



# **Communication Light-Scattering Properties for Aggregates of Atmospheric Ice Crystals within the Physical Optics Approximation**

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Abstract: This paper presents the light-scattering matrices of atmospheric-aggregated hexagonal ice particles that appear in cirrus clouds. The aggregates consist of the same particles with different spatial orientations and numbers of these particles. Two types of particle shapes were studied: (1) hexagonal columns; (2) hexagonal plates. For both shapes, we studied compact and non-compact cases of particle arrangement in aggregates. As a result, four sets of aggregates were made: (1) compact columns; (2) non-compact columns; (3) compact plates; and (4) non-compact plates. Each set consists of eight aggregates with a different number of particles from two to nine. For practical reasons, the bullet-rosette and the aggregate of hexagonal columns with different sizes were also calculated. The light scattering matrices were calculated for the case of arbitrary spatial orientation within the geometrical optics approximation for sets of compact and non-compact aggregates and within the physical optics approximation for two additional aggregates. It was found that the light-scattering matrix elements for aggregates depend on the arrangement of particles they consist of.

Keywords: atmospheric particles; ice aggregates; light scattering; geometrical optics; physical optics



Citation: Timofeev, D.; Kustova, N.; Shishko, V.; Konoshonkin, A. Light-Scattering Properties for Aggregates of Atmospheric Ice Crystals within the Physical Optics Approximation. *Atmosphere* **2023**, *14*, 933. https://doi.org/10.3390/ atmos14060933

Academic Editor: Evgenij S. Zubko

Received: 26 April 2023 Revised: 23 May 2023 Accepted: 24 May 2023 Published: 26 May 2023



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## 1. Introduction

Atmospheric ice particles, which generally appear in cirrus clouds, are an important component in atmospheric research such as remote sensing and radiation transfer. They are observed at altitudes of 7–10 km with a hexagonal shape and size of 10–1000  $\mu$ m, in general. The density of particles in cirrus clouds is low in comparison with other types of clouds, but they have hard to predict scattering properties because of the specific geometry of particles. These properties are actively studied within international projects, and different methods are used: in situ aircraft measurements, remote sensing from ground and space, etc. A substantial amount of work has been done to study the properties of clouds. At the same time, many features of the light-scattering problem for ice particles are still poorly studied [1–14].

There are direct and remote methods for studying cirrus clouds. Direct measurements include contact measurements from aircraft [15], and remote studies include the monitoring of the atmosphere by lidar networks and photometers. Since direct methods are limited in time and financial resources, in practice, remote methods are more useful. For interpretation of lidar data, it is necessary to solve the inverse problem of light scattering for monochromatic laser radiation. However, one needs a database of light-scattering matrices and the corresponding microphysical properties of cloud particles [15–21]. For solving this problem, numerical methods are usually used [22–27].

Cirrus cloud particles can be distinguished by microphysical structure in two types: single particles (hexagonal columns, plates, bullet, etc.) and aggregates consisting of several particles. According to data of in situ measurements, atmospheric ice aggregates take a significant part of particles in cirrus clouds [28,29]. However, the proper information about

their scattering properties is absent in existing databases. In general, crystals in clouds are arbitrarily oriented. Additionally, it is expected that light scattering by aggregates consisting of the same crystals and light scattering by a single crystal are similar in special cloud. In this case, it is possible to calculate the light-scattering matrix for aggregates using the dependence of light-scattering matrix elements on the number of particles in aggregates. However, if particles in the aggregate are compactly packed, then the direction of scattering light might be changed. Additionally, the distribution of light from a single particle might be different from the distribution for an aggregate.

The purpose of the research is to define the dependence of scattering–matrix elements on the number and arrangement of particles in the aggregate.

#### 2. Materials and Methods

For the calculation of the light-scattering matrix, we use the physical optics approximation method [30]. This method is the most applicable for this problem because of its capability to calculate particle parameter size x > 10 and precise results in the backscattering direction [31]. It was also used for solving problems related to the remote sensing of atmospheric ice crystals [32,33]. In this method, the particle consists of facets with multiple vertices. The method is based on the Beam-splitting algorithm [34], which is similar to the Ray-tracing algorithm [35], but it works with plane-parallel optical beams. In this algorithm, the particle that scatters light consists of facets, which consist of vertices with three-dimensional coordinates. The algorithm splits the light incident on facets into beams. These beams propagate in the particle and could be divided into refracted and reflected beams multiple times before they leave the particle and scatter by their cross-section frame.

The physical optics method calculates the scattering field in the near zone, within the geometrical optics approximation, and in the far zone, within both the geometrical and the physical optics approximation. However, the calculation of diffraction for each scattered beam is a very expensive operation, especially for the case of randomly oriented particles in a cloud. Most of all, the calculation time of the physical optics method increases with the number of facets in the particle. That is why, first-order, we calculated the light-scattering matrix for aggregates within the geometrical optics approximation.

By aggregate, we mean the object consisting of several particles that are attached to each other at one or several points. The positions of these particles are fixed, but when the spatial orientation of the aggregate is changed, all particles change their position simultaneously. In this work, we use regular shapes of crystals for cirrus clouds as the basic particles for aggregates: hexagonal column; hexagonal plate [36] (see Figure 1).



Figure 1. Geometrical shapes of the particles for aggregates: (a) hexagonal column; (b) hexagonal plate.

Based on the assumption that compactly packed aggregates change the direction of scattered light, two types of particle arrangements in aggregates were chosen: compact and non-compact. To determine the difference between these arrangements, we add the compactness index (C):

$$C = \frac{L_{avg}}{R_{min}} \tag{1}$$

$$L_{avg} = \frac{\sum_{i=1}^{N} L_i}{N},$$
(2)

where  $L_i$  is the distance between the geometrical aggregate's center and the center of a single particle *i*; *N* is the number of particles in the aggregate. Coordinates of the center of the aggregate and single particle are calculated by summing the coordinates of all vertices of the aggregate and single particle, respectively, and dividing it by the number of vertices. Based on Equation (1), the most compact aggregate will have C = 1.

In this work, aggregates are created according to the following principle: 9 particles with the same shape, size, and coordinates are generated in the center of the coordinate system. It means that they are located inside each other in this stage. Then, each particle (except the first one) is rotated by a random angle and moved away from the center in a random direction until the particles were not intersecting each other. The final aggregate consists of nine particles, which are attached to each other. This procedure was done 100 times for two shapes (see Figure 1), and 100 aggregates with different arrangements were made. Then, the most compact and the most non-compact aggregates were chosen according to Equation (1). Finally, we made sets of aggregates with *N* from two to nine by removing particles from chosen aggregates one by one. Models of aggregates of nine particles for each type are shown in Figure 2. The following dimensions of basic particles were used: for column height 100  $\mu$ m, base diameter 69.6  $\mu$ m; plates height 15.97  $\mu$ m, base diameter 100  $\mu$ m. The particle geometry corresponds to the model in [37]. The dependencies of *L*<sub>avg</sub> and *C* on *N* in the aggregate are shown in Figure 3.



**Figure 2.** Models of aggregates of 9 particles: (a) compact columns; (b) non-compact columns; (c) compact plates; (d) non-compact plates.



**Figure 3.** (a) Dependence of *C* on *N*; (b) Dependence of  $L_{avg}$  on *N*.

#### 3. Calculation Results within the Geometrical Optics Approximation

For the created aggregates, light-scattering matrices were calculated for all scattering angles within the geometrical optics approximation at the wavelength of 0.532  $\mu$ m with the refractive index of particles being 1.3116 (ice water) [38]. Since the aggregates in a

cloud are assumed to be randomly oriented, the calculation was carried out for one million orientations for each aggregate. As an example, the  $M_{11}$  and  $M_{22}$  elements of the light-scattering matrix are shown in Figures 4 and 5. The value of  $M_{11}$  for the scattering angle of 0° is removed from the plots because it is too high to display. The element  $M_{22}$  is normalized over  $M_{11}:m_{22} = M_{22}/M_{11}$ .



**Figure 4.** The  $M_{11}$  element vs. scattering angle ( $\theta$ ) for the following aggregates: (**a**) non-compact columns; (**b**) compact columns; (**c**) non-compact plates; (**d**) compact plates.



**Figure 5.** The  $m_{22}$  element vs. scattering angle ( $\theta$ ) for the following aggregates: (**a**) non-compact columns; (**b**) compact columns; (**c**) non-compact plates; (**d**) compact plates.

The main interest is the study of the dependence of light-scattering matrix elements on the number of particles in the aggregate (*N*). It is better to start the study with the element  $M_{11}$ , which defines the intensity of light in the scattering angle ( $\theta$ ) for unpolarized incident light. Variability of the other elements is not critical.

Because of the fact that the scattering efficiency is equal to two within the physical optics approximation, the most informative parameter is the  $M_{11}$  divided by the average geometric shadow area ( $G_A$ ). This area can be calculated using the geometry of the aggregate without solving the light-scattering problem. The result shows that the  $M_{11}/G_A$  very slightly changes with the number of particles in the aggregate, except in the case of an aggregate of compact plates.

Let us examine the case of a compact aggregate of plates (see Figure 6d). For convenience, we separately plotted  $M_{11}/G_A$  for the forward scattering direction for these aggregates vs.  $N(M_{11}*/G_A \text{ in Figure 7})$ . In this case, the peak of intensity in the forward scattering direction is re-scattered by another plate appearing right behind the first one with an increasing number of particles. That peak is created by the forward transmission of light through the large plane-parallel facets of a plate particle. While in the case of a non-compact aggregate of plates, the particles do not overlap each other, so the light moves freely in the forward scattering direction (see Figure 8).



**Figure 6.**  $M_{11}/G_A$  for the following aggregates: (**a**) non-compact columns; (**b**) compact columns; (**c**) non-compact plates; (**d**) compact plates.



**Figure 7.**  $M_{11}^*/G_A$  for compact aggregate of plates vs. *N*.



Figure 8. Comparison of the compact (a) and non-compact (b) aggregates of three plates.

One of the typical non-compact aggregate particles in cirrus clouds is a bullet-rosette. We calculated light-scattering matrices for bullet-rosette aggregates with a number of bullets from two to six (Figure 9). Unlike previous aggregates, the arrangement of particles in the bullet-rosette remains orthogonal. Calculation parameters were the same as before. The sizes of every bullet in the aggregates are as follows: the height of the hexagonal part is 100 µm; the base diameter is 42 µm; and the pike angle is 19.7°. The dependences of the elements of the light-backscattering matrix on the scattering angle ( $\theta$ ) are presented in Figure 10, and  $M_{11}/G_A$  in Figure 11. The normalized elements  $m_{12}$ ,  $m_{22}$ , and  $m_{44}$  are also shown in Figure 11. The results show that the  $M_{11}/G_A$  and normalized elements  $m_{12}$ ,  $m_{22}$ , and  $m_{44}$  almost do not change with the number of particles in the bullet-rosette. It means that the optical properties of the bullet-rosette can be evaluated from the optical properties of one bullet.



Figure 9. Models of bullet and bullet-rosette aggregates.



**Figure 10.** Dependences of elements of the light-backscattering matrix on scattering angle ( $\theta$ ): (**a**)  $M_{11}$ ; (**b**)  $m_{12}$ ; (**c**)  $m_{22}$ ; (**d**)  $m_{44}$ .



**Figure 11.**  $M_{11}/G_A$  for bullet-rosette aggregates (colors are the same as in Figure 10).

### 4. Calculation Results within the Physical Optics Approximation

For the lidar application, only the physical optics solution is of practical interest because geometrical optics cannot resolve the backscattering peak of intensity of hexagonal particles. However, the physical optics solution is much more demanding on computing resources, so we examine only two aggregates of ice crystals: bullet-rosette and an aggregate of eight hexagonal columns with different sizes [39]. The geometry for these particles is shown in Figure 12.



Figure 12. Models of aggregates: (a) aggregate of 8 columns; (b) bullet-rosette.

For the aggregate of eight columns, the width *D* and length *L* for each hexagonal column are dimensionless quantities. The center of the column in the particle system is denoted by three coordinates ( $x_0$ ,  $y_0$ ,  $z_0$ ). Then, they are scaled to obtain a proper dimension for an aggregate in calculation. The orientation of the single hexagonal column is specified by three Euler angles  $\alpha^0$ ,  $\beta^0$ , and  $\gamma^0$ , where  $\alpha^0$  defines rotation of the column about the vertical direction;  $\beta^0$  is the angle between the vertical direction and the crystal main axis; and  $\gamma^0$  describes the column rotation about the main axis. The main axis of the hexagonal column is assumed to pass through the centers of the hexagonal facets. Table 1 lists the values of initial geometric parameters for an aggregate composed of eight hexagonal columns. Note that these characteristics are slightly different from the characteristics presented in the paper by P. Yang [39], in order to avoid self-intersections. For convenience, we define the aggregate size through its maximal size  $D_{max}$ .

Table 1. Geometric parameters of aggregate of 8 columns.

No.	D, [μm]	<i>L</i> , [μm]	$\gamma^0$	$eta^0$	α <sup>0</sup>	<i>x</i> <sub>0</sub>	<b>y</b> 0	$z_0$
1	92	158	23	50	-54	0	0	0
2	80	124	16	81	156	15.808	107.189	-60.108
3	56	78	5	57	94	-26.691	73.005	49
4	96	126	13	76	130	-88	-39.19	-11.643
5	106	144	11	29	-21	106.532	33.08	27.801
6	38	54	8	62	-164	35.923	-51.5	-37.533
7	68	102	29	41	60	40.11	-57.227	112.5
8	86	138	19	23	-122	-9.7524	-132.57	57.131

For calculations within the physical optics approximation, the bullet-rosette is composed of six bullets of the same size. The bullet shape is defined by the diameter *D*, the length of the hexagonal part *l*, and the angle of a tip, which is equal to 28° for all sizes. The relationship between the length *L* and the width *D* is  $D = 2.31L^{0.63}$ . Sizes are given in  $\mu$ m.

Our previous study shows that the elements of the scattering matrix obey the power laws over the particle size [40]. Base on the conclusion of the previous section, we can assume that the scattering matrix for aggregates of particles also obey the power laws. These power laws for the aggregate can be obtained from the power law of a single particle by multiplying by the averaging geometrical cross-section of the aggregate. This fact can significantly reduce the calculation time. Since one calculation for an aggregate with the maximum dimension  $D_{max} = 670 \ \mu m$  for the wavelength of 0.355  $\mu m$  takes about 18 days on a modern server with 2 Xeon E5-2660 v2 processors (40 threads), we can precisely calculate only a few of them.

The results are presented in Figure 13. The precise calculation for the aggregate is marked as dots, the solid lines correspond to the light-backscattering matrix of a single particle, and the dashed lines correspond to the evaluation of the light-backscattering matrix of the aggregate. We can see that the light-scattering matrix of the aggregate can be obtained from the matrix of a single particle with good accuracy.



**Figure 13.**  $M_{11}$  and  $M_{22}$  vs.  $D_{max}$  at two wavelengths of incident light: (a) for aggregate of 8 columns; (b) for bullet-rosette.

The power laws are presented in Table 2 and Figure 13.

Particle		$\lambda = 0.355 \ \mu m$	$\lambda = 1.064 \ \mu m$
Aggregate of 8 columns, 30 $\mu$ m $\leq D_{max} \leq 2500 \mu$ m	$M_{11} M_{22}$	$\begin{array}{c} 0.00089 \cdot D_{max}^{3.022} \\ 0.00055 \cdot D_{max}^{3.008} \end{array}$	$\frac{0.00022 \cdot D_{max}^{3.025}}{0.00014 \cdot D_{max}}^{3.017}$
Bullet rosettes, $40 \mu\mathrm{m} \le D_{max} \le 2500 \mu\mathrm{m}$	M <sub>11</sub> M <sub>22</sub>	$\begin{array}{c} 0.0146 \cdot D_{max}{}^{2.184} \\ 0.0068 \cdot D_{max}{}^{2.201} \end{array}$	$\begin{array}{c} 0.00587 {\cdot} D_{max}{}^{2.129} \\ 0.0028 {\cdot} D_{max}{}^{2.155} \end{array}$

**Table 2.** The power laws for the light-backscattering matrices  $(M_{11})$ .

Now it is possible for us to compare  $M_{11}$  for bullet-rosette (6 bullets) and for a single bullet using the existing database. Since  $D_{max}$  for aggregate and for a single particle is different, we use the dependence of  $M_{11}$  on the length of a single bullet ( $L^{bul}$ ). Then,  $M_{11}$  for a single bullet was multiplied by the total scattering cross-section for bullet-rosette. The result is presented in Figure 14.



**Figure 14.** Dependence of  $M_{11}$  on  $L^{bul}$  for bullet-rosette and for single bullet.

#### 5. Discussion

Light-scattering matrix calculations within the geometrical optics approximation for aggregates consisting of hexagonal columns and plates with different arrangements, show quasi-linear dependencies of the first element of the light-scattering matrix ( $M_{11}$ ) on the number of particles (N) in the scattering-angle range of 20–180° (Figure 6). The scattering matrix can be obtained by multiplying the scattering efficiency of a single particle by the

geometrical cross-section of an aggregate. As far as the physical optics solution is obtained from the geometrical optics solution, it should also be slightly changed with increasing *N*. However, this effect does not work with compactly packed-plate aggregates because of their specific geometry. This is a very important conclusion that allows us to extend the light-scattering database of a single particle to the case of aggregates of particles.

Otherwise,  $M_{11}$  for column aggregates shows an unpredictable distribution at angles of 0–20°. This fact can be explained by the decreasing energy at the angle of 0° (forward-scattering direction intensity peak) and redistribution of it to different directions. This energy peak is caused by light that falls orthogonal to the surface of facets and propagates through the particle without refraction. However, it can be refracted in the case of an aggregate of two or more particles when the light that leaves one particle is redirected by falling on another particle. In the case of plate aggregates, this effect is insignificant because of the similar, spatial orientation of plate particles in aggregates.

It is important to note that the calculation was carried out for two cases of individual arrangement of particles in the aggregate, and the results cannot predict the exact values of the elements of the light-scattering matrix for different aggregates. For example, the distribution of  $M_{11}$  in the angular range of 0–20° for a column aggregate with a different arrangement may be different. However, the main dependencies are consistent with the initial assumptions. Further studies should consider more examples of aggregates to obtain satisfactory statistics. It is also necessary to calculate the backscattering matrices in the physical optics approximation with the absorption coefficient.

The  $M_{11}$  for bullet-rosette shows a more predictable dependence on the number of particles. It can be obtained by multiplying  $M_{11}$  for a single bullet of the same size by the total scattering cross-section for bullet-rosette both within the geometrical and the physical optics approximation.

Author Contributions: Conceptualization, D.T. and A.K.; methodology, D.T., N.K. and A.K.; software, D.T.; validation, N.K. and V.S.; formal analysis, V.S.; investigation, D.T.; resources, V.S.; data curation, V.S.; writing—original draft preparation, D.T.; writing—review and editing, N.K.; visualization, D.T.; supervision, A.K.; project administration, A.K.; funding acquisition, N.K. All authors have read and agreed to the published version of the manuscript.

**Funding:** This research was funded by the Russian Science Foundation, grant number 21-77-00083, https://rscf.ru/project/21-77-00083/.

Institutional Review Board Statement: Not applicable.

Informed Consent Statement: Not applicable.

Data Availability Statement: Not applicable.

Acknowledgments: The authors gratefully acknowledge the computing time support provided by the IOA SB RAS supercomputer "Felix-C". Alexander Konoshonkin acknowledges the support of the CAS PIFI (2021VTA0009).

Conflicts of Interest: The authors declare no conflict of interest.

#### References

- 1. Liou, K.N. Influence of cirrus clouds on the weather and climate process: A global perspective. *Mon. Weather Rev.* **1986**, *114*, 1167–1199. [CrossRef]
- Stephens, G.L.; Tsay, S.C.; Stackhouse, P.W., Jr.; Flatau, P.J. The relevance of the microphysical and radiative properties of cirrus clouds to climate and climatic feedback. J. Atmos. Sci. 1990, 47, 1742–1754. [CrossRef]
- Takano, Y.; Liou, K.N. Solar radiative transfer in cirrus clouds. Part I. Single scattering and optical properties of hexagonal ice crystals. J. Atmos. Sci. Papers 1989, 46, 3–19. [CrossRef]
- 4. Sassen, K.; Benson, S. A midlatitude cirrus cloud climatology from the Facility for Atmospheric Remote Sensing: II. Microphysical properties derived from lidar depolarization. *J. Atmos. Sci. Papers* **2001**, *58*, 2103–2112. [CrossRef]
- Prigarin, S.M. Monte Carlo simulation of the effects caused by multiple scattering of ground-based and spaceborne lidar pulses in clouds. *Atmos. Ocean. Opt.* 2017, 32, 79–83. [CrossRef]

- 6. Samoilova, S.V. Simultaneous reconstruction of the complex refractive index and the particle size distribution function from lidar measurements: Testing the developed algorithms. *Atmos. Ocean. Opt.* **2019**, *32*, 628–642. [CrossRef]
- Berry, E.; Mace, G.G. Cloud properties and radiative effects of the Asian summer monsoon derived from A-Train data. J. Geophys. Res. Atmos. 2014, 119, 9492–9508. [CrossRef]
- Kokhanenko, G.P.; Balin, Y.S.; Klemasheva, M.G.; Nasonov, S.V.; Novoselov, M.M.; Penner, I.E.; Samoilova, S.V. Scanning polarization lidar LOSA-M3: Opportunity for research of crystalline particle orientation in the ice clouds. *Atmos. Meas. Tech.* 2020, 13, 1113–1127. [CrossRef]
- Marichev, V.N. Combined method for optical sensing of the lower and middle atmosphere. *Atmos. Ocean. Opt.* 2016, 29, 348–352. [CrossRef]
- Russkova, T.V.; Zhuravleva, T.B. Optimization of sequential code for simulation of solar radiative transfer in a vertically heterogeneous environment. *Atmos. Ocean. Opt.* 2017, 30, 169–175. [CrossRef]
- 11. Kumar, A.; Singh, N.; Anshumali; Solanki, R. Evaluation and utilization of MODIS and CALIPSO aerosol retrievals over a complex terrain in Himalaya. *Remote Sens. Environ.* **2018**, 206, 139–155. [CrossRef]
- Pauly, R.M.; Yorks, J.E.; Hlavka, D.L.; McGill, M.J.; Amiridis, V.; Palm, S.P.; Rodier, S.D.; Vaughan, M.A.; Selmer, P.A.; Kupchock, A.W.; et al. Cloud-Aerosol Transport System (CATS) 1064 nm calibration validation. *Atmos. Meas. Tech.* 2019, 12, 6241–6258. [CrossRef] [PubMed]
- 13. Yang, P.; Liou, K.-N. Single-scattering properties of complex ice crystals in terrestrial atmosphere. *Contr. Atmos. Phys.* **1998**, 71, 223–248.
- Khademi, F.; Bayat, A. Classification of aerosol types using AERONET version 3 data over kuwait city. *Atmos. Environ.* 2021, 265, 118716. [CrossRef]
- Heymsfield, A.J.; Bansemer, A.; Field, P.R. Observations and parameterization of particle size distributions in deep tropical cirrus and stratiform precipitating clouds: Results from in-situ observations in TRMM field campaigns. *J. Atmos. Sci.* 2002, 59, 3457–3491. [CrossRef]
- 16. Reichardt, J.; Wandinger, U.; Klein, V.; Mattis, I.; Hilber, B.; Begbie, R. RAMSES: German Meteorological Service autonomous Raman lidar for water vapor, temperature, aerosol, and cloud measurements. *Appl. Opt.* **2012**, *51*, 8111–8131. [CrossRef]
- 17. Aerosol Robotic Network (AERONET) Homepage. Available online: https://aeronet.gsfc.nasa.gov (accessed on 11 January 2023).
- Marichev, V.N.; Bochkovsky, D.A.; Elizarov, A.I. Optical Aerosol Model of the Western Siberian Stratosphere Based on Lidar Monitoring Results. *Atmos. Ocean. Opt.* 2022, 35 (Suppl. S1), S64–S69. [CrossRef]
- 19. Samoilova, S.V.; Balin, Y.S.; Kokhanenko, G.P.; Nasonov, S.V.; Penner, I.E. Aerosol Layers in the Troposphere: Peculiarities of Variations in Aerosol Parameters at a Change in the Advection Direction. *Atmos. Ocean. Opt.* **2020**, *33*, 347–361. [CrossRef]
- Grynko, Y.; Shkuratov, Y.; Förstner, J. Light scattering by irregular particles much larger than the wavelength with wavelengthscale surface roughness. Opt. Lett. 2016, 41, 3491. [CrossRef]
- 21. Zubko, E.; Videen, G.; Zubko, N.; Shkuratov, Y. Reflectance of micron-sized dust particles retrieved with the Umov law. J. Quant. Spectrosc. Radiat. Transfer. 2017, 190, 1–6. [CrossRef]
- Zubko, E.; Shmirko, K.; Pavlov, A.; Sun, W.; Schuster, G.L.; Hu, Y.; Stamnes, S.; Omar, A.; Baize, R.R.; McCormick, M.P.; et al. Active remote sensing of atmospheric dust using relationships between their depolarization ratios and reflectivity. *Opt. Lett.* 2021, 46, 2352–2355. [CrossRef] [PubMed]
- Purcell, E.M.; Pennypacker, C.R. Scattering and absorption of light by nonspherical dielectric grains. *Astrophys. J.* 1973, 186, 705–714. [CrossRef]
- 24. Yurkin, M.A.; Moskalensky, A.E. Open-source implementation of the discrete-dipole approximation for a scatterer in an absorbing host medium. *J. Phys. Conf. Ser.* 2021, 2015, 12167. [CrossRef]
- 25. Sun, B.; Yang, P.; Kattawar, G.W.; Zhang, X. Physical-geometric optics method for large size faceted particles. *Opt. Express* **2017**, 25, 24044–24060. [CrossRef]
- Yang, P.; Ding, J.; Panetta, R.L.; Liou, K.-N.; Kattawar, G.; Mishchenko, M.I. On the Convergence of Numerical Computations for Both Exact and Approximate Solutions for Electromagnetic Scattering by Nonspherical Dielectric Particles (Invited Review). *Prog. Electromagn. Res.* 2019, 164, 27–61. [CrossRef]
- 27. Liu, J.; Yang, P.; Muinonen, K. Dust-aerosol optical modeling with Gaussian spheres: Combined invariant-imbedding T-matrix and geometric-optics approach. J. Quant. Spectrosc. Radiat. Transfer 2015, 161, 136–144. [CrossRef]
- 28. Kajikawa, M.; Heymsfield, A.J. Aggregation of ice crystals in cirrus. J. Atmos. Sci. 1989, 46, 3108–3121. [CrossRef]
- Um, J.; McFarquhar, G.M.; Hong, Y.P.; Lee, S.-S.; Jung, C.H.; Lawson, R.P.; Mo, Q. Dimensions and aspect ratios of natural ice crystals. *Atmos. Chem. Phys.* 2015, 15, 3933–3956. [CrossRef]
- Borovoi, A.; Konoshonkin, A.; Kustova, N. The physical-optics approximation and its application to light backscattering by hexagonal ice crystals. J. Quant. Spectrosc. Radiat. Transfer 2014, 146, 181–189. [CrossRef]
- Borovoi, A.; Konoshonkin, A.; Kustova, N. Backscatter ratios for arbitrary oriented hexagonal ice crystals of cirrus clouds. *Opt. Lett.* 2014, 39, 5788–5791. [CrossRef]
- Wang, Z.; Shishko, V.; Kustova, N.; Konoshonkin, A.; Timofeev, D.; Xie, C.; Liu, D.; Borovoi, A. Radar-lidar ratio for ice crystals of cirrus clouds. Opt. Express 2021, 29, 4464–4474. [CrossRef]
- Kustova, N.; Konoshonkin, A.; Shishko, V.; Timofeev, D.; Tkachev, I.; Wang, Z.; Borovoi, A. Depolarization Ratio for Randomly Oriented Ice Crystals of Cirrus Clouds. *Atmosphere* 2022, 13, 1551. [CrossRef]

- 34. Konoshonkin, A.V.; Kustova, N.V.; Borovoi, A.G. Beam Splitting Algorithm for the Problem of Light Scattering by Atmospheric Ice Crystals. Part 1. Theoretical Foundations of the Algorithm. *Atmos. Ocean. Opt.* **2015**, *28*, 441–447. [CrossRef]
- 35. Macke, A.; Mueller, J.; Raschke, E. Single scattering properties of atmospheric ice crystal. J. Atmos. Sci. **1996**, 53, 2813–2825. [CrossRef]
- 36. Yang, P.; Stegmann, P.; Tang, G.; Hioki, S.; Ding, J. Improving scattering, absorption, polarization properties of snow, graupel, and ice aggregate particles from solar to microwave wavelengths in support of the CRTM. *JCSDA Q.* **2018**, *59*, 8–14.
- 37. Mitchell, D.L.; Arnott, W.P. A model predicting the evolution of ice particle size spectra and radiative properties of cirrus clouds. Part II. Radiation. *J. Atmos. Sci.* **1994**, *51*, 817–832. [CrossRef]
- 38. Warren, S.G. Optical constants of ice from the ultraviolet to the microwave. Appl. Opt. 1984, 23, 1206–1225. [CrossRef]
- Yang, P.; Bi, L.; Baum, B.A.; Liou, K.-N.; Kattawar, G.W.; Mishchenko, M.I.; Cole, B. Spectrally consistent scattering, absorption, and polarization properties of atmospheric ice crystals at wavelengths from 0.2 to 100 μm. *J. Atmos. Sci.* 2013, 70, 330–347. [CrossRef]
- 40. Konoshonkin, A.; Borovoi, A.; Kustova, N.; Reichardt, J. Power laws for backscattering by ice crystals of cirrus clouds. *Opt. Express* **2017**, *25*, 22341–22346. [CrossRef]

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