

# Article The Spiderweb Structure of Stratocumulus Clouds

## Georgios Matheou <sup>1</sup>, Anthony B. Davis <sup>2</sup> and João Teixeira <sup>2</sup>

- <sup>1</sup> Department of Mechanical Engineering, University of Connecticut, Storrs, CT 06269, USA; matheou@uconn.edu
- <sup>2</sup> Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA 91109, USA

Version June 23, 2020 submitted to Atmosphere

- Abstract: Stratocumulus clouds have a distinctive structure composed of a combination of lumpy
- <sup>2</sup> cellular structures and thin elongated regions, resembling canyons or slits. The elongated slits are
- <sup>3</sup> referred to as "spiderweb" structure to emphasize their interconnected nature. Using very high
- <sup>4</sup> resolution large-eddy simulations (LES) it is shown that the spiderweb structure is generated by
- <sup>5</sup> cloud-top evaporative cooling. Analysis of liquid water path (LWP) and the cloud liquid water
- <sup>6</sup> content shows that cloud-top evaporative cooling generates relatively shallow slits near the cloud
- 7 top. Because most of the liquid water mass is concentrated near the cloud top, these regions of clear
- air have a large impact on the entire-column LWP. When the evaporative cooling is suppressed in the
- LES, LWP exhibits cellular lumpy structure without the elongated low LWP regions. Even though the
- <sup>10</sup> spiderweb signature on the LWP distribution is negligible, the cloud-top evaporative cooling process
- significantly affects integral boundary layer quantities, such as the vertically integrated turbulent
- kinetic energy, mean liquid water path, and the entrainment rate. In a pair of simulations driven only
- <sup>13</sup> by cloud-top radiative cooling, evaporative cooling nearly doubles the entrainment rate.

Keywords: stratocumulus clouds; cloud holes; cloud-top evaporative cooling; buoyancy reversal;
 large-eddy simulation

## 16 1. Introduction

Stratocumulus (Sc) clouds form near the surface, covering about 20% the Earth's surface. Sc 17 have a large effect on the Earth's energy balance. Small variations in the Sc area coverage can 18 produce energy-balance changes comparable to those due to greenhouse gases [e.g., 1-4]. Sc have 19 a distinctive structure composed of a combination of lumpy cellular structures and thin elongated 20 regions, resembling canyons or slits. See, for instance, observations in Figure 1 and additional 21 observations in Fig. 1 of [5] and Figs. 5 and 6 of [6]. This characteristic Sc structure is also reproduced 22 in some large-eddy simulations (LES) [7-9]. The cloud structure registers in the radiance fields in 23 observations (Figure 1) and liquid water path (LWP) in model data. In the present study, of primary 24 interest are the thin elongated regions. We refer to these structures as "spiderweb" to emphasize their 25 interconnected nature <sup>1</sup>. 26

The objective of the present study is to understand the physical processes that create the Sc spiderweb structure. Sc radiative properties depend on the liquid water spatial structure. In turn, the cloud liquid water spatial structure is a result of a turbulent flow whose dynamics are modulated by the various physical processes, such as shear, buoyancy, phase changes, cloud microphysics, etc. The

Spiderweb types vary greatly by spider spices. Even though spiral orb webs are often depicted, webs can be irregular. The stratocumulus cloud-top slits loosely resemble the internal structure of webs made by hackledmesh weavers, i.e., members of the spider family Amaurobiidae. The term "spiderweb" is used in a broad sense without implying true representation of either part or all of any web produced by a spider.

<sup>31</sup> present study aspires to create direct links between the atmospheric boundary layer dynamics and

cloud radiative properties by linking the effects of individual physical processes to the cloud liquidstructure.

In situ observations [e.g., 10-12] and high resolution LES [7,9] show a complex three dimensional 34 cloud structure. Presently, we simplify the exploration of the cloud spatial structure by considering 35 the cloud liquid water path – integrating the cloud depth cloud-depth dimension to form a two 36 dimensional field. Consideration of the LWP is not a significant limitation because of the stratiform 37 nature of Sc. Presently, only Only cloud macrophysical effects are considered and variations of LWP are 38 only related to covariances of total water content, pressure, and temperature. For non precipitating and 39 non drizzling Sc and model resolutions of 1–10 m, this approximation is expected to result in sufficiently 40 representative LWP spatial structure [9]. At smaller scales (centimeters), cloud microphysical effects 41 can affect the local cloud liquid distribution. For instance, regions of low droplet concentration (and 42 consequently low cloud liquid mixing ratios) can be created because of droplet inertial effects [13]. 43 We hypothesize that two main mechanisms control the spatial LWP structure: 44 (a) boundary-layer-deep convective motions, which create the cloud lumpy cellular structure; 45 and (b) evaporative cooling near the cloud top, which creates the spiderweb structure. The hypothesis 46 is based on observations of convection organization confined under an inversion [e.g., 14,15] and 47 visualizations of stratocumulus top turbulence in fine-scale process-level model [e.g., 16, Fig. 3] 48 and [e.g., 17, Fig. 5] models, e.g., [16, Fig. 3] and [17, Fig. 5]. Evaporative cooling and the resulting buoyancy reversal instability (BRI) create shallow groves on the cloud top. 50 The working hypothesis has two important implications for Sc physics: (a) self-similarity of cloud 51 liquid spatial structure may not hold across all scales because two different processes with different 52 length and time scales modulate the cloud liquid distribution. For instance, Davis et al. [18] and Ma et 53 al. [19] report a scale break in observations of Sc liquid water content at 2-1-5 m. (b) attribution of cloud liquid structure to different physical processes. The importance of convective motions driven 55 by surface buoyancy, cloud top radiative cooling, and evaporative cooling in determining structure 56 of the stratocumulus-topped boundary layer has been extensively studied and, presently, we are not 57 introducing any new mechanisms of turbulence generation and cloud liquid modulation. However, 58 we aim to elucidate and, to the extent possible, to decompose the effects of cloud top radiative and evaporative cooling. In the past, rather general terms such as "entrainment," "radiative cooling," and 60 "cloud holes" have been used with somewhat indefinite meanings and, in many cases, interchangeably 61 [e.g., 8,20–23]. 62 For non linear systems with a very large number of degrees of freedom (e.g., some of the present 63

For nonlinear systems with a very large number of degrees of freedom (e.g., some of the present
 simulations utilize computational grids with 20 billion grid cells), it is challenging to attribute outcomes
 to specific physical processes. Thus, a series of perturbation numerical experiments is carried out
 to observe the impact to different physical processes on the stratocumulus-topped boundary layer.
 Thus, the The present hypothesis is assessed by performing Sc LES without accounting for latent heat
 exchange. Additional LES are carried out to control for other physical parameters.
 The present study is enabled by recent improvements in high-resolution model fidelity and

computing power, which results in realistic validated simulations of the *entire* boundary layer [9,24]
 The observations, numerical model, and numerical experiments are described in §2. Simulations
 are based on the DYCOMS II RF01 nocturnal non-precipitating Sc case [25] because of the relatively
 simple Sc physics and the availability of extensive previous LES runs and validation data. Results
 are presented in §3 where the effects of cloud top radiative cooling are examined and support for the
 working hypothesis is discussed. Summary and conclusions are presented in §4.

## 76 2. Methodology

## 77 2.1. Observations



**Figure 1.** Radiance fields from six observed stratocumulus scenes during the ORACLES campaign. All images are nadir views of the 450 nm band aquired by the Airborne Multiangle SpectroPolarimetric Imager (AirMSPI) on 22 September 2016 off the coast of Namibia. In spite of variation in cloud cover, the characteristics stratocumulus structure composed of cellular blobs and thin spiderweb-like slits is visible.

The large-eddy simulations are informed by the Sc radiance structure captured in images from 78 the Airborne Multi-angle Spectro-Polarimetric Imager (AirMSPI) on NASA's Airborne ER-2 Platform 79 [26]. A sample of the AirMSPI images is shown in Figure 1. All images correspond to radiance fields 80 of nadir views of the 450 nm band observed during the ORACLES campaign on 22 September 2016 81 off the coast of Namibia. Further details about the instrument and campaign can be found in Xue 82 et al. [6]. The pixel resolution is 10 m and the scenes in Figure 1 are about  $11 \times 11$  km wide. In 83 spite of variations of cloud cover and intensity, the characteristic spiderweb structure is present in all 84 images. The spiderweb is not present in the corresponding coarser resolution (25-m pixels) retrieved 85

cloud properties images, suggesting that fine spacial resolution – less than about 10 m – is critical in
 discerning the spiderweb Sc structure.

88 The present study is enabled by recent improvements in high-resolution model fidelity and

<sup>89</sup> computing power, which results in realistic validated simulations of the entireboundary layer [9,24]

- . The model and numerical experiments are described in §2. Simulations are based on the DYCOMS
- 91 II RF01 nocturnal non-precipitating Sc case [25] because of the relatively simple Sc physics and the
- <sup>92</sup> availability of extensive previous LES runs and validation data. Results are presented in §3 where
- <sup>93</sup> the effects of cloud top radiative cooling are examined and support for the working hypothesis is
- <sup>94</sup> discussed. Summary and conclusions are presented in §4.

### 95 3. Methodology

- 96 The LES
- 97 2.1. Model

The LES model of Matheou & Chung [27] is used. The details of the model formulation, including 98 details of the model setup for the present stratocumulus cases, are described in Refs. [9,24]. The LES model numerically integrates the anelastic approximation of the Navier–Stokes equations [28] on an 100 f-plane using a doubly periodic domain in the horizontal directions. Fully-conservative fourth-order 101 (centered) finite-differences [29,30] and the Quadratic Upstream Interpolation for Convective 102 Kinematics (QUICK) scheme [31] are used for momentum and scalar advection, respectively. The 103 buoyancy adjusted stretched vortex subgrid-scale turbulence model [32–36] is used to account for the 104 effects of unresolved turbulence motions. The third-order Runge-Kutta method of Spalart et al. [37] is 105 used for time integration. All grids are uniform  $\Delta x = \Delta y = \Delta z$ . 106

The simulations are based on the DYCOMS II RF01 case [38], which is a non-precipitating 107 nearly stationary nocturnal stratocumulus-topped marine boundary layer. The flow is driven by 108 prescribed uniform surface latent and heat fluxes, the geostrophic wind,  $\mathbf{u}_{q}$ , and cloud-top radiative 109 and evaporative cooling. The case-specific parameterization of [38] is used for the net longwave 110 radiative flux, which results in strong cooling in a thin layer below the cloud top and small heating 111 near the cloud base. A uniform large-scale horizontal divergence D is used to represent the effects of 112 the large-scale subsidence on the evolution of the boundary layer. We refer to simulations that follow 113 the DYCOMS II RF01 case as "full physics" simulations. Validation of the "full physics" simulations 114 and further details of the present model configuration are described in Refs. [9,24]. 115

In all simulations, the mass of cloud liquid water condensate is diagnosed based on the local saturation water mixing ratio, using the values of pressure and temperature at the center of each grid cell. Thus, no partially saturated air is allowed in each grid cell. Moreover, microphysical effects are not taken into account, such as drizzle, droplet sedimentation and droplet inertial effects [e.g., 13].

A modified definition of buoyancy is used in the model to suppress latent heat exchange (including evaporative cooling)and the resulting BRI. Following Matheou & Teixeira [24], in the standard LES model, buoyancy is defined proportional to deviations of virtual potential temperature  $\theta_v$  from its instantaneous horizontal average  $\langle \tilde{\theta}_v \rangle$ ,

$$b' = g\rho_0 \frac{\tilde{\theta}_v - \langle \tilde{\theta}_v \rangle}{\theta_0},\tag{1}$$

where *g* is the acceleration of gravity,  $\rho_0(z)$  the basic-state density in the anelastic approximation, and  $\theta_0$  the basic-state potential temperature. The virtual potential temperature is

$$\theta_v = \theta \left[ 1 - \left( \frac{R_m}{R} - 1 \right) r - r_l \right], \tag{2}$$

where  $\theta$  is the potential temperature, r and  $r_l$  is the water vapor and liquid water mixing ratios, and  $R = 287.04 \text{ J kg}^{-1} \text{ K}^{-1}$  and  $R_m = 461.5 \text{ J kg}^{-1} \text{ K}^{-1}$  are the gas constants of dry air and water vapor. To suppress latent heat exchange,  $\theta_v$  is modified,

$$\theta_{v,\text{mod}} = \theta_l \left[ 1 - \left( \frac{R_m}{R} - 1 \right) r \right], \tag{3}$$

which is similar to its definition for air without condensate with  $\theta_l$  in the place of  $\theta$ .

The artificial modification of buoyancy suppresses not only evaporative cooling at the cloud top but everywhere in the cloud. Equation (3) is applied everywhere in the cloud to avoid introducing additional parameters and uncertainty. This is a limitation of the current methodology and its effects are analyzed in the next section.

Table 1 summarizes the simulations. Model runs A–E are the same as in Matheou & Teixeira 134 [24] and follow the same naming convection. Runs L3 and M1 are counterparts of runs E3 and A1, 135 respectively, without evaporative cooling. Based on observations, the grid resolution is  $\Delta x = 5$  m for 6 136 runs and  $\Delta x = 1.25$  m for runs A1 and M1. Simulations A3, B3, C3, E3, and L3 are initialized with 137 uniform fields and ran for four hours. Simulations A1 and M1 are initialized from run A1 of [24] at 138 139 §.3.1. Case A1 is merely a continuation of a "full physics" simulation for additional 10 minutes. Case 140 M1 is a short boundary layer evolution without latent heat exchange. 141

Case B3 controls for the effects of radiation on a "full physics" run. Cases E3 and L3 do not include
surface buoyancy fluxes, convection is only driven by cloud-top negative buoyancy production. Case
E3 includes both cloud-top radiative and evaporative cooling and Case L3 only radiative cooling.

#### 145 3. Results

#### 146 3.1. Liquid water path spatial structure

The working hypothesis is qualitatively evaluated by examining LWP and vertical planes of cloud liquid mixing ratio,  $r_l$ . The goal is to contrast simulations with respect to the presence of spiderweb structure in LWP fields and slits of clear air near the cloud top in  $r_l$  vertical planes. Figures 2–4 show LWP from all 5-m resolution runs. Figure 5 shows LWP for the high resolution cases,  $\Delta x = 1.25$  m.

In Figures -2 and 4, two time instances of LWP are shown at t = 2 and 4 h. Cloud cover and LWP significantly decrease with respect to time in the case without radiation (B3), see [24]. Thus,

additionally, LWP is shown in Figure 3 at t = 1 h. All panels in Figures 2–4 exhibit the characteristic

**Table 1.** Summary of the cases simulated. The grid spacing is denoted by  $\Delta x$ . For all runs the grid is homogeneous  $\Delta x = \Delta y = \Delta z$ . The number of horizontal and vertical grid points are  $N_x = N_y$  and  $N_z$ , respectively. "Wind" corresponds to forcing with the geostrophic wind  $\mathbf{u}_g$  or no wind (i.e., no mean surface shear), and *D* is the large-scale divergence. The case-specific parameterization of [38] is used for the net longwave radiative flux, except Case B3 which has null radiative flux at all model levels. Cases C3, L3 and M1 use a modified buoyancy variable (Eq. 3). Surface sensible and latent heat fluxes are denoted by "prescribed" when non-zero.

Run	$\Delta x$ (m)	$L_x$ (km)	N <sub>x</sub>	$N_{z}$	Wind	$\begin{array}{c} D\times 10^{-6} \\ (\mathrm{s}^{-1}) \end{array}$	Radiation	Buoyancy	Surface fluxes
A1	1.25	5.12	4096	1200	ug	3.75	Yes	multi-phase	prescribed
A3	5	5.12	1024	300	$\mathbf{u}_{g}$	3.75	Yes	multi-phase	prescribed
B3	5	5.12	1024	300	ug	3.75	No	multi-phase	prescribed
C3	5	5.12	1024	300	ug	3.75	Yes	modified	prescribed
E3	5	5.12	1024	300	0	0	Yes	multi-phase	$\overline{w\theta}_v = 0$
L3	5	5.12	1024	300	0	0	Yes	modified	$\overline{w\theta}_v = 0$
M1	1.25	5.12	4096	1200	$\mathbf{u}_g$	3.75	Yes	modified	prescribed



**Figure 2.** Liquid water path for the run with full physics, Case A3 (**a**, **b**), run without evaporative cooling, Case C3 (**c**, **d**), and the run without radiation, Case B3 (**e**, **f**). Left column panels (**a**, **c**, **e**) correspond to t = 2 h and right column panels (**b**, **d**, **f**) to t = 4 h.

Sc lumpy structure. However, the spiderweb structure is absent from the LWP plots of cases without cloud-top evaporative cooling (Figure 2 panels **c** and **d**, and Figure 4 panels **c** and **d**).

The contrast with respect to the spiderweb structure is higher in the Cases E3 and F3, which is driven only by cloud-top radiative cooling (Figure 4). As will be quantitatively discussed in the following sections, LWP spatial variability is still present in the cases without evaporative cooling and locations of nearly zero LWP can be observed in Figure 4 panels **c** and **d**. However, these locations of very low LWP are not thin and elongated as, in Figure 4 panels **a** and **b**, but are broader and a few

<sup>161</sup> circular cloud holes are present, similar to the observations in [5].



**Figure 3.** Liquid water path for a run without radiation, Case B3, at t = 1 h.



**Figure 4.** Liquid water path for the runs without surface fluxes (only driven by radiation). Top panels (**a b**) include the effects of evaporative cooling. Bottom panels (c, d) correspond to the run without evaporative cooling. Left column panels (a, c) correspond to t = 2 h and right column panels (b, d) to t = 4 h.

In spite of some evidence of spiderweb structure in Figure 3, the contrast is not as strong as in Figure 4. The lack of a homogenous and high LWP cloud, compared to other cases, may contribute to the reduced contrast.

To remove some of the effects of different boundary layer physics and evolution dynamics, LWP from Cases A1 and M1 is compared in Figure 5. Cases A1 and M1 correspond to a 10-minute evolution, about half the convective time scale, of the boundary layer with and without latent heat exchange.

7 of 16

The boundary layer convective time scale  $t_c = z_i (\overline{ww})^{-\frac{1}{2}} \approx 23$  minutes, where  $z_i = 846$  m is the boundary layer depth and  $\overline{ww}$  the depth-averaged vertical velocity turbulent flux. Thus, the large-scale motions remain well-correlated in Figure 5, since their time-correlation is expected to scale with  $t_c$ . Conversely, the spiderweb dissipates in the simulation without evaporative cooling. The signature of the spiderweb is at places visible in Figure 4 panel **b**, however these regions have higher LWP compared to Figure 4 panel **a**.

#### 174 3.2. Cloud liquid and LWP distributions

Figures 6 and 7 show cross sections of the cloud liquid water mixing ratio for Cases A1 and M1, respectively. The  $r_l$  cross sections correspond to vertical lines in the axes of Figure 5 passing through x = -1 km. The contrast between Figures 6 and 7 is stronger than the corresponding LWP of Figure 5 and provides a clearer indication of the effects of evaporative cooling near the cloud top.

As shown in previous modeling studies [16,39,40], the cloud-top slits because of evaporative cooling do not extend to the entire cloud depth, but are rather concentrated near the top (see also cloud-top boundary distributions in [9]).



**Figure 6.** Cloud liquid mixing ratio on a vertical plain at x = -1 km for Case A1. The elongated domain is partitioned into two panels. Only the cloudy region is shown. Evaporative cooling at the cloud top creates clear-air slits at the cloud top.



**Figure 5.** Adjustment of liquid water path structure to the lack of evaporative cooling. Panel (**a**) shows LWP from a full physics simulation, Case A1, at t = 2.16 h, and panel (**b**) shows Case M1 LWP. Both simulations were initialized from a full-physics LES at t = 2 h and ran for 10 minutes. Because of the relatively short time lapse from the common initial condition, the large-scale LWP structure is similar. The spiderweb LWP structure has shorter time scale and has dissipated in (**b**).



**Figure 7.** Cloud liquid mixing ratio on a vertical plain at x = -1 km for Case M1. The elongated domain is partitioned into two panels. Only the cloudy region is shown. The cloud-top slits are absent in Case M1 (c.f., Figures 6).



Figure 8. Liquid water path of the top-half of the cloud: a Case A1, and b Case M1.



Figure 9. Liquid water path of the bottom-half of the cloud: a Case A1, and b Case M1.

A key question from the observation of Figures 5 and 6 is why the relatively shallow cloud-top slits significantly affect the LWP structure of the entire cloud depth? Sc have most liquid content near the cloud top, thus any modification of the cloud top liquid distribution has a significant impact on the entire column.

Figure 8 shows LWP of the top-half of the cloud (z > 722 m) and Figure 9 shows LWP for the bottom-half of the cloud for Cases A1 and M1. In other words, the sum of panel-panels (**a**) of Figures 8 and 9 equals the LWP contours of Figure 5 panel (**a**). It can be observed that a large fraction of LWP is contributed from the cloud top region. Thus, the LWP structure, including the spiderweb, is because of variations of a relatively thin region near the cloud top.

Figure 9 shows the effects of suppressing latent heat exchange in the lower part of the cloud were evaporative cooling is not expected to prevalent(z < 722 m). Similar to Figures 5 and 8, in Case M1 LWP increases in the low LWP regions of Case A1. The lower part of the cloud has a more classical random turbulent structure without significant differences with respect to the presence of evaporative cooling (c.f., Figure 4).

Cloud liquid mixing ratio on a vertical plain at x = -1 km for Case A1. The elongated domain is partitioned into two panels. Only the cloudy region is shown. Evaporative cooling at the cloud top creates clear-air slits at the cloud top. Cloud liquid mixing ratio on a vertical plain at x = -1 km for Case M1. The elongated domain is partitioned into two panels. Only the cloudy region is shown. The cloud-top slits are absent in Case M1 (c.f., Figures 6).

Figure 10 quantifies the differences in LWP between Cases A1 and M1. For each (x, y) LWP 201 column, the pairs of LWP from of Cases A1 and M1 are recorded, i.e.,  $LWP(x, y)_{A1}-LWP(x, y)_{M1}$ . 202 Then joint (two-dimensional) histograms of the LWP pairs are constructed in Figure 10 for the full 203 column LWP, cloud top and cloud bottom LWP cloud-top LWP (z > 722 m), and cloud-bottom LWP 204 205 (z < 722 m). The histograms lay across the diagonal for perfect correlations. The histograms are not symmetric about the diagonal and spread towards the higher values of Case M1. This is consistent 206 to the observation in with the LWP contours of Figures 5–9 were LWP increases has higher values 207 in Case M1 at the locations of low LWP of compared to Case A1 - This tendency at the same (x, y)208 location. This change in LWP is observed in all LWP values but it is larger in low LWP regions since 209 the contours are broader for lower *x*-axis values in 10. The effects are also present in the lower half 210





**Figure 10.** Correlations between LWP distributions between Cases A1 and M1. The contours correspond to the normalized joint two-dimensional histograms of LWP. The legends correspond to probability per  $g^2 m^{-4} \times 10^3$ . The left panel corresponds to the full-column LWP, upper-half of the cloud is shown in the middle panel and the lower-half of the cloud is shown in the right panel.

of the cloud, however, the amount of cloud liquid is a small fraction in this region. We can conclude that even though the effects of modified buoyancy are present in the entire cloud, any impact of modification of the dynamics of the lower half of the cloud on the present conclusions is likely limited because of the small cloud liquid content and the more random nature of the liquid structure.

Figure 11 shows effects of suppressing latent heat exchange on the LWP Probability Density 215 Functions (PDF) for Case pairs E3–L3 and A1–M1. The LWP PDFs are compared at t = 4 h for 216 Cases E3–L3 and at t = 2.16 h for Cases A1–M1. The PDFs are essentially the normalized (integrate 217 to unity) histograms of LWP. In Cases A1 and M1 the mean LWP is approximately equal, thus the 218 x-axis corresponds to LWP. Cases E3 and L3 have different cloud evolutions (see also next section and 219 Figure 12), thus the *x*-axis is shifted by the location of the PDF mode. In both case pairs the suppression 220 of latent heat exchange affects the left "tail" of the LWP distribution by increasing the occurrences of 221 low LWP columns. Taking into account only the observed differences in the PDFs, we cannot conclude 222 that the change in the PDFs is because of the spiderweb structure. However, the LWP and cloud liquid 223 comparisons show that the spiderweb structure corresponds to low LWP cloud regions and the cases 224 without evaporative cooling show higher liquid water content in the spiderweb region. Therefore, it is 225 likely that the changes of the left PDF "tail" are mostly contributed by the spiderweb Sc structure. 226

#### 227 3.3. Entrainment rate and turbulence

Even though evaporative cooling occurs primarily at the cloud top, it affects the bulk boundary 228 layer dynamics. Comparison of time traces in Figure 12 of simulations driven only by radiative cooling 229 (Cases E3 and L3) shows significant changes to the vertically integrated turbulent kinetic energy 230 (TKE), mean LWP, and entrainment rate. In the cases driven only by radiative cooling, the cloud-top 231 evaporative cooling and the resulting buoyancy reversal instability (BRI) enhances the entrainment 232 rate. The entrainment rate is  $w_e = 0.0017 \text{ m s}^{-1}$  when evaporative cooling is suppressed (Case L3) 233 and nearly doubles to  $w_e = 0.003 \text{ m s}^{-1}$  in simulations with evaporative cooling. A comparison of the 234 Case E3 time traces with the standard DYCOMS II RF01 and other physics-perturbation experiments 235 is documented in [24]. 236

The reduced entrainment in Case L3 results in a cloud with more liquid water content and more radiative cooling at the cloud top. The increased radiative forcing results in more vigorous turbulence (higher TKE). Interestingly, the increase in TKE for Case L3 is not able to compensate for the lack of **BRI induced evaporative-cooling-generated (negative buoyancy)** motions in the entrainment process. The present results suggest that entrainment is mainly affected by the nature of the cloud-top motions

<sup>242</sup> rather than bulk boundary layer properties.





**Figure 11.** Probability Density Functions (PDF) of LWP for Cases E3 and L3 (left) and A1 and M1. The LWP (*x*-axis) in the left panel is shifted by the location of the mode of the PDF.

#### 243 4. Conclusions

<sup>244</sup> Observations (Figure 1 and Refs. [6,12]) show that stratocumulus clouds (Sc) have a distinctive <sup>245</sup> structure composed of a combination of lumpy cellular structures and thin elongated regions, <sup>246</sup> resembling canyons or slits. We refer to the elongated slits as "spiderweb" structure. Using very high <sup>247</sup> resolution ( $\Delta x = 1.25$  and 5 m) large-eddy simulations (LES) of a simple established case of a Sc deck, <sup>248</sup> we show that the spiderweb structure is caused by cloud-top evaporative cooling<del>and the resulting</del> <sup>249</sup> buoyancy reversal instability (BRI).

The effects of evaporative cooling are studied using simulations with a modified buoyancy 250 definition, which does not account for latent heat exchange. However, cloud liquid is diagnosed in the 251 model and used to calculate the parameterized radiative heating/cooling. The results are studied by 252 qualitatively contrasting simulations with and without cloud-top evaporative cooling with respect to 253 the presence of the spiderweb liquid water path (LWP) structure. Analysis of LWP of the entire cloud 254 depth, LWP of fractions of the cloudy column, and instantaneous cloud liquid vertical planes show 255 that cloud-top evaporative cooling generates relatively shallow slits near the cloud top. However, 256 because most of the liquid water mass is concentrated near the cloud top, these regions of clear air 257 have a large impact on the entire-column LWP. 258

Liquid water path of the bottom-half of the cloud: **a** Case A1, and **b** Case M1.

The cellular Sc structure is present in simulations without latent heat exchange, suggesting that the lumpy cloud structure and nearly circular cloud holes is because of are generated by boundary-layer-deep convective motions. These lumpy structures are present when the boundary layer convection is driven from the top by radiative cooling, from the surface, or by a combination of the two.

In the present LES, the spiderweb structure dissipated in about ten minutes (about half the boundary layer convective time scale), suggesting that the BRI-spiderweb process has motions related to cloud-top slits generated by evaporative cooling have short time scales.

The effects of the spiderweb structure on the LWP distributions are small, and discernible only on at the left "tails" of the distributions. This is likely because of the small area fraction of the spiderweb. Even though the spiderweb signature on the LWP distribution is negligible, the cloud-top evaporative cooling process significantly affects integral boundary layer quantities, such as the vertically integrated turbulent kinetic energy, mean liquid water path, and the entrainment rate.

Liquid water path of the top-half of the cloud: **a** Case A1, and **b** Case M1.



**Figure 12.** Time traces of liquid water path (left panel), vertically integrated turbulent kinetic energy (middle panel), and cloud base and cloud top height for Cases E3 and L3.

Funding: This research was funded by NSF-AGS-1916619, the Office of Naval Research, Marine Meteorology
 Program, the NASA MAP Program, the NOAA/CPO MAPP Program, and by the DOE Office of Biological and
 Environmental Research, Earth and Environmental System Modeling Program.

Acknowledgments: The AirMSPI data were obtained from the NASA Langley Research Center Atmospheric 279 Science Data Center. Computational resources supporting this work were provided by the NASA High-End 280 Computing (HEC) Program through the NASA Advanced Supercomputing (NAS) Division at Ames Research 281 Center. The research presented in this paper was supported by the systems, services, and capabilities provided 282 by the University of Connecticut High Performance Computing (HPC) facility. We acknowledge informative 283 discussions with Dr. Jerone Povner, a spider expert. Part of this research was carried out at the Jet Propulsion 284 Laboratory, California Institute of Technology, under a contract with the National Aeronautics and Space 285 Administration. 286

**287** Conflicts of Interest: The authors declare no conflict of interest.

#### 288 References

- Hartmann, D.L.; Ockert-Bell, M.E.; Michelsen, M.L. The effect of cloud type on Earth's energy balance:
   Global analysis. J. Climate 1992, 5, 1281–1304.
- Bretherton, C.S. Convection in stratocumulus-topped atmospheric boundary layers. In *The Physics and Parameterization of Moist Atmospheric Convection*; Springer, 1997; pp. 127–142.
- 3. Stevens, B. Atmospheric moist convection. Annu. Rev. Earth Planet. Sci. 2005, 33, 605–643.
- 4. Wood, R. Stratocumulus clouds. *Mon. Weather Rev.* **2012**, 140, 2373–2423.
- Haman, K.E. Simple approach to dynamics of entrainment interface layers and cloud holes in stratocumulus
   clouds. *Q. J. R. Meteorol. Soc.* 2009, 135, 93–100.
- 297 6. Xu, F.; van Harten, G.; Diner, D.J.; Davis, A.B.; Seidel, F.C.; Rheingans, B.; Tosca, M.; Alexandrov, M.D.;
  298 Cairns, B.; Ferrare, R.A.; others. Coupled Retrieval of Liquid Water Cloud and Above-Cloud Aerosol
- Properties Using the Airborne Multiangle SpectroPolarimetric Imager (AirMSPI). J. Geophys. Res. 2018, 123, 3175–3204.
- 7. Yamaguchi, T.; Randall, D.A. Cooling of entrained parcels in a large-eddy simulation. J. Atmos. Sci. 2012,
   69, 1118–1136.
- Yamaguchi, T.; Feingold, G. On the size distribution of cloud holes in stratocumulus and their relationship
   to cloud-top entrainment. *Geophys. Res. Lett.* 2013, 40, 2450–2454.
- 9. Matheou, G. Turbulence structure in a stratocumulus cloud. *Atmosphere* **2018**, *9*, 392.
- Nicholls, S. The structure of radiatively driven convection in stratocumulus. *Q. J. R. Meteorol. Soc.* 1989, 115, 487–511.
- Gerber, H.; Frick, G.; Malinowski, S.P.; Brenguier, J.L.; Burnet, F. Holes and entrainment in stratocumulus.
   *J. Atmos. Sci.* 2005, *62*, 443–459.

- Haman, K.E.; Malinowski, S.P.; Kurowski, M.J.; Gerber, H.; Brenguier, J.L. Small scale mixing processes at
  the top of a marine stratocumulus–A case study. *Q. J. R. Meteorol. Soc.* 2007, 133, 213–226.
- Karpińska, K.; Bodenschatz, J.F.; Malinowski, S.P.; Nowak, J.L.; Risius, S.; Schmeissner, T.; Shaw, R.A.;
   Siebert, H.; Xi, H.; Xu, H.; others. Turbulence-induced cloud voids: observation and interpretation. *Atmos. Chem. Phys.* 2019, *19*, 4991–5003.
- Schmidt, H.; Schumann, U. Coherent structure of the convective boundary layer derived from large-eddy
   simulations. J. Fluid Mech. 1989, 200, 511–562.
- Sullivan, P.G.; Patton, E.G. The effect of mesh resolution on convective boundary layer statistics and
   structures generated by large-eddy simulation. *J. Atmos. Sci.* 2011, *68*, 2395–2415.
- 16. Mellado, J.P. Cloud-top entrainment in stratocumulus clouds. Annu. Rev. Fluid Mech. 2017, 49, 145–169.
- Mellado, J.P.; Stevens, B.; Schmidt, H. Wind shear and buoyancy reversal at the top of stratocumulus. J.
   *Atmos. Sci.* 2014, 71, 1040–1057.
- Bavis, A.B.; Marshak, A.; Gerber, H.; Wiscombe, W.J. Horizontal structure of marine boundary layer clouds
   from centimeter to kilometer scales. *J. Geophys. Res.-Atmos.* 1999, 104, 6123–6144.
- Ma, Y.F.; Malinowski, S.P.; Karpińska, K.; Gerber, H.E.; Kumala, W. Scaling analysis of temperature and
   liquid water content in the marine boundary layer clouds during POST. J. Atmos. Sci. 2017, 74, 4075–4092.
- vanZanten, M.C.; Duynkerke, P.G. Radiative and evaporative cooling in the entrainment zone of stratocumulus The role of longwave radiative cooling above cloud top. *Boundary-Layer Meteorol.* 2002, 102, 253–280.
- Malinowski, S.P.; Andrejczuk, M.; Grabowski, W.W.; Korczyk, P.; Kowalewski, T.A.; Smolarkiewicz, P.K.
   Laboratory and modeling studies of cloud–clear air interfacial mixing: anisotropy of small-scale turbulence
   due to evaporative cooling. *New J. Phys.* 2008, *10*, 075020.
- Petters, J.L.; Harrington, J.Y.; Clothiaux, E.E. Radiative–dynamical feedbacks in low liquid water path
   stratiform clouds. *J. Atmos. Sci.* 2012, 69, 1498–1512.
- de Lozar, A.; Mellado, J.P. Evaporative cooling amplification of the entrainment velocity in radiatively
   driven stratocumulus. *Geophys. Res. Lett.* 2015, 42, 7223–7229.
- Matheou, G.; Teixeira, J. Sensitivity to physical and numerical aspects of large-eddy simulation of
   stratocumulus. *Mon. Weather Rev.* 2019, 147, 2621–2639.
- Stevens, B.; Lenschow, D.H.; Vali, G.; Gerber, H.; Bandy, A.; Blomquist, B.; Brenguier, J.; Bretherton, C.;
   Burnet, F.; Campos, T.; others. Dynamics and chemistry of marine stratocumulus DYCOMS-II. *Bull. Amer. Meteor. Soc.* 2003, *84*, 579–593.
- 26. Diner, D.J.; Xu, F.; Garay, M.J.; Martonchik, J.V.; Rheingans, B.E.; Geier, S.; Davis, A.; Hancock, B.R.;
  Jovanovic, V.M.; Bull, M.A.; Capraro, K.; Chipman, R.A.; McClain, S.C. The Airborne Multiangle
  SpectroPolarimetric Imager (AirMSPI): a new tool for aerosol and cloud remote sensing. *Atmospheric Measurement Techniques* 2013, 6, 2007.
- Matheou, G.; Chung, D. Large-eddy simulation of stratified turbulence. Part II: Application of the
   stretched-vortex model to the atmospheric boundary layer. J. Atmos. Sci. 2014, 71, 4439–4460.
- <sup>347</sup> 28. Ogura, Y.; Phillips, N.A. Scale analysis of deep and shallow convection in the atmosphere. *J. Atmos. Sci.*<sup>348</sup> 1962, 19, 173–179.
- Morinishi, Y.; Lund, T.S.; Vasilyev, O.V.; Moin, P. Fully Conservative Higher Order Finite Difference
   Schemes for Incompressible Flow. *J. Comput. Phys.* 1998, 143, 90–124.
- 351 30. Matheou, G.; Dimotakis, P.E. Scalar excursions in large-eddy simulations. J. Comput. Phys. 2016, 327, 97–120.
- 353 31. Leonard, B.P. A stable and accurate convective modelling procedure based on quadratic upstream
   interpolation. *Comput. Methods in Appl. Mech. Eng.* 1979, 19, 59–98.
- 355 32. Lundgren, T.S. Strained spiral vortex model for turbulent fine structure. *Phys. Fluids* 1982, 25, 2193–2203.
- 33. Misra, A.; Pullin, D.I. A vortex-based subgrid stress model for large-eddy simulation. *Phys. Fluids* 1997, 9, 2443–2454.
- 358 34. Pullin, D.I. A vortex-based model for the subgrid flux of a passive scalar. *Phys. Fluids* 2000, *12*, 2311–2316.
- 359 35. Voelkl, T.; Pullin, D.I.; Chan, D.C. A physical-space version of the stretched-vortex subgrid-stress model
   for large-eddy simulation. *Phys. Fluids* 2000, *12*, 1810–1825.
- 36. Chung, D.; Matheou, G. Large-Eddy Simulation of Stratified Turbulence. Part I: A Vortex-Based
   Subgrid-Scale Model. J. Atmos. Sci. 2014, 71, 1863–1879.

- 363 37. Spalart, P.R.; Moser, R.D.; Rogers, M.M. Spectral methods for the Navier–Stokes equations with one infinite
   and two periodic directions. J. Comput. Phys. 1991, 96, 297–324.
- 365 38. Stevens, B.; Moeng, C.H.; Ackerman, A.S.; Bretherton, C.S.; Chlond, A.; De Roode, S.; Edwards, J.; Golaz,
- J.C.; Jiang, H.L.; Khairoutdinov, M.; Kirkpatrick, M.P.; Lewellen, D.C.; Lock, A.; Muller, F.; Stevens,
- D.E.; Whelan, E.; Zhu, P. Evaluation of large-eddy simulations via observations of nocturnal marine
   stratocumulus. *Mon. Weather Rev.* 2005, 133, 1443–1462.
- 369 39. Mellado, J.P. The evaporatively driven cloud-top mixing layer. J. Fluid. Mech. 2010, 660, 5–36.
- 40. de Lozar, A.; Mellado, J.P. Mixing driven by radiative and evaporative cooling at the stratocumulus top. J.
   Atmos. Sci. 2015, 72, 4681–4700.

<sup>372</sup> © 2020 by the authors. Submitted to *Atmosphere* for possible open access publication <sup>373</sup> under the terms and conditions of the Creative Commons Attribution (CC BY) license <sup>374</sup> (http://creativecommons.org/licenses/by/4.0/).