



Article

Impacts of National Highway G214 on Vegetation in the Source Area of Yellow and Yangtze Rivers on the Southern Qinghai Plateau, West China

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Abstract: Engineering corridors on the Qinghai–Tibet Plateau have substantially modified the regional ecosystem functions and environment, resulting in changes in the alpine ecosystem. In addition, the building and operation of these engineering corridors have led to rapid permafrost degradation, which in turn has impacted local vegetation along these corridors. This study investigated vegetation changes and their driving factors by the methods of coefficient of variation, correlation analysis, and GeoDetector in a 30 km wide buffer zone at each side along the National Highway G214 (G214) at the northern and southern flanks of the Bayan Har Mountains in part of the source area of the Yellow and Yangtze rivers on the southern Qinghai Plateau, West China. The following results were obtained: (1) The Normalized Difference Vegetation Index in Growing Season (NDVI_{gs}) rose slightly in 2010–2019, with an average annual change rate of 0.006/a. Patterns of NDVI_{gs} along the G214 exhibited “low at the northern flank and high at the southern flank of the Bayan Har Mountains”. (2) Spatially, average NDVI_{gs} increased from the first buffer zone at the distance of 0–10 km from the highway centerline to the second buffer zone at 20–30 km perpendicularly away from the G214. Furthermore, the first buffer zone had the lowest coefficient of variation, possibly due to a low vegetation recovery as a result of the greatest influence of the G214 on NDVI_{gs} at 0–10 km. (3) Furthermore, annual precipitation (AP) was the dominant factor for significantly ($p < 0.01$) and positively influencing the variations in NDVI_{gs} ($R = 0.75$, $p < 0.01$). Additionally, NDVI_{gs} was more strongly influenced by the two combined factors than any single one, with the highest q -value (0.74) for the interactive influences of AP and annual average air temperature (AAAT) and followed by that of the AP and mean annual ground temperature (MAGT) at the depth of zero annual amplitude (15 m). Evidently, the construction and operation of the G214 have directly and indirectly affected vegetation through changing environmental variables, with significant impacts on NDVI_{gs} extended at least 20 km outwards from the highway. This study helps better understand the environmental impacts along the engineering corridors in elevational permafrost regions at mid and low latitudes and their management.



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1. Introduction

Climate warming and human disturbances can directly affect the thermal state of permafrost or indirectly impact permafrost through changes in vegetation, snow cover, land use and cover change, and in many other hydroclimatic aspects [1,2]. A gradual intensification of climate warming might lead to increased droughts and floods, reduced vegetation cover and biodiversity, and warming and thawing of permafrost, with profound impacts on natural ecosystems and human societies [3–7]. Found between the atmosphere and soil or surface waters [8], vegetation is important for global mass, energy, and carbon cycles [9]. With rapidly advancing technology in remote sensing, the accumulation of abundant multi-temporal, multi-band, and higher-resolution satellite remote sensing data [10,11] has provided powerful tools and data support for the monitoring of large-scale vegetation dynamics [12–14]. This has facilitated the study of longtime series of vegetation change, which plays a crucial role in our responses and adaptation to rapidly changing climate, prudent management of the ecological environment, and future harmonious coexistence of humans and nature [15–17].

As a result of climate change and anthropogenic pressures, vegetation cover has changed substantially at global and regional scales [18–20]. Using the Normalized Difference Vegetation Index (NDVI), numerous studies have documented the rapidness and complexity of vegetative responses to climate change [21–23], and as a handy and effective tool, NDVI could quantitatively reflect the dynamics of vegetation cover [24–28]. Globally, annual precipitation (AP) and air temperature are key climate controllers and important environmental variables for NDVI dynamics and their changes [29,30]. Regionally, Growth Season NDVI (NDVI_{gs}) is highly correlated with growth season precipitation, and the NDVI tends to increase significantly [31,32]. For example, in alpine regions, changes in NDVI are also influenced by ground thermal regimes, snow cover, and altitudinal climate zonation [33–35]. In plains, hills, mountains, and grasslands, changes in NDVI are inevitably and extensively affected by the magnitude and intensity of human activities (e.g., extensive urbanization, massive construction projects, intensive cattle grazing, and rapid population growth) [36–38]. Overall, these studies have improved our understanding of the responsive shifts of vegetation to hydroclimatic and environmental changes at different spatiotemporal scales.

The responses of vegetation to environmental changes are extremely complex [39,40] and highly variable [41,42]. Related studies have shown that the effects of precipitation, air temperature, permafrost conditions, and human activities on vegetation are spatially heterogeneous [43–45]. Meanwhile, there are certain lags in related vegetation changes in response to shifts in hydroclimate patterns [46,47]. Furthermore, many scholars have analyzed the correlations between NDVI and its drivers, using correlation, regression trends, and geographically weighted regression, e.g., [48–50]. However, analyses with a single method may not be adequate for reflecting the changes or evolution of the interactive hydroclimatic and environmental processes or trends. Therefore, in this study, the methods of spatial autocorrelation, coefficient of variation (CV), and Geographical Detector (GeoDetector) have been adopted for more systematic and more in-depth investigation of the influences of single and multiple drivers on changes in NDVI in alpine permafrost regions.

Since the 21st century, with rapidly developing highway networks, engineering activities have increased. The construction, operation, and maintenance activities of highways have affected vegetation in permafrost regions [51,52]. Despite an overall increased trend in alpine vegetation, the deterioration of alpine vegetation was found in the Qinghai–Tibet Engineering Corridor [53]. In addition, another study showed that the impacts of the

Qinghai–Tibet Engineering Corridor were extended to about 1.5 km on each side [54]. However, there is insufficient information regarding the dominant influencing factors on alpine vegetation. Additionally, there is a paucity of information regarding elevational permafrost and its degradation along the highway and whether this is associated with alpine vegetation and its changes. The impacts of climate change on the Qinghai–Tibet Plateau (QTP), a region of key ecological significance, have been well-reported, with only little attention paid to the smaller corridor regions along major lifelines, such as the Qinghai–Tibet highway and railway [55,56].

The source area of the Yellow and Yangtze rivers (SAYYR) is on the southern Qinghai Plateau (SQP), West China. Its unique geological and hydroclimatic environment has long nurtured relatively stable alpine vegetation ecosystems extremely sensitive and responsive to hydroclimatic and environmental changes [57–59]. Under the dual action of climate change and human disturbances, alpine ecology and the periglacial environment along the National Highway 214 (G214) in the SAYYR have changed substantially, and vegetation cover and land use have also been modified accordingly [58,60]. The G214 is an important part of the Qinghai–Kang (West Sichuan Province) Engineering Corridor, and vegetation and changing permafrost have irreplaceable roles in maintaining basin-wide and regional ecological security and sustainable development of human society. In this study, along the G214 in the SAYYR, the spatiotemporal patterns of alpine vegetation have been characterized; the impacts of the G214 on the NDVI_{gs} have been evaluated, and their environmental drivers have been identified. This study could help better understand the impacts of highway and hydroclimate changes on alpine vegetation in elevational permafrost regions and, accordingly, planning and management.

2. Materials and Methods

2.1. Study Area

In the SAYYR on the SQP, permafrost interacts intricately and intensively with alpine hydroclimate and ecological environment. The continental high-plateau or alpine climate of the study area is characterized by long and severely cold winters but short and mild summers, large diurnal temperature differences, small AP, major rainfall in the warm season, and distinct dry and wet seasons [54]. During the period of 2010–2019, average annual air temperature (AAAT) was -5.7°C , and AP was 326.2–618.9 mm in the study area. Precipitation varied significantly from season to season. The temperature and precipitation decreased from southeast to northwest in the SAYYR [26].

Increasing human activities have resulted in permafrost degradation in the SAYYR [58,60,61]. The construction and operation of the G214 have directly or indirectly affected alpine vegetation and changed soil hydrology and nutrients along the engineering corridor, resulting in warming and/or thawing of the underlying active layer and permafrost and modified or altered alpine ecosystems in the right-of-way along the engineering corridor [55,62–65].

The chosen study area is a segment of the G214 from the Changshitoushan Mountain Pass to Qingshui'he Riverside crossing the northern and southern flanks of the Bayan Har Mountains; it is a 60 km wide buffer zone (a 30 km buffer zone at each side) with elevation at 3761–5767 m a. s. l. (Figure 1). This 30 km buffer zone was divided into three subzones based on the resolution of the data we used. The southern flank of the Bayan Har Mountains in the Yangtze River Basin is characteristic of a semi-humid and cold climate, while the northern flank in the Yellow River Basin is characteristic of a semi-arid and cold climate, with strong influences from the summer monsoons on the southern flank and the upper winter westerlies on the northern flank [65]. They result in more precipitation on the windward southern flank and less precipitation on the leeward northern flank of the Bayan Har Mountains, significantly impacting the spatial variations in vegetation features [60]. Sedge-dominated alpine (marsh) meadows prevail in the study area.

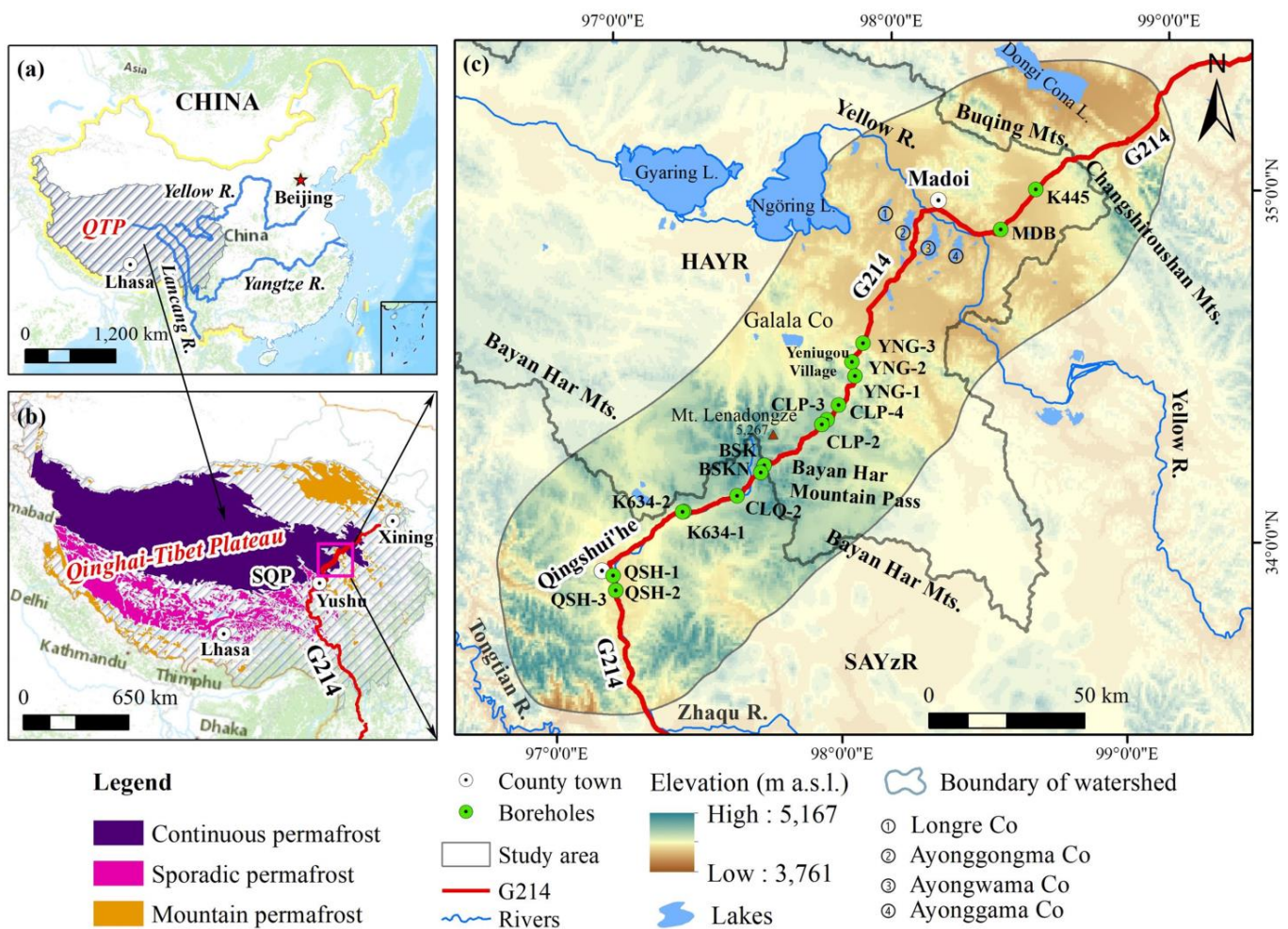


Figure 1. Location map of the study area on northeastern Qinghai–Tibet Plateau in West China. Notes: (a) Relative location of the Qinghai–Tibet Plateau in China; (b) National Highway 214 (G214) (red line) across the southern Qinghai Plateau (SQP) and the key study segment of G214 across the Bayan Har Mountains (pink box). Permafrost distribution on the upper panel is derived from a global permafrost zonation index [66]; (c) Permafrost boreholes (green points), county towns (white points), rivers and lakes (blue line and areas), and study area (highlighted black-bordered box, a 60 km wide buffer zone mapped along each side of the G214).

2.2. Data Sources

Growth season (from May to September) NDVI ($NDVI_{gs}$) data in the study area for each year from 2010 to 2019 were downloaded from the Data Center for Resource and Environmental Sciences, Chinese Academy of Sciences (URL: <https://www.resdc.cn/data.aspx?DATAID=342>, accessed on 5 March 2023), with a spatial resolution of 1.0 km. Elevation data were downloaded from the Japan Aerospace Exploration Agency (JAXA URL: <https://www.eorc.jaxa.jp/ALOS/en/aw3d30/data/index.htm>, accessed on 9 March 2023); the slope data were generated by DEM, with a spatial resolution of 30 m (released in March 2017). Meteorological data were collected from the annual spatial interpolation dataset of Chinese meteorological elements (URL: <http://www.resdc.cn>, accessed on 5 March 2023), and in this paper, a 1.0 km spatial resolution was obtained for the datasets of AP and AAAT. Snow cover data were acquired from the day-by-day cloudless MODIS Normalized Difference Snow Index (NDSI) product for Asian high mountains, with a spatial resolution of 0.5 km (URL: <https://cstr.cn/18406.11.Cryos.tpcd.272836>, accessed on 9 March 2023) [67] (Table 1). The data of mean annual ground temperatures (MAGT) at the depth of zero annual amplitude (15 m) were manually obtained from boreholes in

July or August in 2010, 2013, 2016, and 2019 along the G214. All ground temperatures are calculated from the electrical resistivity measured in the field using the Fluke meter (accuracy of 0.05%, model: 289, Fluke Corporation, Everett, Washington, DC, USA). More detailed information about boreholes can be found in Luo et al. [54].

Table 1. List of the data in this study.

Category	Data	Year	Spatial Resolution
Vegetation index	NDVI	2010, 2013, 2016, 2019	1 km
Snow index	NDSI	2010, 2013, 2016, 2019	0.5 km
Meteorological factors	AP and AAAT	2010, 2013, 2016, 2019	1 km
Terrain factors	Elevation and slope	2017	30 m
Permafrost	MAGT	2010, 2013, 2016, 2019	N/A

Note: NDVI, Normalized Difference Vegetation Index; NDSI, Normalized Difference Snow Index; AP, annual precipitation; AAAT, average annual air temperature; MAGT, mean annual ground temperature, which was obtained at the depth of zero annual amplitude at 15 m in depth in July or August from the 16 boreholes along the G214 in the study area.

2.3. Methods

2.3.1. Pearson Correlation Analysis

Pearson correlation method can be applied to detect the correlation between each factor. In this case, positive and negative correlation coefficients represent the effect of one variable on another variable and are used to calculate the correlation between different factors [68,69]. The calculation equation is shown as follows.

$$R_{xy,z} = \frac{R_{xy} - R_{xz}R_{yz}}{\sqrt{(1 - R_{xz})^2} \sqrt{(1 - R_{yz})^2}} \quad (1)$$

where x , y , and z represent three factors. R_{xy} , R_{xz} , and R_{yz} represent the correlations between the pairs of factors x , y , and z , respectively. $R_{xy,z}$ represents the correlation between the variables xy and z after excluding the interference associated with factor z .

2.3.2. Coefficient of Variation (CV)

The CV can reflect the fluctuating characteristics of the time series data [70]. This study aims to investigate the large variation of NDVI in response to changes in regional climate (i.e., AAAT and AP). The calculation equation is shown as follows.

$$CV = \frac{1}{\overline{NDVI}} \sqrt{\frac{1}{n-1} \sum_{i=1}^n (NDVI_i - \overline{NDVI})^2} \quad (2)$$

where CV is the coefficient of variation of $NDVI$ values; $NDVI_i$ is the $NDVI$ in the year i , and; $\overline{NDVI}$ is the multi-year average of annual $NDVI_i$ of the study region from 2010 to 2019.

The stability of CV was classified into five categories according to the natural interruption point method [71]: Minimum (0.02, 0.12), Low (0.13, 0.30), Medium (0.31, 0.70), High (0.71, 1.54), and Maximum fluctuations (1.55, 4.12). A large value of CV means greater $NDVI$ variation and poorer stability; the smaller the value of CV , the higher the $NDVI$ stability.

2.3.3. Geographical Detector

The Geographical Detector (GeoDetector) is used to detect spatial stratified heterogeneity, and thus to reveal the drivers behind it. These include factor, interaction, risk, and ecological detectors [72]. The factor detector is used to detect the spatial differentiation of the dependent factor Y and how much the changes in the independent factor X explain

the spatial differentiation of the attribute Y [73]. The performance of the GeoDetector is generally valued by the function of q-statistic with its range of (0 to 1). A value of 1 indicates the factor completely controls the spatial patterns of the geographical phenomenon, and a value of 0 indicates no relationship between the dependent and independent factors.

$$q = 1 - \frac{\sum_{h=1}^L N_h \sigma_h^2}{N \sigma^2} \quad (3)$$

where $h = 1, L$ is the strata of variable Y or factor X (that is, classification or partitioning), and N_h and N are the numbers of units in the layer h and the whole region, respectively. The variables σ_h^2 and σ^2 are variances of the Y values of the h^{th} layer and the entire area, respectively.

The interaction detector measures the value of q-statistic for the two explanatory variables after their interaction $q(X_1 \cap X_2)$ (Table 2). The interaction is used to investigate whether there is interaction between two factors and the variation in the degree of influence on the attribute Y under the interaction [74]. Ecological probing is used to compare whether there is a significant difference in the effect of factor–factor interactions on the spatial differentiation of Y [45,75].

Table 2. Types of interaction between two covariates.

Criteria of Interval	Interaction
$q(X_1 \cap X_2) < \min[q(X_1), q(X_2)]$	Nonlinear weakening
$\min[q(X_1), q(X_2)] < q(X_1 \cap X_2) < \max[q(X_1), q(X_2)]$	Single-factor nonlinear weakening
$q(X_1 \cap X_2) > \max[q(X_1), q(X_2)]$	Dual-factor enhancement
$q(X_1 \cap X_2) = q(X_1) + q(X_2)$	Independence
$q(X_1 \cap X_2) > q(X_1) + q(X_2)$	Nonlinear enhancement

3. Results

3.1. Spatiotemporal Patterns of NDVI_{gs} along the G214

The results of interannual variation analysis showed that regional NDVI_{gs} fluctuated between 0.53 to 0.58 in the past decade (2010–2019) and increased slightly with a linearly regressed increase rate of 0.006/a. Compared with a trend of slight decline in NDVI_{gs} at the northern flank of the Bayan Har Mountains in 2010–2019, NDVI_{gs} at the southern flank increased instead. Spatial analysis exhibited a northward decreasing pattern of NDVI_{gs} and “low at the northern flank and high at the southern flank of the Bayan Har Mountains” (Figure 2a). To facilitate a comparison of NDVI_{gs} patterns, the spacing method was applied to divide the NDVI_{gs} into five classes, namely Lower (0.10, 0.20), Low (0.20, 0.40), Medium (0.40, 0.60), High (0.60, 0.80), and Higher (0.80, 0.90). During the period of 2010 to 2019, the areal extent of the Higher NDVI_{gs} class increased by 14.9%; that of the High class increased by 0.3%; that of the medium class increased by 1.6%; that of the Low class decreased by 5.7%; and that of the Lower class decreased by 11.1% (Figure 2b). This shows that the NDVI_{gs} class in the study area changed significantly, and the NDVI_{gs} change in the south flank showed an increasing trend compared with the north flank, while the NDVI_{gs} change trend in the north flank had greater fluctuations but showed an overall increasing trend. In terms of G214 highway, the NDVI_{gs} change characteristics showed a decreasing feature from south to north along G214 in the Bayan Har Mountains. Therefore, there were more areas with the High and the Higher NDVI_{gs} in the southern flank, and most areas of the Low and the Lower NDVI_{gs} were in the northern flank. Overall, NDVI_{gs} displayed a slightly rising trend in the study area from 2010 to 2019. Spatially, a northward decline was found in NDVI_{gs} across the study area, characterized by “low at the northern flank and high at the southern flank”.

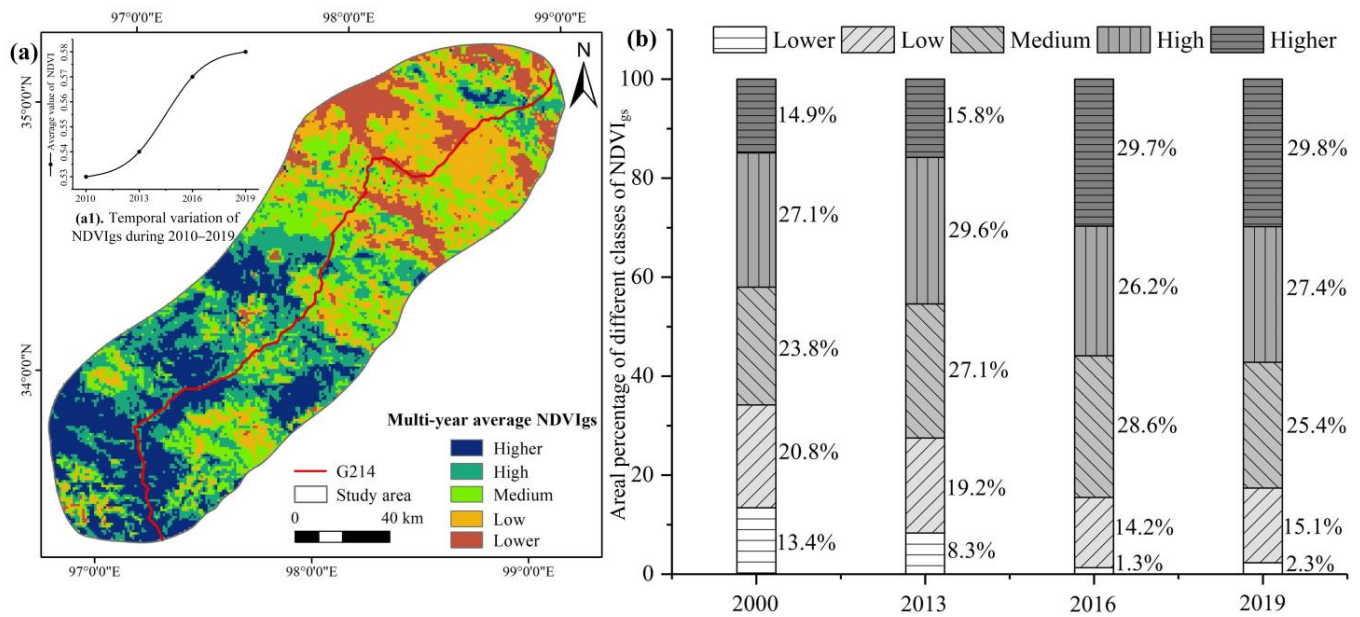


Figure 2. Distributive features of five classes of growth season Normalized Difference Vegetation Index (NDVI_{gs}) along the National Highway G214 from the Changshitoushan Mountain Pass to Qingshui'he Riverside at the northern and southern flanks of the Bayan Har Mountains, southern Qinghai, China. Notes: (a) Spatiotemporal variation of NDVI_{gs}, averaged by NDVI_{gs} during 2010–2019; (b) Areal percentage of different classes of NDVI_{gs} in the study area in 2010, 2013, 2016, and 2019.

3.2. NDVI_{gs} Variations in the Three Buffer Zones along the G214

The average NDVI_{gs} increased gradually with the increasing perpendicular distance away from the G214, i.e., 0–10, 10–20, and 20–30 km symmetrically with the highway centerline as the axis. Among them, the average NDVI_{gs} was greatest in the buffer zone of 20–30 km with a value of 0.57 and lowest at the zone of 0–10 km with a value of 0.50 (Figure 3). Average NDVI_{gs} over the three buffer zones exhibited a similar overall increasing trend from 2010 to 2019. In addition, there was a slight increasing trend from 2010 to 2013, with the same pattern of increase or decrease in each buffer zone. However, NDVI_{gs} gradually decreased in each buffer zone from 2013 to 2016. Then, NDVI_{gs} increased steadily from 2016 to 2019, reaching the maxima in all three buffer zones in 2019. In summary, the NDVI_{gs} of each buffer zone decreased from 2013 to 2016, with fluctuations, and gradually increased with some fluctuations from 2016 to 2019. This indicates the gradually ameliorating vegetation of each buffer zone over the study period.

The average of coefficients of variation (CV) was 0.23 from 2010 to 2019 in the three buffer zones. To study the influences of the G214 on the stability of NDVI_{gs}, the engineering corridor along the G214 was divided into equally spaced buffer zones at an interval of 10 km. They are: the first buffer zone at distances of 0–10 km perpendicularly away from the highway centerline, the second at 10–20 km, and the third at 20–30 km, symmetrically at both sides of the G214, to analyze the variability of NDVI_{gs} (Figure 4). The variability was relatively low in the first outgoing (0–10 km) buffer zone, with minimum and low fluctuation areas accounting for 96.3% in the three buffer zones, while the areas of the medium, high, and maximum fluctuation account for only 3.7%. Therefore, from 2010 to 2019, the NDVI_{gs} in the 0–10 km buffer zone had not changed much and had the best stability. In the second outgoing (10–20 km perpendicularly away from the G214) buffer zone, the magnitude of changes in NDVI_{gs} was relatively larger than that of the first (0–10 km) buffer zone, and the areas with medium and greater fluctuations in NDVI_{gs} account for 3.9% in all three buffer zones. Among them, the patchy areas with the maximum fluctuations account for only 0.1% of the total area. In the third buffer zone (20–30 km

perpendicularly away from the G214), the area with medium fluctuations was 3.4%, and that of high and maximum fluctuations was 1.4%. This suggests the best stability of CV at the first buffer zone (0–10 km) followed by the second (10–20 km), while the third buffer zone (20–30 km) had the least stability.

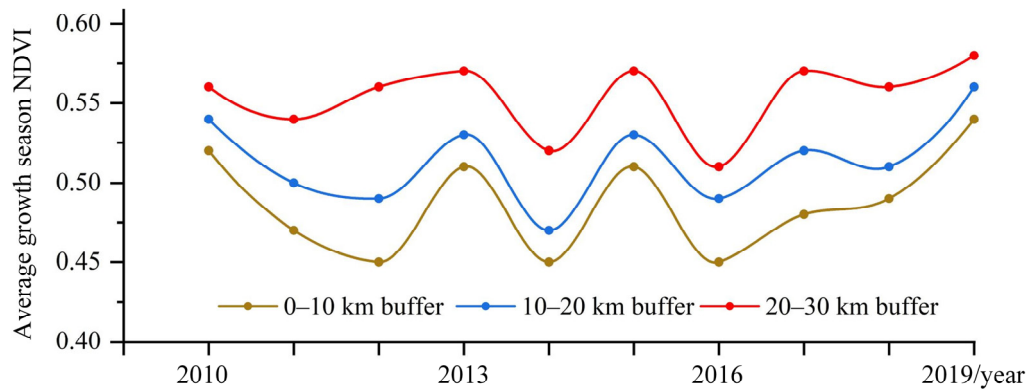


Figure 3. Changes in average growth season Normalized Difference Vegetation Index ($NDVI_{gs}$) in three buffer zones along the National Highway G214 from the Changshitoushan Mountain Pass to the Qingshui'he Riverside at the northern and southern flanks of the Bayan Har Mountains on the southern Qinghai Plateau, West China, from 2010 to 2019.

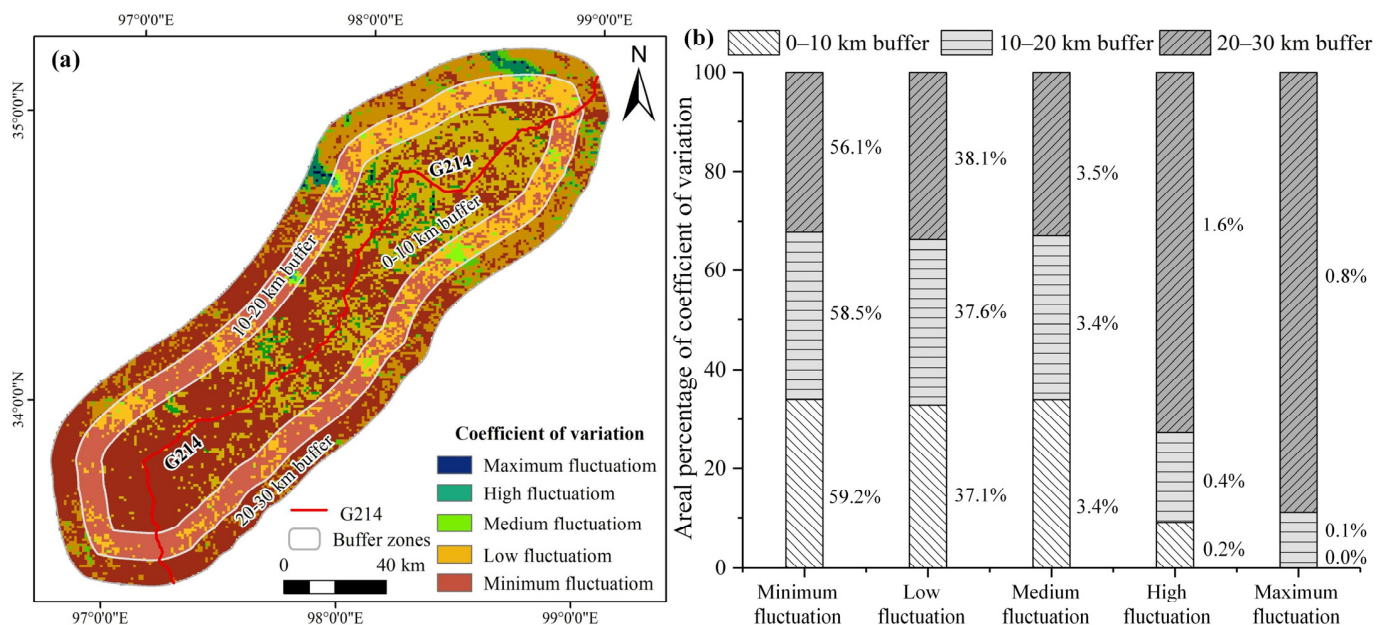


Figure 4. Coefficients of variation (CV) of growth season Normalized Difference Vegetation Index ($NDVI_{gs}$) within the three buffer zones along National Highway G214 from Changshitoushan Mountain Pass to Qingshui'he Riverside at the northern and southern flanks of the Bayan Har Mountains, southern Qinghai, China, from 2010 to 2019. Notes: (a) Spatial stability distribution of coefficients of variation within the 0–30 km buffers; (b) Areal percentage of coefficient of variation in the three buffer zones.

3.3. Correlation and Sensitivity Analysis of Drivers for $NDVI_{gs}$

The GeoDetector was applied to examine the environmental drivers for $NDVI_{gs}$. In this study, the q -values of different factors were calculated by the GeoDetector (Table 3). In Table 3, AP had the highest q -value (0.52), followed by AAAT (0.45) and elevation (0.31). In addition, Pearson correlation also showed a higher positive correlation of $NDVI_{gs}$ with AP

($R = 0.75$), followed by that of AAAT ($R = 0.63$). Meanwhile, there was a significant negative correlation between AAAT and NDSI ($R = -0.59$) (Figure S1). The q -value of MAGT was 0.20, that of NDSI was 0.08, and that of the slope was the lowest (0.04), indicating insignificant effects of NDSI and slope on $NDVI_{gs}$. Overall, AP and AAAT were the main drivers for variations in $NDVI_{gs}$. The contributions of each driver to the spatial variations in $NDVI_{gs}$ decline in the order of AP, AAAT, elevation, MAGT, NDSI, and slope.

Table 3. The q -statistics of the contribution of each factor to $NDVI_{gs}$ along National Highway G214 from the Changshitoushan Mountain Pass to the Qingshui'he Riverside at the northern and southern flanks of the Bayan Har Mountains, southern Qinghai, China, from 2010 to 2019.

Types	Elevation	Slope	NDSI	AP	AAAT	MAGT
q -statistic	0.31	0.04	0.08	0.52	0.45	0.20
p -value	<0.01	<0.01	<0.01	<0.01	<0.01	<0.01

Note: NDSI, Normalized Difference Snow Index; AP, annual precipitation; AAAT, average annual air temperature. Mean annual ground temperature (MAGT) was obtained at the depth of zero annual amplitude at 15 m in depth from the 16 boreholes along the G214 in the study area.

The interaction detection analysis demonstrated that the effects of the two interactive factors were greater than those of a single factor. The interactions of all influencing factors showed a two-factor enhancement, among which the q -value of $AP \cap AAAT$ was the highest at 0.74 (Figure 5). This suggests that the interactive AP and MAAT were the two key factors for driving the spatial variations in $NDVI_{gs}$ along the G214. Additionally, the q -values of $AP \cap MAGT$, $AP \cap$ elevation, elevation $\cap$ MAGT, and $AP \cap$ NDSI were all greater than 0.6, indicating the strong interactive effects of these four factors on $NDVI_{gs}$.

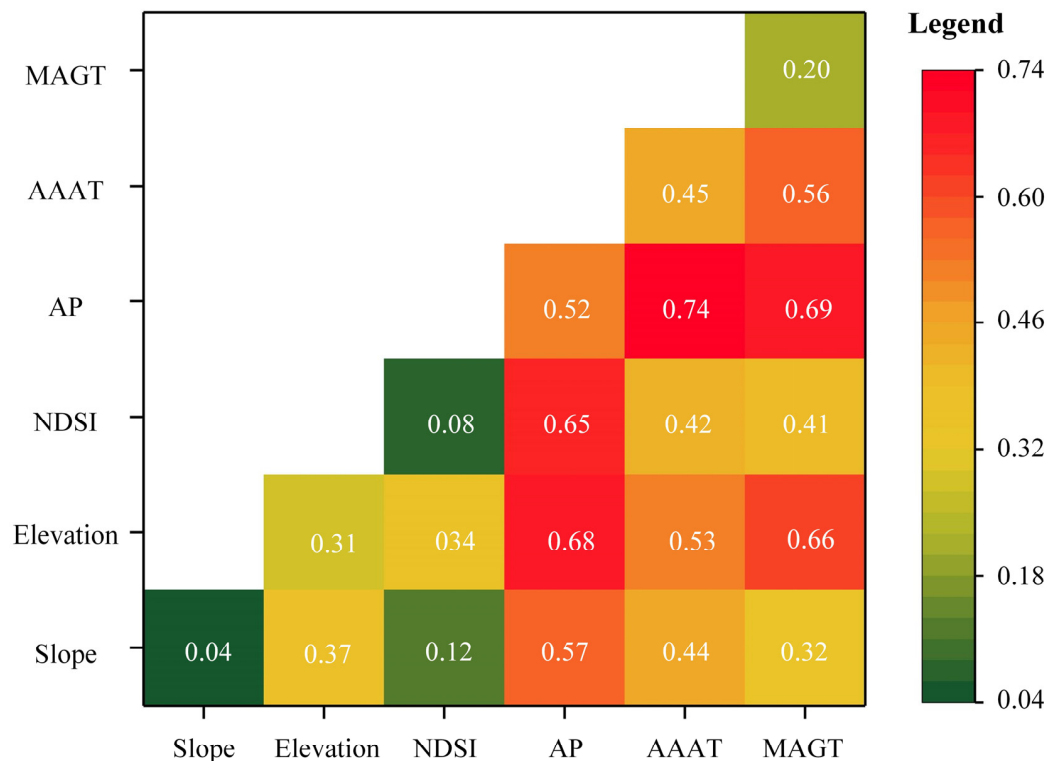


Figure 5. Effects of two interactive factors on Normalized Difference Vegetation Index (NDVI) along the National Highway G214 from the Changshitoushan Mountain Pass to Qingshui'he Riverside at the northern and southern flanks of the Bayan Har Mountains, southern Qinghai, China, from 2010–2019. Note: NDSI, Normalized Difference Snow Index; AP, annual precipitation; AAAT, average annual air temperature; MAGT, mean annual ground temperature.

4. Discussion

4.1. Vegetation Change in the Source Area of the Yellow and Yangtze Rivers (SAYYR)

The NDVI_{gs} along the G214 in our study area showed a northward decreasing pattern, possibly because of the upward decline trend in hydrothermal regimes from southeast to northwest in the SAYYR with rising elevation [76,77]. Furthermore, NDVI_{gs} was “low at the northern flank and high at the southern flank of the Bayan Har Mountains”. This indicates a higher vegetation cover at the southern flank in the source area of the Yangtze River than at the northern flank in the Headwater Area of the Yellow River along the G214 in the study area due to the linear relationship between vegetation cover and NDVI_{gs} [78]. There could be two main explanations for this. One is that on the northern flank, there are lower elevations, gentler slopes, more lakes and rivers, and lower vegetation cover. Therefore, the NDVI_{gs} at the northern flank of the Bayan Har Mountains are mostly areas of low NDVI_{gs}, with milder fluctuations. Second, there is more AP at high elevations at the southern flank and thus better vegetation. Therefore, this leads to more areas of high NDVI_{gs} at the south flank, with large fluctuations.

From 2010 to 2019, NDVI_{gs} displayed an overall upward trend in the study area, ranging between 0.53 and 0.58 and with an increasing rate of 0.006/a. This trend is consistent with previous studies on changes in NDVI in the SAYYR [79–81]. A study using multiple high-resolution satellite data showed a slightly increasing pattern, and the NDVI ranged from 0.23 and 0.27 in the Headwater Area of Three Rivers (Yangtze, Yellow, and Lancang Rivers) on the northeastern QTP from 1982 to 2015 [24]. The increasing rate of NDVI in the study area (0.0002/a from 1980 to 2018 [59]) was higher than that for the whole SAYYR (0.0020/a from 1998 to 2016 [82]).

4.2. Impacts of the Highway on Vegetation in Permafrost Regions

Spatially, NDVI_{gs} showed an upward trend from the buffer zones at the distance of 0–10 to 20–30 km perpendicularly away from the G214, suggesting that the impacts of the highway on vegetation weaken with increasing outward distance. NDVI_{gs} varied substantially in the first buffer zone and stabilized in the second zone. Therefore, we may preliminarily conclude that the G214 have a significant effect at a distance extending up to 10 km from the highway. This result is comparable with findings of previous studies [54,83].

The former G214 was about 4 m wide. The construction of the approximately 6 m wide new G214 (the Gonghe–Yushu Expressway), almost parallel and adjacent to the former one, was started in 2011, and it was put into operation in 2017. During the construction, soils were collected from the nearby or designated quarries and filled into the subgrade, generally about 1–5 m higher than the original natural ground. This process destroyed the pristine alpine grasslands and created new patches or stripes of bare ground along the highway. In addition, similar bare patches and stripes were created by the temporary stacking of construction materials and facilities. Although revegetation was carried out after the highway’s construction, the recovery rate of alpine vegetation was slow or even impossible because of the simple plant community structure and the harsh periglacial and arid environment [51,63]. This could be the main reason for the most stable zone of 0–10 km. In addition, there are substantially less important local, branch, and access roads and trails and small- to mid-size towns and villages along the highway, which have also modified and damaged the alpine grasslands to different extents, creating some new patches and stripes of bare ground.

It was unexpected that the influences of the G214 on the SQP extended as far as 10 km, since NDVI_{gs} still showed similar changing patterns with the zones close to the highway. This is inconsistent with some results from other permafrost regions on the interior QTP, which documented that the impacts of the Qinghai–Tibet Engineering Corridor (both highway and railway) from Golmud to Lhasa, Tibet Autonomous Region, Southwest China were limited to about 1.5 km on each side [53]. This may be because of the different vegetation indicators used in the study [84]. In addition, the source and resolution of data and the unique location of the research area, in addition to regional hydroclimate

and alpine environments, could be reasons for this discrepancy. In our research, a spatial resolution of 1.0×1.0 km NDVI data was used, while a spatial resolution of 250×250 m was used in the other study [86]. The influences of data resolution on the study of highway effects, in addition to the criteria for judging the “affected zones,” need further systematic and in-depth studies in the future.

4.3. Interactive Impacts of Geocryological, Topographical, Hydroclimatic, and Other Factors on NDVI

Changing vegetation patterns along the G214 in the buffer zone were mainly driven by AP, AAAT, MAGT, and construction and operation of the G214. Pearson correlation analysis show that AP, AAAT, and MAGT have significant positive effects on NDVI_{gs}. This is supported by earlier studies of the QTP, showing that AP and AAAT are important for vegetation dynamics [24,53,83,85]. On the QTP, precipitation and temperature are key factors for plant growth, and due to the nature of the arid environment, precipitation is more limiting than other driving factors for vegetation growth [83]. However, other studies reported that the upward trend of NDVI was more associated with recently rising air temperature on the QTP [86]. This is associated with the hydrothermal conditions of the study site. For example, vegetation is more related to air temperature in the Headwater Area of the Three Rivers on the northeastern QTP, where the regional climate is characterized of relatively low AAAT and relatively higher AP, and thus the warmth has more of the limiting effect for vegetation growth [87]. This is also the main reason for the NDVI patterns of “low at the northern flank and high at the southern flank” in the study area striding across the Bayan Har Mountains along the G214.

Although precipitation and temperature dominated in driving NDVI changes in alpine regions, ground temperature was also a significant driver for NDVI changes [88], suggesting that the presence of permafrost could be a major driver for better vegetation growth on the QTP. This result agrees well with other studies on the QTP, e.g., [53,63,82].

Permafrost, as an aquitard, could inhibit the unfrozen water in the surface soil from percolating downwards and stored soil moisture in the active layer in the growing season or slightly above the permafrost table, thus facilitating vegetation growth [64]. In this study, permafrost was gradually warming from 2010 to 2019 according to the data from the borehole along the G214, and the average rate was 0.01 °C/a (Table S1). It has been widely reported that the plateau permafrost has been degrading extensively and rapidly, as characterized by shrinkage of the area extent of permafrost, increasing of the active layer thickness, ground temperature, and the extent of taliks, e.g., [89,90]. These effects of permafrost degradation could enhance the upward evaporation and downward infiltration of soil waters, subsequently adversely impacting the growth of alpine vegetation [53,63].

Furthermore, the results of GeoDetector analysis showed a stronger interactive effect of two factors on NDVI_{gs} than a single one. This agrees well with other similar observations, e.g., [83,91,92]. Our study showed that NDVI_{gs} was largely influenced by the interactions between $AP \cap AAAT$ and $AP \cap MAGT$. Additionally, the interactions between $AP \cap$ elevation, $AP \cap NDSI$, and $MAGT \cap$ elevation had relatively stronger effects on NDVI. This is because AP, AAAT, and MAGT vary along the elevational gradient and north–south flanks of the Bayan Har Mountains, affecting the spatial distribution and growth of vegetation [93].

Unfortunately, this study is unable to separate the individual influences of the G214 and environmental factors. The construction of the highway locally and directly destroyed the alpine vegetation along the highway, while the impacts of the operating highway on alpine vegetation could be indirectly exerted through changing the environmental variables [54]. For example, several studies have documented permafrost degradation along the major engineering corridors on the QTP, e.g., [62,90,94] and at high latitudes, e.g., [44,95–97]. This may further alter the hydrothermal conditions and landscapes along the highways and accelerate vegetation changes under a warming climate.

4.4. Implications and Limitations

This research used NDVI from MODIS data to study the impacts of G214 on vegetation changes, and the main environmental drivers were detected. We found that MAGT was one of the major influencing factors for alpine vegetation. However, due to the data's limitations, only data from 16 boreholes along the G214 at the northern and southern flanks of the Bayan Har Mountains were used to address the permafrost changes and their impacts on NDVI. Furthermore, it is challenging to separate the influences of the G214 and permafrost degradation on NDVI, although we have theoretically concluded that the vegetation is affected by the G214 and permafrost degradation. More systematically designed permafrost boreholes away from the highway and systematic methods to analyze the data are deemed necessary to investigate the influences of highway on vegetation in permafrost regions.

The impacts of the G214 on vegetation were explored in this study. However, it is challenging to define the extents of the impacts of the G214 in the study area, because the criteria for the impacted zone could be very complex and not all variables are systematically measured over long periods. In addition, the impacted outward distance is tremendously influenced by the sources of datasets and the areal extent of the study region. More studies in different spatial scales are needed to further explore and understand the influences of highways on vegetation in alpine permafrost regions.

5. Conclusions

The spatiotemporal patterns of NDVI_{gs} along the G214 at the southern and northern flanks of the Bayan Har Mountains were reported in this paper. Furthermore, the impacts of the G214 on NDVI_{gs} of alpine vegetation along the engineering corridor therewith were studied. Finally, the links between NDVI_{gs} and environmental variables were systematically analyzed. The main conclusions are as follows: (1) NDVI_{gs} showed a slight upward tendency (0.006/a) from 2010 to 2019. The spatial distribution of NDVI_{gs} along the G214 showed a pattern of “low at the northern flank and high at the southern flank of the Bayan Har Mountains”. NDVI_{gs} declined by 16.8% in the Low and Lower classes and increased by 15.2% in the High and Higher classes. (2) The average NDVI_{gs} gradually increased with the increasing outward distance from the G214, indicating the greatest influence in the first buffer zone at the perpendicular distances of 0–10 km away from the highway. The average CV in the engineering corridor was 0.23, and CV in the first buffer zone of 0–10 km was relatively small, with the highest stability of NDVI_{gs} and with 96.3% of the area showing minimum and low fluctuations. (3) AP, AAAT, and elevation were the major drivers for the NDVI patterns, with the highest *q*-value of AP (0.52). Furthermore, the interaction analysis showed that the two-factor interactions had a stronger effect than that of one single factor. The interactive *q*-value of AP $\cap$ AAAT was the highest (0.74), suggesting that these interactive climate variables dominated in driving the distribution of NDVI_{gs} and its changes.

Supplementary Materials: The following supporting information can be downloaded at: <https://www.mdpi.com/article/10.3390/rs15061547/s1>, Figure S1: Correlation between hydroclimatic and environmental drivers and NDVI on north and south flanks of the Bayan Har Mountains along the G214 on the northeastern Qinghai-Tibet Plateau during 2010–2019. Note: NDSI, Normalized Difference Snow Index; AP, annual precipitation; AAAT, average annual air temperature. MAGT, mean annual ground temperature; Table S1: Warming rate of ground temperature during 2010–2019 in boreholes on north and south flanks of the Bayan Har Mountains along the National Highway G214 from the Changshitoushan Mountain Pass to the Qingshui'he Riverside at the northern and southern flanks of the Bayan Har Mountains, southern Qinghai, China from 2010 to 2019.

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