

Article

Characteristics of Supergene Gold of Karst Cavities of the Khokhoy Gold Ore Field (Aldan Shield, **East Russia)**

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Abstract: Typomorphic features of supergene gold in karst cavities were studied in the recently discovered Au-Te-Sb-Tl deposit within the Khokhoy gold ore field of the Aldan-Stanovoy auriferous province (Aldan shield, East Russia). Two morphological types of supergene gold, massive and porous, are recognized there. The first type is represented by gold crystals and irregular mass, with the fineness ranging from 835 to 1000‰. They are closely associated with goethite, siderite, unnamed Fe, Te, and Tl carbonates, Tl tellurites/tellurates and antimonates, as well as avicennite with a Te impurity. The second type is represented by mustard gold of two types with different internal structure: microporous and dendritic. The supergene gold is characterized by persistently high fineness. Along with Ag, it invariably contains Hg (up to 5.78 wt%) and Bi, and, rarely, Pb, Cu, and Fe. The supergene gold is chemically homogeneous, and its particles are all marked by high fineness, without any rims or margins. The obtained characteristics made it possible to prove the existence of two genetic types of supergene gold. Mustard microporous gold is the result of the decomposition of the associated minerals-goethite, Tl oxides, tellurium, Fe, Mn and Tl carbonates and antimonates, containing microinclusions of gold. Massive gold and dendrites are newly formed. The decomposition, remobilization, and reprecipitation of residual gold nanoparticles and their aggregation led to the formation of dendrites, and with further crystal growth and filling of pores, to gold of massive morphology. In terms of morphology, internal structure, fineness, and trace element composition, supergene gold of the Khokhoy gold ore field is comparable to gold from the Kuranakh deposit (Russia) and the Carlin-type gold deposits. It also is similar to spungy and mustard gold from Au–Te and Au–Sb deposits, weathering crusts, and placers. Its main characteristic feature is a close paragenesis with Tl minerals.

Keywords: gold ore field; karst cavities; monolithic and porous gold; Tl oxides (avicennite); Tl carbonates; Tl tellurates and antimonates

1. Introduction

The importance of gold-bearing supergene zones increased significantly in relation to the discovery and commercial exploitation of the Carlin-type gold deposits (USA) [1–5]. In Russia, assigned to this type are deposits localized in karst cavities such as the Kuranakh deposit within the Central Aldan district of Yakutia and the Vorontsovskoe deposit in the Urals [6–9]. The question of native gold formation in supergene conditions, including in karst cavities, remains are as yet poorly investigated. Supergene gold in karst cavities has specific features, exhibits various textural and chemical characteristics that can help in elucidating the process of formation of supergenic gold.

That's why studies on the typomorphic features of the recently discovered supergene gold from the Khokhoy gold ore field within the Aldan-Stanovoy auriferous province are quite topical now.



druse quartz lenses, often with chalcedony rim.

The Khokhoy gold ore field is located in the Verkhneamginskaya auriferous zone, in the up-stream basin of the river Khokhoy, a right tributary of the river Amga [10]. The territory is a part of the Aldan-Stanovoy shield, in the basemen subsidence and sedimentary cover enlargement areas (Aldan shield, East Russia) (Figure 1). The majority of the sediments are those of Lower Cambrian terrigenous-carbonate sediments (predominantly dolomitic ones) with stratigraphic mismatches of closed Lower Jurassic terrigenous sandstone sediments. Mesozoic magmatism resulted in little stratified intrusions, monzonite lakkolithes and syenite-porphyries as well as in alkaline gabbroides. The ore field is structured by a vast faulted area with the North-Eastern strike. The faults' kinematics implies them being normal faults and slip faults with the amplitude of more than 100 m with echelon faults of the North-Western strike forming a ladder-shaped block structure of the ore field. Mineralization takes place where the echelon strippings intersect each other within the tectonized contact of Cambric and Jurassic sediments. Ores are accumulated in the karst cavities forming a vast zone with the North-Western strike of more than 10 km crossing the ore field from the South to the North. Almost the whole anomalous zone consists of karst holes of various size and mineralization extent, open as well as closed with sandstones. The latitude of the karst cavities opened during the mine working ranges from 5–15 to 50 more than 50 m, the depth is 15–45 m or more. Karst cavities are formed of intensive limonitized argillaceous-sandy fulvous matter with primary ores and enclosing rocks fragments of various size. According to X-ray phase and thermal analyses, karst include quartz, potassium feldspars, muscovite, goethite, and less often calcite, hematite, and fluorite. Clinochlor, Fe-clinochlor, lepidocrocite, jarosite and kaolinite also attested. Primary ores exist as fragments of jasperoids, pyrite-adular-quartz metasomatites (Figure 2). The metasomatites contain 93.53–94.45% of SiO₂, 1.41–1.91% of Al₂O₃ and 0.88–1.56% of K₂O. The primary ores almost exclusively consist of fine-grain quartz with some adular in the form of little, frequently idiomorphic interpositions with late

The geochemical association of the Khokhoy gold ore field elements is the following Au, Sb, Te and Tl. The main minerals of karst cavities are of fine grain quartz, chalcedony, opal, adular, sericite, calcite, barite, fluorite, goethite, limonite, hematite. Unnamed Fe, Mn, Te, and Tl carbonates, galena, weissbergite, berthierite, arsenopyrite, chalcocite, unnamed sulfide Re and W, avicennite, hollandite, chalcopyrite, acanthite, chlorargyrite, fine grain native gold and silver are rarely. Gold mineralization of the Khokhoy gold ore field has a hypogene-supergene nature. The loose gold-bearing rocks here are secondary formations, which originated from oxidation, disintegration, and redeposition in the karst cavities of primary ores such as pyrite-adular-quartz metasomatites that resulted from silicic-potassic metasomatism of carbonate rocks. Strong Tl enrichment occurs during potassic metasomatism, in the fault zones, as is well seen in the Khokhoy gold ore field. The Au–Tl–As–Sb–Te–Ba geochemical profile of the mineralization and its low-temperature formative conditions suggest the epithermal origin of the ores. With regard to structural-morphological and mineralogical-geochemical parameters, gold mineralization of the Khokhoy gold ore field is comparable to the Kuranakh-type gold deposits of the Central Aldan district of Yakutia, representing unique supergene, shallow-depth, friable ores with free gold localized in karst cavities. They are characterized by large reserves with a relatively low gold grade.



Figure 1. Structure of the Khokhoy gold ore field (Aldan shield, East Russia).



Figure 2. Ore bearing karst, (a) general look, (b) pyrite-adular-quartz jasperoids, (c) section of b.

2. Materials and Methods

We used grab samples and friable materials of argillaceous and sandy fragments, taken from the surface mines and the core drilling holes of the Khokhoy gold ore field karst (Table 1). The grab samples of 0.3–0.5 kg taken from the karst fragments were used for making polished slides. The friable materials of 3 kg, on the other hand, were hydroseparated and to get the heavy residue. All samples were investigated with the use of binocular microscope, then we detected individual gold lumps and the associated minerals, which we forced together under pressure, made solid with epoxide and polished. All the polished samples were investigated with the use of Jenavert microscope in the reflected light, photographed and made ready for microprobe analysis. In order to investigate the chemical composition of the native gold and the associating minerals (ore, vein, and supergene minerals) as well as to detect unknown minerals we used microprobe analysis. The samples were analyzed with Camebax microanalyzer (Cameca, Courbevoie, France). We investigated the composition of gold making 3–5 probes in the center and on the edges of the gold lumps (the analysis was performed by N. Khristoforova). The majority of the samples were analyzed with the use of scanning electron microscope JEOL JSM-6480LV and energy spectrometer by Oxford (JEOL, Tokyo, Japan), which was used for taking the picture (the analysis was performed by S. Popova and S. Karpova). We made the quantitative XPP analyses using OXFORD INCA ENERGY 350 (Oxford Instruments, UK). The analysis conditions are as follows, the accelerating voltage of 20 kV, the measuring current flow is 1.08 nA, the measurement time is 10 s. The photographing conditions are as follows, voltage is 20 kV, the current flow is 17 nA. The analytic lines are Cu, Fe, Zn–K α ; Sb, S–L α . The standards are gold 750‰—Au, Ag; HgTe (coloradoite)—Hg, Te; CuSbS₂ (chalcostibite)—Cu, Sb, S; ZnS (sphalerite)—Zn; CuFeS₂ (chalcopyrite)—Fe; PbS (galena)—Pb; FeAsS (arsenopyrite)—As; BaSO₄ (barite)—Ba; ZrSiO₄ (zircon)—Zr; manganese 100%—Mn.

Na	Commite Labor	G	rid			Secondary Gold		
NO.	Sample Laber -	Easting	Northing	No. Grains	Frimary Gold	Massive	Porous 0	
1	4004	6555710	21514040	2	2	0	0	
2	3237	6550789	21510138	1	1	0	0	
3	K-21	6556300	21513025	1	1	0	0	
4	12014	6556295	21501333	1	2	0	0	
	5-15-53, 54, 57-60, 64, 65, 67-74, 76-78, 80-82,							
5	84-87, 90, 94-95, 97-98, 100-103, 105-108, 110-113, 115-118, 6-15	6555705	21514030	82	0	72	10	
6	1-17-1-7	6556046	21513843	7	0	0	7	
7	1-17-8-14	6555603	21514221	7	0	0	7	
8	1-17-15	6552872	21511211	1	0	0	1	
9	1-17-16-21	6556022	21513960	6	0	0	6	

Table 1. Details of samples studied.

Limits of element detection (wt%) X-ray spectral microprobe analysis: Au 0.117, Ag 0.061, Hg 0.083, Cu 0.031, Fe 0.019, Pb 0.066, Bi 0.095. Limits of element detection (wt%) scanning electron microscope equipped with energy spectrometer: Au 1.84, Ag 0.96, Hg 1.6, Cu 1.22, Fe 1.04, Pb 1.78, Bi 2.7.

3. Results

3.1. Supergene Gold of the Khokhoy Gold Ore Field

Primary Native Gold

Visually native gold in primary ores is extremely rare [10]. It is assumed that there is invisible finely divided gold, as evidenced by the gold content in the ore. Native gold is largely attested in porous oxidized pyrite relics or in quartz associated with barite and galena as submicroscopic cloddy effusions of the size up to 5 μ m, which is why we could not make conditional analysis. The native gold on the whole barely contains any impurities (970–999‰), which is consistent with the ratio of Au/Ag in ores 100/1. Very rarely, grains with up to 15 wt% of silver are attested.

3.2. Supergene Gold of the Karst Cavities

The physicochemical processes that have taken place during the development of the karst have influenced the typomorphic characteristics of gold. The karst cavities gold is most frequently attested in loose condition. At the same time, it is bigger in its size than that of primary ores. The Khokhoy gold ore field gold is represented by two types: massive and porous.

3.2.1. Massive Gold

Massive gold is represented by crystals and irregular mass (Figures 3 and 4). Crystals are rarely attested in karst holes and make 1–2% of the total mass. They are represented by individuals of dodecahedral (Figure 3a) and octahedral (Figure 3b) shapes with clear cut facets, and rarely with smooth edges. On the surface of the facets some dints and dents indicating sliding processes are attested. The size of the crystals and irregular mass of gold is reaches 0.2 mm. The fineness of the gold is very high, namely 950–1000‰ (Table 2; Nos. 1 and 2).

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Figure 3. The dodecahedral (**a**) and octahedral (**b**) crystals of Khokhoy ore field. Mounted polished sections, scanning electronic microscope in back-scattering electron mode. The numbers in microphotographs are analitical spots, here and in other figures.



Figure 4. Massive gold. (a) gold with Hg alloy, (b,c) pure gold, (d) gold with Ag alloy. Mounted polished sections, scanning electronic microscope in back-scattering electron mode.

No.	Sample *	Sp.	Ag	Au	Hg	Total	Fineness
1	IZ 10	1	-	98.31	-	98.31	1000
2	K_12	2	4.73	94.28	-	99.01	952
3		1	-	94.01	3.61	97.62	963
4	5_15-64	2	-	95.90	2.92	98.82	970
5	- 1 4	1	-	100.48	-	100.48	1000
6	5_15-74	2	0.64	99.11	-	99.75	993
7	- 4- 04	1	-	98.16	-	98.16	1000
8	5_15-84	2	-	98.39	-	98.39	1000
9		1	4.42	95.13	-	99.55	956
10	5_15-85	2	-	101.40	-	101.40	1000
11		1	-	101.78	-	101.78	1000
12		2	-	98.71	-	98.71	1000
13	1_17-5	3	-	99.08	-	99.08	1000
14		4	-	99.35	-	99.35	1000
15		1	-	95.56	4.57	100.13	954
16		3	0.28	96.64	3.54	100.47	962
17		4	-	100.29	-	100.29	1000
18		5	-	100.91	-	100.91	1000
19	1 17-15	6	0.37	96.37	2.85	99.59	967
20		7	-	95.50	4.74	100.24	953
21		8	-	96.75	3.08	99.83	969
22		9	-	98.85	1.56	100 41	984
23		10	-	93.91	4.33	98 24	956
24		10	_	98.45	-	98.45	1000
25		1	_	100.57	_	100 57	1000
26		5	_	99 77	_	99 77	1000
20	1 17-13	6		97.90		97.90	1000
28	1_17 10	7	-	97.20	_	97.90	1000
20		8	_	101.06	_	101.06	1000
30		9		00 11		99.11	1000
31		1	-	99.01	_	99.01	1000
22		2	5.08	92.01	-	08 10	020
22		2	5.90	92.21 100.1 2	-	100.12	1000
24	1_17-18	4	-	95 59	-	08.36	072
25		5	2.77	97.49	-	97.49	1000
33 26		6	-	97.49	-	97.49	028
27		0	0.00	92.30 0E 41	-	90.04	930
37 20		2	1.25	95.41	2.57	96.57	900
20	1_17-7	0	1.25	90.70	2.69	100.62	962
39 40		0	0.95	97.10 05.51	2.00	100.81	904
40		9	2.05	93.31	5.04 1.54	100.38	949
41		1	1.97	97.07	1.34	100.18	975
42		2	-	100.08	-	100.00	1000
43 44	1 17 14	3	- 21/	77./0	- 2.01	77.70 08.10	027
44	1_1/-14	4	5.14	92.52	5.01	90.40	937
43 46		3 7	2.05	100.20	-	100.20	1000
40 47		/	2.93	93.4Z	1.17	77.54 00.20	708 054
4/ 10		ð	- E 40	95.05	4.33	77.38 08.92	936
48		2	5.40	93.43	-	98.83	945
49 50	1 17 00	4	3.99	97.98	-	101.97	961
50	1_17-20	5	3.13	96.10	-	99.24	968
51		6	2.18	98.61	-	100.78	978
52		7	2.54	96.86	-	99.40	974
53		3	-	98.02	-	98.02	1000
54	1_17-6	6	-	98.74	-	98.74	1000
55		7	-	98.78	-	98.78	1000

Table 2. The composition (in wt%) and fineness (in ‰) of the supergene gold shown in Figures 3, 4, 8 and 10.

Notes: * Nos. 1–10—massive gold; 1–2—crystals; 3–10—irregular mass gold; 11–55—porous gold. Sp.—analysis spot in figures. The analyzes were performed on a scanning electron microscope equipped with energy spectrometer.

99.82

_

99.82

1000

56

8

The most widespread are cloddy gold particles of massive structure (Figure 4). In terms of morphology, they belong to the irregularly-shaped type, and associate closely with various minerals formed from the weathering of primary ores (Figure 5). The supergene gold often occurs in association with unnamed Fe, Te, and Tl carbonates (Figure 5a, Table 3; Nos. 1–8). Aggregates of intimately intergrown unnamed tellurates of Tl and native gold are occasionally observed along cracks in the gold particles (Figure 5b, Table 3; Nos. 9–15). In the interstices between massive gold there are found dendritic gold particles closely interwoven with goethite (Figure 5c, Table 3, Nos. 16–23). The most intricate forms of gold are seen in a very rare mineral avicennite (Tl_2O_3), an oxide of thallium with a Te impurity (Figure 5d, Table 3; Nos. 24–39). Sometimes massive gold occurs in assemblage with siderite (Figure 5e,f, Table 3; Nos. 40–43) and goethite (Figure 6, Table 3, Nos. 14–15). At a higher magnification, micron gold inclusions are visible in associating minerals (Figure 5f, Table 3, Nos. 44–55; Figure 6, Nos. 9–13). At a higher magnification, micron-sized gold grains are seen in siderite (Figure 5f, Table 3; Nos. 44–55). Fineness of the monolithic gold varies from 835 to 1000‰, with the high-fineness particles prevailing (Tables 2 and 3). Along with Ag, trace elements include mercury (up to 3.61 wt%), bismuth, and, more rarely Fe, Cu, Zn, and Pb (Table 4).



Figure 5. Massive gold in association with various supergene minerals. (**a**) massive gold in association with unnamed carbonates Te, Fe, Tl, (**b**) massive gold with inclusions of unnamed tellurates Tl, (**c**) honeycomb gold at the edges, sponge gold in the center in close association with goethite, (**d**) inclusion of massive and sponge gold in avicennite, (**e**) massive gold in association with siderite, (**f**) enlarged fragment of e. Mounted sections, scanning electron microscope, image in the back-scattering electron mode.

No.	Sample	Sp.	Au	Ag	Hg	Te	Fe	Tl	Si	0	Total
1	5_15_73	1	99.61	-	1.48	-	-	-	-	-	101.09
2	5_15_73	2	101.61	-	-	-	-	-	-	-	101.61
3	5_15_73	6	96.87	1.38	2.6	-	-	-	-	-	100.85
4	5_15_73	7	-	-	-	13.73	11.23	22.78	1.53	17.78	67.05
5	5_15_73	8	-	-	-	10.97	12.29	15.58	0.92	15.1	54.86
6	5_15_73	9	-	-	-	13.13	12.44	20.99	2.51	17.26	66.33
7	5 15 73	10	-	-	-	16.08	14.58	24.42		16.08	71.16
8	5 15 73	11	-	-	-	14.93	11.82	20.9	1.39	14.51	63.55
9	5 15 101	1	50.27	3.62	-	-	23.15	-	-	23.06	100.1
10	5_15_101	2	50.45	2.56	-	-	20.81	-	-	25.94	99.76
11	5 15 101	3	52.45	2.6	-	-	20.57	-	-	23.26	98.88
12	5 15 101	4	53.02	3.95	-	-	21.93	-	-	19.83	98.73
13	5 15 101	5	51.43	1.64	-	-	24.96	-	-	30.32	108.35
14	5 15 101	6	92.07	7.39	-	-	-	-	-	-	99.46
15	5 15 101	7	93.01	6.74	-	-	-	-	-	-	99.75
16	1 17 11	1	93.41	5.64	-	-	-	-	-	-	99.05
17	1 17 11	2	98.17	-	-	-	-	-	-	-	98.17
18	1 17 11	3	94.18	6.00	-	-	-	-	-	-	100.18
19	1 17 11	4	93.28	-	-	-	1.68	-	-	4.01	98.97
20	1 17 11	5	98.79	-	_	-	-	-	-	-	98.79
21	1 17 11	6	97 45	2.33	-	-	_	-	-	-	99.78
22	1 17 11	7	-	-	-	4 67	60 42	-	-	35 45	100.54
23	1_{17}_{11}	8	_	-	-	5.48	58 33	_	_	35.23	99.04
20	5 15 3	1	99 39	-	-	-	-	_	_	-	99 39
25	5 15 3	2	99 39	-	-	_	_	_	_	-	99 39
26	5 15 3	3	94 79	6.28	_	_	_	_	_	_	101.07
27	5 15 3	4	93.91	6.42	_	_	_	_	_	_	101.07
28	5 15 3	5	93 39	<u> </u>	_	_	_	_	_	_	97.83
20	5 15 3	6	91 78	5 77	_	_	_	_	_	_	97.55
30	5 15 3	7	93.87	5.91	_	_	_	_	_	_	99 78
31	5 15 3	8	93.13	4 56	_	_	_	_	_	_	97.69
32	5 15 3	9	-		_	4 93	_	75 65	_	17 25	97.83
33	5_15_5 5_15_3	10	_	_	_	3.96	_	81.6	_	13.65	99.21
34	5 15 3	11	_	_	_	75	_	79.55	_	13.00	100 54
35	5 15 3	12	_	_		5.15	_	83.25	_	13.47	100.54 101.77
36	5 15 3	12	_	_		5.13	_	78 79	_	13.07	97 73
37	5_15_3	13	-	-	-	3.85	-	81 00	-	13.41	97.73
38	5_15_3	15	-	-	-	774	-	76 21	-	15.04	97.90
30	5_15_3	15	-	-	-	1.74	-	20.21 81.01	-	16.16	102.04
39 40	5 15 115	10	- 02.00	-	-	4.07	-	01.91	-	10.10	00.06
40	5 15 115	2	93.99	6.95	-	-	-	-	-	-	99.00
41	5 15 115	2	92.51	0.95	-	-	45.22	-	3 20	20.85	70.70
42	5 15 115	4	-	-	-	-	45.22	-	2.20	30.24	77.27
43	5_15-115 5_15_115f	1	- 00.63	7.02	-	-	45.20	-	2.23	50.24	07.65
44	5_15-1151 E 1E 11Ef	1	90.03	6.80	-	-	-	-	-	-	97.03
43	5_15-1151 E 1E 11Ef	2	91.13	6.65	-	-	-	-	-	-	97.93
40	5_15-1151	3	92.30	0.03	-	-	-	-	-	-	99.13
47	5_15-1151	4	90.37	7.36	-	-	10 70	-	-	- 20/10	90.15 70.00
40	5_15-1151	5	-	-	-	-	40.20	-	2.24	20.40	79.00
49 50	5_15-115f	07	-	-	-	-	47.20	-	2.10	∠7.3U 20.04	70.8U
50	5_15-115f	/	-	-	-	-	40.93	-	2.27	30.94	0U.00
51	5_15-115f	ð	-	-	-	-	40.23	-	2.32	30.48	01.03 70.04
52 52	5_15-115f	9 10	-	-	-	-	47.95	-	2.42	20.00	/9.00
53 E 4	5_15-115f	10	-	-	-	-	48.55	-	2.07	30.76	01.38 05.94
54	5_15-115f	11	15.33	2.87	-	-	38.46	-	1.89	37.28	95.84
55	5_15-115t	12	15.66	-	-	-	41.85	-	2.23	36.90	96.64

Table 3. The composition of the massive gold and associated minerals (in wt%) shown in Figures 5 and 6.

Note: The analyzes were performed on a scanning electron microscope.



Figure 6. Massive gold with goethite: (a) general view, (b) enlarged goethite fragment with microinclusions of gold in backscattered electrons, (**c**–**f**) in X-rays of O, Fe, Au and Ag.

Table 4. The composition (in wt%) and fineness (in ‰) of the massive gold (in wt%) shown in Figure	e 11.
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No.	Sample	Fe	Cu	Pt	Ag	Au	Hg	Pb	Bi	Total	Fineness
1	5-15-53	0.01	-	-	1.29	100.15	0.01	0.10	0.17	101.73	984
2	5-15-54	0.05	-	-	3.18	94.70	0.10	0.03	-	98.06	965
3	5-15-57	0.06	0.00	0.11	0.28	97.43	0.08	0.21	0.19	98.35	990
4	5-15-58	0.04	-	0.07	1.56	97.17	0.10	0.08	0.12	99.14	980
5	5-15-59	0.03	-	0.08	4.71	93.17	-	0.10	0.18	98.27	947
6	5-15-60	0.02	-	0.01	1.59	98.26	0.02	0.07	0.08	100.05	981
7	5-15-64	0.01	-	0.01	1.40	96.88	0.02	0.07	0.18	98.57	982
8	5-15-65	-	-	-	0.93	99.64	0.10	0.10	0.13	100.89	987
9	5-15-67	0.01	-	0.04	1.58	97.67	0.07	0.02	0.06	99.45	981
10	5-15-68	-	-	-	1.13	97.37	0.14	0.10	0.23	98.97	983
11	5-15-69	0.01	0.00	0.09	0.83	96.84	0.14	0.10	0.18	98.19	985
12	5-15-70	0.01	0.02	-	1.48	99.47	0.08	0.20	0.22	101.48	980
13	5-15-71	-	0.01	0.16	1.11	99.05	-	0.11	0.21	100.66	984
14	5-15-72	0.01	-	0.01	1.02	100.11	0.08	0.06	0.12	101.39	987
15	5-15-73	-	0.01	0.05	0.73	99.35	0.03	0.05	0.15	100.38	990
16	5-15-74	0.01	-	-	0.60	100.57	0.01	0.06	0.12	101.38	992
17	5-15-76	-	0.00	0.14	1.33	96.65	0.05	0.03	0.10	98.30	982
18	5-15-77	-	-	-	1.46	97.81	0.13	0.03	0.06	99.48	983
19	5-15-78	0.00	0.01	-	0.71	99.74	-	0.12	-	100.58	992
20	5-15-80	0.01	0.01	0.13	0.92	97.97	0.11	0.05	0.19	99.38	986
21	5-15-81	0.02	0.02	0.13	0.90	97.48	0.10	-	0.14	98.78	986
22	5-15-82	0.04	0.02	0.05	0.91	98.35	0.03	0.12	0.17	99.68	987
23	5-15-84	0.00	0.02	0.15	0.88	98.13	0.01	-	0.19	99.38	987
24	5-15-85	-	0.01	0.01	7.46	90.65	0.04	-	0.03	98.20	923
25	5-15-86	0.07	0.02	0.05	5.13	92.93	0.26	0.06	0.03	98.55	942
26	5-15-87	0.01	0.01	0.04	1.00	98.02	0.19	0.13	0.09	99.47	985
27	5-15-90	-	0.01	0.02	1.45	96.79	0.02	0.05	0.19	98.53	982
28	5-15-94	-	0.01	-	1.29	99.27	0.12	0.08	0.25	101.02	983
29	5-15-95	0.12	-	0.05	8.14	92.21	-	0.05	0.02	100.60	917
30	5-15-97	0.01	0.11	0.08	2.99	97.31	0.09	0.03	0.10	100.79	965
31	5-15-98	-	0.13	0.06	2.62	98.32	-	0.00	-	101.20	972
32	5-15-100	-	0.09	0.10	3.20	94.78	0.04	0.02	0.06	98.30	964
33	5-15-101	-	0.08	0.08	7.27	90.71	0.11	-	0.05	98.35	922
34	5-15-102	0.00	0.03	0.01	7.41	90.18	0.01	0.07	0.04	97.88	921
35	5-15-103	0.00	0.08	0.04	4.12	93.65	0.05	0.07	0.10	98.13	954
36	5-15-105	-	0.07	0.07	7.39	91.03	-	0.02	0.23	98.84	921
37	5-15-106	-	0.06	0.09	7.16	91.45	-	-	-	98.79	926
38	5-15-107	0.15	0.01	0.08	15.95	81.64	-	0.05	0.01	97.93	834
39	5-15-108	0.00	0.09	0.14	3.77	94.80	-	0.04	0.17	99.09	957
40	5-15-110	0.11	0.04	0.17	8.49	87.04	0.16	0.77	0.17	96.96	898
41	5-15-111	-	0.04	0.06	7.17	91.49	0.06	0.04	0.20	99.08	923
42	5-15-112	0.00	0.05	0.01	6.88	92.73	-	0.02	0.14	99.82	929
43	5-15-113	0.01	0.01	-	7.36	92.62	0.06	0.06	0.15	100.28	924
44	5-15-115	-	0.08	0.05	3.68	97.45	-	0.01	0.15	101.42	961
45	5-15-116	-	0.08	0.03	3.37	97.24	-	0.04	0.09	100.86	964
46	5-15-117	0.00	0.07	0.07	7.09	91.55	-	-	0.14	98.96	925
47	5-15-118	0.11	0.01	0.05	6.47	93.29	0.06	-	-	100.04	933

Note: Analyzes performed on an X-ray microanalyzer.

3.2.2. Porous Gold

Besides the massive gold, porous gold particles can be found. Prevalent are irregular mass, with lesser flattened forms. Both varieties are characterized by microporosity and a mustard color, indicating a wide occurrence of mustard gold in the Khokhoy gold ore field (Figure 7). The term "mustard gold" was got in by W. Lindgren [11]. Typical features of mustard gold are low reflectivity, porous or colloidal texture and rusty, reddish, orange-red and brown-yellow colors in reflected light. It is characteristic of gold-telluride and gold-antimony deposits.





Figure 7. Native gold of the Khokhoy gold ore field, (**a**) microporous spongy-mustard gold, (**b**) framework gold with massive gold in the center, (**c**,**d**) dendritic mustard gold. Mounted sections, ore microscope, image in reflected light.

"Mustard" gold in the ore field according to the internal structure is microporous and dendritic. Porous irregular mass and flattened gold are friable aggregates of fulvous color. The porosity of gold is clearly reflected in the investigation of gold in a scanning electron microscope. The closeness of mustard and sponge gold is often observed (Figure 8a,b). Sometimes, massive gold grains occur with the development of hollow spongy and mustard gold closer to the center (Figure 8c). Less commonly observed is microporous mustard gold with elongated massive gold fibers in the central part (Figure 8d). The mustard gold micropores can be hollow or filled with various chemical elements Fe, Te, Cu, Mn, Sb, and Tl.





Figure 8. Mustard and sponge gold of microporous structure: (**a**,**b**) closely located mustard and sponge gold, (**c**) massive gold with spongy and mustard gold close the center, (**d**) microporous mustard gold with elongated fibers of massive gold in the center. Mounted sections, scanning electron microscope, image in the backscattering electron mode.

Mustard gold of dendritic structure is more widespread (Figure 9). Dendrites are represented in the center by dendritic-branched lumps of gold surrounded by spongy mustard gold (Figure 9a) in close association with iron hydroxides. Sometimes there are looped massive gold grains interspersed with mustard-sponge gold (Figure 9b). Varieties of mustard twiggy gold prevail, the pores of which are mostly hollow (Figure 9c,d).

According to microprobe analyses, in individual grains one can assume the presence of iron oxides (hydroxides) in the pores. Figure 10 shows the internal structure of the enlarged fragments of dendritic mustard gold.



Figure 9. Mustard gold of dendritic structure, (**a**,**b**) dendritic gold in the center, meandering massive golds, mustard porous gold along the edges in close association with iron hydroxides, (**c**–**f**) varieties of twiggy gold. Mounted sections, scanning electron microscope, image in the backscattering electron mode.



Figure 10. The internal structure of mustard-spongy (**a**,**b**,**h**) and dendritic mustard (**c**–**g**) gold. Fragments of the gold shown in Figures 6–8. (**a**,**b**) brain structure; (**c**,**d**) massive golds are cemented with microporous mustard gold; (**e**,**f**) tortuous gold in the cement of microporous mustard-sponge gold; (**g**) streaky-meandering golds in the center are bordered with brain-shaped golds; (**h**) a lump of gold of the brain-structure. Mounted sections, scanning electron microscope, back-scattering electron mode.

Sponge and mustard gold fineness is above >900‰. Among impurities, in addition to silver, Hg (up to 5.48 wt%) and Bi (up to 0.42 wt%) are constantly present, the remaining elements Fe, Zn, Pb, Cu, and Pt are found sporadically (Table 2, Nos. 11–56, Table 5).

No.	Sample	Sp.	Fe	Cu	Zn	Ag	Au	Hg	Pb	Bi	Total	Fineness
1	1-17-1	1	-	0.02	-	4.79	94.29	0.15	-	0.15	99.40	949
2	1-17-1	2	0.06	-	0.01	1.01	96.85	0.07	0.08	0.10	98.17	986
3	1-17-2	1	0.05	0.02	-	8.86	89.44	-	0.08	0.05	98.50	908
4	1-17-6	1	0.05	0.02	-	0.852	97.14	0.02	0.01	0.16	98.25	989
5	1-17-6	2	0.03	0.02	0.02	0.717	99.28	0.17	-	0.22	100.46	988
6	1-17-6	3	0.03	0.01	-	2.52	94.86	0.22	0.01	0.08	97.73	971
7	1-17-7	1	-	0.04	-	0.752	98.68	0.09	0.03	0.15	99.74	989
8	1-17-7	2	-	0.02	-	0.092	99.74	0.25	0.08	0.16	100.34	994
9	1-17-7	3	0.05	0.01	0.01	0.194	99.6	0.26	0.04	0.26	100.42	992
10	1-17-7	4	0.06	0.02	-	0.627	97.98	0.14	-	0.12	98.945	990
11	1-17-8	1	-	-	-	5.17	95.24	-	0.02	0.10	100.54	947
12	1-17-8	2	-	-	-	5.67	95.96	0.07	0.04	0.09	101.84	942
13	1-17-8	3	-	0.02	-	6.07	94.35	0.27	-	0.07	100.77	936
14	1-17-8	4	-	0.02	0.01	5.09	91.64	0.18	0.04	0.12	97.1	944
15	1-17-8	5	-	-	-	4.89	95.79	0.29	0.02	0.14	101.13	947
16	1-17-9	1	-	-	-	6.04	92.1	0.27	-	0.01	98.42	936
17	1-17-9	2	0.01	0.01	-	3.28	96.79	0.17	-	0.14	100.40	964
18	1-17-10	1	-	0.01	-	1.07	98.75	0.02	0.11	0.16	100.11	986
19	1-17-10	2	-	0.01	-	1.16	99.94	0.32	0.09	0.13	101.66	983
20	1-17-10	3	-	-	-	1.19	99.93	0.10	-	0.28	101.50	985
21	1-17-11	1	-	-	-	1.42	98.56	0.24	0.02	0.25	100.49	981
22	1-17-11	2	-	0.01	0.012	3.99	96.46	0.19	-	0.17	100.83	957
23	1-17-11	3	-	0.01	-	3.95	96.14	0.35	0.03	0.18	100.66	955
24	1-17-12	1	-	-	-	2.56	97.78	0.20	0.07	0.02	100.63	972
25	1-17-13	1	-	0.03	-	-	99.62	-	-	0.24	99.89	997
26	1-17-14	1	0.07	-	-	0.79	98.82	0.1	0.05	0.22	100.04	988
27	1-17-15	1	0.03	-	0.03	0.30	96.62	0.30	0.10	0.24	97.62	990
28	1-17-17	1	-	-	0.02	0.29	99.72	0.27	0.01	0.24	100.56	992
29	1-17-18	1	-	-	-	3.87	90.87	4.63	-	0.18	99.55	913
30	1-17-19	1	0.04	-	-	6.78	91.49	2.97	-	0.066	101.35	903
31	1-17-21	1	-	-	-	1.88	97.46	0.22	0.03	0.42	100.01	975
32	1-17-21	2	0.02	0.02	0.01	4.00	93.29	0.04	-	0.18	97.56	956

Table 5. The composition (in wt%) and fineness (in ‰) of the porous gold shown in Figure 11.

Note: Analyzes performed on an X-ray microanalyzer.

All gold varieties are chemically homogeneous and characterized by high fineness of the gold particles without any rims or margins. The distribution patterns of fineness in both types of native gold massive and porous, do not differ essentially. The high-fineness gold (>950‰) makes up 64%, the amount of particles of lower fineness (800–950‰) comes to only 36% of the total. Along with Ag, the porous native gold invariably contains mercury (up to 5.78 wt%) and bismuth, and, less frequently, lead, copper, and iron.

Cumulative diagrams clearly show high concentrations of Pb and Cu in the massive gold, and of Hg in the porous one (Figure 11). The amount of Fe and Bi impurities in both types of supergene gold is similar.



Figure 11. Cumulative plots of trace element ((**a**) Au, (**b**) Pb, (**c**) Cu, (**d**) Fe, (**e**) Bi, (**f**) Hg) vs. silver concentrations in massive (green triangle) and porouse (blue rombus) gold of the Khokhoy gold ore field.

4. Discussion

Two genetic types of supergene gold, massive and porous, are recognized there.

The first type is represented by gold crystals and irregular mass, with the fineness ranging from 835 to 1000‰. They are closely associated with goethite, siderite, unnamed Fe, Te, and Tl carbonates, Tl tellurites/tellurates and antimonates, as well as avicennite with a Te impurity.

The second type is represented by mustard gold with different internal structure: microporous and dendritic.

Evidentially, mustard gold is characteristic of gold-telluride [12–15] and gold-antimony [16–18] deposits, weathering crust [19–21] and placers [22,23]. Recently, mustard gold, sometimes with an admixture of Tl (up to 6.34 wt%), was found in the Oleninsky Au–Ag deposit (Kola Peninsula, Russia) [24]. In all the cases, mustard gold was developed in the oxidation zone as a result of the decomposition of tellurides, antimonides, sulfides, bismuthides Au (Ag) and low-fineness gold [14]. In the Kuranakh deposit, mustard gold is the product of the decomposition of tellurates [25]. Researchers consider the occurrence of mustard gold a result of hypogenous as well as supergenous processes. The proponents of its hypogenic origin believe that mustard gold, Au tellurates/antimonates, and

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complex gold oxides were formed from hydrothermal low-temperature solutions with a high oxidation potential [16]. Supergene processes explain the origin of the secondary mustard gold of high fineness of placers in the North-East of Russia [23].

The formation of mustard gold at the Dongping mines (Hebei province, China) has been related to decomposition of calaverite by selective leaching of tellurium while leaving the gold alloy in the cavity formed by the alteration reaction [12,13]. This type of pseudomorphic alteration was also documented by Palache et al. [26]. The occurrence of microporous gold has also been observed under cold climatic conditions, such as at the Aginskoe low-sulfidation ephithermal deposit in Central Kamchatka, Russia [27]. In this deposit, calaverite is the main Au telluride mineral and it has been partially replaced by porous gold. By comparing the textures of microporous gold from this natural occurrence with those obtained experimentally via the dealloying of gold–(silver) tellurides [28–30], Okrugin et al. [31] confirmed that natural microporous gold can form via the replacement of telluride minerals and assessed the role that hydrothermal fluids may play in the formation of microporous gold. The formation of the secondary high-fineness mustard gold from placers in northeast Russia is interpreted to be due to supergene processes. Although no antimonides, nor tellurides, or gold bismuthides were found in the ores of the Khokhoy gold ore field, their presence is assumed by the minerals associated with supergenic gold and the microchemical composition. They are represented by Fe, Mn and Tl tellurates and carbonates, as well as avicennite, Tl antimonates, goethite, limonite, hydrogoethite, and siderite of which the breakdown contributes to the formation of microporous gold particles of spungy habit.

So, what is the formation mechanism of dendrites? Interesting data on the formation of native gold dendrites in epithermal ores are presented by Saunders [32–34]. Author showed a scanning electron microscope (SEM) image of electrum dendrites that appeared to have formed from aggregation of smaller colloid-sized particles. Metallic nanoparticles appear to form from supersaturated solutions and can form dendrites by the self-assembly and aggregation of the nanoparticles. These dendrites are typically an intermediary stage to more traditional crystal formation as the infilling of branches of the dendrites occurs. In epithermal ores, the dendrites of electrum appear to be preserved due to the infilling of other nanoparticles between the branches, such as silica nanoparticles. The 'fractal' electrum dendrites have been observed in many the Tertiary bonanza epithermal ores in northern Nevada, and have been interpreted to be evidence of nanoparticle nucleation and aggregation in ore formation. More recently, similar textures and genetic interpretations have been made from ores from the southeastern USA [33] and Bulgaria [35]. Finally, disseminated electrum nanoparticles have recently been discovered in the epithermal Round Mountain (Nevada) deposit and have been proposed to be precursors for coarser electrum crystals there [36,37].

We believe that this is a more suitable mechanism of formation for the supergene gold of the Khokhoy gold ore field. The colloidal gold transfer is thought to be possible not only in hydrothermal conditions but in supergene zones too [38]. In the oxidation zone of low-sulfidation ores, to which group the Khokhoy ores belong, the role of gold colloids could be significant. Nanoparticles or colloids of gold with nanoparticles of silica (opal, chalcedony) could form dendrites by the self-assembly and aggregation of the nanoparticles. These dendrites were an intermediary stage to the formation of massive gold particles and crystals by the infilling of the dendrite branches. The massive gold type occurs in the most hypsometrically high levels of karst formations of the Khokhoy gold ore field.

The style of mineralization is among the main factors defining the microchemical composition of native gold, particularly the placer gold [39]. In the Khokhoy gold ore field, the main ore minerals associated with primary gold are pyrite, hematite, galena, and chalcosine. Supergene gold occurs in paragenesis with goethite, siderite, oxides, carbonates, tellurates, and antimonates of thallium. The set of trace elements (Cu, Pb, Fe), both in the massive and porous gold, corresponds to main elements of the associated minerals. The constant presence of Bi in both gold types suggests an intrusive source for them. The preferential accumulation of Hg in porous gold may be explained by its friable texture.

The lack of gold placers in the Khokhoy gold ore field may be due to that microaggregates of mustard gold are instable in supergene process because of friable texture, thus having a weak potential for placer formation.

The fact that native gold is unvisible and fine-grained in primary ores and visible in the karst cavities indicates that it grew larger in size in the oxidation zone of karst formations.

5. Conclusions

Relationships of massive gold with unnamed tellurates and carbonates of thallium and with avicennite (Te_2O_3) are first described for the Khokhoy gold ore field. Along with massive gold closely associated with Tl minerals, there are abundant porous particles, the so called spungy and mustard gold.

The research has shown that the Khokhoy gold ore field, according to its mineralogical and geochemical features, should be classified as a deposit of Au–Te–Sb–Tl mineral composition localized in karst cavities. A typical gold ore deposit in karst cavities is the Kuranakh deposit of Central Aldan district of South Yakutia [6], with which our data is being compared. It should be noted that in the Kuranakh deposit one of the typomorphic geochemical elements is thallium, but its mineral form has not been identified. In the Khokhoy ore field, a diverse spectrum of thallium minerals is attested. This fact makes closer to Carlin-type gold deposits in the West of the USA [1–5], Alshar epithermal Au–As–Hg–Tl deposit in Macedonia [40,41], Vorontsovskoe deposit in the Urals, Russia [7–9,42] etc. The main difference between the Khokhoy gold ore field and these deposits at this stage of research is the absence of As and Hg minerals—realgar, auripigment, cinnabar. In addition to thallium minerals, Sb and Te minerals (weissbergite, antimonite, berthierite, and unnamed antimonates and tellurates of thallium) are widespread in the ores.

Supergene gold in karst cavities has specific features: (1) morphology—massive (gold crystals and cloddy particles of monolithic structure) and porous (microporous and dendritic); (2) associated minerals—goethite, limonite, avicennite, hydrogoethite, and siderite, Fe, Mn and Tl tellurates and carbonates, Tl antimonates; (3) persistently high fineness and chemically homogeneous; (4) microchemical elements Hg, Bi, Fe. The obtained characteristics made it possible to prove the existence of two genetic types of supergene gold. Mustard microporous gold is the result of the decomposition of the associated goethite minerals, Tl oxides, tellurates, carbonates and antimonates Fe, Mn and Tl containing microinclusions of gold. Massive gold and dendrites are newly formed. The decomposition, remobilization, and reprecipitation of residual gold nanoparticles and their aggregation led to the formation of dendrites, and with further crystal growth and filling of pores, to gold of massive morphology.

The supergene gold of the Khokhoy gold ore field is comparable in its typomorphic characteristics to that of the Kuranakh deposit, as well as with to that of Carlin-type gold deposits in the West of the USA, to that of the Alshar epithermal deposit in Macedonia, to that of the Vorontsovskoe deposit in the Urals, Russia, and also have similarities with those of sponge and mustard gold of gold-telluride and gold-antimony deposits, weathering crusts and placers.

However, at the same time, the gold has certain uniqueness, namely the paragenesis with thallium minerals. To date, no such relationship has been cited in the literature.

Author Contributions: The idea of the research belongs to G.S.A., who has then designed it, analyzed the results and wrote the paper. L.A.K. was responsible for the fieldwork in the auriferous zone; she also supported the research with valuable ideas. V.N.K. was also responsible for the fieldwork in the auriferous zone; moreover she made the mineralgraphic description and took the photographs of the polished sections. All authors have read and agreed to the published version of the manuscript.

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References

- 1. Hofstra, A.H.; Cline, J.S. Characteristics and models for Carlin-type gold deposits. *Econ. Geol.* **2000**, *13*, 163–220.
- 2. Cline, J.S.; Hofstra, A.H.; Muntean, J.L.; Tosdal, R.M.; Kenneth, K.A. Carlin-Type Gold Deposits in Nevada: Critical Geologic Characteristics and Viable Models. *Econ. Geol.* **2005**, *100*, 451–484.
- 3. Volkov, A.V.; Sidorov, A.A. The geological-genetic model of Carlin type gold deposits. *Litosphere* **2016**, *6*, 145–165. (In Russian)
- 4. Zhong, H.R.; Chao, S.W.; Wu, B.X.; Zhi, T.G.; Hofstra, A.H. Geology and geochemistry of Carlin-type gold deposits in China. *Miner. Depos.* **2002**, *37*, 378–392. [CrossRef]
- 5. Xia, Y.; Su, W.; Zhang, Z.; Liu, J. Geochemistry and Metallogenic Model of Carlin-Type Gold Deposits in Southwest Guizhou Province, China. In *Geochemistry Earth's System Processes*; INTECH: London, UK, 2012.
- 6. Rodionov, S.M.; Fredericksen, R.S.; Berdnikov, N.V.; Yakubchuk, A.S. The Kuranakh epithermal gold deposit (Aldan Shield, East Russia). *Ore Geol. Rev.* **2014**, *59*, 55–65. [CrossRef]
- 7. Sazonov, V.N.; Murzin, V.V.; Grigor'ev, N.A. Vorontsovsk gold deposit: An example of carlin-type mineralization in the Urals, Russia. *Geol. Ore Depos.* **1998**, *40*, 139–151.
- 8. Murzin, V.V.; Naumov, E.A.; Azovskova, O.B.; Varlamov, D.A.; Rovnushkin, M.Y.; Pirajno, F. The Vorontsovskoe Au–Hg–As ore deposit (Northern Urals, Russia): Geological setting, ore mineralogy, geochemistry, geochronology and genetic model. *Ore Geol. Rev.* **2017**, *85*, 271–298. [CrossRef]
- 9. Vikentyev, I.V.; Tyukova, E.E.; Vikent'eva, O.V.; Chugaev, A.V.; Dubinina, E.O.; Prokofiev, V.Y.; Murzin, V.V. Vorontsovka Carlin-style gold deposit in the North Urals: Mineralogy, fluid inclusion and isotope data for genetic model. *Chem. Geol.* **2019**, *508*, 144–166. [CrossRef]
- 10. Anisimova, G.S.; Kondratieva, L.A.; Sokolov, E.P.; Kardashevskaya, V.N. Gold mineralization of the Lebedinsky and Kuranakh types in the Verkhneamginsky district (South Yakutia). *Otech. Geol.* **2018**, *5*, 3–13. (In Russian)
- 11. Lindgren, W. *Mineral Deposits*; McGraw-Hill Book Company, Inc.: New York, NY, USA; London, UK, 1933; p. 930.
- 12. Petersen, S.B.; Makovicky, E.; Li, J.L.; Rose-Hansen, J. Mustard gold from the Dongping Au–Te deposit, Hebei province, People's Republic of China. *Neues Jahrb. Mineral. Mh.* **1999**, *8*, 337–357.
- 13. Li, J.; Makovicky, E. New studies on mustard gold from the Dongping Mines, Hebei Province, China: The tellurian, plumbian, manganoan and mixed varieties. *Neues Jahrb. Mineral. Abh.* **2001**, *176*, 269–297.
- 14. Tolstykh, N.D.; Palyanova, G.A.; Bobrova, O.V.; Sidorov, E.G. Mustard Gold of the Gaching Ore Deposit (Maletoyvayam Ore Field, Kamchatka, Russia). *Minerals* **2019**, *9*, 489–506. [CrossRef]
- Tolstykh, N.; Vymazalova, A.; Tuhy, M.; Shapovalova, M. Conditions of formation of Au–Se–Te mineralization in the Gaching ore occurrence (Maletoyvayam ore field), Kamchatka, Russia. *Mineral. Mag.* 2018, 82, 649–674. [CrossRef]
- 16. Gamyanin, G.N.; Nekrasov, I.; Zhdanov, J.J.; Leskova, I.V. Auroantimonate—A new natural compound of gold. *Dokl. Akad. Nauk SSSR* **1988**, *301*, 947–950. (In Russian)
- 17. Amuzinsky, V.A.; Anisimova, G.S.; Zhdanov, Y.Y. *Sarylakhskoe and Sentachanskoe Gold Antimony Deposits*; Nauka/Interperiodika: Moscow, Russia, 2001; 218p. (In Russian)
- 18. Zacharias, J.; Nemec, M. Gold to aurostibite transformation and formation of Au–Ag–Sb phases: The Krásná Hora deposit, Czech Republic. *Mineral. Mag.* **2017**, *81*, 987–999. [CrossRef]
- 19. Kalinin, Y.A.; Kovalev, K.R.; Naumov, E.A.; Kirillov, M.V. Gold in the weathering crust the Suzdal' deposit (Kazakhstan). *Russ. Geol. Geogphys.* **2009**, *50*, 174–187. [CrossRef]
- 20. Kalinin, Y.A.; Palyanova, G.A.; Bortnikov, N.S.; Naumov, E.A.; Kovalev, K.R. Aggregation and Differentiation of Gold and Silver during the Formation of the Gold-Bearing Weathering Crusts (on the Example of Kazakhstan Deposits). *Dokl. Earth Sci.* **2018**, *482*, 1193–1198. [CrossRef]

- Kalinin, Y.A.; Palyanova, G.A.; Naumov, E.A.; Kovalev, K.R.; Pirajno, F. Supergene remobilization of Au in Au-bearing regolith related to orogenic deposits: A case study from Kazakhstan. Ore Geol. Rev. 2019, 109, 358–369. [CrossRef]
- 22. Reith, F.; Stewart, L.; Wakelin, S.A. Supergene Au transformation: Secondary and nano-particulate Au from southern New Zealand. *Chem. Geol.* **2012**, *320–321*, *32–45*. [CrossRef]
- 23. Litvinenko, I.S.; Shilina, L.A. Hypergene gold neomineralization in placer deposits of Nizhne-Myakitsky ore-placer field, North-East Russia. *J. Ores Met.* **2017**, *1*, 75–90. (In Russian)
- 24. Kalinin, A.A.; Savchenko, Y.E.; Selivanova, E.A. Mustard Gold in the Oleninskoe Gold Deposit, Kolmozero–Voronya Greenstone Belt, Kola Peninsula, Russia. *Minerals* **2019**, *9*, 786–800. [CrossRef]
- 25. Yablokova, S. About a new morphological variety of gold and its origin. *Dokl. Akad. Nauk SSSR* **1972**, 205, 936–939. (In Russian)
- 26. Palache, C.; Berman, H.; Frondel, C. Dana's System of Mineralogy I; Wiley: New York, NY, USA, 1944.
- Andreeva, E.D.; Matsueda, H.; Okrugin, V.M.; Takahashi, R.; Ono, S. Au–Ag–Te mineralization of the low-sulfidation epithermal Aginskoe deposit, Central Kamchatka, Russia. *Resour. Geol.* 2013, 63, 337–349. [CrossRef]
- 28. Xu, W.; Zhao, J.; Brugger, J.; Chen, G.; Pring, A. Mechanism of mineral transformations in krennerite, Au3AgTe8, under hydrothermal conditions. *Am. Miner.* **2013**, *98*, 2086–2095. [CrossRef]
- 29. Zhao, J.; Xia, F.; Pring, A.; Brugger, J.; Grundler, P.V.; Chen, G. A novel pre-treatment of calaverite by hydrothermal mineral replacement reactions. *Miner. Eng.* **2010**, *23*, 451–453. [CrossRef]
- Zhao, J.; Brugger, J.; Xia, F.; Ngothai, Y.; Chen, G.; Pring, A. Dissolution-reprecipitation vs. solid-state diffusion: Mechanism of mineral transformations in sylvanite, (AuAg)₂Te₄, under hydrothermal conditions. *Am. Miner.* 2013, *98*, 19–32. [CrossRef]
- 31. Okrugin, V.M.; Andreeva, E.; Etschmann, B.; Pring, A.; Li, K.; Zhao, J.; Griffiths, G.; Lumpkin, G.R.; Triani, G.; Brugger, J. Microporous gold: Comparison of textures from Nature and experiments. *Am. Miner.* **2014**, *99*, 1171–1174. [CrossRef]
- 32. Saunders, J.A.; Schoenly, P.A. Boiling, colloid nucleation and aggregation, and the genesis of bonanza gold mineralization at the Sleeper Deposit, Nevada. *Miner. Depos.* **1995**, *30*, 199–211. [CrossRef]
- 33. Saunders, J.A. Textural evidence of episodic introduction of metallic nanoparticles into bonanza epithermal ores. *Minerals* **2012**, *2*, 228–243. [CrossRef]
- 34. Saunders, J.A.; Burke, M. Formation and aggregation of gold (electrum) nanoparticles in epithermal ores. *Minerals* **2017**, *7*, 163–173. [CrossRef]
- 35. Marinova, I.; Ganev, V.; Titorenkova, R. Colloidal origin of colloform-banded textures in the Paleogene low-sulfidation Khan Krum gold deposit, SE Bulgaria. *Miner. Depos.* **2014**, *49*, 49–74. [CrossRef]
- 36. Burke, M. An Electron Microscopy Investigation of Gold and Associated Minerals from Round Mountain, Nevada. Master's Thesis, Miami University, Oxford, OH, USA, 2017.
- 37. Burke, M.; Rakovan, J.; Krekeler, M.P.S. A study by electron microscopy of gold and associated minerals from Round Mountain, Nevada. *Ore Geol. Rev.* **2017**, *91*, 708–717. [CrossRef]
- 38. Petrovskaya, N.V. Native Gold; Nauka: Moscow, Russian, 1973; 347p. (In Russian)
- 39. Chapman, R.J.; Mortensen, J.K.; LeBarge, W.P. Styles of lode gold mineralization contributing to the placers of the Indian River and Black Hills Creek, Yukon Territory, Canada as deduced from microchemical characterization of placer gold grains. *Miner. Depos.* **2011**, *46*, 881–903. [CrossRef]
- 40. Volkov, A.V.; Serafimovski, T.; Kochneva, N.T.; Tomson, I.N.; Tasev, G. The Alshar epithermal Au–As–Sb–Tl deposit, southern Macedonia. *Geol. Ore Depos.* **2006**, *48*, 175–192. [CrossRef]
- 41. Palinkaš, S.S.; Hofstra, A.H.; Percival, T.J.; Šoštari'c, S.B.; Palinkaš, L.; Bermanec, V.; Pecskay, Z.; Boev, B. Comparison of the Allchar Au–As–Sb–Tl Deposit, Republic of Macedonia, with Carlin-Type Gold Deposits. *Rev. Econ. Geol.* **2018**, *20*, 335–363.
- Kasatkin, A.V.; Makovicky, E.; Plášil, J.; Škoda, R.; Agakhanov, A.A.; Karpenko, V.Y.; Nestola, F. Tsygankoite, Mn₈Tl₈Hg₂(Sb₂₁Pb₂Tl)_{Σ24}S₄₈, a New Sulfosalt from the Vorontsovskoe Gold Deposit, Northern Urals, Russia. *Minerals* 2018, *8*, 218. [CrossRef]



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