



Article

Early-Warning Signals of Drought-Flood State Transition over the Dongting Lake Basin Based on the Critical Slowing Down Theory

Hao Wu 1,2,3, Wei Hou 4,*, Dongdong Zuo 5, Pengcheng Yan 6 and Yuxing Zeng 1,2

- ¹ Hunan Climate Center, China Meteorological Administration, Changsha 410118, China
- ² Hunan Key Laboratory of Meteorological Disaster Prevention and Reduction, Changsha 410118, China
- ³ Yueyang National Climatic Observatory, China Meteorological Administration, Yueyang 414000, China
- ⁴ National Climate Center, China Meteorological Administration, Beijing 100081, China
- 5 School of Mathematics and Physics, Yancheng Institute of Technology, Yancheng 224000, China; dongdz_vow@163.com
- 6 Institute of Arid Meteorology, China Meteorological Administration, Lanzhou 730000, China; yanpc@iamcma.cn
- * Correspondence: houwei@cma.gov.cn

Abstract: In this study, the standardized precipitation index (SPI) data in Hunan Province from 1961 to 2020 is adopted. Based on the critical slowing down theory, the moving t-test is firstly used to determine the time of drought-flood state transition in the Dongting Lake basin. Afterwards, by means of the variance and autocorrelation coefficient that characterize the phenomenon of critical slowing down, the early-warning signals indicating the drought-flood state in the Dongting Lake basin are explored. The results show that an obvious drought-to-flood (flood-to-drought) event occurred around 1993 (2003) in the Dongting Lake basin in recent 60 years. The critical slowing down phenomena of the increases in the variance and autocorrelation coefficient, which are detected 5–10 years in advance, can be considered as early-warning signals indicating the drought-flood state transition. Through the studies on the drought-flood state and related early-warning signals for the Dongting Lake basin, the reliabilities of the variance and autocorrelation coefficient-based early-warning signals for abrupt changes are demonstrated. It is expected that the wide application of this method could provide important scientific and technological support for disaster prevention and mitigation in the Dongting Lake basin, and even in the middle and lower reaches of the Yangtze River.

Keywords: drought-flood state transition; critical slowing down; early-warning signals; SPI

Citation: Wu, H.; Hou, W.; Zuo, D.; Yan, P.; Zeng, Y. Early-Warning Signals of Drought-Flood State Transition over the Dongting Lake Basin Based on the Critical Slowing Down Theory. *Atmosphere* **2021**, *12*, 1082. https://doi.org/ 10.3390/atmos12081082

Academic Editors: Hanbo Yang, Songjun Han and Bing Gao

Received: 31 July 2021 Accepted: 20 August 2021 Published: 23 August 2021

Publisher's Note: MDPI stays neutral with regard to jurisdictional claims in published maps and institutional affiliations.



Copyright: © 2021 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (http://creativecommons.org/licenses/by/4.0/).

1. Introduction

The Dongting Lake basin is located in the middle reaches of the Yangtze River and the north of the southern Yangtze River, mainly covering the Hunan Province, which is known as a fertile land of fish and rice in China. Historical data show that the Dongting Lake basin has been deeply affected by drought and flood disasters [1]. Particularly since the mid-20th century, flood and drought disasters frequently occur in the Dongting Lake basin under the background of significant global warming. Moreover, the phenomenon of drought-flood state transition (critical transition) has become more and more frequent [2,3]. The duration for the climate maintained in a drought or flood state over the Dongting Lake basin is getting shorter and shorter, that is, the transition cycle of drought-flood state significantly shortens [1,4,5].

The drought-flood state transition often leads to regional flood, drought, or deterioration of water environment [6]. If effective measures of early-warning can be carried out before the drought-flood state transition, local government and relevant departments will have sufficient time for management planning to minimize disaster losses as much as possible. Therefore, it is very important to explore the prediction and early-warning methods for drought-flood state transition in the Dongting Lake basin.

However, as the drought-flood state transition involves complex nonlinear processes, the existing climate model has poor capability in predicting possible drought-flood transition events [7-11]. At present, research into predicting regional abnormal drought or flood events caused by a drought-flood state transition are mainly based on the establishment of models taking the external forcings (sea temperature, sea ice, etc.) as a bridge. However, the results have shown that it is often difficult to overcome the limitations of the linear method adopted in the prediction model, which only contains weak mechanisms due to excessive simplification. Some of the studies have demonstrated that the model's prediction results over the Dongting Lake basin in 1999 and 2007 are almost opposite to the observations [12,13]. In summary, the accurate prediction of regional drought or flood events caused by the drought-flood state transition is a long-term challenge for meteorological departments. Therefore, it is still an extremely difficult task to accurately predict the drought-flood state transition based on the existing prediction theories and technical levels. Previous studies have shown that many complex nonlinear systems have critical thresholds (also called the tipping points, where the behavior or structure of the system changes dramatically) at which the system shifts abruptly from one state to another [14,15]. This critical transition phenomenon widely exists in various fields of nature and human society, such as the abrupt changes of climate system, the outbreak of diseases, the sudden collapse of real estate markets, the sudden paralysis of networks, the sudden extinction of species and the sudden disappearance of oasis [16-18]. Such transitions are often catastrophic. For example, the abrupt changes of the climate system may result in the decline of the civilization or the rise and fall of dynasties [18-20]. However, it is difficult to accurately predict such phenomena based on existing technologies. Therefore, it is urgent to investigate the mechanisms and early-warning signals for critical transitions of nonlinear systems.

According to the definition of the United Nations International Strategy for disaster reduction [21], the early-warning system is "the set of capacities needed to generate and disseminate timely and meaningful warning information to enable individuals, communities and organizations threatened by a hazard to prepare and to act appropriately and in sufficient time to reduce the possibility of harm or loss". A generally accepted viewpoint on early-warning system (EWS) is that the EWS is a social process aiming to address the need to avoid harm due to hazards [22-25]. EWS embraces the technical aspect, this technology can give warning signals to people when hazards are approaching [26,27], and depends on how people define the hazard or the type of hazard they will face. The start time of hazard is a key professional problem that needs to be solved in EWS, because EWS does not play its role until the hazard has appeared [27]. Some hazards have obvious external characteristics and can be quickly detected by us, such as tsunami or tornado caused by earthquake, we need to warn them in a few minutes; hurricane warnings are on the order of hours to weeks; but other hazards continue to develop at a very slow speed, such as drought, we will issue early-warning a few months in advance according to the prediction of precipitation [28].

A large amount of research has already indicated that some hazard will occur more frequently over a period of time, especially in the context of climate warming [29], such as the floods that occurred frequently in Europe between 1985 and 2009 [30–32]. Yunnan and other provinces in Southwest China have experienced frequent droughts since 2009 [33–36]. If we can warn people in advance that droughts or floods will occur frequently in

the next few years, so people can more reasonably arrange the industrial/agricultural production and take defensive measures to minimize the loss or impact of disasters. Therefore, it is necessary to find the early-warning signals longer in advance.

At the beginning of the 21st century, scientists found that the critical slowing down theory has important application potential in the study of early-warning signals for complex dynamical systems [14,37,38]. The critical slowing down is a concept in statistical physics, which refers to the phenomenon that when the dynamic system approaches the critical point before its transition from one phase to another, especially at the critical point, there will be dispersed fluctuations conducive to the formation of new phases. This dispersed fluctuation presents not only increasing amplitude, but also longer duration, slower recovery rate after the perturbation and decreased ability to return to the previous phase. This phenomenon is called "slowing down" [39]. Research on several real systems such as the climate system and the ecosystem has revealed that when the system approaches to the critical point, there will be a critical slowing down phenomenon featured by the slowing down in recovery rate, and the increases in the variance and autocorrelation coefficient after the system perturbation which provides a new idea for investigating the early-warning signals of abrupt changes in complex dynamical systems [14]. Based on the critical slowing down theory, Yan et al. [40] processed the data of radon concentrations in water and revealed that the early-warning signals of increasing variances and autocorrelation coefficients appeared before the Wenchuan earthquake in 2008. In recent years, several scholars [41-45] applied this theory to climate data, such as the regional temperature in China, the Pacific Decadal Oscillation index and the intensity of Aleutian low, aiming to study the early-warning signals of abrupt climatic changes. On the basis of existing early-warning signals, He et al. [46-48] made some improvements and developed a set of quantitative early-warning signals and methods for abrupt climatic changes. In view of the limitations of the linear method in previous research on the drought-flood state transition and the complex nonlinear processes involved in the drought-flood state transition over the Dongting Lake basin, this kind of state transition can be regarded as a systematic mutation. Therefore, previous research on the early-warning signals of abrupt changes based on the critical slowing down theory can provide effective methods for the research on the early-warning signals of the drought-flood state transition over the Dongting Lake basin.

Considering this, this study investigates the early-warning signals of the drought-flood state transition over the Dongting Lake basin based on the critical slowing theory. Firstly, the moving *t*-test (MTT) is used to detect the drought-flood state transition over the Dongting Lake basin. Then, by analyzing the variances and autocorrelation coefficients that characterize the phenomenon of critical slowing down, the early-warning signals for drought-flood state transition are explored which provides important reference for the prediction of drought and flood events over the Dongting Lake basin, thereby offering scientific and technological support for disaster prevention and mitigation over the basin.

2. Data and Methods

2.1. Data

The data used in this paper are the monthly precipitation during 1961–2020 from 97 national meteorological stations released by the National Meteorological Information Center of China Meteorological Administration. After the stations with missing data were excluded, 82 stations were finally selected. On this basis, the standardized precipitation index (SPI) based on the monthly precipitation data is calculated at each station. The SPI proposed by McKee et al. [49] describes the intensity of moisture. It is calculated by referring to [50], and the intensity grades are shown in Table 1. The SPI calculation program is derived from the National Drought Mitigation Center of the United States (https://drought.unl.edu/droughtmonitoring/SPI/SPIProgram.aspx,accessed on 5 July 2021). In this study, the annual

SPI is adopted to investigate the drought-flood state transition and its early-warning signals over the Dongting Lake basins.

Table 1. Drought-flood grades of the standardized precipitation index (S)	5PI	.).
--	-----	-----

Grades	Category	SPI Value
1	Extremely wet	2.00 ≤ SPI
2	Severely wet	$1.50 \le SPI \le 1.99$
3	Moderately wet	$1.00 \le SPI \le 1.49$
4	Mildly wet	$0.00 \le SPI \le 0.49$
5	Normal	$-0.49 \le SPI \le 0.49$
6	Mild drought	$-0.99 \le SPI \le -0.49$
7	Moderate drought	$-1.49 \le SPI \le -1.00$
8	Severe drought	$-1.99 \le SPI \le -1.5$
9	Extreme drought	SPI ≤ -2.00

2.2. Methods

2.2.1. Variance and Autocorrelation Coefficient

In this study, the moving variance and autocorrelation coefficient of the sequence are calculated to explore the early-warning signals for the drought-flood state transition in terms of the SPI sequence. The detailed method is shown in Figure 1.

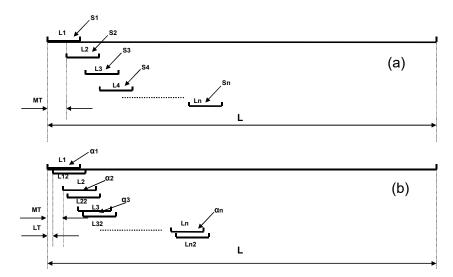


Figure 1. We calculated the variance and autocorrelation coefficient by sliding the window. (a) We calculated variance by sliding the window, L1, L2, L3, . . ., Ln denote windows of the same length(ML), S1, S2, S3, . . ., Sn represent the variances of the corresponding windows,L is the total length of the sequence, and MT is the sliding step; (b) we calculated the autocorrelation coefficient by sliding the window, L1(L12), L2(L22), L3(L32), . . ., Ln(Ln2)represent windows of the same length, α 1 denotes the autocorrelation coefficients of L1 and the L12, α 2 denotes the autocorrelation coefficient of the L2, while L22, the α n refer to the autocorrelation coefficients of the Ln and Ln2, LT represents the lag time, and L, ML and MT have the same meaning as those in (a).

Atmosphere **2021**, 12, 1082 5 of 15

2.2.2. Relation of the Critical Slowing down with the Increasing Autocorrelations and Increasing Variances

Critical slowing down often leads to an increase in the autocorrelation and variance of the fluctuations in a stochastically forced system approaching a bifurcation at a threshold value of a control parameter [14,37]. It is assumed that there is a repeated disturbance of the state variable after each period of Δt (that is, additive noise). Between disturbances, the return to equilibrium is approximately exponential with a certain recovery speed of λ . In a simple autoregressive model, this can be described as follows:

$$x_{n+1} = e^{\lambda \Delta t} x_n + s \varepsilon_n \,, \tag{1}$$

where x_n is the deviation of the state variable x from the equilibrium state, \mathcal{E}_n is a random quantity satisfying the standard normal distribution, and s is the mean square error. If λ and Δt are independent of x_n , this model can be written as a first-order autoregressive (AR(1)) process:

$$X_{n+1} = \alpha X_n + S \mathcal{E}_n \,, \tag{2}$$

where the autocorrelation $\alpha = e^{\lambda \Delta t}$ is zero for white noise and close to one for red noise.

The variance of an AR(1) is as follows:

$$Var(x_{n+1}) = E(x_n^2) + (E(x_n))^2 = \frac{s^2}{1 - \alpha^2}$$
 (3)

Generally, when the system approaches the critical point, the return of the disturbance with small amplitude to equilibrium gets slower and slower [51,52], that is, λ approaches zero and the autocorrelation α is close to one. Thus, the variance tends to infinity (Equation (3)). Therefore, the increasing variance and autocorrelation coefficient can be used as the early-warning signals for the system approaching the critical point. In the calculation of this study, s represents the variance of the whole sequence, which is a constant value, while the Var value in Equation (3) changes with the window size and moving step.

2.2.3. Moving t-Test Method

The MTT is used to test the abrupt change in a sequence by examining the significance of the difference between the averages of two samples. The basic idea is as follows. The problem of whether there is a significant difference between the mean values of two subsequences in a climatic sequence can be regarded as the test of the significance of the population means between two samples. If the difference between the mean values of two subsequences exceeds a certain confidence level, it is considered that the mean value has changed qualitatively and thus the abrupt change occurs. Note that in order to avoid the drift of the tipping point caused by the artificially selected subsequence length in the MTT method, the subsequence length can be repeatedly changed for experimental comparison in specific application to improve the reliability of calculation results. In this study, this method is used to detect the abrupt change of sequences.

3. Results and Analysis

Since the mid-20th century, the drought-flood state transition over the Dongting Lake basin occurs more and more frequently under significant global warming, and correspondingly, the duration for the climate maintained in a drought or flood state is getting shorter and shorter. The drought-flood state transition often leads to regional flood, drought or deterioration of water environment [1,6]. Therefore, it is of great significance to study the early-warning signals for the drought-flood state transition over the Dongting Lake basin. In this section, the early-warning signals for the drought-flood state transition over the Dongting Lake basin are explored based on the critical slowing down theory.

3.1. Detection of the Drought-Flood State Transition over the Dongting Lake Basin

Firstly, the SPI values over the Dongting Lake basin are calculated through the grid-area-weighted average method [53]. This section mainly investigates the decadal variation characteristics of the SPI and the characteristics of the drought-flood state transition (abrupt change) over the Dongting Lake basin.

Figure 2 shows that the SPI sequence over the Dongting Lake basin in recent 60 years is featured by obvious decadal changes. The trend of the moving averages reveals that there are obvious more precipitation-to-less precipitation (less precipitation-to-more precipitation) events in 1983 and 2003 (1993 and 2016). All these events are collectively referred to as the drought-flood state transition events (abrupt change). Studies have shown that there are obvious decadal changes for the drought and flood conditions over the Dongting Lake basin. In particular, the less precipitation-to-more precipitation events around 1993 are widely concerned [4,5].

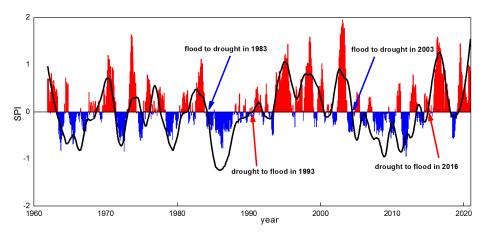
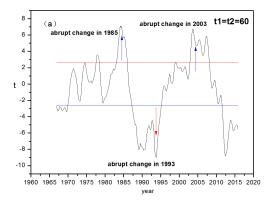


Figure 2. Temporal variation of the SPI sequence over the Dongting Lake basin in recent 60 years. The bar represents the SPI value (red bar indicates positive SPI and blue bar indicates negative SPI), and the black curve represents the trend extracted from 51month moving average.

Figure 3 shows the detection curve for the abrupt change in the SPI sequence over the Dongting Lake basin based on the MTT method. When the size t of the sliding window is set as 5 years, multiple abrupt changes will be found in the SPI sequence, which are especially significant in 1985, 1993 and 2003. When t is set as 10 years, two obvious abrupt changes in 1993 and 2003 can also be found. Therefore, it can be considered that the SPI sequence over the Dongting Lake basin has two obvious abrupt changes (drought-flood state transition) in 1993 and 2003. In the following section, the early-warning signals for these two abrupt changes are explored based on the critical slowing down theory.



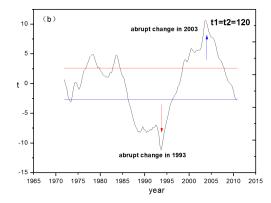


Figure 3. Detection of the abrupt change in the SPI sequence over the Dongting Lake basin based on the moving t-test (MTT) method with the size of the sliding window being. (a) 5 years (60 months) and (b) 10 years (120 months), respectively.

Figure 3 shows that the SPI sequence over the Dongting Lake basin has obvious decadal varaitions, and there are two obvious abrupt changes (drought-flood state transition) in 1993 and 2003. Therefore, the state of the drought or flood over the Dongting Lake basin can be divided into three periods: 1961–1992, 1993–2002 and 2003–2020, all of which experience the state transition of less precipitation-more precipitation-less precipitation (Figure 4). Specifically, during 1961–1992, the drought intensity in central and southeast Hunan is stronger; during 1993–2002, the waterlogging in western and northern Hunan is more obvious; during 2003–2020, the drought intensity in the east of central Hunan and the waterlogging in northern and western Hunan are stronger. The spatial distribution of the SPI values demonstrates that there are significant differences in the spatial distributions of the drought-flood conditions before and after the transition in 1993.

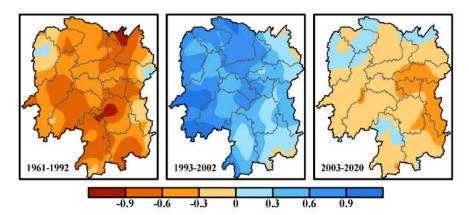
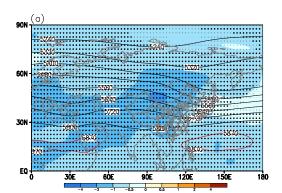


Figure 4. Spatial distributions of the SPI values in different decades over the Dongting Lake basin.

3.2. Circulation Background of the Drought-Flood Transition over the Dongting Lake Basin

The above analysis shows that the Dongting Lake basin has experienced less precipitation—more precipitation—less precipitation state transitions in three periods of 1961–1992, 1993–2002 and 2003–2020. In order to explore the circulation backgrounds for drought-flood state transitions over the Dongting Lake basin under different decadal backgrounds, the 500-hPa geopotential height and 200-hPa zonal wind are analyzed in this section.

The low-pass filtering is carried out on the 500-hPa geopotential height to filter out the inter-annual components below 10 years. Figure 5 shows the fields of 500-hPa geopotential height and the composite anomalies corresponding to different decades over the Dongting Lake basin. In the first stage of the drought-flood transition (Figure 5a), the intensity of the polar vortex in Asian is strong, and there are significant negative anomalies over the Ural Mountains and the Okhotsk Sea, which is unfavorable to the development of the blocking highs over the Ural Mountains and Okhotsk Sea. For the mid-latitude region over the Asia continent, there are also significant negative anomalies, resulting in the relatively straight westerly circulation, which is unfavorable for the southward movement of cold air. Meanwhile, the western Pacific subtropical high (WPSH) is significantly weak, which is not beneficial to the transportation of the warm and humid airflow from the western Pacific and the South China Sea to southern China. This circulation configuration leads to less rainfall in this decade over the Dongting Lake basin. In the second stage of the drought-flood transition (Figure 5b), the field of 500-hPa geopotential height and corresponding composite anomalies show that the polar vortex is slightly weaker, and the Eurasian mid-high latitudes present a distribution of two-trough and one-ridge. Specifically, the Ural Mountains and its western area as well as the Okhotsk Sea are controlled by significant negative anomalies of geopotential height, indicating weak blocking highs over the Ural Mountains and the Okhotsk Sea. The Lake Baikal and its southern area are dominated by the ridge with obvious meridional circulations, which is beneficial to the southward movement of the cold air to affect China. Meanwhile, the westward extension and enhancement of the WPSH are more significant compared with that in the previous stage, and the Dongting Lake basin is located in the southward water vapor conveyor belt to the west side of the WPSH, which are favorable to more precipitation in this decade over this region. In the third stage of the drought-flood transition over the Dongting Lake basin (Figure 5c), the Eurasian mid-high latitudes are controlled by positive anomalies of geopotential height. The blocking highs over the Ural Mountains and the Okhotsk Sea significantly strengthen, which is not conducive to the southward movement of cold air. This is also consistent with the overall background of higher decadal temperature. At this stage, the WPSH is obviously stronger with larger area and westward position, and the ridge line is obviously northward to its normal position. The main rain belt is located in northern China, while the weather to the south of the Yangtze River is sunny, hot with less rain under the control of the WPSH.



Atmosphere 2021, 12, 1082 9 of 15

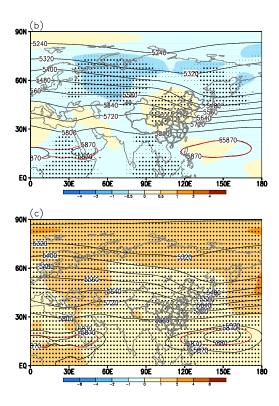


Figure 5. 500-hPa geopotential height field (black solid line, unit: gpm) and the composite anomaly field (shading, unit: gpm) at the stages of (a) 1961–1992, (b) 1993–2002 and (c) 2003–2020 over the Dongting Lake basin. The red solid line represents the climatology of 5870-gpm isoline from 1981 to 2010, and the black dots indicate the area with the values exceeding the confidence level at 95%.

Figure 6 shows the 200-hPa zonal wind and the composite anomalies corresponding to different decades over the Dongting Lake basin. In the first stage (Figure 6a), the 200hPa zonal wind over the mid-latitude region in eastern Asia presents a significant "positive-negative" distribution, where positive zonal wind anomalies appear over the northern region of China, and negative zonal wind anomalies over the southern region. This reveals that the location of the westerly jet is more northerly and easterly than its normal position. Therefore, it is difficult for the cyclonic disturbance in the westerlies to affect the southern region of China, leading to less rainfall in the Dongting Lake basin. In the second stage (Figure 6b), the circulation is obviously opposite to that of the previous stage. Negative zonal wind anomalies prevail over the northern region of China, while positive zonal wind anomalies over the southern region, which implies that the location of westerly jet is more south than its normal location, thus leading to significantly more-than-normal precipitation over the Dongting Lake basin. In the third stage (Figure 6c), the distribution of 200-hPa zonal wind over the eastern part of the mid-latitude in the Asia continent is similar to that in the first stage, posing a "positive and negative" distribution pattern from north to south. The location of westerly jet stream is slightly norther than normal, causing less precipitation over the Dongting Lake basin, but not as significant as that in the first stage. The results are consistent with the conclusions in previous studies [4,54].

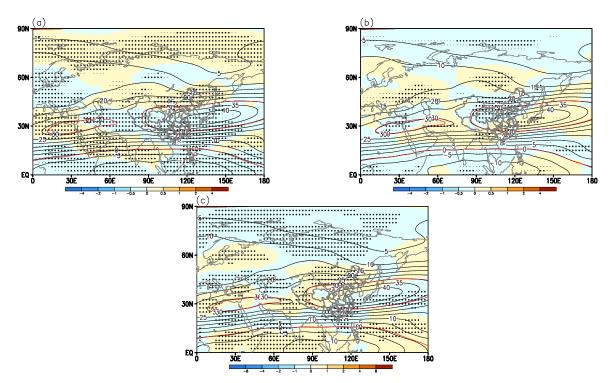


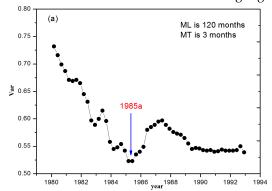
Figure 6. Same as Figure 5, (a) 1961–1992, (b) 1993–2002 and (c) 2003–2020 over the Dongting Lake basin. but for 200-hPa zonal wind (contours, unit: m·s⁻¹), where the red solid lines represent the climatology of wind speed isolines of 0 m·s⁻¹ and 30 m·s⁻¹ from 1981 to 2010.

3.3. Early-Warning Signals for the Drought-Flood State Transition over the Dongting Lake Basin Based on the Variance

At the beginning of the 21st century, scientists found that the critical slowing down theory has important potentials in studying the early-warning signals for complex dynamical systems. This section explores the early-warning signals for the drought-flood state transition over the Dongting Lake basin by the variance that characterizes the phenomenon of critical slowing down.

Figure 7 shows the detection of the abrupt changes in the SPI sequence over the Dongting Lake basin based on the variance signal. Here, ML in this study is set to 120 months (10 years), and the MT is set to 3 months. For the variance of the sequence, the window width and the MT, please refer to Figure 1. Figure 7a shows the variance-based detection of the SPI sequence for the abrupt change in 1993. As shown in Figure 1, by sliding forward a fixed-width window with fixed MT in the SPI sequence, a new sequence is obtained in the fixed-width window with the new location. The variance is further calculated for this new sequence. The arrow in Figure 7a illustrates that the variance starts to increase around 1985. The above theoretical analysis shows that the critical slowing down leads to the decrease in the internal change rate of the drought-flood state, and the similarity between the states at any time and the previous state in the complex nonlinear system as well. Thus, the autocorrelation coefficient is close to one, and the variance is close to infinity according to Formula (3). Therefore, the critical slowing down phenomenon of increasing variance when the complex nonlinear system approaches to the critical point might be an early-warning signal for the abrupt change in complex nonlinear system. It is considered that the early-warning signal for the drought-flood state transition over the Dongting Lake basin occurs around 1985. It can be seen that the early-warning signal is about 8 years ahead of the abrupt change in the SPI sequence, indicating that the variance acts as a good indicator for the prediction of the drought-flood state transition. Similarly, the variance signal for the abrupt change of the SPI sequence in 2003 over the

Dongting Lake basin appears in 1998 (in Figure 7b). The critical slowing down based on increasing variance is a potential early-warning signal for the drought-flood state transition over the Dongting Lake basin.



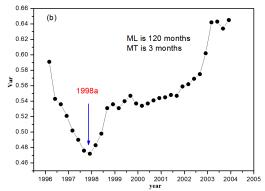


Figure 7. Variance signal-based detection for abrupt changes in (a) 1993 and (b) 2003 in the SPI sequence over the Dongting Lake basin.

3.4. Early-Warning Signals for the Drought-Flood State Transition over the Dongting Lake Basin Based on the Autocorrelation Coefficient

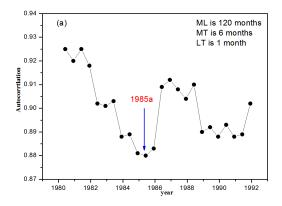
According to the above studies, the critical slowing down phenomenon based on the increasing variance can be used as an early-warning signal for the drought-flood state transition over the Dongting Lake basin. In this section, another quantity that characterizes the critical slowing down theory, the autocorrelation coefficient, is used to explore the early warning signals for the drought-flood state transition over the Dongting Lake.

Figure 8 shows the autocorrelation coefficient signal-based detection of the abrupt changes in the SPI sequence over the Dongting Lake basin. Similarly, ML in this study is set to 120 months (10 years), the MT is set to 6 months, and the lag time (LT) is set to 1 month. The variance of the sequence, window width and MT are shown in Figure 1. As shown in Figure 1, by sliding forward a fixed-width window with one fixed LT in the original SPI sequence, another sequence with the same length as the sequence in the original window is obtained. The correlation between the obtained sequence and the original sequence is calculated, thus obtaining the lag autocorrelation for the sequence itself. The MT is the same as described in the variance signal section. Figure 8a displays the autocorrelation coefficient-based detection for the abrupt change of the SPI sequence in 1993. The autocorrelation coefficient begins to increase from year marked with an arrow in Figure 8a. As the critical slowing down leads to the decrease of the internal change rate in the drought-flood state and the state of complex nonlinear system at any time is becoming more and more similar to its previous state, the autocorrelation coefficient of the SPI sequence is close to one. It is considered that the early-warning signal for the drought-flood state transition in 1993 appears around 1985. Therefore, the early-warning signal of the increasing autocorrelation coefficient appears eight years before the drought-flood state transition over the Dongting Lake basin in 1993, which serves as a good indicator for the prediction of the drought-flood state transition over the Dongting Lake basin. Similarly, the autocorrelation coefficient signal for the abrupt change over the Dongting Lake basin in 2003 appears in 1998 in the SPI sequence. Therefore, in view of the two drought-flood state transitions of the SPI sequence over the Dongting Lake basin, the early-warning signal of increasing autocorrelation coefficient for abrupt changes appears 5–10 years before the state transition. The critical slowing down phenomenon based on the increasing autocorrelation coefficient is a potential early warning signal for the drought-flood state transition over the Dongting Lake basin.

Note that when the data amount remains constant, the larger the window width and the longer the MT, the more stable the result will be. Variations in window width and LT

do not influence the occurrence of the signal, but only affect the signal stability. Related studies are detailed in Wu et al. [25–28], which will not be covered here.

In summary, when the drought-flood system approaches the critical point, the phenomena of critical slowing down will occur, featured by increasing variance and increasing autocorrelation coefficient, indicating the occurrence of a drought-flood state transition. Based on the variance and autocorrelation coefficient, the occurrence times of the early-warning signals for the drought-flood state transition are basically the same, which further proves the feasibility to explore the early warning signals for the drought-flood state transition based on the critical slowing down theory.



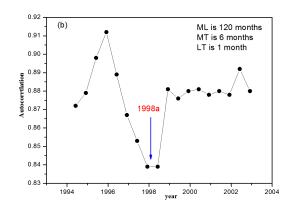


Figure 8. Autocorrelation coefficient signal-based detection for the abrupt changes in (a) 1993 and (b) 2003 of the SPI sequence over the Dongting Lake basin.

4. Conclusions and Discussions

In view of the complex nonlinear characteristics of the drought-flood state transition, the theories and methods related to the nonlinear science are adopted for investigation in this study. First, the MTT method is performed on the SPI sequence over the Dongting Lake basin in recent 60 years for detecting the drought-flood state transitions. Afterwards, the physical basis and statistical significance of the critical slowing down phenomenon are further introduced. On this basis, the theory of critical slowing down is used to explore the early-warning signals for abrupt changes in the SPI sequence over the Dongting Lake basin. The conclusions are as follows.

Detection on the SPI sequence over the Dongting Lake basin reveals that there are two significant drought-flood state transitions, namely the drought-to-flood event in 1993 and the flood-to-drought event in 2003.

- (1) By analyzing the 500-hPa geopotential height and 200-hPa zonal wind, it is found that the circulation backgrounds during 1961–1992 and 2003–2020 are favorable for less precipitation and drought over the Dongting Lake Basin. While the circulation backgrounds during 1993–2002 results in more precipitation and even floods over the Dongting Lake basin.
- (2) The critical slowing down phenomena with the increases in variances and autocorrelation coefficients appear 5–10 years before the two drought-flood state transitions, indicating the feasibility of exploring the early-warning signals for drought-flood state transitions based on the critical slowing down theory.
- (3) The critical slowing down phenomena with the increases in variances and autocorrelation coefficients appear 5–10 years before the two drought-flood state transitions, indicating the feasibility of exploring the earlywarning signals for drought-flood state transitions based on the critical slowing down theory.

Through the study on the drought-flood state and related early-warning signals over the Dongting Lake basin, the reliability of the early-warning signals of the variance and

autocorrelation coefficient characterizing the critical slowing down phenomenon in detecting abrupt changes is demonstrated. Besides, the introduction of the critical slowing down theory is of important practical significance and scientific value for deep understanding of drought-flood state transition and for exploring the early warning signals of its state transition based on the SPI sequence. Furthermore, the extensive application of this method is expected to provide important scientific and technological support for disaster prevention and mitigation over the Dongting Lake basin and even in the middle and lower reaches of the Yangtze River. Note that although the the results of the present manuscript show that the increasing variances and autocorrelation coefficients caused by the critical slowing down phenomenon are potential early warning signals, and more robust conclusions require more experiments and applications to confirm. The confidence probability of early warning signals, the circulation characteristics, the characteristics of external forcing changes, and the relationship between the critical slowing down phenomenon and the intensity/amplitude of abrupt changes before and after the transitions still need to be further studied.

This paper studies a method that can identify the early warning signal of Drought-Flood state transition in a long period of time in the future, which strictly belongs to the category of climate shift. The climate shift is not necessarily a hazard per se, but it will have a great impact on the frequency, intensity and duration of hazards. The early warning signal of climate shift can warn the change and remind people to be prepared for the possible consequences of this change. For example, we can study some specific disasters and determine the threshold of precipitation which can lead to disaster, such as flood and landslide [55]; in addition, the high-risk locations of disasters can be determined in the region with reference to the disasters that have occurred in the previous more Rainy period or less rainy period and their impacts, and the probability of possible disasters in the future can be estimated in combination with the current exposure and vulnerability.

EWS involves six aspects [27]: (1) What is happening with respect to the hazards and/vulnerabilities? (2) When are the impacts likely to occur? (3) Where are at risk? (4) Who are at risk? (5) What causes the risk? (6) How can EWS play a better role, whether it is for a specific hazard or guiding a long-term risk response process? For drought or flood disasters, the early warning signal of Drought-Flood state transition can answer the first question about hazards and the second question, as well as help to solve the sixth question. On the premise that the accuracy of climate model prediction needs to be further improved, the method in this paper can use the existing observation data to provide a new technology for early warning of climate transition. In addition, it can also analyze the influencing factors such as atmospheric circulation and sea surface temperature in time, so as to provide comparison and reference for improving the climate model.

Author Contributions: Methodology H.W., W.H. and P.Y.; writing original draft preparation, H.W. and Y.Z.; writing review and editing, H.W. and W.H.; visualization, D.Z. and P.Y. All authors have read and agreed to the published version of the manuscript.

Funding: This work was funded by the National Natural Science Foundation of China (Grant: 42005058,41775078,41675092), the Hunan Provence Natural Science Foundation of China (Grant: 2020JJ5298), the Natural Science Research Project of Higher Education in Jiangsu Provinces (Grant: 20KJB170004), the school-level research projects of Yancheng Institute of Technology (Grant: xjr2019052).

Institutional Review Board Statement: Not applicable.

Informed Consent Statement: Not applicable.

Data Availability Statement: The data of 97 stations can be downloaded from China Meteorological Data Center (https://data.cma.cn,accessed on 10 July 2021), further inquiries can be directed to the corresponding author/s.

Conflicts of Interest: The authors declare no conflict of interest.

References

1. Huang, Q.; Sun, Z.; Opp, C.; Lotz, T.; Jiang, J.; Lai, X. Hydrological drought at dongting lake: Its detection, characterization, and challenges associated with three gorges dam in central Yangtze, China. *Water Resour. Manag.* **2014**, *28*, 5377–5388, doi:10.1007/s11269-014-0807-8.

- 2. Grubler, A. IPCC fifth assessment report. Weather 2014, 68, 310.
- 3. Tang, G.; Ding, Y.; Wang, S.; Ren, G.; Liu, H.; Zhang, L. Comparative analysis of the time series of surface air temperature over China for the last 100 years. *Adv. Clim. Chang. Res.* **2010**, *1*, 11–19.
- 4. Zhang, J.; Liao, Y.; Wu, H.; Zhang, J.; Zhao, H. Characteristics of atmospheric circulation anomalies and drought in summer and autumn in Hunan province. *J. Arid. Meteorol.* **2018**, *36*, 353–364. (In Chinese)
- 5. Wu, H.; Zhang, J.; Yan, P.; Zeng, Y.; Duan, L. The research on drought characteristics of SPI at different time scales in Hunan province. *Adv. Meteorol. Sci. Technol.* **2021**, *11*, 139–147. (In Chinese)
- 6. Yang, S.; Wu, B.; Zhang, R.; Zhou, S. Relationship between an abrupt drought-flood transition over mid-low reaches of the Yangtze River in 2011 and the intraseasonal oscillation over mid-high latitudes of East Asia. *Acta Meteorol. Sin.* **2013**, 27, 129–143, doi:10.1007/s13351-013-0201-0.
- 7. Ding, R.; Li, J. Relationships between the limit of predictability and initial error in the uncoupled and coupled Lorenz models. *Adv. Atmos. Sci.* **2012**, *29*, 1078–1088, doi:10.1007/s00376-012-1207-8.
- 8. Wan, S.; Feng, G.; Dong, W.; Li, J.; Gao, X.; He, W. On the climate prediction of nonlinear and non-stationary time series with the EMD Method. *Chin. Phys. B* **2005**, *14*, 628–633.
- 9. Ding, R.; Liu, B.; Gu, B.; Li, J.; Li, X. Predictability of ensemble forecasting estimated using the Kullback-Leibler divergence in the Lorenz Model. *Adv. Atmos. Sci.* **2019**, *36*, 837–846, doi:10.1007/s00376-019-9034-9.
- 10. Feng, G.-L.; Yang, J.; Zhi, R.; Zhao, J.-H.; Gong, Z.-Q.; Zheng, Z.-H.; Xiong, K.-G.; Qiao, S.-B.; Yan, Z.; Wu, Y.-P.; et al. Improved prediction model for flood-season rainfall based on a nonlinear dynamics-statistic combined method. *Chaos Solitons Fractals* **2020**, *140*, 110160, doi:10.1016/j.chaos.2020.110160.
- 11. Hu, S.; Chou, J.; Cheng, J. Three-pattern decomposition of global atmospheric circulation: Part I—Decomposition model and theorems. Clim. Dyn. 2018, 50, 2355–2368.
- 12. Fan, M.; Jiang J. Study of heavy causes contrasting 1999 to 1998 summer over Changjiang river basin. *Meteorol. Mon.* **2001**, 27, 38–41. (In Chinese)
- 13. Gao, H.; Wang Y. Sea surface temperature and the general circulation in 2007 and their influences on the climate of China. *Meteorol. Mon.* **2008**, *34*, 107–112. (In Chinese)
- 14. Scheffer, M.; Bascompte, J.; Brock, W.A.; Brovkin, V.; Carpenter, S.R.; Dakos, V.; Held, H.; van Nes, E.; Rietkerk, M.; Sugihara, G. Early-warning signals for critical transitions. *Nature* **2009**, *461*, 53–59, doi:10.1038/nature08227.
- 15. Fisher, L.; Scheffer, M. Critical transitions in nature and society. Am. J. Psychol. 2011, 124, 365–367.
- Gopalakrishnan, E.A.; Sharma, Y.; John, T.; Dutta, P.S.; Sujith, R.I. Early warning signals for critical transitions in a thermoacoustic system. Sci. Rep. 2016, 6, 35310, doi:10.1038/srep35310.
- 17. Lenton, T.M.; Livina, V.N.; Dakos, V.; van Nes, E.; Scheffer, M. Early warning of climate tipping points from critical slowing down: Comparing methods to improve robustness. *Philos. Trans. R. Soc. Math. Phys. Eng. Sci.* **2012**, 370, 1185–1204, doi:10.1098/rsta.2011.0304.
- 18. Alley, R.; Marotzke, J.; Nordhaus, W.; Overpeck, J.; Peteet, D.; Pielke, R.; Pierrehumbert, R.; Rhines, P.; Stocker, T.; Talley, L. Abrupt climate change. *Science* **2003**, 299, 2005–2010.
- 19. He, W.; Xie, X.; Mei, Y.; Wan, S.; Zhao, S. Decreasing predictability as a precursor indicator for abrupt climate change. *Clim. Dyn.* **2021**, *56*, 3899–3908, doi:10.1007/s00382-021-05676-1.
- 20. Severinghaus, J.P. Abrupt climate change at the end of the last glacial period inferred from trapped air in polar ice. *Science* **1999**, 286, 930–934, doi:10.1126/science.286.5441.930.
- 21. UNISDR Terminology. UNISDR (United Nations International Strategy for Disaster Risk Reduction), Geneva, 2012. Available online: http://www.unisdr.org/we/inform/terminology (accessed on 22 August 2021).
- 22. Gruntfest, E.; Downing, T.; White, G. Big Thompson flood exposes need for better flood reaction system to save lives. *Civ. Eng.* **1978**, *48*, 72–73.
- 23. Lewis, J. Development in Disaster-Prone Places: Studies of Vulnerability; Intermediate Technology Publications: London, UK, 1999.
- 24. Cyr, J.F.S. At risk: Natural hazards, people's vulnerability, and disasters. *J. Homel. Secur. Emerg. Manag.* **2005**, 2, doi:10.2202/1547-7355.1131.
- 25. Wisner, B.; Gaillard, J.; Kelman, I. Handbook of Hazards and Disaster Risk Reduction; Routledge: Abingdon, London, 2012.
- 26. Mileti, D.; Ebrary, I. Disasters by design—A reassessment of natural hazards in the United States. Ameaas 1999, 8, 699.
- 27. Kelman, I.; Glantz, M.H. Early Warning Systems Defined; Springer: Dordrecht, Germany, 2014.
- 28. Hoekstra, S.; Klockow, K.; Riley, R.; Brotzge, J.; Brooks, H.; Erickson, S. A preliminary look at the social perspective of warn-on-forecast: Preferred tornado warning lead time and the general public's perceptions of weather risks. *Weather Clim. Soc.* **2011**, *3*, 128–140, doi:10.1175/2011wcas1076.1.
- 29. IPCC, Climate Change 2014: Synthesis Report. Contribution of Working Groups I, II and III to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change 2014.https://www.ipcc.ch/(accessed on 22 August 2021).
- 30. Kundzewicz, Z.; Pińskwar, I.; Brakenridge, G. Large floods in Europe, 1985–2009. Hydrol. Sci. J. 2013, 58, 1–7.

31. EM-DAT: The CRED/OFDA International Disaster Database. Available online: http://www.emdat.be (accessed on 20 January 2020).

- 32. DFO. Available online: http://floodobservatory.colorado.edu/Archives/index.html (accessed on 22 August 2021).
- 33. Zhang, D.-D.; Yan, D.-H.; Lu, F.; Wang, Y.-C.; Feng, J. Copula-based risk assessment of drought in Yunnan province, China. *Nat. Hazards* **2014**, *75*, 2199–2220, doi:10.1007/s11069-014-1419-6.
- 34. Wang, L.; Wang, S.; Wang, X.; Zhang, L.; Wang, F. Feasibility study of five drought indices for use in Yunnan Province. *J. Irrig. Drain.* **2017**, *36*, 117–124.
- 35. Yu, Y.; Wang, J.; Li, X. Monitoring the occurrence of drought in Central Yunnan Province based on MODIS data. *J. Irrig. Drain.* **2018**, *37*, 91–98.
- 36. Yu, Y.; Wang, J.; Cheng, F.; Chen, Y. Soil moisture by remote sensing retrieval in the tropic of cancer of Yunnan Province. *Pol. J. Environ. Stud.* **2020**, *29*, 1981–1993, doi:10.15244/pjoes/110203.
- 37. Carpenter, S.; Brook, W. Rising variance: A leading indicator of ecological transition. Ecol. Lett. 2006, 9, 311–318.
- 38. Guttal, V.; Jayaprakash, C. Changing skewness: An early warning signal of regime shifts in ecological systems. *Ecol. Lett.* **2008**, *11*, 450–460.
- 39. Yu, L.; Hao, B. Phase Transitions and Critical Phenomena; Science Press: Beijing, China, 1984. (In Chinese)
- 40. Yan, R.; Jiang, C.; Zhang, L. Study on critical slowing down phenomenon of radon concentrations in water befer the Wenchuan Ms 8.0 earthquake. *Chin. J. Geophys.* **2011**, *54*, 1817–1826. (In Chinese)
- 41. Wu, H.; Feng, G.; Hou, W.; Yan, P. The early warning signals of abrupt climate change in different regions of china. *Acta Phys. Sin.* **2013**, *62*, 059202. (In Chinese)
- 42. Wu, H.; Hou, W.; Yan, P. Using the principle of critical slowing down to discuss the abrupt climate change. *Acta Phys. Sin.* **2013**, 62, 039206. (In Chinese)
- 43. Wu, H.; Hou, W.; Yan, P.-C.; Zhang, Z.-S.; Wang, K. A study of the early warning signals of abrupt change in the Pacific decadal oscillation. *Chin. Phys. B* **2015**, *24*, 089201, doi:10.1088/1674-1056/24/8/089201.
- 44. Tong J L, Wu H, Hou W; et al. The early warning signals of abrupt temperature change in different regions of China over recent 50 years. *Chin. Phys. B* **2014**, 23, 049201.
- 45. Yan, P.C.; Feng, G.L.; Hou, W. A novel method for analyzing the process of abrupt climate change. *Nonlinear Process. Geophys.* **2015**, 22, 249–258, doi:10.5194/npg-22-249-2015.
- 46. He, W.; Zhao, S.; Liu, Q.; Jiang, Y.; Deng, B. Long-range correlation in the drought and flood index from 1470 to 2000 in Eastern China. *Int. J. Clim.* 2015, 36, 1676–1685, doi:10.1002/joc.4450.
- 47. He, W.-P.; Liu, Q.-Q.; Gu, B.; Zhao, S.-S. A novel method for detecting abrupt dynamic change based on the changing Hurst exponent of spatial images. *Clim. Dyn.* **2016**, 47, 2561–2571, doi:10.1007/s00382-016-2983-0.
- 48. Xie, X.-Q.; He, W.-P.; Gu, B.; Mei, Y.; Zhao, S.-S. Can kurtosis be an early warning signal for abrupt climate change? *Clim. Dyn.* **2018**, *52*, 6863–6876, doi:10.1007/s00382-018-4549-9.
- 49. Mckee, T.; Doesken, N.; Kleist, J. The relationship of drought frequency and duration to time scales. In Proceedings of the 8th Conference on Applied Climatology, Hanover, Germany, 17–22 January 1993.
- 50. Belayneh, A.; Adamowski, J. Standard precipitation index drought forecasting using neural networks, wavelet neural networks, and support vector regression. *Appl. Comput. Intell. Soft Comput.* **2012**, 2012, 1–13, doi:10.1155/2012/794061.
- 51. Dakos, V.; Scheffer, M.; van Nes, E.; Brovkin, V.; Petoukhov, V.; Held, H. Slowing down as an early warning signal for abrupt climate change. *Proc. Natl. Acad. Sci. USA* **2008**, *105*, 14308–14312, doi:10.1073/pnas.0802430105.
- 52. Held, H.; Kleinen, T. Detection of climate system bifurcations by degenerate fingerprinting. *Geophys. Res. Lett.* **2004**, 31, doi:10.1029/2004gl020972.
- 53. Jones, P.D.; Groisman, P.Y.; Coughlan, M.; Plummer, N.; Wang, W.-C.; Karl, T.R. Assessment of urbanization effects in time series of surface air temperature over land. *Nature* **1990**, *347*, 169–172, doi:10.1038/347169a0.
- 54. Li, S.; Feng, G.; Hou, W.; Cheng, J. Characteristics of atmospheric circulation patterns over East Asia and their impacts on precipitation in summer. *Clim. Res.* **2019**, *78*, 117–133, doi:10.3354/cr01544.
- 55. Luca, D.; Versace, P. Diversity of rainfall thresholds for early warning of hydro-geological disasters. Adv. Geosci. 2017, 44, 53–6