

In our paper, the thermometer and Oxygen fugacity are estimated using the trace element concentrations in zircon and melt composition (represented by whole-rock). The detail methods are described as following.

Blundy and Wood [1] showed a model to predict the mineral melt partition coefficient for a cation (i) by the lattice strain energy, which created by substituting a cation whose ionic radius (r_i) differs from the optimal value for that site (r_0) according the equation:

$$\ln D_i = \ln D_0 - \frac{4\pi EN_A}{RT} \left(\frac{r_i}{3} + \frac{r_0}{6} \right) (r_i - r_0)^2 \quad (1)$$

Where D_i is partition coefficient of element i, D_0 is the strain compensated partition coefficient, N_A is Avogadro's number, R is the gas constant, and T is the temperature (K). When D_0 , N_A , R and T are treated as constants, the $\ln D_i$ versus $(r_i/3 + r_0/6)(r_i - r_0)^2$ yields a linear relation.

Based on the lattice strain model, the zircon $\text{Ce}^{4+}/\text{Ce}^{3+}$ ratios could use following formula [2].

$$\left(\frac{\text{Ce}^{4+}}{\text{Ce}^{3+}} \right)_{\text{zircon}} = \frac{\text{Ce}_{\text{melt}} - \frac{\text{Ce}_{\text{zircon}}}{D_{\text{zircon/melt}}^{\text{Ce}^{3+}}}}{\frac{\text{Ce}_{\text{zircon}}}{D_{\text{zircon/melt}}^{\text{Ce}^{4+}}} - \text{Ce}_{\text{melt}}} \quad (2)$$

Where Ce_{melt} and $\text{Ce}_{\text{zircon}}$ are the concentrations of Ce in melt and zircon, respectively. $D_{\text{Ce}^{3+}}$ and $D_{\text{Ce}^{4+}}$ are partial coefficients calculated from the linear fit of trivalent cations (REE^{3+}), and tetravalent cations (such as, Th^{4+} , U^{4+} and Hf^{4+}), respectively (Figure S1).

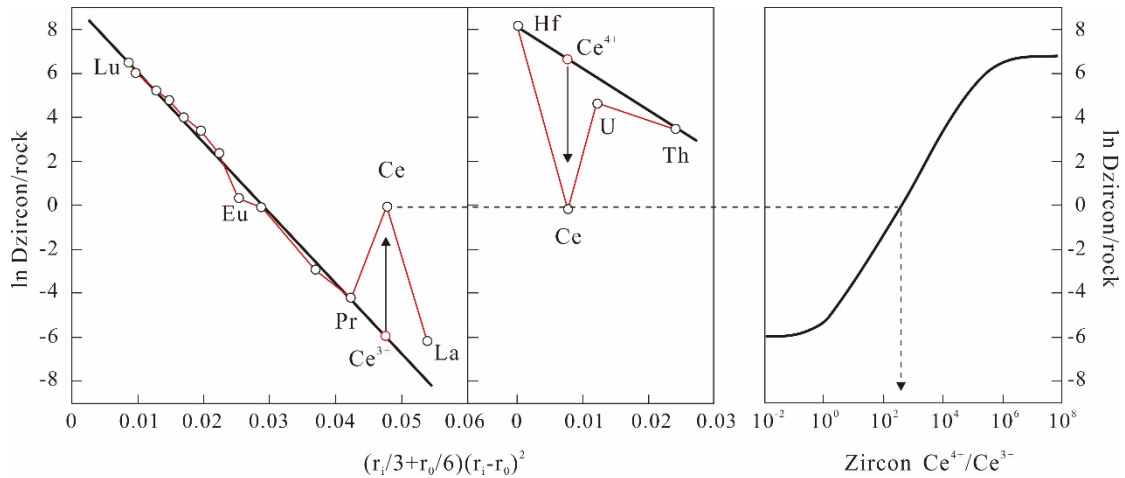


Figure S1. Diagrammatic representation of zircon/rock distribution coefficients by a lattice-strain parameter for trivalent and tetravalent cations[2]. The r_i represent ionic radius of a cation i , and r_0 is ionic radius of the optimal cation.

Partition coefficients for $D_{\text{Ce}^{3+}}$ and $D_{\text{Ce}^{4+}}$ can be determined by interpolation. Due to La and Pr are present very low concentration levels in natural zircons, and Eu, U are multivalent elements, these elements should be excluded in the calculation procedure [3]. Values of r_0 used for 3+ and 4+ cations were 0.93 Å (determined by regression) and 0.84 Å (8-fold coordinated Zr^{4+}), and ionic radii for r_i were from Shannon [4]. Smythe and Brennan [3] described a following equality for estimating $\text{Ce}^{4+}/\text{Ce}^{3+}$ in melt:

$$\left(\frac{\text{Ce}^{4+}}{\text{Ce}^{3+}} \right)_{\text{melt}} = 1.04877 * \left[\frac{\text{Ce}_{\text{zircon}} - (\text{Ce}_{\text{melt}} * D_{\text{Ce}^{3+}}^{\text{zircon/melt}})}{(\text{Ce}_{\text{melt}} * D_{\text{Ce}^{4+}}^{\text{zircon/melt}}) - \text{Ce}_{\text{zircon}}} \right] \quad (3)$$

Where $D_{\text{Ce}^{3+}}$ and $D_{\text{Ce}^{4+}}$ can be obtained from Equation (2). The proportions of Ce^{3+} and Ce^{4+} are related to $f\text{O}_2$ of melt which can be obtained through the following equation by Smythe and Brenan [3]:

$$\ln\left(\frac{\text{Ce}^{4+}}{\text{Ce}^{3+}}\right)_{\text{melt}} = \frac{1}{4} \ln f\text{O}_2 + \frac{13136(\pm 591)}{T} - 2.064(\pm 0.011) \frac{\text{NBO}}{T} - 8.878(\pm 0.112) * \text{xH}_2\text{O} - 8.955(\pm 0.091) \quad (4)$$

Where NBO/T is the proportion of non-bridging oxygens to tetrahedrally coordinated cations calculated on an anhydrous basis that can be calculated by the bulk element composition [5], and xH_2O is the mole fraction of water dissolved in the melt. $\text{Ce}^{4+}/\text{Ce}^{3+}$ in melt can be obtained from Equation (2) and T is temperature degrees in K determined using the Ti-in-zircon thermometer as described below.

The Ti-in-zircon thermometer has been widely used to identify magma forming or metamorphism temperatures. The experimentally calibrated thermometer is provided using the expression [6]:

$$\log(Ti_{\text{zircon}}) = (6.01 \pm 0.03) - \frac{5080 \pm 30}{T(\text{K})} \quad (5)$$

Where Ti_{zircon} denote the concentration of Ti in zircon in ppm, and T is temperature degrees in K.

The concentration of trace elements for melts are represented by average values of two samples from Guojialiang monzonites in Huyanshan complex by Ying et al. [7] presented in Table S1.

Table S1. Major and trace element data for the Guojialiang monzonites from Huyanshan complex.

Trace Element (ppm)	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	
XML08-2	36.5	64.1	7.74	28.6	4.98	1.66	4.45	0.67	3.95	0.8	
XMI08-3	22.1	52.8	6.84	25.2	4.69	1.32	4	0.6	3.57	0.69	
Average	29.30	58.45	7.29	26.90	4.84	1.49	4.23	0.64	3.76	0.75	
Trace element (ppm)	Er	Tm	Yb	Lu	Hf	Th	U	Ta	Zr	Y	
XML08-2	2.27	0.35	2.28	0.35	4.3	4.2	0.9	0.45	156	23	
XMI08-3	1.95	0.3	1.92	0.29	4.2	3.2	0.6	0.43	149	19	
Average	2.11	0.33	2.10	0.32	4.25	3.70	0.75	0.44	152.50	21.00	
Major element (wt. %)	SiO ₂	TiO ₂	Al ₂ O ₃	TFe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	LOI
XML08-2	62	0.29	19	3.29	0.1	0.69	1.86	5.22	5.32	0.11	2.07
XMI08-3	61.5	0.26	18.4	3.13	0.13	0.31	2.65	6.74	4.76	0.11	1.98
Average	61.75	0.28	18.70	3.21	0.12	0.50	2.26	5.98	5.04	0.11	2.03

References

1. Blundy, J.; Wood, B. Prediction of crystal-melt partition coefficients from elastic-moduli. *Nature* **1994**, *372*, 452–454.
2. Ballard, J.R.; Palin, M.J.; Campbell, I.H. Relative oxidation states of magmas inferred from Ce(IV)/Ce(III) in zircon: application to porphyry copper deposits of northern Chile. *Contrib. Mineral. Petrol.* **2002**, *144*, 347–364.
3. Smythe, D.J.; Brenan, J.M. Magmatic oxygen fugacity estimated using zircon-melt partitioning of cerium. *Earth Planet. Sci. Lett.* **2016**, *453*, 260–266.
4. Shannon, R.D. Revised effective ionic radii and systematic studies of interatomic distances in halides and Chalcogenides. *Acta Crystallogr. Sect. A* **1976**, *32*, 751–767.
5. Smythe, D.J.; Brenan, J.M. Cerium oxidation state in silicate melts: Combined $f\text{O}_2$, temperature and compositional effects. *Geochim. Et Cosmochim. Acta* **2015**, *170*, 173–187.

6. Watson, E.B.; Wark, D.A.; Thomas, J.B. Crystallization thermometers for zircon and rutile. *Contrib. Mineral. Petrol.* **2006**, *151*, 413–433.
7. Ying, J.F.; Zhang, H.F.; Tang, Y.J. Crust-mantle interaction in the central North China Craton during the Mesozoic: Evidence from zircon U–Pb chronology, Hf isotope and geochemistry of syenitic-monzonitic intrusions from Shanxi province. *Lithos* **2011**, *125*, 449–462.