



Article Distinguishing the Impacts of Human Activities and Climate Change on the Livelihood Environment of Pastoralists in the Qinghai Lake Basin

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Abstract: Grassland vegetation is the largest terrestrial ecosystem in the Qinghai Lake Basin (QLB), and it is also the most important means of production for herders' livelihoods. Quantifying the impact of climate change and human activities on grassland vegetation changes is an essential task for ensuring the sustainable livelihood of pastoralists. To this end, we investigated vegetation cover changes in the QLB from 2000 to 2020 using the normalized difference vegetation index (NDVI), meteorological raster data, and digital elevation and used residual analysis of multiple linear regression to evaluate the residuals of human activities. The residual analysis of partial derivatives was used to quantify the contribution of climate change and human activities to changes in vegetation cover. The results showed that: (1) The vegetation coverage of the QLB increased significantly (0.002/a, p < 0.01), with 91.38% of the area showing a greening trend, and 8.62% of the area suffering a degrading trend. The NDVI decreased substantially along the altitude gradient (-0.02/a, p < 0.01), with the highest vegetation coverage at 3600-3700 m (0.37/a). The vegetation degraded from 3200-3300 m, vegetation greening accelerated from 3300-3500 m, and vegetation greening slowed above 3500 m. (2) The contribution of climate change, temperature (T), and precipitation (P) to vegetation cover change were 1.62/a, 0.005/a, and 1.615/a, respectively. Below 3500 m, the vegetation greening was more limited by P. Above 3500 m, the vegetation greening was mainly limited by T. (3) Residual analysis showed that the contribution of human activities to vegetation cover was -1.618/a. Regarding the altitude gradient, at 3300–3500 m, human activities had the highest negative contribution to vegetation coverage (-2.389/a), and at 3200-3300 m, they had the highest positive contribution (0.389/a). In the past 21 years, the impact of human activities on vegetation coverage changed from negative to positive. Before 2009, the annual average NDVIres value was negative; after 2010, the average yearly NDVIres value turned positive. In general, the vegetation greening of the QLB depends on climate warming and humidification. The positive impact of human activities over the past decade was also essential for vegetation greening. These findings deepen our understanding of the QLB vegetation changes under climate change and human activities.

Keywords: NDVI; climate change; human activities; residual analysis; Qinghai Lake Basin

1. Introduction

As a significant component of the land-atmosphere system [1,2], vegetation fundamentally regulates the material cycle and energy flow on the earth's surface [3,4]—its changes



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Copyright: © 2022 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). may also affect climate through surface hydrothermal regulation and biological evolution [5–8]. In addition, vegetation is also the basis of economic and social development in some regions [9], especially in the Qinghai-Tibet Plateau, where animal husbandry is the primary livelihood [10]. Therefore, studying the relationship between vegetation cover changes, climate, and human factors is of great significance in evaluating the ecological security and sustainable livelihood of farmers and herders in the Qinghai-Tibet Plateau.

As the primary vegetation coverage monitoring method used in the past (sampling method [11], instrument method [12]), field measurement can hardly meet the needs of obtaining vegetation cover monitoring data covering an extensive range. With the development of satellite remote sensing technology, RVI (rainfall variability index) [13], DVI (difference vegetation index) [14], PVI (perpendicular vegetation index) [15], NDVI (normalized difference vegetation index) [16], SAVI (soil-adjusted vegetation index) [17], and other vegetation indices obtained from vegetation reflectance spectral characteristics are widely used for large-scale vegetation monitoring. Among these, NDVI takes into account the interference of the terrain and the vegetation canopy, eliminating the error caused by radiation, and can well reflect the biomass and greenness of vegetation. It has been proved to have high application value in response to environmental changes [18]. Therefore, NDVI has become a standard indicator for monitoring vegetation growth and coverage density [19].

Vegetation change and its attribution to the Qinghai-Tibet Plateau have always been hot topics for scholarly research. In the last 30 years, the NDVI of the Qinghai-Tibet Plateau has increased significantly at a rate of 0.001–0.002/a [20,21]. The alpine grassland and alpine meadow have been improved [22]. Climate change was considered one of the reasons for vegetation cover changes [23]. Wu et al. found that temperature (T), precipitation (P), and radiation energy explained 66.2% of the changes in the alpine grasslands on the Qinghai-Tibet Plateau [24]. Liu et al. believed that P and active photosynthetic radiation were the main factors affecting the NDVI variation of different grassland types on the Tibetan Plateau [25]. In other regions within the Qinghai-Tibet Plateau, the impact of climatic factors on NDVI exhibits noticeable spatial and temporal differences. In the semi-arid and cold areas of the upper reaches of the Yarlung Zangbo River, herbs and shrubs were susceptible to changes in P and T [26], while NDVI in the lower reaches was significantly and positively correlated with surface T [27]. Chen et al. found that the difference in climate change in the Qinghai-Tibet Plateau determined the spatial difference in vegetation response. The correlation between the northern and southern parts of the plateau was opposite to *P* and *T*, respectively. In the south part of the plateau, the greening trend slowed down due to increased cooling and humidification, and some areas even deteriorated [28]. The IPCC Sixth Assessment Report pointed out that at the current rate of carbon dioxide and other greenhouse gases emissions, the global T increase would reach or exceed 1.5 $^{\circ}$ C over the next 20 years [29]. The response of vegetation change to climate change in the Qinghai-Tibet Plateau would also persist. Fan et al. predicted that the future climate change intensity would directly affect the rate of vegetation change on the Qinghai-Tibet Plateau, especially regarding the altitude gradient, where the change of vegetation types in low altitude and high cold areas could reach 7.54%/10 a-11.32%/10 a, respectively [30].

In addition, human activities are also an essential factor in the vegetation change of the Qinghai-Tibet Plateau. On the one hand, the impact of human activities on vegetation in the Qinghai-Tibet Plateau was smaller than that of climate factors [31], and the effect on vegetation also changed from limiting to promoting. On the other hand, in the southern, eastern [32], and northeastern [28] regions of the Qinghai-Tibet Plateau, human activities, such as grazing, engineering construction, and the influx of tourists, have caused significant disturbance to vegetation [32–34], resulting in a decline in NDVI values, alpine vegetation degradation, and other issues. In summary, the NDVI and climatic factors of the Qinghai-Tibet Plateau show spatial non-stationarity and scale dependence [35], which requires us to pay attention to the dynamics of vegetation changes in different regions within the plateau.

The human–land relationship is the focus of regional sustainability research [36]. Grasslands are not only the primary carrier of the human-land relationship on the Qinghai-Tibet Plateau, but are also the most vulnerable to disturbance by human activities. Therefore, a comprehensive multi-perspective study on the evolution of the grassland resources environment and its interaction with climate and humans is important to maintain the coordinated development of the human-nature-economic system. The Qinghai Lake Basin (QLB) is located in the northeastern part of the Qinghai-Tibet Plateau, and it is a transitional zone between the alpine region of the Qinghai-Tibet Plateau, the arid and semi-arid region of the northwest, and the arid region of the Loess Plateau area. However, the specific changes that have occurred in the vegetation of the QLB, especially the degree to which climate change and human activities have affected the changes in vegetation cover, and the main factors affecting the changes in the vegetation cover of the QLB have not been studied quantitatively. To this end, based on the MODIS NDVI dataset, this research evaluates the distribution and trend of vegetation cover in the QLB from 2000 to 2020, reveals the residual trend of changes in vegetation cover caused by human activities over the past 21 years, and quantifies the contribution of climate change and human activities to vegetation cover dynamics from the two dimensions of space and altitude. The research results will promote our understanding of the dynamic changes of the alpine steppe in the northeastern Qinghai-Tibet Plateau and are significant for the targeted implementation of regional ecological conservation.

2. Materials and Methods

2.1. The Study Area

The QLB is a closed inland basin (Figure 1), which is located in the northeastern Qinghai-Tibet Plateau ($36^{\circ}15'-38^{\circ}20'$ N, $97^{\circ}50'-101^{\circ}20'$ E), with an area of about 2.96×10^4 km². China's largest saltwater lake—Qinghai Lake—is located here. The topographical altitude of the QLB decreases from northwest to southeast [37], and the average elevation is above 3000 m. The QLB is distinguished by a typical continental plateau climate, with intense solar radiation. In most areas, the average annual *T* is below 0 °C, and the average annual *P* is below 400 mm [38].



Figure 1. Location of the QLB.

The vegetation types in the QLB show the coexistence of temperate vegetation and alpine vegetation. The main vegetation types are alpine grassland, alpine meadow, sandy

vegetation, halophytic meadow, marsh meadow, etc.. Among these, the alpine meadow is the most widely distributed area, mainly at the altitude of 3000–4000 m, accounting for 70.26% of the size of the QLB. In addition, the vertical zonal distribution of vegetation in the QLB is apparent. At lower altitudes, the lake basin and river valley are dominated by grassland vegetation such as *Achnatherum splendens* L., *Stipa capillata* L., and *Agropyron cristatum* L. In contrast, alpine steppe, alpine shrub, and alpine meadow are the primary vegetation types in the surrounding mountains and areas at higher altitudes [39].

Relying on the abundant grassland resources, the QLB has become an important pastoral area on the Qinghai-Tibet Plateau, and grazing is the main livelihood of residents. By the end of 2019, the total population of the QLB was 108,639, and the animal husbandry population was 73,937, accounting for 12.03% of the total regional population [40].

2.2. Data Sources and Processing

We used NDVI to indicate the vegetation cover of QLB, which is derived from the MOD13Q1 data product released by NASA (National Aeronautics and Space Administration) (https://search.earthdata.nasa.gov, accessed on 15 November 2021), with a temporal resolution of 16 days and a spatial resolution of 250 m. Due to its higher resolution and lower uncertainty, the MOD13Q1 data product has been widely used to study vegetation change on the Qinghai-Tibet Plateau [41]. The digital elevation (DEM) data was sourced from the China Geospatial Data Cloud Platform (http://www.gscloud.cn/, accessed on 12 January 2022), with a spatial resolution of 90 m. T and P were selected as climatic elements affecting NDVI, which were obtained from the China's National Earth System Science Data Center (http://www.geodata.cn, accessed on 23 December 2021), with a spatial resolution of 1 km. The population and GDP spatial distribution kilometer grid dataset were obtained from the Resource and Environmental Science and Data Center of the Institute of Geographical Sciences and Natural Resources Research, Chinese Academy of Sciences (https://www.resdc.cn/, accessed on 8 February 2022), with a resolution of 1 km. The livestock data per unit area are derived from the statistical data of the towns in Haibei, Hainan, and Haixi.

The above data were preprocessed using Python 3.5 software (Python Software Foundation, Beaverton, OR, USA), written to complete splicing, cropping, and data format conversion. At the same time, all raster data were resampled to a spatial resolution of 250 m \times 250 m. The contributions of climate change and human activities to changes in vegetation cover were calculated by MatLab R2021b (MathWorks, Natick, MA, USA). All visualizations in this research were processed by ArcGIS Pro 2.8.0 (Esri, Berkeley, CA, USA).

2.3. Research Methods

2.3.1. Trend Analysis of Vegetation Coverage and Climate Factors

The linear trend of *T*, *P*, and NDVI from 2000 to 2020 was estimated using the least squares regression method. The calculation formula of this method is as follows [42]:

$$Slope = \frac{n \times \sum_{i=1}^{n} i \times j_{i} - \sum_{i=1}^{n} i \sum_{i=1}^{n} i}{n \times \sum_{i=1}^{n} i^{2} - \left(\sum_{i=1}^{n} i\right)^{2}}$$
(1)

where *slope* is the regression equation slope and *n* is the length of the time series, which is 21 in this study. If *slope* > 0, vegetation *T*, *P*, and NDVI increase; if *slope* < 0, *T*, *P*, and NDVI decrease.

2.3.2. Quantifying the Impact of Climate Change and Human Activities on Vegetation Cover

In this research, the residual analysis method based on multiple linear regression was used to evaluate the influence trend of the QLB human activities on vegetation cover. This method has been widely used in assessing vegetation cover changes since it was first proposed in 2004 [43,44]. The residual analysis first needs to establish a multiple linear regression model of *T*, *P*, and NDVI at the pixel scale to predict the predicted value of NDVI ($NDVI_{pre}$) under the influence of *T* and *P*. The difference between the observed values of NDVI ($NDVI_{obs}$) and $NDVI_{pre}$ is the residual ($NDVI_{res}$), that is, the degree of influence of human activities on vegetation. The calculation formula is:

$$NDVI_{pre} = a \times P + b \times T \tag{2}$$

$$NDVI_{res} = NDVI_{obs} - NDVI_{pre}$$
(3)

Among these, *a* is the regression coefficient between NDVI and *T*, and *b* is the regression coefficient between vegetation cover and *P*; if $NDVI_{res} < 0$, it means that human activities have caused vegetation degradation; if $NDVI_{res} > 0$, it means that human activities have promoted vegetation growth; if $NDVI_{res} = 0$, the vegetation cover changes are attributed to climate change.

Although residual analysis based on multiple linear regression can distinguish the effects of climate change and human activities on vegetation cover, it cannot be quantitatively assessed. To this end, this research used residual analysis of partial derivatives to quantify the contribution of climate change and human activities to changes in vegetation cover. At the same time, the altitudes of the QLB were reclassified into 17 categories according to the 100 m interval. We then extracted the contribution rates of climate change and human activities on different altitude gradients. The residual analysis method based on partial derivatives was proposed by Roderick et al. in 2007 [45] and has been widely used in many studies [46]. The calculation formula is:

$$NDVI_{slope} \approx C_{_con} + H_{_con} = T_{_con} + P_{_con} + H_{_con}$$

= $\frac{\partial NDVI}{\partial T} \times \frac{dT}{dt} + \frac{\partial NDVI}{\partial P} \times \frac{dP}{dt} + H_{_con}$ (4)

Among these, $NDVI_{slope}$ is the interannual variation slope of NDVI; $C_{_con}$, $H_{_con}$, $T_{_con}$ and $P_{_con}$ represent the contribution of climate change, human activities, T, and P to vegetation covering NDVI, respectively; The sum of $T_{_con}$ and $P_{_con}$ is $C_{_con}$; the residual of $NDVI_{slope}$ and $C_{_con}$ is approximately equal to $H_{_con}$; $\frac{\partial NDVI}{\partial T}$ and $\frac{\partial NDVI}{\partial P}$ represent the partial correlation coefficients of T, P, and NDVI, respectively (excluding the interference of Pand T, respectively); $\frac{dT}{dt}$ and $\frac{dP}{dt}$ are the interannual variation slopes of T and P in the time variable t, respectively. Among these, partial correlation analysis is an effective method to study the linear relationship between two factors, while eliminating the interference of other factors. The calculation formula is [47]:

$$R_{xy,z} = \frac{R_{xy} - R_{xz} \times R_{yz}}{\sqrt{(1 - R_{xz}^2) \times (1 - R_{yz}^2)}}$$
(5)

where $R_{xy,z}$ is the partial correlation coefficient between x and y under the condition that when the influence of variable z is excluded, R_{xy} , R_{xz} and R_{yz} are the simple correlations between x and the other two variables y and z, respectively. The significance test of the partial correlation coefficient is completed by the t test, and its calculation formula is [48]:

$$t = \frac{R_{xy}}{\sqrt{1 - R_{xy}^2}} \sqrt{n - m - 1}$$
(6)

Among these, *n* is the number of samples (the time series is 2000–2020, that is, n = 21), and *m* is the number of independent variables.

3. Results

3.1. Spatial Distribution Characteristics and Variation Trend of NDVI

From 2000 to 2020, the annual average NDVI value for the QLB was 0.28, fluctuating and rising between 0.26 and 0.32 (Figure 2), showing a significant increasing trend (0.017/10a, p < 0.01). In 2004, the annual average NDVI value was the minimum (0.26), and in 2018, it was the maximum (0.32).



Figure 2. Variation of NDVI in the QLB from 2000 to 2020.

The mean NDVI across the QLB decreased from southeast to northwest (Figure 3a). The areas with the most significant NDVI values were distributed in the northern and southern parts of Qinghai Lake, while the areas with the smallest NDVI values were mainly distributed in the upper Buha River in the northwest and the Shadao area on the east bank of Qinghai Lake. The changing trend of the calculated NDVI can be seen (Figure 3b). Pixels with a positive inter-annual variation slope of NDVI accounted for 91.38% of the entire study area, and pixels with a negative inter-annual variation only accounted for 8.62%, indicating that the overall vegetation coverage of the QLB in the past 21 years was dominated by improvement. More specifically, the trend of NDVI value increasing from the north to the south of the basin was obvious. The areas with the most vegetation improvement were mainly concentrated on both sides of the Buha River Valley and the northern part of Qinghai Lake (0.03–0.13/10a). The areas with the most severe vegetation degradation were mainly distributed around the shores of Qinghai Lake (north bank, east bank, west bank) and the upper reaches of the Shaliu River.



Figure 3. Spatial distribution (a) and variation trend (b) of NDVI in the QLB from 2000 to 2020.

3.2. Spatial Distribution Characteristics and Variation Trends of Climate Factors

T and *P* determined the total input of heat and moisture and were the main climatic factors for vegetation growth. The spatial distribution and variation trend of the QLB climate factors are shown in Figure 4. It can be seen that the annual average *T* of the QLB ranged from -13.37 °C to 1.51 °C (Figure 4a), and the distribution was consistent with the altitude (Figure 1). The average annual *T* was higher in the areas around Qinghai Lake and the middle and lower reaches of the Buha River in the south, and lower in the high-altitude areas in the north. The annual average *P* showed an increasing spatial distribution pattern from the west (318.98 mm) to the east (629.56 mm) (Figure 4b). During the 21 years, the pixels with reduced *T* only accounted for 5.19% of the watershed area (Figure 4c) and were mainly distributed in the northwest of the study area. The pixels with elevated *T* accounted for 94.81% of the watershed area, and the areas with the fastest warming were mainly around Qinghai Lake (0.02–0.06 °C/a). From 2000 to 2020, the annual average *P* of the QLB generally increased at a rate of 3.08–7.08 mm/a (Figure 4d) and showed a spatially symmetrical distribution pattern, with the slope of *P* gradually increasing along the center to the southeast and northwest.

3.3. Residual Trends of Human Activities

Figure 5 shows the changing trend of $NDVI_{res}$ from 2000 to 2020. The annual average $NDVI_{res}$ value of the QLB ranged from -0.02 to 0.03 (Figure 5a), showing a significant upward trend ($\mathbb{R}^2 = 0.48$, p < 0.01). From 2000 to 2009, $NDVI_{res}$ was mainly negative, and only in 2006 was $NDVI_{res}$ positive. After 2010, $NDVI_{res}$ was mainly positive, which indicated that the adverse effects of human activities on vegetation coverage had gradually weakened in the past 21 years, and the favorable effects on vegetation coverage have continued to increase. Further analysis of the spatial variation of $NDVI_{res}$ can be seen (Figure 5b). A total of 89.74% of the pixel $NDVI_{res}$ increased, mainly distributed in the Buha River basin and along the periphery of Qinghai Lake. In comparison, 10.22% of the pixel $NDVI_{res}$ decreased, distributed primarily in the northeast of the basin.



Figure 4. Spatial distribution (**a**,**b**) and variation trend (**c**,**d**) of *T* and *P* in the QLB from 2000 to 2020.



Figure 5. Interannual variation (a) and trend (b) of NDVI_{res} in the QLB from 2000 to 2020.

3.4. Contribution of Climate Change and Human Activities to Vegetation Cover

To determine the relationship between vegetation cover and climate factors, we analyzed the partial correlation between *T*, *P*, and NDVI (Figure 6). Throughout the QLB, the partial correlation coefficients of *T*, *P*, and NDVI were 0.21 (p > 0.05) and 0.37 (p > 0.05), respectively. As shown in Figure 6a, 80.68% of the regional annual *T* in the study area was positively correlated with NDVI. Only 15.65% of the pixels reached the significance level of p < 0.05 (Figure 6b), while the area where *P* was positively correlated with NDVI reached 91.88% (Figure 6c), and 44.47% of the areas showed p < 0.05 (Figure 6d).



Figure 6. Spatial distribution and significance (*p*-value) of partial correlation coefficients between climatic factors and NDVI in the QLB from 2000 to 2020. (**a**,**b**) is the partial correlation coefficient and significance of *T* and NDVI(*p*-value), respectively; (**c**,**d**) is the partial correlation coefficient and significance of *P* and NDVI (*p*-value), respectively.

Figure 7 shows the residual analysis based on partial derivatives. The contribution of $T(T_{con})$ to the interannual variation of NDVI was 0.005/a, and T_{con} values ranged from -0.03-0.04/a (Figure 7a). The area with $T_{con} > 0$ accounted for 82.78% of the QLB area, mainly distributed around Qinghai Lake and the middle and lower reaches of the Buha River; the area with T_{con} value < 0 accounted for only 17.72% of the QLB area, mainly in the northwest of the study area. The contribution of $P(P_{con})$ to the interannual variation of NDVI was 1.615/a, with P_{con} values ranging from -4.019-4.983/a (Figure 7b). The area with $P_{con} > 0$ accounted for 91.88% of the basin area, and the highest values were mainly distributed in the eastern part of the study area and the middle and upper reaches of the Buha River. The area with P_{con} value < 0 only accounted for 8.12% of the watershed area, and was mainly distributed on the edge of the study area and the north, east, and west shores of Qinghai Lake.

Based on the results of $T_{_con}$ and $P_{_con}$, the spatial distribution contribution of climate change ($C_{_con}$) and human activities ($H_{_con}$) to the interannual variation of NDVI is obtained (Figure 8). From 2000 to 2020, the average annual $C_{_con}$ value was 1.62/a, ranging from -4.023-4.988/a (Figure 8a). The area with $C_{_con} > 0$ accounted for 91.91% of the watershed area, and the area with $C_{_con} < 0$ only accounted for 8.09% of the watershed area. Compared with $C_{_con}$, human activities showed a strong negative impact on NDVI changes (Figure 8b): the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 91.9% of the watershed area, and the area with $H_{_con} < 0$ accounted for 8.1% of the watershed area. Overall, in 2000–2020, $H_{_con}$ was -1.618/a, ranging from -4.985-4.021/a.



Figure 7. Spatial distribution of T_{con} (**a**) and P_{con} (**b**) in the QLB from 2000 to 2020.



Figure 8. Spatial distribution of C_{con} (**a**) and H_{con} (**b**) in the QLB from 2000 to 2020.

3.5. Distribution and Greening of Vegetation along the Altitude Gradient

Figure 9 shows the distribution of NDVI along the altitude gradient and the greening trend in the QLB. NDVI decreased significantly along the elevation gradient ($R^2 = 0.73$, p < 0.01), and NDVI decreased by 0.02 for every 100m of increase in altitude. Among these, NDVI rises rapidly below 3600 m, reaches the highest value (0.37/a) at 3600–3700 m, and then decreases continuously. Above 4600 m, NDVI is less than 0.1, and almost no vegetation grows. This can be seen from the trend of vegetation greening according to the altitude gradient. Vegetation at 3200–3300 m shows a trend of degradation, vegetation greening at 3300–3500 m is accelerated, and vegetation greening above 3500 m is slowed down.



Figure 9. The distribution of NDVI along the altitude gradient and the greening trend in the QLB.

Figure 10 shows the change slopes of the climatic factors and the distribution of their contributions to vegetation cover changes along the altitudinal gradient. The mean values of $T_{_con}$ and $P_{_con}$ over the altitude gradient are 0.0032/a and 1.296/a, respectively. For every 100m increase in altitude, $T_{_con}$ and $P_{_con}$ decreased by 0.0792/a ($R^2 = 0.27$, p < 0.01) and 0.0006 ($R^2 = 0.27$, p < 0.05), respectively. In terms of the changing trend, from 3200 to 3300 m, the *T* rises faster, and the *P* increases less, and $T_{_con}$ and $P_{_con}$ were -0.006/a and -0.39/a, respectively. Water limitation may be the main reason for vegetation growth. From 3300 to 3400 m, the obvious climate warming and humidification led to a rapid increase in $T_{_con}$ and $P_{_con}$. From 3400 to 3500 m, *T* rises slowly, $T_{_con}$ remains at the maximum value (0.009/a), *P* continues to increase, $P_{_con}$ increases to the maximum value (2.383/a), and the increase in precipitation may be the main reason for vegetation greening. Above 3500 m, *P* increased rapidly, but temperature decreased rapidly, and $T_{_con}$ and $P_{_con}$ continued to decrease. *T* was the main reason for the limiting of vegetation greening.



Figure 10. The change slope of *T* and *P* and the distribution of T_{con} and P_{con} along the altitude gradient of the QLB.

Figure 11 shows the distribution of $C_{_con}$ and $H_{_con}$ over the altitude gradient. At 3200–3300 m, $C_{_con}$ and $H_{_con}$ reached the lowest value (-0.39/a) and the highest value (0.389/a), respectively. From 3300 m to 3500 m, $C_{_con}$ rises rapidly, and $H_{_con}$ falls quickly, reaching the maximum value (2.392/a) and the minimum value (-2.389/a), respectively. Above 3500 m, the importance of $C_{_con} > 0$ and $H_{_con} < 0$ continued to decrease, and the absolute value of the former was greater than that of the latter, indicating that vegetation greening was mainly affected by climate change. Overall, the favorable impact of climate change on vegetation cover ($C_{_con} = 1.299/a$) along the altitude gradient was higher than the unfavorable impact of human activities ($H_{_con} = -1.298/a$).



Figure 11. Distribution of *C*_{_con} and *H*_{_con} along the altitude gradient of the QLB.

4. Discussion

4.1. Three-Dimensional Distribution Pattern of Vegetation Cover

A multi-dimensional assessment of vegetation cover changes in the Qinghai-Tibet Plateau is essential for understanding the sustainability of the livelihoods of the pastoralists. This study explored the changes in vegetation cover in the QLB from the three dimensions of temporal, spatial, and altitudinal aspects and obtained many interesting results. From the temporal dimension, the vegetation coverage of the QLB increased significantly, consistent with the research results of Xiong and Han et al. [21,49], further confirming the greening trend of the QLB vegetation. From the spatial dimension, the vegetation has only been degraded in the northern part of the QLB and a small part of the lakeshore over the past 21 years. Compared with the areas with degraded vegetation, the area of greenery is much larger. Interestingly, the areas with low vegetation coverage in the QLB did not degrade significantly, but showed a greening trend, and the spatial heterogeneity of different vegetation coverage degradation levels also existed in other areas of the Qinghai-Tibet Plateau [50]. The QLB is the same as the Qinghai-Tibet Plateau from the altitudinal dimension. The altitude determines the vegetation distribution pattern [51], which is also the main reason for the heterogeneity of vegetation greening along the altitude gradient [52]. Therefore, analyzing the influence of altitudinal factors on vegetation changes is of great significance to better understand the interaction mechanism of vegetation-climate-human on the Qinghai-Tibet Plateau.

4.2. Methods to Quantitatively Assess Changes in Vegetation Caused by Climate Change and Human Activities

Quantitatively assessing the relative contributions of climate change and human activities from complex long-term changes in vegetation is challenging. Hence, this study selected NDVI as the evaluation index of vegetation coverage. It used multiple regression and partial derivative residual analysis to determine the main factors of vegetation change in alpine regions. We found that using one of the residual methods alone to determine the factors contributing to changes in vegetation cover is insufficient, and the resulting data may "lie". The vegetation coverage of the QLB showed extensive and significant growth from 2000 to 2020. From the partial derivative residual analysis results (Figure 8), human activities did not seem to contribute positively to the greening of the QLB. The residual analysis based on multiple regression shows that the time difference of the intensity of human activities on vegetation coverage is the main reason for this "wrong" result. Yin et al. [53] also confirmed that unreasonable human activities were the main reason for the decline in vegetation cover in the QLB in the first decade of the 21st century. Still, on the whole, the positive impact of human activities on vegetation cover is increasing [28]. Therefore, the combined use of multiple regression and partial derivative residual analysis is superior to the results obtained by using a single method because it can quantify the contribution of influencing factors and monitor the impact of influencing factors on vegetation changes, but also avoid neglecting the impact of human activities on vegetation cover change in temporal and spatial dimensions in research. The effect of human activities on vegetation cover changes over time and space provides a new attempt to guide the sustainable development of the human-land relationship in grassland ecosystems.

4.3. The Impact of Climate Change on Changes in Vegetation Cover

As the "magnifying glass" of global climate change, the effects of *T* and *P* changes on vegetation coverage in the Qinghai-Tibet Plateau have been confirmed in many studies [54,55]. Previous studies have shown that climate warming may alter vegetation phenology, leading to an earlier vegetation growth season and promoting vegetation growth in high-latitude cold and wet areas [56]. However, we found that the increase in *T* of the QLB is limited, and the $T_{_con}$ to the growth in vegetation cover is not significant. $T_{_con}$ was positive only in regions where *T* was higher and became warmer (Figure 4a,c). On the contrary, due to the cold and rainy climate characteristics of the QLB [57], $C_{_con}$ humidification contribution to the increase in vegetation cover is higher, with $P_{_con}$ more than 300 times that of $T_{_con}$. Previous studies also came to the same conclusion that the alpine steppe in the northeastern Qinghai-Tibet Plateau showed a strong response to *P* changes [58,59]; water availability was the main factor limiting vegetation growth of the QLB [44], and *P* increase was the main climatic factor of greening in the QLB [38].

The vegetation coverage in the QLB shows a degradation trend in the area with the lowest altitude (3200–3300 m), and the negative contribution of P(-0.389/a) is much higher than that of T(-0.006/a), which may be due to the high water consumption of vegetation. While precipitation increases but still cannot meet the water demand of vegetation, an increase in temperature instead leads to a decrease in available water for vegetation. Wang et al. [37] showed that the water consumption of the QLB vegetation decreases significantly along the altitude gradient and is higher at 3200–3300 m (more than 300 mm). In addition, it is generally believed that the coastal gradient greening of vegetation will be subject to more low-temperature restrictions, and the hydrothermal conditions required for its growth will also change significantly [60]. However, we found that climate warming caused the limitation of vegetation greening along the altitude gradient of T in the QLB to be narrowed, with vegetation greening displayed above 3300 m above sea level ($NDVI_{slope} > 0$). This phenomenon also exists in other areas of the Qinghai-Tibet Plateau [52,61], mainly due to the increase in T rise in the alpine region, which increases the melting of permafrost, ice, and snow [62].

4.4. The Impact of Human Activities on Changes in Vegetation Cover

As the external driving force for the change in the QLB grassland ecosystem, human activities showed a change process of "continuous-deterioration long-term fluctuationbenign maintenance" over the 21 years (Figure 5a). From 2000 to 2004, the population density of QLB increased rapidly, from 2.97 persons/km² to 6.94 persons/km² (Figure 12). Herders had to raise more livestock to ensure their livelihoods, resulting in the highest number of animals per unit area in the past 21 years (Figure 12). All these may be the reasons for the continuous deterioration of vegetation cover. Wang et al. showed that the population of the QLB increased rapidly before 2005, resulting in unreasonable land use, increased grazing intensity, desertification, environmental pollution, and other serious problems becoming more prominent, which caused the vegetation cover to deteriorate continuously [63]. From 2005 to 2014, the Chinese government successively implemented ecological, environmental protection, and comprehensive management projects, such as the return of grazing land to grassland, wetland protection, degraded grassland management, and seasonal grazing prohibition. The population also declined rapidly (Figure 12), and vegetation degradation began to decrease. However, due to factors such as industrial structure, livelihood habits, and ecological engineering periodicity, the H_{con} was < 0 most of the time. From 2015 to 2020, the continuous decline in population and livestock numbers reduced direct pressure on vegetation. On the other hand, with the income from grassland subsidies, grazing prohibition subsidies, tourism development, etc., the livelihood structure of QLB herders is gradually diversifying, which is also conducive to the continuous greening of vegetation. By 2021, the proportion of the primary industry in the QLB had dropped to 24.62%, while the proportion of the tertiary industry, dominated by eco-tourism, had reached 51.27%, completing the structural transformation of livelihoods in the QLB. In general, once human activities adversely affect vegetation cover, the speed of recovery will be prolonged. This has important implications for further understanding the fragility of the QLB ecosystem.



Figure 12. Variation of population density, livestock density, and GDP in the QLB from 2000 to 2020.

Spatially and altitudinally, the vegetation greening of the QLB under the influence of human activities seems to be fragile. The favorable impact of climate warming and wetting on vegetation greening is almost equivalent to the adverse effects of human activities (C_{con} and H_{con} are 1.62/a and -1.618/a, respectively). In other words, if the intensity of climate warming and wetting is weakened, the vegetation cover of the QLB may deteriorate rapidly, even if the intensity of human activities remains unchanged, which will affect

regional ecological security and herders' livelihoods. Furthermore, our study found that $H_{_con}$ reached the highest value (0.389/a) at 3200–3300 m, while these areas are located around the shore of Qinghai Lake and belong to the core protected areas of the QLB, where ecological engineering and grazing prohibition may be the main reasons for increased vegetation cover. From 3300–3500 m, $H_{_con}$ drops rapidly to reach the minimum (–2.389/a). There are four towns in these areas (Figure 1), two of which are county towns (Gangcha County and Tianjun County), and the population is nearly three times that of low-altitude areas (below 3300 m, Figure 13a). Studies have shown that the rapid development of cities and towns, and the population increase over the past 21 years, are the main reasons for the lowest $H_{_con}$ value [64]. Above 3500 m, the altitude increase in the QLB neither restricts human activities, nor reduces their intensity. For example, in the middle and upper reaches of the Buha River, where population density is low (Figure 13b), the negative contribution of human activities remains high, which is mainly determined by the livelihood patterns of local pastoralists, where pasture and water distribution are the main factors in the choice of grazing location.



Figure 13. Population density over the spatial (**a**) and altitudinal gradients (**b**) of the QLB from 2000 to 2020.

As in other parts of the world, the vegetation of the QLB is facing dual pressures from climate change and human activities [54,65]. However, the factors affecting vegetation cover are multifaceted. This research selected only T and P as the variable climate factors, ignoring other natural factors, such as solar radiation [46], snow-covered areas [66], and the impact of human activities on vegetation cover. Additionally, the impact of the rising lake level of Qinghai Lake on vegetation coverage cannot be ignored. Recent studies have shown that the lake area of Qinghai Lake increased by 156.31 km² from 2000 to 2014, the total length of the shoreline increased by 8.01 km, and the maximum advancing distance of the shoreline reached 2.5 km [67], especially in the east, north, and west areas of the lakeshore, which is essential for the vegetation degradation around Qinghai Lake over the past 21 years (Figure 3b). All of the above need to be further quantified in future research.

5. Conclusions

From 2000 to 2020, the NDVI of the QLB increased significantly (0.002/a, p < 0.01), with 91.38% of the area showing a greening trend, and only 8.62% of the area showing a degrading trend. T_{con} and P_{con} were positive to vegetation greening, and P_{con} was more than 300 times that of T_{con} . C_{con} and H_{con} were 1.62/a and -1.618/a, respectively. There are obvious differences in the distribution and change of vegetation cover along the altitude gradient. Vegetation coverage decreased significantly along the altitude gradient (-0.02/a, p < 0.01), with the highest vegetation coverage at 3600–3700 m (0.37/a). C_{con} , H_{con} , T_{con} ,

and P_{con} were 1.299/a, -1.29/a, 0.003/a, and 1.296/a, respectively. From 3200 to 3300 m, the negative contribution of climate change was the highest (-0.39/a), and the positive contribution of human activities was the highest (0.389/a). At 3400–3500 m, the positive contribution of climate change and the negative contribution of human activities were the highest, 2.392/a and -2.389/a, respectively.

In general, the temporal difference in the intensity of human activities was the main reason for the lower H_con value over the past 21 years. The greening of vegetation in the QLB depends not only on warming and humidification caused by climate change, but also on the increase in the positive impact of human activities after 2010.

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