1.63 | −0.96 |
| JL-U-04 | 432 | 0.0562 | 0.0006 | 0.002283 | 0.000024 | 0.282378 | 0.000013 | −13.93 | −5.08 | 1.28 | 1.74 | −0.93 |
| JL-U-05 | 428 | 0.0314 | 0.0005 | 0.001303 | 0.000021 | 0.282305 | 0.000014 | −16.52 | −7.48 | 1.35 | 1.89 | −0.96 |
| JL-U-06 | 424 | 0.0417 | 0.0006 | 0.001740 | 0.000021 | 0.282424 | 0.000015 | −12.30 | −3.46 | 1.19 | 1.63 | −0.95 |
| JL-U-07 | 423 | 0.0374 | 0.0006 | 0.001508 | 0.000026 | 0.282260 | 0.000015 | −18.10 | −9.22 | 1.42 | 1.99 | −0.95 |
| JL-U-08 | 429 | 0.0482 | 0.0007 | 0.001979 | 0.000024 | 0.282324 | 0.000018 | −15.84 | −6.96 | 1.34 | 1.85 | −0.94 |
| JL-U-09 | 425 | 0.0332 | 0.0006 | 0.001387 | 0.000025 | 0.282336 | 0.000022 | −15.43 | −6.48 | 1.31 | 1.82 | −0.96 |
| JL-U-10 | 429 | 0.0337 | 0.0022 | 0.001377 | 0.000088 | 0.282347 | 0.000015 | −15.03 | −5.98 | 1.29 | 1.79 | −0.96 |
| JL-U-11 | 420 | 0.0423 | 0.0007 | 0.001754 | 0.000023 | 0.282501 | 0.000017 | −9.57 | −0.81 | 1.08 | 1.46 | −0.95 |
| JL-U-12 | 422 | 0.0394 | 0.0015 | 0.001636 | 0.000059 | 0.282366 | 0.000017 | −14.37 | −5.55 | 1.27 | 1.76 | −0.95 |
| JL-U-13 | 417 | 0.0362 | 0.0010 | 0.001481 | 0.000038 | 0.282323 | 0.000014 | −15.88 | −7.12 | 1.33 | 1.85 | −0.96 |
| JL-U-14 | 432 | 0.0373 | 0.0011 | 0.001540 | 0.000041 | 0.282430 | 0.000015 | −12.09 | −3.03 | 1.18 | 1.61 | −0.95 |
| JL-U-15 | 424 | 0.0530 | 0.0021 | 0.002127 | 0.000089 | 0.282398 | 0.000016 | −13.23 | −4.50 | 1.24 | 1.69 | −0.94 |
| JL-U-16 | 431 | 0.0462 | 0.0006 | 0.001863 | 0.000022 | 0.282238 | 0.000034 | −18.88 | −9.94 | 1.46 | 2.04 | −0.94 |
| JL-U-17 | 432 | 0.0304 | 0.0007 | 0.001295 | 0.000029 | 0.282264 | 0.000016 | −17.98 | −8.85 | 1.41 | 1.98 | −0.96 |

5. Discussion

5.1. Genetic Type and Magma Source

Granite's genesis and geodynamic mechanisms are closely related to rock types [72]. Whalen, et al. [73] proposed that A, I, S, and M-type granites can be distinguished according to rocks' geochemical characteristics, such as FeO^*/MgO and the Y vs. $10,000 \times \text{Ga}/\text{Al}$ discriminant diagram. In the discrimination diagram, all samples fall into I, S, and M regions, indicating that they are not A-type granites (Figure 6a,b). I-type granites are generally calcium-rich and aluminum-poor, with a high $\text{Na}_2\text{O}/\text{K}_2\text{O}$ ratio, and are mostly meta-aluminous ($A/\text{CNK} < 1.1$). I-type granites' dark minerals mostly contain clinopyroxene or amphibole. S-type granites are rich in aluminum but low in calcium, with a low $\text{Na}_2\text{O}/\text{K}_2\text{O}$ ratio. They are peraluminous ($A/\text{CNK} > 1.1$), and contain metamorphic minerals such as sillimanite, cordierite, garnet, and andalusite [74]. Jilongjie plutonic rocks are meta-aluminous to weakly peraluminous ($A/\text{CNK} = 0.80\text{--}1.04$) and contain amphiboles (Figure 2), which are similar to I-type granites. Moreover, the Jilongjie pluton falls into the I-type granite region in the Rb/Zr vs. SiO_2 plot (Figure 6c), as well as the I-type granite trend in the P_2O_5 vs. SiO_2 plot (Figure 6d). Therefore, its petrological and petrochemical characteristics indicate that Jilongjie granite is an I-type granite.

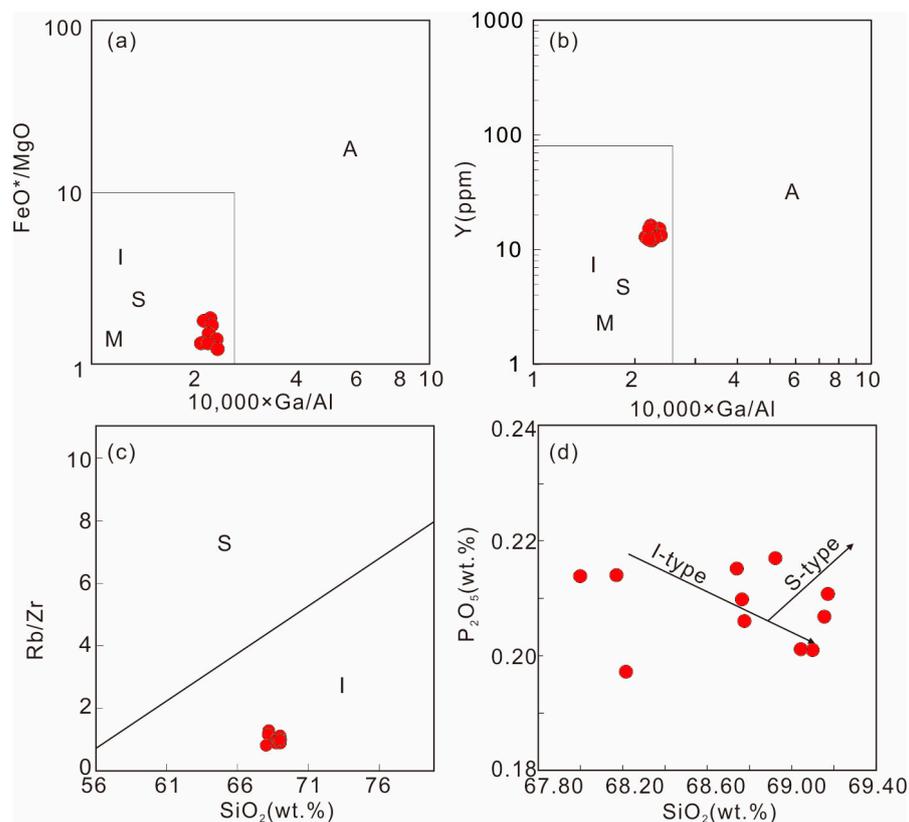


Figure 6. Chemical discrimination diagrams for I-, S-, M-, and A-type granites; (a) $10,000 \times \text{Ga}/\text{Al}$ vs. FeO^*/MgO [75]; (b) $10,000 \times \text{Ga}/\text{Al}$ vs. Y [75]; (c) Rb/Zr vs. SiO_2 [76]; and (d) P_2O_5 vs. SiO_2 [77].

5.2. Magma Source

Generally, the following models could account for the generation of I-type granite: (1) partial melting of metamorphic intermediate-mafic volcanic rocks [78–80]; (2) fractionation from the mantle-derived mafic magma [81,82]; (3) partial melting of juvenile crust induced by asthenosphere underplating [83,84]; and (4) mixing between mantle-derived mafic- and crust-derived felsic magma [85–88].

I-type granites generated by the fractional crystallization of mantle-derived mafic magma ordinarily have the following characteristics [81,82,89]: (1) massive ultramafic and

mafic lavas exposed around the study area; (2) samples with obvious negative Eu and Sr anomalies, indicating that magma formation was the result of the fractional crystallization of plagioclase from ultramafic and mafic melts; (3) the occurrence of mafic enclaves; and (4) enrichment with Sr-Nd-Pb isotopic features. The geological and geochemical characteristics of the Jilongjie pluton rule out fractional crystallization of mantle-derived mafic magma.

Zircon grains from Jilongjie granite samples had negative $\varepsilon_{\text{Hf}}(t)$ values ranging from -9.9 to -0.7 (Figure 7); this rules out partial melting of metamorphic intermediate-mafic volcanic rocks and fractionation from mantle-derived mafic magma (Figure 7). Based on these negative $\varepsilon_{\text{Hf}}(t)$ values and the model age of 1.04–1.46 Ga, the most direct explanation is that they originated from the anatexis or remelting of ancient crustal materials [90]. However, zircon Hf isotope data show obvious inhomogeneity (its variation range was several ε units); this required an open system to cause remarkable changes in the $^{176}\text{Hf}/^{177}\text{Hf}$ ratio in the melt (Figure 7) [91]. As zircon Hf isotope ratios hardly change with partial melting or fractional crystallization, the heterogeneity of zircon Hf isotopes likely indicates the interaction of mantle- and crust-derived magmas [92]. Therefore, similar to the heterogeneity of zircon Hf isotopes observed in other parts of the world, the Jilongjie pluton with similar characteristics is also interpreted as the result of the mixing of mantle- and crust-derived magmas [77,91–94]. Previous studies indicated that igneous rock with a high transition metal content are generally interpreted as direct melting from mantle peridotite or mixed melting from crust and mantle materials [95,96]. The high content of transition metals in Jilongjie granite was most likely derived from the mixed melting of crustal and mantle materials (Table 4). Moreover, in the $(\text{La}/\text{Yb})_{\text{N}}$ diagram, the Jilongjie pluton falls into a mixed mantle and crust region (Figure 7c).

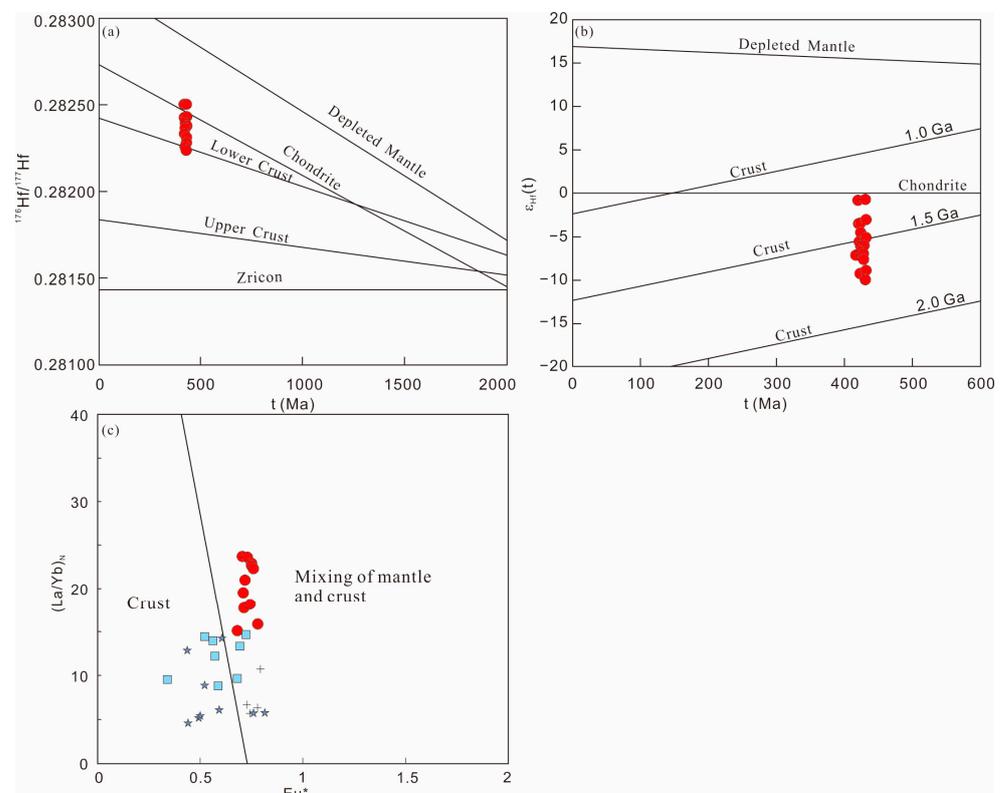


Figure 7. (a) $^{176}\text{Hf}/^{177}\text{Hf}$ vs. t (Ma); (b) $\varepsilon_{\text{Hf}}(t)$ vs. t (Ma) [97]; (c) $(\text{La}/\text{Yb})_{\text{N}}$ vs. Eu^*/Eu^+ [98]. The data source is the same as that in Figure 4. The legends are the same as Figure 4.

However, it is difficult to explain the whole-rock major and trace geochemical characteristics of Jilongjie pluton based solely on the mixing of mantle- and crust-derived material.

The depletion of Ba, Nb, Ta, Sr, and Ti indicate that its parent magma has undergone significant fractional crystallization (Figure 5). For example, the depletion of Nb, Ta, and Ti indicates the fractional crystallization of titanium-rich mineral phases (such as ilmenite and/or rutile) and Ca-amphibole, and the strong depletion of Sr and Ba indicates the fractional crystallization of plagioclase and potassium feldspar. Moreover, the Rb vs. Sr and Ba vs. Sr diagrams also suggest fractional crystallization of plagioclase and potassium feldspar (Figure 8a,b). Fractional crystallization of zircon and K-feldspar are also indicated in Zr vs. SiO₂ and Ba vs. SiO₂ diagrams, respectively (Figure 8c,d).

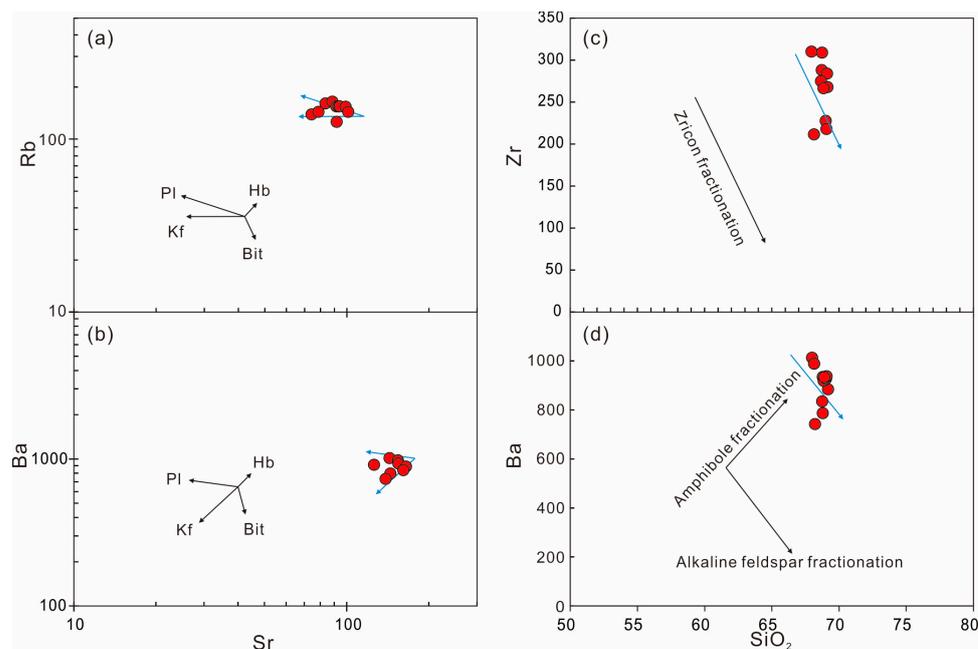


Figure 8. (a) Rb vs. Sr; (b) Ba vs. Sr; (c) Zr vs. SiO₂; and (d) Ba vs. SiO₂ diagrams of granite in the Jilongjie area (after Zong, et al. [99]). Mineral abbreviations are the same as Figure 2.

Mantle material may have played an important role in the SCB's Triassic magmatism, including the Shangmushui, Daning, and Wanyangshan plutons [45–47]. The geochemical characteristics of Jilongjie granite are similar to those of the above plutons (Figures 4 and 5). Therefore, based on geological, major and trace element geochemical, and Hf isotope data, we suggest that the parent magma of Jilongjie granite originated from mixed crust-derived felsic and mantle-derived mafic magmas, followed by fractional crystallization during its ascent or in the emplacement level.

5.3. Tectonic Implications

According to previous studies, granites and granitic gneisses dated from the Late Ordovician to the Early Devonian are considered the significant products of early Paleozoic magmatism in the SCB [37–39,44,100–102]. The early Paleozoic Wuyi–Yunkai orogeny in the SCB was the first extensive tectonothermal event since the Neoproterozoic break-up of the Rodinia supercontinent, roughly synchronized with the Caledonian orogeny in Europe [1,103,104]. However, the tectonic–magmatic evolution of the early Paleozoic remains controversial. Some scholars suggest that the Wuyi–Yunkai orogeny was the result of continental collisions or arc collisions caused by the subduction and closure of the Huanan Ocean between the Cathaysian and Yangtze blocks during the Caledonian period [105–107]. Others argue that the Yangtze and Cathaysian blocks were still continuous in the early Paleozoic, and thus the Wuyi–Yukai orogeny represents an intracollisional collision [7,14,37,108–110]. The following petrological and sedimentary geological features indicate that the Wuyi–Yunkai orogeny was more likely an intraplate orogeny rather than a subduction-related orogenic event: (1) Using statistics from previous research, we

found that massive granites are widely distributed, which is inconsistent with the linear distribution of magmatic rocks in the subduction mode (Figure 1a; [6]). (2) Early Paleozoic ophiolitic suites, arc andesites, and calc-alkaline volcanic rocks related to the closure of the Huanan Ocean (mentioned above) are absent [12,14,109]. (3) The paleoecological and biostratigraphical evolution in the Cathaysia and Yangtze blocks are related and continuous [111]. (4) The age spectra of detrital zircon from lower Paleozoic sandstones of the Cathaysia and Yangtze blocks have similar characteristics [109,112].

Although the Jilongjie pluton is a significant distance from the Wuyi–Yunkai orogeny’s core area, according to its temporal and spatial distribution characteristics, it is likely to be the westward extension of the orogenic belt or the product of early Paleozoic granitic magmatism [31,113]. Therefore, it should have been generated in the same tectonic setting. However, it is controversial whether the tectonic setting of the early Paleozoic magmatic rocks (especially 460–400 Ma) in the South China Plate was syn-collisional [43,113] or post-collisional [26,27,37,39,44,100,101]. As shown in Figure 9, Jilongjie granites plot in the field of syn-collisional to post-collisional granites.

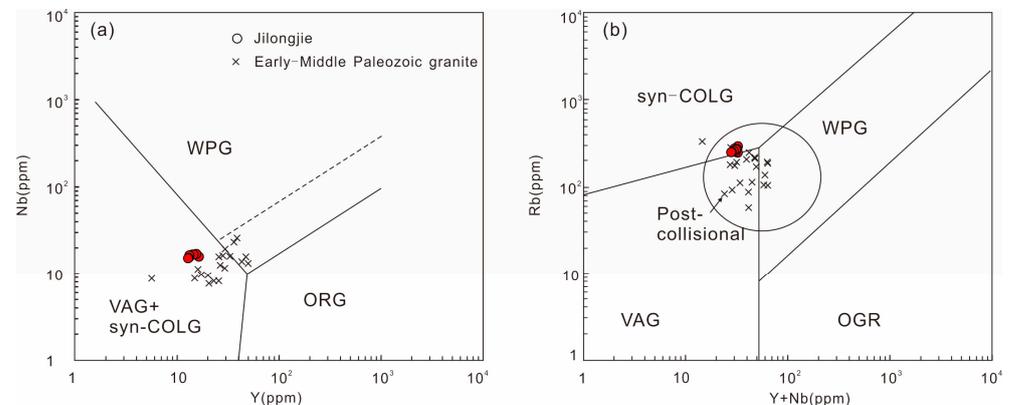


Figure 9. (a) Y vs. Nb and (b) (Y + Nb) vs. Rb [114]. The data source is the same as Figure 4.

Increasing evidence indicates that extensional mechanisms related to the post-collision stage were responsible for the generation of magmatic rocks in the SCB after 435 Ma, as follows: (1) Massive granites are widely distributed in the SCB, which is inconsistent with the characteristics of a small amount of migmatite and leucogranite generated in the syn-collisional orogenic stage [115]. (2) The high-magnesium basalt in the Wuyi–Yunkai orogenic belt indicates that the potential temperature of mantle melting at that time exceeded 1300 °C, which is significantly different from the syn-collision extrusion regime [42]. (3) The mafic intrusive rocks and contemporaneous granitic rocks generated by the decompression melting of the mantle in the SBC due to lithospheric extension are characterized by a bimodal pattern, which is consistent with post-collision extensional magmatism rather than a compression regime [26,101]. (4) Recent studies of deformation and advanced metamorphism indicate a prograde metamorphism associated with synorogenic crustal thickening and a retrograde metamorphism with postorogenic rapid exhumation at 460–435 and 435–400 Ma, respectively [26,27,100].

In summary, we suggest that the tectonic–magmatic evolution model of the Early Paleozoic (460–400 Ma) in the SCB can be summarized as follows: (1) During the syn-collision period (460–435 Ma), the crust was significantly shortened and thickened, causing high-temperature crustal anatexis and generating granitic rocks, accompanied by thickening of the lithospheric root (Figure 10a) [7,26,37,39]; (2) between 435 and 400 Ma, as the lithosphere was denser than the underlying asthenosphere, delamination caused part of the lithosphere root to be removed [116]. The subsequent upwelling of the asthenosphere provided heat to melt the lithospheric mantle [117]. The partial melting of the depleted lithosphere produced mafic magma, which intruded into the middle and upper crust, forming a huge magma chamber [118]. The intrusive mafic magma promoted massive crustal melting and

generated granitic melts [117]. Subsequently, the depleted mantle-derived material mixed with the granitic parent magma, generating intermediate granites, including the Jilongjie pluton (Figure 10b).

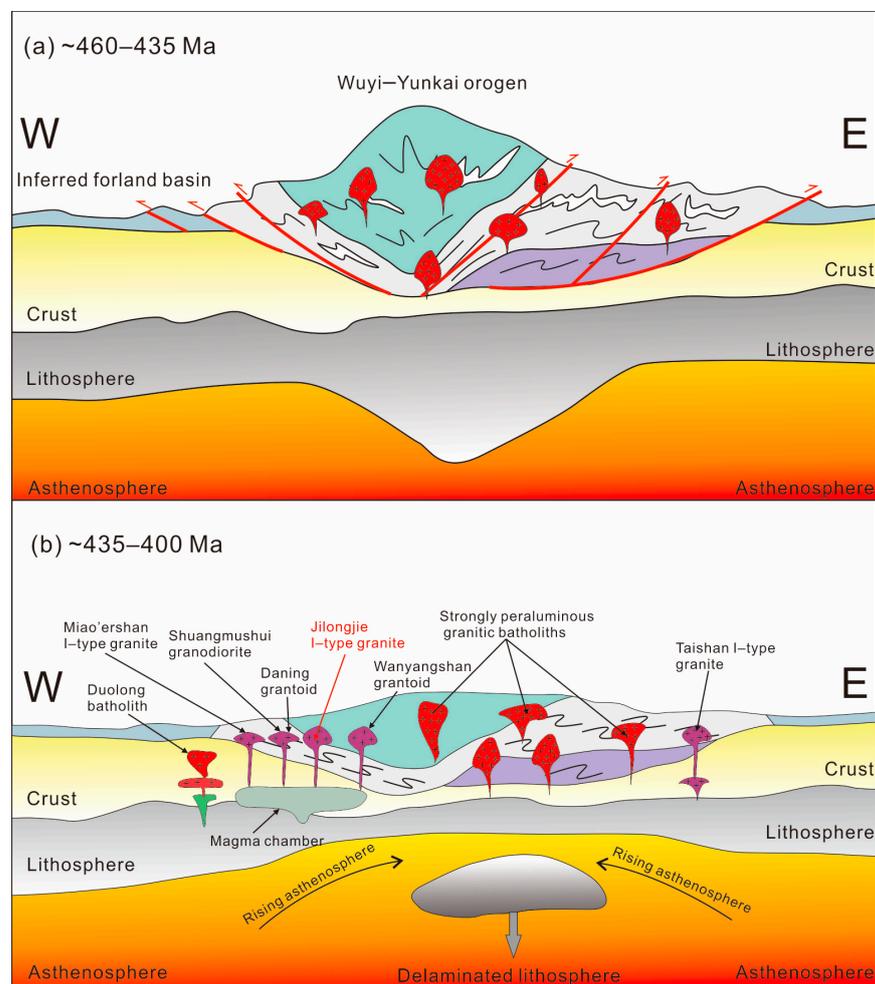


Figure 10. Simplified diagrams showing (a) ~460–435 Ma, crustal thickening and (b) ~435–400 Ma, lithospheric delamination in central SCB (modified after [101]). Data sources: Duolong batholith [44]; Miao'ershan I-type granite [27]; Taishan I-type granite [41]; strongly peraluminous granitic batholiths [37,39,119].

6. Conclusions

- (1) The Jilongjie pluton was emplaced at ~426 Ma and displays shoshonite and metaluminous characteristics.
- (2) Jilongjie granites' parent magma originated from a mixing of crust-derived felsic and mantle-derived mafic magmas, and then underwent fractional crystallization during its ascent and/or emplacement.
- (3) The post-collisional extensional mechanism associated with the delamination of the thickened lithosphere was responsible for post-435 Ma igneous rock (including the Jilongjie pluton) in the SCB.

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