



Article Solubility Product of Vanadinite Pb₅(VO₄)₃Cl at 25 °C—A Comprehensive Approach to Incongruent Dissolution Modeling

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Abstract: Although vanadinite (Pb₅(VO₄)₃Cl) occurs in abundance in various terrestrial geochemical systems of natural and anthropogenic origin and is seriously considered as a potential nuclear waste sequestering agent, its actual application is severely limited by a lack of understanding of its basic thermodynamic parameters. In this regard, the greatest challenge is posed by its incongruent dissolution, which is a pivotal hurdle for effective geochemical modeling. Our paper presents an universal approach for geochemical computing of systems undergoing incongruent dissolution which, along with unique, long-term experiments on vanadinites' stability, allowed us to determine the mineral solubility constant. The dissolution experiments were carried out at pH = 3.5 for 12 years. Vanadinite has dissolved incongruently, continuously re-precipitating into chervetite (Pb₂V₂O₇) with the two minerals remaining in mutual equilibrium until termination of the experiments. The empirically derived solubility constant $K_{sp,V,298} = 10^{-91.89 \pm 0.05}$ of vanadinite was determined for the first time. The proposed modeling method is versatile and can be adopted to other mineral systems undergoing incongruent dissolution.

Keywords: lead apatite; chervetite; solubility product; dissolution experiment; thermodynamics; geochemical modeling

1. Introduction

Vanadium is one of the most important metals in modern technology. It is the twentysecond most common element in the Earth's crust and the fifth most common transition metal. Small amounts of vanadium are essential for metabolism of several bacteria, algae, fungi and lichens (e.g., [1–5]). It also partakes in several bodily functions and life processes [6]. Vanadinite $Pb_5(VO_4)_3Cl$ is a secondary mineral, usually found in the oxidation zones of lead deposits and as a product of corrosion of infrastructure. In nature, it is often associated with high V content in mica and clay sediments subjected to hydrothermal processes, occurring under neutral or slightly alkaline conditions [7-9]. However, the mineral was also found in lead pipe water systems and it occurs in many already closed lead mines around the world [10,11]. Overall, there are more than 3000 worldwide locations, where vanadinite is present along with its phosphate and arsenate counterparts: pyromorphite $Pb_5(PO_4)_3Cl$ and mimetite $Pb_5(AsO_4)_3Cl$ [12], with which it forms solid solution series [13-15]. The minerals belong to the group of lead apatites [16], the composition of which is expressed by the general formula $Pb_5(BO_4)_3X$, where BO_4 represents the coordinated tetrahedron oxyanions (e.g., PO4³⁻, AsO4³⁻, VO4³⁻) and X represents the monovalent electronegative anions (e.g., F⁻, Cl⁻, Br⁻, I⁻, OH⁻) [17–19]. The hexagonal structure of most of the minerals in this group is consistent with the apatite structure



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Copyright: © 2021 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). $Ca_5(PO_4)_3F$ (spatial group P63/m) [14,17,19–24]. Many mineral phases of the apatite group are promising in the stabilization and recycling of industrial and nuclear waste and are being intensively studied in this respect [25–35]. The lead apatites are particularly interesting in this matter due to their relatively low solubility products, e.g., [36–39]. For this reason, induced precipitation of pyromorphite (logK_{sp} ~ -79) by phosphates amendments to soil or waste is a proposed and applied method of in situ remediation of lead-contaminated sites ([40–43] and the literature cited therein). Pentavalent vanadium in the form of VO_4^{3-} anion, one of the building blocks of vanadinite, is the most toxic and mobile form of the element that bears the potential of having a major impact on human health in the near future [44,45]. Furthermore, the level of vanadium in the environment is increasing alarmingly and its spread due to the combustion of hydrocarbon fuels, industrial and mining activities has been of concern in recent decades [46,47]. Therefore, the increasing number of studies on vanadinites as sequestering agents is not surprising [26,48-51]. However, the stability of vanadinites is, to date, poorly understood [10,52] and both the missing solubility and thermodynamic data are of high importance for a wide range of remediation strategies geared towards reducing the effects of potentially toxic substances containing Pb, V, As and radioactive elements [27]. Determination of the solubility product of vanadinite $(K_{sp,V})$ poses significant challenges and, to date, no reliable data on this matter have been presented. Gerke et al. [10] in their studies of groundwater in the Pb-containing pipes, proposed a log $K_{sp,V}$ value of -86, but their approach to estimate the solubility of vanadinite failed to consider fundamentals of thermodynamic modeling regarding congruency of dissolution and solution equilibrium. As such, the authors themselves recommended continued investigations into the matter. Unlike orthophosphate or orthoarsenate ions, orthovanadate groups tend to polymerize in aqueous solution (e.g., [53]). Hence, the difficulty in determining $K_{sp,V}$ stems from the fact that it is prone to re-precipitation into its polymerized form—chervetite ($Pb_2(V_2O_7)$)—under a wide range of conditions [54]. During the incongruent dissolution of vanadinite—as the system approaches thermodynamic equilibrium, a key factor in determining the properties of the minerals—chervetite is formed. Recrystallization is relatively slow and occurs at different rates, which highly depend on the pH and temperature; however, this process is still poorly understood in the context of vanadinite solubility [55]. In the carbonaceous environments, the amount of adsorbed (or possibly precipitated as chervetite) vanadium is greatly affected by the dissolved organic compounds (humic substances, organic acids) because of reduction of vanadate V(V) to vanadyl (V(IV)) VO^{2+} ions [56–59]. Furthermore, thermodynamic modelling of incongruent dissolutions poses significant challenges and, even though it is relevant to a large number of mineralogical systems, including other apatites, few approaches have been proposed to date, and no universal method can be referenced [60–63].

The growing importance of lead vanadate phases for industrial and environmental applications, along with the obvious lack of reliable thermodynamic data for vanadinite as the endmember of Pb-apatites solid solutions series, has served as an inspiration for the presented unconventional, 12-year long experiment on its solubility and stability. Our aim was to supplement thermodynamic data and to provide a reliable basis for further studies on Pb-V phases. We have faced the challenge of incongruent dissolution and, in this work, we demonstrate an approach to data processing that allowed us to experimentally determine the solubility product of vanadinite for the first time, and which is also sufficiently universal to be applied to other systems.

2. Materials and Methods

2.1. Synthesis of Vanadinite

Synthesis of vanadinite was carried out at 90 °C and pH 3.5 adjusted using 0.1 M HNO₃ (MERCK, supra pure) by dropwise mixing of the aqueous solutions of Pb(NO₃)₂ (POCH.SA, pure for analysis), KCl (CHEMPUR, pure for analysis), and NH₄VO₃ (POCH.SA, pure for analysis), in molar proportions based on the stoichiometry of the vanadinite (Pb:V:Cl = 5:3:1 molar ratio). The temperature and the pH of the synthesis were indicated by preliminary

geochemical modeling [64,65] as the conditions favorable for the VO₄^{3–} presence in the solution, thus favorable for vanadinite precipitation. Suspension of forming precipitate was stirred using a mechanical stirrer for 5 h and then allowed to settle. After 2 weeks, the precipitate was separated from the solutions by decantation, washed thoroughly several times with redistilled water and acetone and dried at 110 °C.

2.2. Solid Characterization

The X-ray powder diffraction (XRD) pattern of the solids was obtained using a Philips PW 3020 X'Pert-APD Diffractometer system (with a Cu anode and a graphite monochromator) using a step scan mode at a step size of 0.02° 2 Θ and a rate of 1 s per step. The scanning electron microscopy with elemental microanalysis (SEM/EDS) was performed using a variable pressure field emission FEI QUANTA 200 FEG microscope on gold-coated and uncoated samples at 20 kV. For the wet chemical analysis, -small portions of synthetic vanadinite (~50 mg) were digested in 0.02 M EDTA (2,2',2'',2'''—(ethane-1,2-diyldinitrilo)tetraacetic acid) (POCH.SA, pure for analysis) and analyzed for Pb, V and Cl, as described in the next section. Compared to other reagents e.g., HNO₃, the EDTA is an effective solvent for mineralization of Pb-apatites yielding relatively quick and complete dissolution of the solid [14,36,37,39,55].

2.3. Solution Analysis

A Total Reflection X-ray Fluorescence was used to determine the Pb and V content in the solutions resulting from dissolution experiments. The samples were inoculated with a nickel stock solution (1000 mg dm⁻³, MERCK, supra pure) as an internal standard, and analyzed at 50 kV and 12 mA with a Nanohunter II spectrometer equipped with a Mo lamp (t = 3000 s). Atomic Absorption Spectroscopy AAS (SavantAA GBC Scientific Equipment) was used to determine the concentration of Pb and V in the solutions resulting from digestion of solids in EDTA for wet chemical analysis of vanadinite. The turbidimetric method, with silver nitrate at 380 nm, was used to determine the concentration of Cl in all solutions with a HITACHI U-1800 UV-Vis Spectrophotometer. The universal meter CX-505 with a glass pH electrode ERH-13-6 was used to measure the pH. All analyses were performed in triplicate.

2.4. Dissolution Experiments

Dissolution experiments were performed in triplicate in 250 mL polycarbonate bottles at 25 °C \pm 1 °C. Equal portions (150 mg) of synthetic vanadinite were washed thoroughly with redistilled water and introduced as a suspension to the bottles containing 250 mL of 0.05 M KNO₃ solution. The KNO₃ (POCH.SA, pure for analysis) was used as the background electrolyte because both, K⁺ and NO₃⁻ ions are not involved in the reactions in the studied experimental system. At this concentration of the background electrolyte, the increase in ion concentration caused by the dissolution of the mineral during the experiment does not cause significant changes in the total ionic strength of the solution, and thus does not affect the experimental conditions. At the same time, the ionic strength is low enough that the Debye-Hückel equation can be used to calculate ion activity. The experiments were performed in triplicate at pH = 3.5, adjusted using 0.1 M HNO₃ (MERCK, supra pure). This pH value was indicated in our previous studies as adequate for the potential thermodynamic modeling of vanadinite stability [55]. The dissolution was carried out for 12 years. For the first 8 months, the bottles were immersed in a closed water bath with thermostatic control and manually stirred at least two times a week. After this time, the bottles were transferred to a room temperature of 22 $^\circ\text{C}\pm$ 2 $^\circ\text{C}$ and kept for 9 years in darkness; the reactors were occasionally mixed and checked for leaks. For the last 3 years of the experiments, the temperature was raised to 25 °C \pm 1 °C and the reactors were manually stirred at least once a month. During the experiments, 4 mL aliquots of the solutions were collected periodically, filtered using a 0.2 µm polycarbonate filter and analyzed in terms of pH and the total concentration of Pb, V and Cl. The evolution

pattern of [Pb], [V] and [Cl] was observed to determine the equilibrium in the suspensions; the system was considered in equilibrium when the results of three consecutive solution analyses were equal at 95% confidence level. The residual solids were syringe-sampled, washed with acetone, air-dried, and characterized with SEM/EDS and XRD.

2.5. Statistics

The statistical software Statistica 13 (StatSoft Polska) was used to process the solution composition data. The normal distribution was verified with the use of the Kolmogorov–Smirnov (K–S) test with Lilliefors correction (K–S–L) and the Shapiro–Wolf test (S–W). The Wald–Wolfowitz, the U Mann–Whitney and the Kołmogorow–Smirnow tests and the Student's t test were applied accordingly to verify whether the elemental contents of the solutions change significantly over time at 95% confidence level.

3. Results

3.1. Synthesized Solid

The synthesis yielded very fine, crystalline vanadinite with crystals smaller than 2 μ m (Figure 1). Lead, vanadium, chlorine, and oxygen were the only elements detected by the SEM/EDS elemental microanalysis (data not shown). No other crystalline or amorphous components were observed. Comparison of XRD patterns with data reported by Dai and Hughes [66] and Trotter and Barnes [67] yielded good agreement, indicating vanadinite Pb₅(VO₄)₃Cl as the sole reaction product (Table A1). No other phases were found within the detection limit of the method (equal to ca. 0.1 wt%). The wet chemical analysis of synthesized solid confirmed the purity of the solid (mass fraction purity = 0.995).



Figure 1. SEM microphotograph of synthetic vanadinite. The precipitate was aged for 2 weeks.

3.2. Dissolution Experiments

The dissolution experiments were carried out over an exceptionally long period— 12 years. Standard empirical thermodynamic studies of apatites last several weeks, as this has been widely accepted to be sufficiently long for the solids to equilibrate with the solutions [36,37,39,68]. However, vanadates tend to polymerize, inducing re-precipitation of vanadinite Pb₅(VO₄)₃Cl into chervetite Pb₂V₂O₇; the system is highly pH and temperature dependent and the kinetics of these reactions are poorly understood [55]. This caused difficulties in the interpretation of dissolution data and in unambiguous determination of the equilibrium of the system with the use of the dissolution–precipitation method (e.g., [37]). Thus, we decided to run the experiments over an exceptionally long period, recognizing the long-term stability of the solutions' composition as the sufficient base to estimate the missing thermodynamic parameters for vanadinite. The analytical methods (TXRF), requiring small amounts of aliquots, as well as the proper sealing of the reactors assured the quality of the long-term experiments, and the effect of potential evaporation was negligible. The SEM image of the dissolution residue after 12 years of aging in the suspension is presented in Figure 2. During the experiments, apart from the very fine, hexagonal particles of primary vanadinite, (the experimental material subject to dissolution; please compare with Figure 1), notably larger (>10 μ m) prismatic crystals with the monoclinic symmetry appeared in the solid fraction. The EDS (data not shown) and the XRD analysis indicated that the prismatic crystals of the secondary mineral phase are chervetite Pb₂V₂O₇. [69,70]. All X-ray diffraction peaks were assigned to either vanadinite or chervetite, confirming beyond doubt the incongruent nature of the dissolution process in which vanadinite re-precipitates partially into chervetite (Table A1). The possible mechanism behind the appearance of chervetite in the solid residue was that dissolving vanadinite released Pb and V ions in concentrations exceeding the solubility product of chervetite inducing its precipitation as the secondary mineral. However, the recrystallization of vanadinite into chervetite due to adsorption of polymerized vanadates or their direct formation on the vanadinite surface cannot be excluded.



Figure 2. SEM microphotographs of the residual solids from the dissolution experiments of vanadinite. Fine, hexagonal, stubby crystals were identified as vanadinite while notably larger, tabular ones were identified as chervetite.

The evolution of Pb, V and Cl concentration in the experimental solutions with time is shown in Table 1 and additionally visualized in Figure A1 (Appendix B) The pH was systematically measured; however, its value remained stable for the whole 12 years, oscillating at 3.5 ± 0.05 , hence the data was not included in the table. The applied statistical analysis yielded that at the 95% confidence level the solution samples collected within the 1st, 2nd, 5th, 8th and 19th week as well as the 116th, 132th and 147th month of the experiment were identical in terms of Cl concentration. This indicated that Cl had reached its close-to-equilibrium level in the solutions relatively quickly, i.e., within the first week of the experiment. In contrast, the concentrations of Pb_{tot} and V_{tot} systematically increased in the solutions and only the last three samples collected during the final 3 years of the experiment were statistically identical at 95% confidence level (Table 1). It is worth mentioning that vanadium concentration in most of the samples collected at the beginning of the experiment was below or at its detection limit (0.5 μ mol dm⁻³). The results suggested that the system had reached significant, long-term steady state conditions: three years of statistically stable solution parameters. Hence, we concluded that the final levels (in μ mol dm⁻³) of Cl = 24.4 \pm 3.2, Pb = 45.9 \pm 7.00 and V = 2.4 \pm 1.6, expressed as the mean value of concentrations determined in the solutions during the last three years of the experiment, adequately represented the equilibrium in the solutions. The final molar ratios of the Pbtot to Cltot in the solutions varied between 1.8 and 1.9, while the ratios of Pb_{tot} to V_{tot} fell in the range between 17 and 20. The proportions were neither typical for vanadinite (Pb:V;Cl = 5:3:1), nor for chervetite (Pb:V = 1:2), indicating dissolution in

which the properties of both involved phases simultaneously affect the composition of the solution.

					Concon	tration	umol de	m = 3				[DL]	T CI *	T TAD \$*	T TADASS
Time		Pb _{tot.}		V _{tot.}		Cl _{tot} .		[Pb] _{tot} /[V] _{tot}	[[Cl] _{tot}	Chervetite	Vanadinite	Vanadinite			
	0.1	5.7	±	0.4	b.d.	±	-	8.3	±	11.0	-	0.7	-	-	-
weeks	0.7	18.5	±	18.6	b.d.	±	-	18.1 c	±	11.1	-	1.0	-	-	-
	2.0	19.6	±	1.4	b.d.	±	-	19.2 c	±	11.2	-	1.0	-	-	-
	4.6	24.0	±	0.7	b.d.	±	-	20.9 c	±	11.2	-	1.1	-	-	-
	7.7	26.7	±	0.8	b.d.	±	-	23.1 c	±	4.0	-	1.2	-	-	-
	13.4	27.5	±	2.7	0.37	±	0.04	14.9	±	6.4	75	1.8	-0.68	-92.43	-94.57
	18.9	28.5	±	2.0	0.50	±	0.04	20.2 c	±	5.2	57	1.4	-0.56	-92.28	-94.05
	23.9	28.4	±	0.7	0.46	±	0.08	18.6	±	5.6	62	1.5	-0.59	-92.32	-94.17
	116	46.3 _a	±	10.6	2.39 b	±	0.86	24.8 c	±	4.0	19	1.9	0.14	-91.87	-91.42
	132	45.6 _a	±	3.0	$2.62_{\rm b}$	±	2.84	23.4 c	±	4.0	17	1.9	0.16	-91.91	-91.39
	147	45.8 a	±	4.4	2.28 b	±	1.61	25.2 c	±	4.5	20	1.8	0.13	-91.87	-91.47

Table 1. Selected parameters of experimental solutions: the raw and calculated data.

* logarithm of Saturation Index for Chervetite in the solution calculated with the use of the geochemical model PHREEQC (Phreeqc Interactive 3.5.0-14000) [64]; for undersaturated solution SI < 0, for saturated solution SI = 0, for supersaturated solution SI > 0. ** Ion Activity Product (IAP) for Vanadinite calculated with the use of PHREEQC [64]—assuming solution equilibrated with Chervetite. *** Ion Activity Product (IAP) for Vanadinite calculated with the use of PHREEQC [64]—neglecting the presence of Chervetite. $_{a, b, c}$ Values within a column followed by the same letters are statistically identical at 95% confidence level. Shaded table rows represent equilibrium.

4. Discussion

Thermodynamic modelling of incongruent dissolutions poses significant challenges. In this regard, few approaches for the apatite group have been proposed [71–74]. In this study, the geochemical speciation model PHREEQC (Phreeqc Interactive 3.5.0-14000) [64] was used to support the calculations of thermodynamic parameters. The MINTEQA2 [70] was chosen as the PHREEQC database containing, in our opinion, the most comprehensive dataset related to vanadium compounds. The database did not include referenced thermodynamic data for vanadinite, since, to date, no reliable report has been proposed in these terms; however, the equilibrium constant of chervetite dissolution reaction (1) was indicated there as $10^{-0.95}$.

$$PbVO_{35} + 3H^{+} = Pb^{2+} + VO^{2+} + 1.5H_{2}O$$
(1)

The compositions of the experimental solutions were proceeded with the use of PHREEQC [64] to verify their possible saturation with chervetite. The results presented as the logarithm of chervetite saturation index (SI) are shown in Table 1. The log SI values indicated supersaturation with chervetite for the last three years of the vanadinite dissolution. Even though the log SI values only slightly exceeded 0, the experimental conditions were favorable for vanadates polymerization as indicated by the PHREEQC modeling [64] and the precipitation of chervetite occurred as confirmed by the SEM/EDS and XRD analysis. Hence, there were two possible approaches to further the thermodynamic modeling of vanadinite's solubility product, K_{sp} : (i) acknowledging the equilibrium with chervetite or (ii) neglecting the presence of a secondary phase as having little effect on the equilibrium [71–74]. Both approaches were exercised using PHREEQC [64].

If the dissolution reaction of vanadinite is written as in the Equation (2),

$$Pb_{5}(VO_{4})_{3}Cl = 5Pb^{2+} + 3VO_{4}^{3-} + Cl^{-}$$
(2)

then the logarithm of the ion activity product (IAP) for the reaction (2) is given by:

$$\log IAP = 5\log\{Pb^{2+}\} + 3\log\{PO_4^{3-}\} + \log\{Cl^{-}\}$$
(3)

where {} denotes activities of the ions. At equilibrium, the Ion Activity Product (IAP) calculated for the dissolution reaction of a mineral is equal to its solubility product, K_{sp} . The activities of aqueous Pb²⁺, VO₄³⁻, and Cl⁻ were calculated considering the two different approaches: (i) and (ii) mentioned above with the use of the geochemical speciation model PHREEQC [64]. All reactions were modeled as open to the air with $P_{CO2} = 10^{-3.5}$ atm, at pH = 3.5 and temperature = 25 °C. The experimental concentrations of Pb_{tot}, V_{tot} and

Cl_{tot} were used as the input data and the calculated log IAPs are listed in Table 1. For all equilibrated solutions, the obtained values were of similar magnitude; however, the IAP values that neglected the presence of chervetite were slightly higher. At equilibrium, the average log IAP for vanadinite calculated based on the approach (i) was: -91.89 ± 0.05 , whereas the average log IAP value obtained using the alternative approach (ii) was: -91.43 ± 0.08 (uncertainties represented by the double standard deviation from three values calculated at equilibrium).

To validate the calculations made upon the "inverse modeling" based on the three-step empirical approach with the use of the geochemical speciation model PHREEQC [64] was applied. In the first step of the inverse approach, the MINTEQA2 [65] thermodynamic database available in the PHREEQC [64] package was modified by the amendment of the dissolution reaction for vanadinite (4):

$$Pb_{5}(VO_{4})_{3}Cl = 5Pb^{2+} + 3VO_{4}^{3-} + Cl^{-}; \log_{k} = X$$
(4)

where "log_k" expresses solubility product of vanadinite (log $K_{sp,V,298}$). In the second step, the value of the vanadinite solubility product (log $K_{sp,V,298}$) expressed by X in the Equation (4) was modified in the database with the use of experimentally determined IAP values calculated at equilibrium (Table 1): X = -90.87, -90.91, -90.42, -90.39 and -90.47. The input file of a modeling assumed (Appendix C):

- The concentration of background electrolyte KNO3 = 0.05 M;
- Constant temperature (24 °C, 25 °C or 26 °C);
- Initial fixed pH = 3.5;
- Forced equilibrium of the solution with the vanadinite and chervetite.

In the third step, the output data of each modeling were compared with the results of the dissolution experiment, in terms of the solution parameters such as: pH, total Pb, Cl and V concentrations and saturation index of other potentially precipitated phases. The obtained set of selected data is presented in Table 2.

Table 2. Comparison of the selected solution parameters determined in dissolution experiments of vanadinite and via computing with geochemical model PHREEQC [64].

м		ining the Date	Equ	Residual			
M	ethod of Obta	lining the Data	Pb (µmol dm ⁻³)	V (µmol dm ⁻³)	Cl (µmol dm ⁻³)	Phases *	
	Experir	nental	$45.9\pm7.0~{}_{\rm a}$	$2.4\pm1.6~{}_{b}$	$24.4\pm3.2~\mathrm{c}$	none	
ling assuming: **	rroach (ii)	Log K _{sp,V} = -90.42	66.5 ± 8.4	$0.9\pm0.2~{}_{\rm b}$	32.8 ± 4.3	none	
del	ddy	$Log K_{sp,V} = -91.47$	64.0 ± 8.0	0.9 ± 0.2 b	31.5 ± 4.1	none	
om	A	$Log K_{sp,V} = -91.39$	68.0 ± 7.7	$0.9\pm0.2~{}_{\rm b}$	33.5 ± 3.9	none	
ochemical	proach (i)	Log K _{sp,V} = -91.87	$47.3\pm7.0~\text{a}$	$1.4\pm0.2~{}_{\rm b}$	$22.9\pm3.6~{\rm c}$	none	
Ge	Apl	$Log K_{sp,V} = -91.91$	$45.9\pm6.7~{}_{\rm a}$	$1.5\pm0.3~{}_{\rm b}$	$22.2\pm3.5~{}_{\rm c}$	none	

* Phases other than Vanadinite and Chervetite; ** Geochemical modeling with the use of PHREEQC [64], in which the MINTEQA2 [65] database was amended with particular values of the solubility product of Vanadinite (log $K_{sp,V}$) listed in the column 3. Approach (i): acknowledging presence of Chervetite. Approach (ii): neglecting presence of Chervetite; \pm Double standard deviation of triplicates (experimental data) or of modeling at three different temperatures (modeling data); _{a, b, c}, Values within a column followed by the same letters are statistically identical at 95% confidence level.

The PHREEQC [64] calculations with the assumed vanadinite solubility products log $K_{sp.V,298}$ equal -90.87 and 90.91 yielded solution parameters that were statistically identical (at 95% confidence level) with the empirical results provided by the dissolution experiment. This suggested that the calculations acknowledging the presence of chervetite were more adequate in this case. Thus, we recommend the value ($K_{sp,V,298} = 10^{-91.89 \pm 0.05}$) as the first empirically determined solubility product of vanadinite which can be used for further thermodynamic calculations, modeling and experimental approximations. The uncertainty (± 0.05) corresponds to a double standard deviation of the IAP values calculated for equilibrium in this study. The presented approach involving "inverse geochemical modeling" is a universal method to validate the assumptions made during thermodynamic modeling. The method is simple and in case of a complex incongruent dissolution systems it may refine uncertainties or even act as the straightforward calculations without intricate raw data processing. The obtained value ($K_{sp,V,298} = 10^{-91.89 \pm 0.05}$) is notably lower than the value ($K_{sp,V,298} = 10^{-86}$) proposed by Gerke et al. [10], however, their data was just an approximation based on presumptions of water parameters in lead-containing pipes and did not consider incongruent nature of vanadinite dissolution, precipitation of chervetite or equilibrium evidence in the solution. Nevertheless, regarding the general tendency of Pb-apatites for extremely low solubility products and the relative value of vanadinite parameter when compared to other Cl-Pb-apatites e.g., mimetite $K_{sp,V,298} = 10^{-79.11 \pm 0.33}$ [36] and pyromorphite $K_{sp,V,298} = 10^{-79.60 \pm 0.33}$ [39], our results were self-consistent.

5. Conclusions

For the first time, the experimentally determined solubility product of vanadinite equal to $K_{sp,V} \ 10^{-91.89 \pm 0.05}$ is presented. Vanadinite is an exceptional mineral among Pb-apatite series; it dissolves incongruently, inducing precipitation of chervetite as the sole secondary phase. The incongruent mechanism of vanadinite's dissolution poses a challenge to the direct experimental determination of the solubility constant; however, a new calculation approach, along with the unique, long term experiments proposed here, made it possible to overcome this difficulty. The approach to modeling with a use of the presented inverse method is novel, universal and can be applied to other systems.

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Conflicts of Interest: The authors declare no conflict of interest.

Appendix A

		v	'anadinite P	b ₅ (VO ₄) ₃	,Cl		Dissolution		Chervetite Pb ₂ V ₂ O ₇				
h k l	Dai and [[66	Hughes	Trotter and Barnes [67]		Synthesis, This Study		This Study		[70]		[69]		h k 1
	d- Spacing (Å)	I/Io	d- Spacing/ (Å)	I/Io	<i>d-</i> Spacing/ (Å)	I/Io	<i>d-</i> Spacing/ (Å)	I/Io	d- Spacing/ (Å)	I/Io	<i>d-</i> Spacing/ (Å)	I/Io	11 K I
110	5.1587	7.83	5.1655	8.91	5.171	7	5.170 4.998 4.949	10 4 15	4.9545 4.9422	6.98 57.75	4.9573 4.9438	7.26 61.74	$\overline{111}$
200	4.4676	24.21	4.4735	25.16	4.476	20	4.475	29	4.0007	27.99	4 2227	42.05	111
111	4.2202	32.9	4.2249	34.62	4.230	27	4.339	30	4.3337	37.88	4.3337	42.95	111
201	3.8159	4.89	3.8203	2.97	3.824	4	4.155 3.828	6 5	4.1469	4.71	4.1460	3.79	201
002	3.6689	7.66	3.6715	6.84	3.678	8	3.680	5	3.6770	1.66	3.6762	1.59	310
102 120.	3.3939	26.22 17.03,	3.3966	26.95 14.39,	3.390	38	3.592 3.450 3.433 3.403	14 22 15 14	3.5886 3.4491 3.4312	60.57 100 66.96	3.5879 3.4487 3.4351	61.34 100 76.92	$211 \\ 120 \\ \overline{2}02$
210	3.3772	18.84	3.3816	19.26			3.381	29	2 2740	10 70	2 27/1	10.00	101
							3.273 3.217 3.188	4 29 4	3.2749 3.2138 3.1815	18.70 85.59 15.91	3.2761 3.2130 3.1840	18.90 83.30 12.88	$401 \\ 400 \\ \overline{1}12 \\ \overline{1}21$
							3.172	10	3.1743, 3.1710	1.48, 32.64	3.1748, 3.1713	2.11, 33.84	021,
121, 211	3.0678	45,76, 54.24	3.0716	46.19, 53.81	3.073, 2.993	94, 100	3.073	100					
112 300	2.9899 2.9784	95.04 53.24	2.9926 2.9823	91.58 55.19	2.841	1	2.996 2.982 2.934 2.863	65 28 6 6	2.9320 2.8607	12.83 30.45	2.9314 2.8633	11.63 35.03	$\frac{410}{312}$
201	2 7507	1.45	2 7621	1 1 6	2 762	1	2.839	4	2.8419	13.05	2.8427	16.40	112
301	2.7397	1.45	2.7031	1.10	2.703	1	2.748 2.724 2.711	8 7 5	2.7475 2.7219 2.7100	28.12 41.21 27.44	2.7471 2.7221 2.7096	27.05 41.71 26.70	320 202 221
220	2.5793	2.89	2.5828	1.99	2.582	3	2.585	5	2.7100	22.11	2.7090	20.70	
122	2.4848	2.01	2.4873	1.96	2.489	3	2.519	6	2.5213	23.04	2.5225	25.58	122
							2.419 2.351	$\frac{4}{4}$	2.4198 2.3525	3.64 10.57	2.4192 2.3538	3.61 11.55	$\frac{510}{322}$
131 302	2.3479 2.3124	1.27 5.55	2.3508 2.3149	$1.40 \\ 4.43$	2.350 2.316	1 6	2.316	6					
						-	2.276	5	2.2766, 2.2704	17.94, 2.40	2.2782, 2.2720	20.30, 2.96	003, 512
							2.237	13	2.2376	23.01, 2.51	2.2394, 2.2373	1.62, 24.71	113, 230
400 113	2.2338 2.2101	$5.48 \\ 10.66$	2.2367 2.2119	4.94 9.69	2.238 2.214	5 9	2.215	7					
							2.184	4	2.1892, 2.1860	6.51, 12.57	2.1893, 2.1859	8.28, 12.40	$\overline{231}, 131$
							2.127	5	2.1259, 2.1258	17.04	2.1262, 2.1252	18.00	421
222	2.1101	35.78	2.1124	34.74	2.114	33	2.114 2.090 2.068	29 8 5	2.1153 2.0918 2.0687	10.67 6.17 2.61	2.1145 2.0930 2.0684	11.02 7.02 3.73	
132, 312	2.0536	3.93, 12 51	2.0559	5.09, 12.68	2.055	18	2.056	15					
100		10.01		15.00			2.006	4	2.0078	21.67	2.0090	23.93	<u>6</u> 12
123, 213	1.9810	17.15, 11.10	1.9828	15.11, 10.38	1.982	28	1.984	13	1.9810	5.89	1.9815	8.14	$\overline{1}32$
140.	1.0.400	8.28.	1.0524	7,87.	1.050	01	1.977	13	1.9774	15.87	1.9772	17.79	322
410	1.9498	16.55	1.9524	14.76	1.952	21	1.952	20					
402	1.9080	12.00	1.9102	20.30	1.711	20	1.911 1.891 1.838	4	1.8906 1.8388	6.31 2.54	1.8907 1.8393	6.84 2.13	$\overline{6}21$ 123
004 322	1.8345 1.7895	12.99 1.69	1.8357 1.7916	12.81	1.838 1.791	2	1.791 1.732	4 5	1.7337	1.36	1.7340	1.52	223

Table A1. Powder X-ray diffraction peaks from synthetic $Pb_5(VO_4)_3Cl$ used in this study compared with data reported in the literature: Miller indices *h k l*, distance *d* between adjacent (*h k l*) planes and the peak intensity ratio *l*/*lo*.

Appendix B



Figure A1. The evolution of Pb_{tot} , Cl_{tot} and V_{tot} concentrations in the experimental solutions with time.

Appendix C

Exemplary input data for PHREEQC simulation used for "Inverse modeling".

```
SOLUTION 1
             25
   temp
            4
   pe
   redox
            pe
   units
            mmol/kgw
   density
            1
   -water
            1 # kg
    Κ
             50
    N(5) 50
PHASES
 pH_Fix
  H+ = H+
   log_K = 0.0
Vanadinite
       Pb5(VO4)3Cl = 5Pb+2 + 3VO4-3 + Cl -
       log_k
               -91.87
EQUILIBRIUM_PHASES 1
   CO2(g)
               -3.5
                      1.0
chervetite 0.0
   Vanadinite 0.0
                      HNO3
                                 10.0
   pH_Fix
              -3.5
     -force_equality
END
```

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