



Article

Examining Storage Capacity Loss and Sedimentation Rate of Large Reservoirs in the Central U.S. Great Plains

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Abstract: Reservoirs created by impounding sediment-laden streams infill over time, reducing storage capacity and altering water quality. Increasing freshwater demand on a limited supply is adding pressure on reservoir water storage and management across much of the western and central U.S. Determining sedimentation rates is helpful to ensure a reliable and sustainable clean water supply for drinking, irrigation and recreation purposes. In the state of Kansas, located in the central Great Plains, bathymetric surveys have been completed recently for many major state and federally constructed reservoirs. In this paper, we examine sediment infill rate and storage capacity loss for all 24 federally operated reservoirs in Kansas. As of 2016, these reservoirs have an average age of 52 years and collectively have lost approximately 17% of their original capacity, with the highest single-reservoir loss of 45%, the highest annual loss rate of 0.84%, and the highest annual sediment yield of 1688 m³/km²/year. Results from this paper provide baseline information pertinent to the development, prioritization and assessment of remediation and mitigation strategies for addressing the sediment infill problem in Kansas, with general implications for other large reservoirs across the Great Plains as well as other climatologically and ecologically similar regions around the world.

Keywords: bathymetry; flood control; Great Plains; impoundment; Kansas; reservoir half-life; sediment management; sediment yield; soil erosion; water supply

1. Introduction

Streamflow has been managed by dam construction on large and small rivers for various purposes throughout modern human history, with dams in operation as old as Lake Homs water supply dam (c. 1300 BC) in Syria [1,2]. In areas with inadequate subsurface water supplies and highly variable surface flows, which describes much of the U.S. Great Plains, impoundments can provide water supply

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stability and also limit flooding by storing floodwater and releasing it gradually, thereby mitigating peak flows downstream.

Dam construction can have negative impacts on upstream and downstream ecosystems, water chemistry, stream flow and wetlands [3–7]. Sediment trapping in these reservoirs can be substantial [8–10], with the estimated mean annual global reservoir infill rate of 0.5–1% resulting in crucial storage loss, particularly in areas prone to severe storms [11]. Disproportionately high amounts of infill sediment are delivered during extreme high flow events typically associated with intense storms [12]. As human population increases, there is an associated higher demand for drinking and irrigation water. This presents a precarious situation in areas that have highly variable weather and extended drought periods such as the Great Plains, where reservoir sedimentation escalates water management concerns [13,14].

Kansas (Figure 1, inset map) is one of several states in the Great Plains region that depends heavily on reservoir-sourced drinking water, with about 80 reservoirs (including 17 of the 24 federal reservoirs examined in this study) providing primary or backup water supply for more than 60% of its residents [15]. To the south, the state of Texas has 188 reservoirs with a normal capacity greater than 6 Mm³ [16], many of which are used for drinking water supply. In Oklahoma, approximately 61 reservoirs have a normal capacity greater than 14 Mm³, all of which help meet the critical need for water in this state [17]. Eastern New Mexico, eastern Colorado, Nebraska and other Great Plains states similarly rely on reservoirs to meet at least part of their water supply needs (Figure 1). In the Great Plains, the majority of large federal reservoirs were designed and built in the 20th century principally to control flooding and secondarily to provide water for a number of human uses. Now well into the 21st century, many communities are faced with greater water demands but less available water due to reservoir sedimentation in conjunction with potentially more common episodic weather extremes, including both flooding and drought [18–24].

Recent efforts in Kansas have focused on estimating the amount of sediment accumulation in larger reservoirs. Bathymetric surveys have been initiated and supported primarily by the Kansas Water Office (KWO), the U.S. Army Corps of Engineers (USACE), and the U.S. Department of Interior—Bureau of Reclamation (BOR) during the last two decades [25,26]. High-quality lake-bottom topography and lake volume data for 24 federal and 55 state and local lakes resulted from these surveys (for convenience, we use "lake" and "reservoir" interchangeably in reference to impoundments, not to be confused with natural lakes). Due to the low detail of the original lake surveys (which, for this study, have generally been derived from pre-impoundment topographic contour maps as shown in several bathymetric reports available from [25]), there is inherent error when conducting infill volume estimates; however, the differencing approach used here has been found to be reasonably accurate [27,28]. An example of a more specific type of error that can enter into this kind of assessment can be seen in the bathymetric report for Clinton Lake [25], where a number of borrow pits along with an elevated road used for material transport during dam construction are visible in the lake bottom elevation change map. These features were present for the recent bathymetric survey but do not appear in the pre-impoundment topographic map. While such data artifacts are fairly common, they are generally of small magnitude relative to the whole-lake sediment volume estimates.

The objective of this study is to examine sedimentation (infill amount and rate) and capacity depletion (storage loss rate) of the 24 federally constructed reservoirs in Kansas (hereafter referred to as KS-FdR). Capacity refers to the volume of the *conservation pool*, which is the body of water below the normal operating level of a reservoir. The bathymetric survey information used in this study originated from data collected between 1998 and 2014 (Table 1), making sediment infill data the most accurate to date. None of the KS-FdR have been dredged prior to their most recent survey. However, a large-scale dredging project on John Redmond Reservoir occurred in 2016 [29], which is described later.

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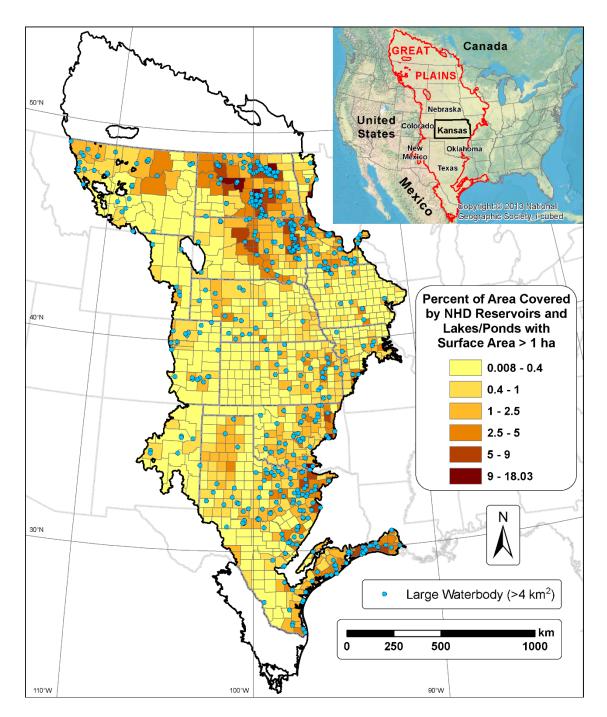


Figure 1. Waterbody summary information from the U.S. Geological Survey National Hydrography Dataset (NHD) is shown for the U.S. Great Plains [30,31]. Aggregation zones used in the main map are county boundaries. 24 federally constructed reservoirs in the central state of Kansas (most of which correspond with blue dots) are the focus of this sedimentation study.

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Table 1. Physical and hydrological characteristics of the 24 federally operated reservoirs in Kansas (KS-FdR).

Reservoir	W (km ²)	t ₀ (year)	V ₀ (Mm ³)	t _r (year)	V_r (Mm ³)	V_c (Mm ³)	L _c (%)	D (%)	Y (m ³ /km ² /year)	t _{1/2} (year)	Bathymetric Data Source *
Big Hill	95	1980	34	2010	29	28	17.3	0.48	1688	2085	USACE
Cedar Bluff	13,794	1950	229	2000	213	208	9.3	0.14	23	2306	BOR
Cheney	2566	1964	206	1998	197	193	6.5	0.12	99	2368	USGS
Clinton	950	1977	159	2009	146	143	9.9	0.26	433	2172	KBS
Council Grove	677	1963	65	2008	54	52	18.9	0.36	341	2104	KBS
El Dorado	634	1979	202	2004	196	193	4.7	0.13	410	2369	OWRB
Elk City	1650	1965	65	2008	46	43	34.2	0.67	263	2040	KBS
Fall River	1433	1948	37	2010	26	24	35.1	0.52	135	2045	KBS
Hillsdale	372	1981	101	2009	95	94	7.3	0.21	571	2220	KBS
John Redmond	6642	1963	126	2014	78	76	39.6	0.75	142	2030	KBS
Kanopolis	6322	1948	90	2007	60	55	38.9	0.57	81	2036	KBS
Keith Sebelius	1799	1964	45	2000	43	42	6.4	0.12	30	2373	BOR
Kirwin	3635	1955	123	2013	117	117	4.8	0.08	26	2597	KBS
Lovewell	894	1957	55	2011	44	43	21.3	0.36	223	2096	KBS
Marion	535	1967	105	2008	99	98	6.0	0.12	243	2371	KBS
Melvern	869	1972	190	2009	185	183	3.6	0.08	182	2575	KBS
Milford	9777	1967	512	2009	460	452	11.8	0.24	125	2177	USACE
Perry	2924	1966	300	2009	247	238	20.6	0.41	424	2088	USACE
Pomona	835	1963	87	2009	68	66	24.7	0.47	488	2071	KBS
Toronto	1854	1959	34	2010	20	19	44.2	0.77	140	2024	KBS
Tuttle Creek	25,070	1962	525	2009	317	286	45.4	0.84	175	2022	USACE
Waconda	6518	1967	299	2001	271	258	13.5	0.28	127	2148	BOR
Webster	2997	1956	95	2011	94	94	1.2	0.02	7	4388	KBS
Wilson	4976	1963	306	2008	294	292	4.5	0.09	53	2548	KBS
Sum	97,817		3990			3298					
Mean	4076		166			137	17.9	0.34	268		
Median	1827		114			96	12.6	0.27	159		
Max	25,070		525			452	45.4	0.84	1688		
Min	95		34			19	1.2	0.02	7		

W is the drainage area, t_0 is the beginning of reservoir operation, V_0 is the original reservoir capacity, t_r is the date of the most recent survey, V_r is the reservoir capacity at the most recent bathymetric survey, V_c is the reservoir capacity in the year 2016, L_c is the loss of capacity to 2016, D is the annual depletion rate, Y is the sediment yield, and $t_{1/2}$ is the year that half of the reservoir will be infilled by sediment. * USACE = U.S. Army Corps of Engineers; BOR = U.S. Bureau of Reclamation; USGS = U.S. Geological Survey; KBS = Kansas Biological Survey; OWRB = Oklahoma Water Resources Board.

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2. Methods and Materials

2.1. Study Area

The study sites include all the KS-FdR (Figure 2, Table 1). Aside from groundwater in the low populated counties of western Kansas, KS-FdR are the most important drinking water sources for the state's citizens [32]. Surface water storage and population density are clearly related, with 15 of the 24 KS-FdR located in the more heavily populated eastern third of the state (Figure 3). This is consistent with the fact that KS-FdR placement was driven principally by flood control needs, which are greater in the eastern part of the state due to increased precipitation (Figure 2).

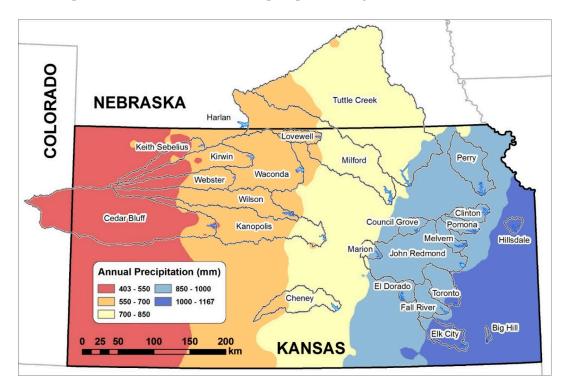


Figure 2. Watershed extents are shown for the KS-FdR, overlaid on an average annual precipitation map developed using PRISM climatological normal data (1981–2010) [33]. Harlan County Lake, a federal reservoir in Nebraska that is not part of the study, serves as the upstream watershed boundary for Milford.

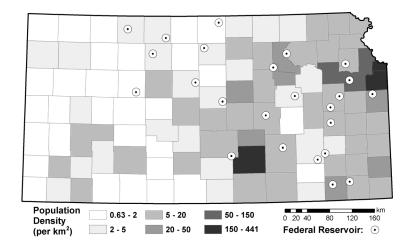


Figure 3. Population density (est. 2010) across Kansas, by county. The 24 KS-FdR locations are also shown.

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Annual rainfall indicates a gradient from west (approximately 500 mm/year) to east (approximately 1000 mm/year), with more frequent and greater extreme precipitation events in the east [34,35]. Thus, eastern Kansas reservoirs usually have more available inflow (both in absolute terms and relative to their respective capacities) than those toward the west.

Reservoir sediment volume has been widely studied in the state, and reservoir storage capacity data are available for all the KS-FdR [15,26,36–38]. Decreasing storage capacity due to sediment infill coupled with an estimated population increase of approximately 17,131 people per year ([39]; growth rate estimated using annual values from 1990–2014, with a determination coefficient (R²) of 0.99) suggests a decreasing per capita available surface water supply (Figure 4). The increase in Kansas' population of approximately 164,294 persons between 2000 and 2010 alone represents an additional demand of nearly 1.8 Mm³ of drinking water per year, based on an average domestic water use of 307 liters (0.307 m³) per person per day [40]. Considering that the KS-FdR collectively are losing capacity at a rate of 13 Mm³/year while demand is increasing by 1.8 Mm³/year, surplus water supply capacity is effectively being reduced by 14.8 Mm³/year.

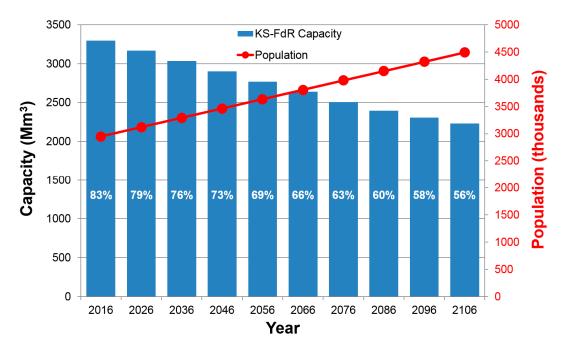


Figure 4. Linearly projected declining KS-FdR total storage capacity is shown against estimated Kansas population increase of 17,131/year described in the text. Percent of original total capacity is noted on the bars.

2.2. Data Sources

Depending on their location in the drainage system and other factors, the KS-FdR are monitored by the Kansas City or Tulsa districts of the USACE or the BOR. KWO, a state agency with responsibilities for state water resource planning and supply management, provided the general characteristics of the reservoirs [25]. The Kansas Biological Survey (KBS, a research center at the University of Kansas and a research agency of the state of Kansas), USACE, BOR, USGS, and Oklahoma Water Resources Board provided the bathymetry data used by the state for reservoir capacity and sedimentation rate analysis [33]. The majority of the surveys (15 of 24) were completed by KBS (Table 1). Climatological precipitation data from 1981–2010 were processed by and downloaded from the Parameter-elevation Regressions on Independent Slopes Model (PRISM) Climate Group [33].

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2.3. Methods

To obtain bathymetric data, KBS operated a Biosonics DT-X echosounding system with a 200 kHz split-beam transducer and a 38-kHz single-beam transducer. Samples were densely collected along transect lines with a typical spacing of approximately 100–200 m (see Figure 5 for example), and depth estimates from these measurements were spatially interpolated to produce a map representing lake bottom topography at the time of survey. Additionally, a number of sediment cores were collected with each KBS survey, which provided information on sediment particle size, bulk density (which is used in this paper), thickness, and composition. Readers are referred to [41] and many similar reports (obtainable from [25]) for more information about the procedures of the KBS and other bathymetric surveys. Details on instruments, number of survey points in each reservoir, pre-survey preparation, survey procedure, and lake level adjustments are among the steps explained in those publications.

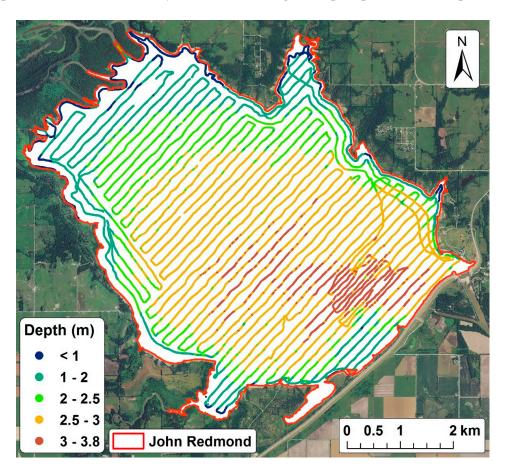


Figure 5. Bathymetric survey points collected in 2014 for John Redmond Reservoir are shown. Sample points appear as continuous lines due to the dense sampling (roughly 4 m between successive points) along the sample paths. Approximately 65,000 depth values were collected during this survey, along paths spaced approximately 150 m apart. Background imagery is from the 2012 U.S. Department of Agriculture National Agricultural Imagery Program (NAIP) dataset.

Drainage area was calculated for each KS-FdR separately; for watersheds that included more than one federal reservoir, the upstream reservoirs' drainage area was calculated individually and excluded from the drainage area of the downstream reservoir (Figure 2). This was the case for four reservoirs (John Redmond, Kanopolis, Milford, and Waconda). While a substantial portion of each KS-FdR's upstream watershed lies behind much smaller dams [30], the sediment trapping and pass-through behavior of these relatively small impoundments during extreme flood events generally is not well known, and these structures predominantly were built without a specific flood control purpose.

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For these reasons, and to be consistent with most other sediment yield studies, we did not exclude smaller impoundment catchments from the KS-FdR watersheds. Consequently, our sediment yield estimates likely are biased low to some degree, considering that smaller upstream impoundments are trapping some sediment that otherwise might have found its way into the downstream KS-FdR. Nonetheless, with KS-FdR found to have >90% sediment trapping efficiency [42–45] and existing at a much larger scale than all other upstream impoundments, this approach to watershed boundary determination is expected to result in a reasonably accurate estimate of sediment yield for the KS-FdR watersheds. From the most recent survey and watershed extent data in addition to the original surveys, we estimated sediment yield, annual depletion rate, and reservoir half-life (when the reservoir conservation pool is half full of sediment) for the KS-FdR (Table 1) [46].

The sediment yield was calculated by dividing the total accumulated sediment in the reservoir by the period between the most recent bathymetric survey and the original survey, divided by the reservoir drainage area:

$$Y = [(V_0 - V_r)/(t_r - t_0)]/W \tag{1}$$

Y is the sediment yield (m³/km²/year), V_0 is the original reservoir capacity (m³), V_r is the reservoir capacity at the most recent bathymetric survey, t_r is the date of the most recent bathymetric survey (year), t_0 is the beginning of reservoir operation, and W is the drainage area (km²). In this study, the sediment yield excludes the trapped sediment in the portion of the flood pool that is above the conservation pool as well as the sediment that is passed through the reservoir.

Percent loss of capacity (L_r) from time t_0 to t_r was calculated by:

$$L_r = 100 \times (V_0 - V_r) / V_0 \tag{2}$$

 V_c denotes current reservoir capacity as of 2016 (= t_c) and was estimated by:

$$V_c = V_0 - \left(\frac{V_0 - V_r}{t_r - t_0}\right) \times (t_c - t_0) \tag{3}$$

The loss of capacity was estimated for 2016 because the most recent surveys were completed in different years. The use of 2016 estimations for all reservoirs allows for more current and temporally consistent comparisons to be made among individual reservoir capacity losses. Also, using 2016 values allows us to avoid any current volume adjustment for John Redmond due to the dredging activity that occurred late that year.

Annual depletion rate (*D*; also a %) was calculated by:

$$D = L_r/(t_r - t_0) \tag{4}$$

The year that half of the reservoir $(t_{1/2})$ will be infilled with sediment is given by:

$$t_{1/2} = t_0 + [V_0/(Y \times W)]/2 \tag{5}$$

3. Results and Discussion

Physical and hydrological characteristics of the KS-FdR are summarized in Table 1. Drainage area ranges from 95 km² (Big Hill) to 25,070 km² (Tuttle Creek), with a total catchment area of 97,817 km² for all the KS-FdR (72,560 km² in Kansas, or 34% of the state; 22,733 km² in Nebraska; and 2524 km² in Colorado) [30]. Three KS-FdR have watersheds that cross into a neighboring state: 75% (18,862 km²) of Tuttle Creek and 40% (3872 km²) of Milford watersheds are in Nebraska, and 18% (2524 km²) of Cedar Bluff watershed is in Colorado. Generally, western and central reservoirs have larger drainage areas to help maintain adequate inflow due to the lower precipitation (and thus lower runoff) rates across their watersheds. The KS-FdR were built between the 1940s and 1980s, with the earliest gate closure (i.e., start of reservoir fill, or start of reservoir operation) occurring in 1948 (Kanopolis) and the

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most recent occurring in 1981 (Hillsdale). Original capacities were estimated primarily using digitized pre-impoundment topographic contour data. The most recent bathymetric surveys (and thus capacity estimates) were completed between 1998 (Cheney) and 2014 (John Redmond).

Annual depletion rate ranges from 0.02% (Webster) to 0.84% (Tuttle Creek) with a mean of 0.34%. Globally, reservoir capacity loss has been estimated to be in the range of 0.5–1% annually [11,47], but a review of the literature indicates that a broad range of capacity losses has occurred. For example, annual capacity loss is 1.5% for Tarbela Dam in Pakistan [11,48], 2% for Manwan Dam in China [49], and 0.2% on average in Ceará, Brazil, based on studying 21 reservoirs [50]. In Central Europe, a study of 19 reservoirs with storage capacity ranging from 1.5 to 226 Mm³ indicated an average annual depletion rate of 0.51% [51]. In the United States, 0.03% annual loss has been reported for Lake Mead upstream of Hoover Dam based on the sediment load in the first 13 years of dam operation [52]. By contrast, there are also cases where a reservoir was fully infilled by sediment only a few decades after dam construction; for example, full capacity loss was documented for Sweasey Dam in California in 1969 because of infilling sediment during 31 years of operation [53].

In a comprehensive reservoir study [54], the following average annual loss rates were found for different size categories: (a) 3.56% for reservoirs with $V_0 < 12.3 \times 10^3$ m³ using data from 190 reservoirs, (b) 2.00% for reservoirs with 12.3 \times 10³ m³ $< V_0 < 123.3 \times 10^3$ m³ using data from 257 samples, (c) 1.02% for reservoirs with 123.3 \times 10³ m³ $< V_0 < 1233 \times 10^3$ m³ using data from 283 samples, (d) 0.81% for reservoirs with 1233 \times 10³ m³ $< V_0 < 12330 \times 10^3$ m³ using data from 176 samples, (e) 0.43% for reservoirs with 12,330 \times 10³ m³ $< V_0 < 123,300 \times 10^3$ m³ using data from 107 samples, (f) 0.23% for reservoirs with 123,300 \times 10³ m³ $< V_0 < 123,300 \times 10^3$ m³ using data from 69 samples, and (g) 0.16% for reservoirs with $V_0 > 1,233,300 \times 10^3$ m³ using data from 23 samples. These results suggest that greater capacity loss rates are observed for smaller reservoirs. The KS-FdR sizes fall in the (b) and (c) ranges from that study and exhibit lower annual loss rates, but they do show a similar general trend with annual loss decreasing as reservoir size increases (also noted in [55]) (Figure 6). An average of 0.22% annual storage loss has been documented for U.S. reservoirs with 24% of the infill sediment originating from cropland erosion [55]. Kansas is an agricultural state with about 47% of the land use within the KS-FdR catchments being classified as cropland [56], which likely contributes to the high sediment concentrations found in the state's streams and reservoirs [57–59].

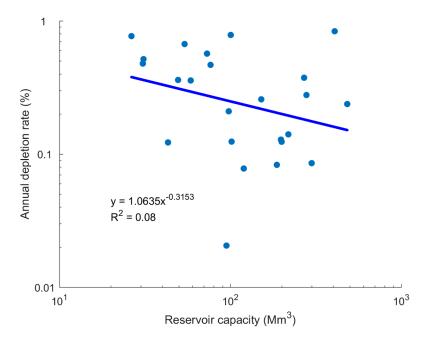


Figure 6. Consistent with other studies, annual depletion rate was found to be (weakly) inversely correlated with original capacity for the KS-FdR.

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Total original capacity for all the KS-FdR is 3990 Mm³ where 692 Mm³ (17.3%) has been lost since the establishment of the reservoirs. Loss of capacity ranges between 1 Mm³ (1.2%) and 238 Mm³ (45.4%) for Webster and Tuttle Creek, respectively, with an average percent loss of 17.9% across all the KS-FdR (Table 1). A majority of the reservoirs with higher percent loss are located in the eastern half of Kansas, and generally the storage capacity is lower for these reservoirs compared to reservoirs in western Kansas (Figures 2 and 7). Precipitation distribution follows an east-west gradient from higher to lower annual values, suggesting higher inflow to the eastern reservoirs. Streamflow is the source of sediment in the reservoirs, and higher sediment yield generally is expected in areas receiving higher precipitation [60]. Higher magnitude and frequency of extreme daily and annual rainfall events has been found in eastern Kansas [24]. Higher runoff carries more sediment, resulting in generally higher capacity loss for eastern reservoirs in Kansas [12]. For most of the state, a shift has been observed in the maximum monthly rainfall from June to May, when soil is also moister due to lower temperatures, potentially suggesting a transition to higher runoff generation earlier in spring.

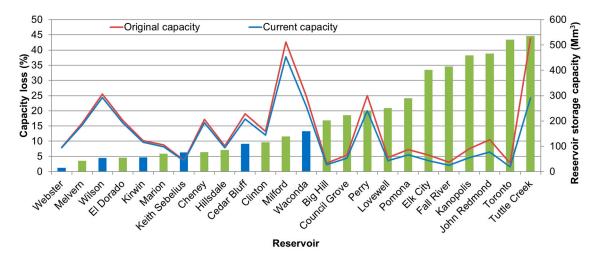


Figure 7. Capacity (lines) and capacity loss (bars) are shown for the KS-FdR. Blue bars denote reservoirs in the western half of the state.

Each KS-FdR should be useful for flood control well beyond the time its conservation pool is predominantly infilled. However, functionality with respect to non-flood-control design purposes is reduced when sediment starts to settle in the conservation pool. Crucial operational conditions may exist by the time the reservoir is half full of sediment [54,61–63] or possibly even sooner depending on factors such as pool fluctuation, lake bottom topography, sediment grain size, and intake and outflow structure design and configuration [64,65]. Furthermore, as the sediment settles in the reservoir, capacity and sediment trapping efficiency decline and the half-life period is not necessarily representative of the time it takes the reservoir to become half-filled with sediment because of shorter residence time and thus increased pass-through of suspended solids [66].

Nonetheless, if we ignore this reduction in sedimentation rate (which will be slow to build), six of the KS-FdR are projected to be half full with sediment by mid-century under the assumption that sediment infilling will continue at current rates (Table 1): Tuttle Creek (2022), Toronto (2024), John Redmond (2030 if we ignore dredging that occurred in 2016, which otherwise pushes this date back to 2033), Kanopolis (2036), Elk City (2040), and Fall River (2045). Five of these reservoirs (except Kanopolis) are located in the eastern third of the state. Originally, these six reservoirs accounted for 22% of the collective KS-FdR volume; as of 2016, this ratio has dropped to 15%, reflective of the fact that these reservoirs are losing their capacity faster than the others. Consequently, sediment upstream and within these six reservoirs may need to be managed prior to the other 18. KS-FdR, through direct extraction or release of stores downstream, are a primary source of water for various uses including drinking water and irrigation, particularly in prolonged drought periods when sufficient quantities of

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natural surface water baseflow are not available. Because of their critical importance for meeting the state's water supply needs, all manner of KS-FdR sediment management techniques and strategies, both upstream and in situ, should be considered to extend their usable lifetime [36].

The state of Kansas has purchased all or some part of the conservation pool available water in 13 of the 24 KS-FdR for municipal, industrial and irrigation use. Four of these contracts are on reservoirs that are projected to be half full with sediment by 2040 (Elk City, John Redmond, Kanopolis, and Tuttle Creek). The other nine of the 13 reservoirs still have water available for purchase, though two are fully committed (Clinton and Hillsdale) [67]. With increasing population, drinking and irrigation water demands continue to increase along with water management concerns. In 2015, the state of Kansas created a long-term (50-year) Water Vision [68] to "provide Kansans with the framework, policy and tools, developed in concert with stakeholders, to manage, secure and protect a reliable, long-term statewide water supply while balancing conservation with economic growth". One objective is to manage reservoirs sustainably while continuing to meet different domestic, agricultural, and industrial demands such as drinking water, irrigation, and recreation. With this Vision, the state hopes to anticipate and address challenges related to water including ensuring access to reliable water supply, protecting reservoirs, and reducing vulnerability to extreme events. Recommendations from the Vision could provide guidance to other regions with similar sedimentation and water challenges.

Sedimentation rate, as a physical quantity, is an absolute measure that depends on two factors, sediment yield (supply rate from the reservoir catchment) and trapping efficiency (retention rate, which is >90% for the KS-FdR). On the other hand, in its commonly used percent form, sedimentation rate provides a relative measure due to its additional dependency on reservoir volume. This is an important distinction, with the former definition appropriate for scientific examination and the latter perhaps more meaningful for reservoir prioritization for sediment management.

KS-FdR watershed sediment yield ranges from 7 m 3 /km 2 /year (Webster) to 1688 m 3 /km 2 /year (Big Hill). The average sediment yield is 268 m 3 /km 2 /year with a median of 159 m 3 /km 2 /year and a standard error about the mean of 71 m 3 /km 2 /year (Table 1). For comparison, watershed sediment yield rates for 66 surveyed reservoirs in California typically were higher than Kansas estimates, with a minimum, maximum, and mean of 8, 5085, and 413 m 3 /km 2 /year, respectively [69]. Similar rates to the KS-FdR were indicated in a different study [70], ranging approximately between 100 m 3 /km 2 /year and 1200 m 3 /km 2 /year for relatively small impoundments (drainage area < 70 km 2) in southwestern Ohio.

Sediment yield has been reported using mass rather than volume by many researchers. In two maps prepared by the U.S. Department of Agriculture on sediment delivered to rivers and streams from sheet and rill erosion [71] and erosion rates on cropland [72], the Northern Plains was among the regions with the highest erosion rates in the U.S. with an average of 1.2×10^6 kg/km²/year on croplands. Volume-based sediment yield values reported in this paper can be compared to the mass-based sediment yields by multiplying by sediment bulk density. An average bulk density of 1.1 t/m^3 was reported [73] from the Reservoir Sedimentation Survey Information System (RESIS) for the conterminous U.S. [74]. However, we assume an average bulk density of 0.8 t/m^3 for Kansas reservoirs based on sediment core samples obtained during recent regional bathymetric surveys (e.g., [75]).

The sediment delivery ratio (percentage of eroded soil amount that reaches the watershed outlet on an annual basis) to water bodies is impacted by various factors including drainage area and overland flow length [76,77]. On average, more time and opportunity are available for eroded soil to deposit in larger watersheds before reaching reservoirs located far downstream in a catchment. Consequently, the sediment delivery ratio and sediment yield are likely to be lower in larger watersheds [78,79]. Sedimentation rate and drainage area for the KS-FdR are correlated as in other regions, with sediment yield generally decreasing as drainage area increases (Figure 8). In an older study [80], sediment load was analyzed in 326 small reservoirs in Kansas with drainage areas ranging from 0.21 km² to 98 km², with sediment yield documented between 24 m³/km²/year and 5792 m³/km²/year. Though the sediment yields in that study with smaller watersheds are higher than the results reported here, the author also found a negative trend between drainage area and sediment yield.

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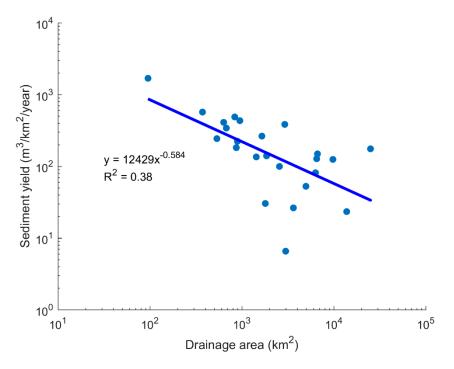


Figure 8. Consistent with other studies, sediment yield was found to be inversely correlated with drainage area for the KS-FdR.

4. Conclusions

Sedimentation and capacity depletion characteristics of 24 large federally operated reservoirs were examined in Kansas, an agricultural state situated in the central U.S. Great Plains. Original capacities range from 34 to 525 Mm³ and drainage areas range from 95 to 25,070 km². Annual depletion rate and sediment yield were estimated using the most recent bathymetric surveys compared to pre-impoundment topographic data. Annual depletion rates range from 0.02% to 0.84%, with generally higher values for smaller reservoirs. Sediment yield was found to be between 7 and 1688 m³/km²/year, with generally larger values for reservoirs with smaller drainage areas. Reservoirs in wetter eastern Kansas typically exhibit higher loss of capacity than those in central and western parts of the state. As of 2016, with an average age of 52 years, Kansas has lost approximately 17% of the original capacity of its federal reservoirs to sedimentation.

Reservoir placement within a watershed typically is optimized for storing water for flood control or water supply. Consequently, finding locations with optimal geography, hydrology, topography, and geology for building new dams in the same region can be challenging [81]. For these reasons, along with other prohibitive logistical factors (e.g., economic, legal, and regulatory), new large impoundment construction currently is considered an unlikely option for Kansas as it seeks to ensure a future water supply for its citizens and industries.

Reservoir management techniques should be identified and assessed for cost effectiveness in light of the documented sedimentation rates and the increasing demands for reservoir-supplied surface water [82,83]. Several options are discussed in [36] that affect different components of the reservoir-watershed system, and a number of these are being enacted or seriously considered in Kansas. For example, in 2013, the USACE approved a 0.61-m raise of the conservation pool at John Redmond to increase available allocable water supply [84], which had the effect of extending the reservoir's half-life by about 15 years. Kansas was compelled to action at John Redmond primarily because of a contractual water supply obligation with a local power company that extends through 2045. State officials initiated this process more than 10 years prior to approval, reflecting the difficulty in enacting conservation pool level changes when working with federal flood control reservoirs. The state's John Redmond dredging

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project, which was anticipated during the pool-raise approval process, is the first such experiment in the country to recover water supply capacity in a large federal reservoir. At a cost of \$20M (US), 2.3 Mm³ of sediment was removed from John Redmond in 2016 [29], which extended its half-life by an additional 2.5 years. The implied annual maintenance dredging cost for this one reservoir alone is \$8M/year, which is not sustainable considering current state budget limitations for water resources management. Alternatively, USACE is exploring possibilities for sediment transport and hydrosuction sediment removal at Tuttle Creek, but as yet no action is planned [85]. With these and other in situ sediment removal options, a good understanding of lake bottom geometry and sediment deposition patterns is necessary to maximize efficiency and effectiveness.

Reducing sediment yield from the upstream watershed offers another approach to slowing sediment infill [86]. In the drainage network above several federal reservoirs, Kansas is investing heavily in streambank stabilization projects [87] to harden streambank erosion hotspots. While techniques are improving and treated lengths are adding up, it is not yet clear if these stationary treatments applied to non-stationary stream channels are effectively and sustainably slowing reservoir sedimentation. As a counter-alternative, holistic watershed management options are also being investigated in an effort to identify adaptable, sustainable, and environmentally beneficial sediment mitigation strategies (e.g., [88,89]).

Sediment management, both upstream and within the reservoir, is a rapidly growing concern as reservoir infill progresses and water retention capacities are diminished [90,91]. Climate change may yet be another factor impacting reservoir water availability for multiple uses including drinking water. More observed and projected frequent and greater flooding across the world [92] may result in greater sediment delivery to reservoirs and thus increased sedimentation rates and capacity loss in some areas, while more frequent and intense droughts will reduce the available water in many impoundments through increased evaporation and reduced storage opportunity between periods of adequate surface water inflow. Results from this study provide baseline information that will help guide and support further investigation of reservoir sedimentation and sediment management, in Kansas as well as in other regions around the world facing similar challenges. Future research directions will involve more detailed geospatial analyses of landscape and climate factors of the KS-FdR reservoir-watershed systems to better understand the sediment yields that are driving reservoir sedimentation.

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