

## Article

# A Multi-Temporal Analyses of Land Surface Temperature Using Landsat-8 Data and Open Source Software: The Case Study of Modena, Italy

Tommaso Barbieri \* , Francesca Despini and Sergio Teggi 

Department of Engineering “Enzo Ferrari”, University of Modena and Reggio Emilia, Via Vivarelli 10, 41125 Modena, Italy; francesca.despini@unimore.it (F.D.); sergio.teggi@unimore.it (S.T.)

\* Correspondence: tommaso.barbieridief@unimore.it; Tel.: +39-059-205-6295

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**Abstract:** The Urban Heat Island (UHI) phenomenon, namely urban areas where the atmospheric temperature is significantly higher than in the surrounding rural areas, is currently a very well-known topic both in the scientific community and in public debates. Growing urbanization is one of the anthropic causes of UHI. The UHI phenomenon has a negative impact on the life quality of the local population (thermal discomfort, summer thermal shock, etc.), thus investigations and analyses on this topic are really useful and important for correct and sustainable urban planning; this study is included in this context. A multi-temporal analysis was performed in the municipality of Modena (Italy) to identify and estimate the Surface Urban Heat Island (SUHI, strictly correlated to the UHI phenomenon) from 2014 to 2017. For this purpose, Landsat-8 satellite images were processed with Quantum Geographic Information System (QGIS) to obtain the Land Surface Temperature (LST) and the Normalized Difference Vegetation Index (NDVI). For every pixel, LST and NDVI values of three regions of interest (ROI, i.e., Countryside, Suburbs, and City Center) were extracted and their correlations were investigated. A maximum variation of 6.4 °C in the LST values between City Center and Countryside was highlighted, confirming the presence of the SUHI phenomenon even in a medium-sized municipality like Modena. The implemented procedure demonstrates that satellite data are suitable for SUHI identification and estimation, therefore it could be a useful tool for public administration for urban planning policies.

**Keywords:** urban heat island; land surface temperature; remote sensing; Landsat-8; semi-automatic classification plugin; QGIS; global warming; urbanization

## 1. Background

About half of the world population lives in urban areas [1]. The global urbanization rate is expected to increase by 70% compared to the current world population [2], both because of the continued emergence of new urban areas [3] and because of the constant population migration from rural to urban and suburban areas [4,5]. It is not therefore surprising that the negative impacts of urbanization are an ever-growing global concern [6–10]. Urbanization has a negative impact on the environment, mainly due to pollution, changes in the physical and chemical properties of the atmosphere, and in the type of cover of the soil surface [11]. These phenomena lead to so-called Urban Heat Islands (UHI), namely urban areas where the atmospheric temperature is significantly higher than that in the surrounding rural areas [12]. The presence of UHIs is an increasing phenomenon studied by the international scientific community because of its dangerous and significant effects. In fact, the temperature increase has effects on the environment (higher temperatures cause higher energy consumption, photochemical smog, and worsening of the air quality), on the climate and on human health [13–16].

Concerning urban planning policies, nowadays, public administrations have to take into account not only socioeconomic perspectives but also environmental sustainability [17]. Obviously, the UHI phenomenon is also a topic that has to be considered in administration decisions because it negatively influences population life quality. It leads to thermal discomfort and thermal shock, especially during the hot seasons [18,19]. The presence of UHI is typical of highly urbanized areas, but there is currently a progressive increase of this phenomenon not only for metropolises but also for medium-sized municipalities.

The main objective of this study is to investigate the presence, extent, and intensity of Surface Urban Heat Islands (SUHIs) in the municipality of Modena through the use of remote sensing images. Environmental remote sensing is a very useful tool to analyze this phenomenon by calculating the Land Surface Temperature (LST); this is why this study deals with SUHI instead of UHI [20–25]. The trend of the LST and the SUHI intensity are highly correlated—an increase of 0.1 °C of the urban LST compared to the rural LST corresponds to an increase of 0.04 °C of SUHI intensity [26]. This result, however, is retrieved only in the Mediterranean area, due to the specific characteristics of the cities of this region, which may not be found in other parts of the world.

Also, SUHI and UHI phenomena are strictly correlated and in particular the SUHI represents an indirect estimate of the UHI [27,28]. In scientific literature, a variety of studies have combined LST, usually used for the SUHI phenomenon, and air temperature data, usually used for the UHI phenomenon [29–33].

Some preliminary assessments on the study area are mandatory to understand and analyze results. First of all, Modena has different characteristics from the American cities that are usually used to define the UHI phenomenon [34]. Modena is a medium-sized municipality, without a high urban density and with a surrounding countryside with a lot of little villages. Instead, American metropolises usually have a very high urban density and the surrounding areas are empty. For these reasons, the difference of temperatures between cities and the countryside are not as high as in other studies in scientific literature dealing with bigger cities.

Furthermore, Modena has many well-distributed green areas that contribute to reduce the UHI phenomenon and to mitigate the temperatures in the areas that have a high urban density. Indeed, green areas are one of the most common strategies used to mitigate UHIs [17,35–38].

The city of Modena therefore represents a different area from those usually studied for the analysis of the SUHI phenomenon in the existing scientific literature. In the current state of the art, in fact, it is possible to find numerous studies on UHI related to large cities [39] both in America and in Asia [40,41]. In the Mediterranean area there are many studies focused on large cities such as Barcelona [42], Athens [43,44], Salonicco [45], and Tel Aviv [46]. In Italy, numerous studies considered the UHI of Rome or Milan [47–50], while recently other studies have been implemented on cities like Bologna [51], Padua [52], and Venice [53]. Concerning the southern part of Italy, currently there is only one study that reports an analysis of the LST of the major southern cities (Bari, Naples, Palermo, and Catania) [54]. In this study, the LST value of the rural area is not present, thus it is not possible to quantify the SUHI.

For the city of Modena, some considerations of the UHI phenomenon have been carried out in the studies of Bonafè and Zauli Sajani [51,55]. Contrary to the methodology presented in this paper, these studies use ground station measurements for UHI intensity estimation. Furthermore, in Reference [51], not the entire area of Modena is considered but only specific districts.

No existing study in the scientific literature uses remote sensing data on a medium-sized city like Modena, placed in a highly polluted area (Po Valley), to identify and estimate SUHIs. In addition, there are only few studies that use Quantum Geographic Information System (QGIS) procedures (open source software) applied to remote sensing data for these purposes [56].

Another observation concerns the climate that has characterized the study area in the last two years. In this period, various climatic anomalies occurred as a consequence of the global climate change. In particular, winter was very humid with high temperatures and no snow. Also, summer was



quite wet, especially in July and between the end of August and the beginning of September. These weather and climate conditions surely have an influence on the results obtained by remote sensing, because evapotranspiration influences the temperatures of urban and rural areas [57–61].

In this framework it is important to highlight that the municipality of Modena, in 2010, put in place a plan of action for sustainable energy (SEAP). This plan is included in the “good practices” for the European Green Infrastructure Strategies because it provides an increase in green infrastructures in agricultural areas, ecological zones, and public green. In particular, an increase in urban forest of 127.5 ha has been planned for implementation between 2011 and 2020 [62,63].

Thus, using the methodology developed in this paper, the city administrators will have a powerful tool to identify SUHIs and their intensity in different years. In this way, urban policies can be directed to those areas where the SUHI phenomenon is still present and creates discomfort in the population. Furthermore, the public administration could evaluate appropriate mitigation measures such as green roofs, cool materials, and green areas [64,65] that lead to a more sustainable life quality. In addition, using this methodology, it is possible to evaluate whether the policies entailed within the framework of the European Green Infrastructure Strategies are producing the desired results.

## 2. Materials and Methods

The analysis of the SUHI phenomenon in these areas was carried out using an experimental plugin for QGIS [66], called the Semi-Automatic Classification Plugin (SCP) [67], that allows use of the free open source software QGIS as a remote sensing software. With this plugin, temperature maps can be obtained from “raw” remote sensing images acquired by the Landsat-8 satellite sensor.

The study area includes the municipality of Modena located in the Emilia Romagna region, Italy. The environmental context is the Po Valley, which is characterized by a high population density and intensive agriculture. The LST was computed using the equation suggested by Weng [23]. Then, additional processing was conducted in order to create three Regions of Interest (ROI) within each image: the City Center region, the Suburbs region, and the Countryside region. For each ROI, LST maps and Normalized Difference Vegetation Index (NDVI) maps were retrieved. Creating these ROIs also allowed the estimation of the temperature difference between Countryside and City Center, and thus allowed us to correlate it with the SUHI phenomenon in the studied period (2014–2017) [68–70]. Finally, NDVI maps were used to analyze the correlation between land cover and LST maps.

### 2.1. Study Area

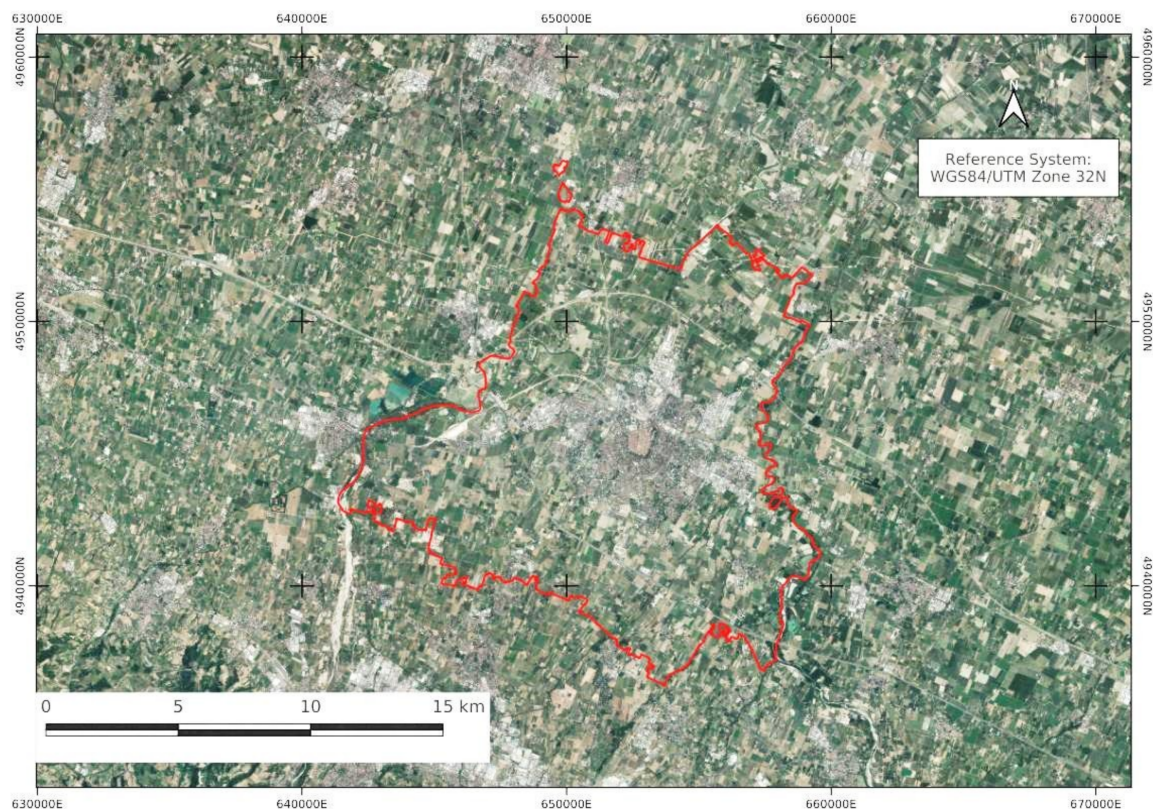
The study area is the municipality of Modena, Italy (Figure 1). Modena has 184,826 inhabitants and an area of 183.19 km<sup>2</sup> [71]. Modena is located in the north of Italy, along the “via Aemilia”, an ancient Roman road running from Rimini to Piacenza on the river Po (Figure 2). The Po Valley is characterized by a high population density and by processes of industrialization and intensive agriculture (it is among the most productive agricultural areas within Europe) [72]. The climate of this region is partially continental: summers are hot with intense heat waves while winters are cold and wet, usually with atmospheric stability conditions and fog. Normally precipitation is concentrated during autumn and spring. Summer and winter are the driest seasons [57,58,72,73]. In order to have a clear climate situation of the study area, Table 1 reports some climate information. The data are divided into two series: an historical series from 1971–2000 and another time series from 2001 to 2017. For each series, Table 1 shows values of seasonal mean temperature, seasonal mean of the maximum daily temperatures, seasonal mean of the minimum daily temperatures, and seasonal mean of precipitation. The data referred to the City Center region were acquired by the Weather Station “Osservatorio Geofisico di Modena”, located in the urban area of Modena (Long: 10°55′47.2″ E, Lat: 44°38′52.9″ N) [74], while the data referred to the Countryside region were acquired by the ARPAE Weather Station of Albareto (Long: 10°57′24″ E, Lat: 44°42′7″ N). In Table 1, all temperature values of the historical series increase in the time series in every season. The increase of the values vary between 1.4 °C and 2 °C. Also, the values of precipitation increase in every season except summer. In Modena,

starting from the 1980s, climate change has exhibited the same trend as the global climate change. The changes cause extreme events such as jumps in temperature or increases in the distribution and amount of precipitation [57].

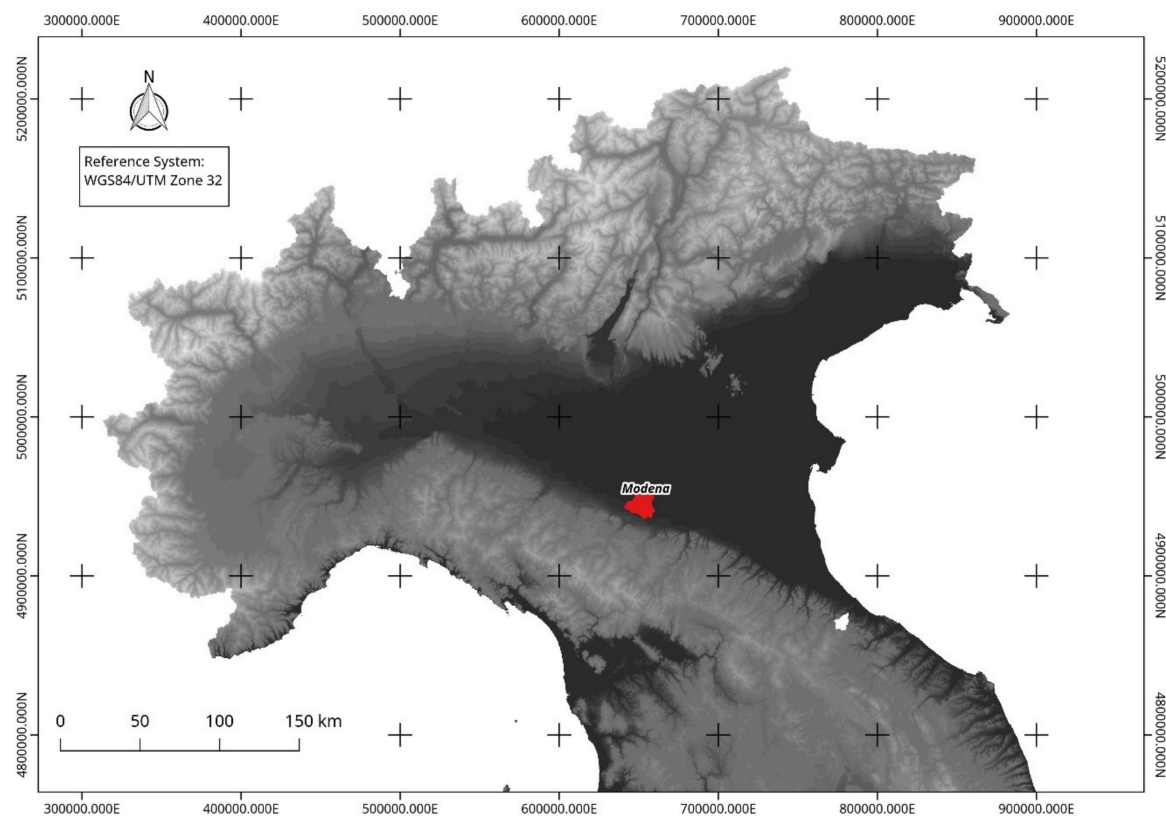
The area of Modena was chosen as a first area of interest for the proposed methodology for its characteristics in terms of:

- Dimensions. Modena is a medium-sized municipality and the scientific literature does not investigate SUHIs of these kind of cities with remote sensing data;
- Climate situation;
- Geographical position. Modena is located in one of the most industrialized areas of Europe, as well as one with the highest levels of air pollutants [75];
- Active urban policies. Analyzing three years of data allowed us to investigate whether urban policies on green infrastructures are effective or not.

These characteristics render Modena an ideal test site for the developed methodology.



**Figure 1.** Ortho image of the city from an aerial photograph of Modena provided by the cartographic archive of the Emilia Romagna region (AGEA 2011). the municipality borders are shown in red (Source: Geodatabase Emilia Romagna region—<http://geoportale.regione.emilia-romagna.it>, image elaborated by the authors).



**Figure 2.** Digital Terrain Model (DTM) of Italy with the location of the city of Modena. (Source: The Ministry of the Environment and Protection of Land and Sea of Italy—[www.pcn.minambiente.it/mattm/](http://www.pcn.minambiente.it/mattm/), image elaborated by the authors).

**Table 1.** Seasonal mean temperature, seasonal mean of the maximum daily temperatures, seasonal mean of the minimum daily temperatures, and seasonal mean of precipitation of Modena between 1971 and 2000 (historical series) and between 2001 and 2017 (Source: Osservatorio Geofisico di Modena).

HISTORICAL SERIES FROM 1971 TO 2000				
Season	$T_M$ (°C)	$TMAX_M$ (°C)	$TMIN_M$ (°C)	$Prec_M$ (mm)
Spring	17.5	21.7	13.3	53.5
Summer	23.1	27.3	18.8	49.5
Autumn	8.8	11.2	6.3	62.2
Winter	5.8	8.7	2.9	40.6
TIME SERIES FROM 2001 TO 2017				
Season	$T_M$ (°C)	$TMAX_M$ (°C)	$TMIN_M$ (°C)	$Prec_M$ (mm)
Spring	19.5	23.7	15.3	60.5
Summer	24.5	28.7	20.3	45.9
Autumn	10.4	12.9	8.0	67.5
Winter	7.2	10.1	4.2	53.8

$T_M$ : seasonal mean temperature;  $TMAX_M$ : seasonal mean of the maximum daily temperatures;  $TMIN_M$ : seasonal mean of the minimum daily temperatures;  $Prec_M$ : seasonal mean of precipitation.

## 2.2. Landsat-8 Data

The Landsat-8 satellite was successfully launched on 11 February 2013 and deployed into orbit with two instruments on board: (1) the Operational Land Imager (OLI), whose spatial resolution is 30 m, with nine bands in the visible (VIS), the near infrared (NIR), and the short-wave infrared (SWIR) spectral regions; and (2) the Thermal Infrared Sensor (TIRS) with two spectral bands in the

long-wave infrared (LWIR) region. The spatial resolution of TIRS data is 100 m with a revisit time of 16 days [75–79] (Table 2).

**Table 2.** Operational Land Imager (OLI) and Thermal Infrared Sensor (TIRS) (Landsat-8) spectral bands.

Spectral Band	Wavelength ( $\mu\text{m}$ )	Spatial Resolution (m)
Band 1—Coastal/Aerosol	0.435–0.451	30
Band 2—Blue	0.452–0.512	30
Band 3—Green	0.533–0.590	30
Band 4—Red	0.636–0.673	30
Band 5—NIR	0.851–0.879	30
Band 6—SWIR	1.566–1.651	30
Band 7—SWIR	2.107–2.294	30
Band 8—Panchromatic	0.503–0.676	15
Band 9—Cirrus	1.363–1.384	30
Band 10—LWIR	10.60–11.19	100 (resampled to 30)
Band 11—LWIR	11.50–12.51	100 (resampled to 30)

The data used for this study are the “raw” remote sensing images acquired from the Landsat-8 satellite sensors, provided free of charge by the United States Geological Survey (USGS) (Table 3).

**Table 3.** Image dataset used in this study.

Acquisition Date	Day of the Week	Season	Acquisition Time—Coordinated Universal Time (UTC)
26 January 2014	Sunday	Cold	10.00
7 April 2014	Monday	Mid	10.00
6 August 2014	Wednesday	Hot	10.00
25 October 2014	Saturday	Mid	10.00
12 December 2014	Friday	Cold	10.00
15 January 2015	Wednesday	Hot	10.00
10 September 2015	Thursday	Mid	10.00
16 January 2016	Saturday	Cold	10.00
21 April 2016	Thursday	Mid	10.00
26 July 2016	Tuesday	Hot	10.00
5 October 2016	Wednesday	Mid	10.00
9 January 2017	Monday	Cold	10.00
17 May 2017	Wednesday	Hot	10.00
4 July 2017	Tuesday	Hot	10.00
22 September 2017	Friday	Mid	10.00

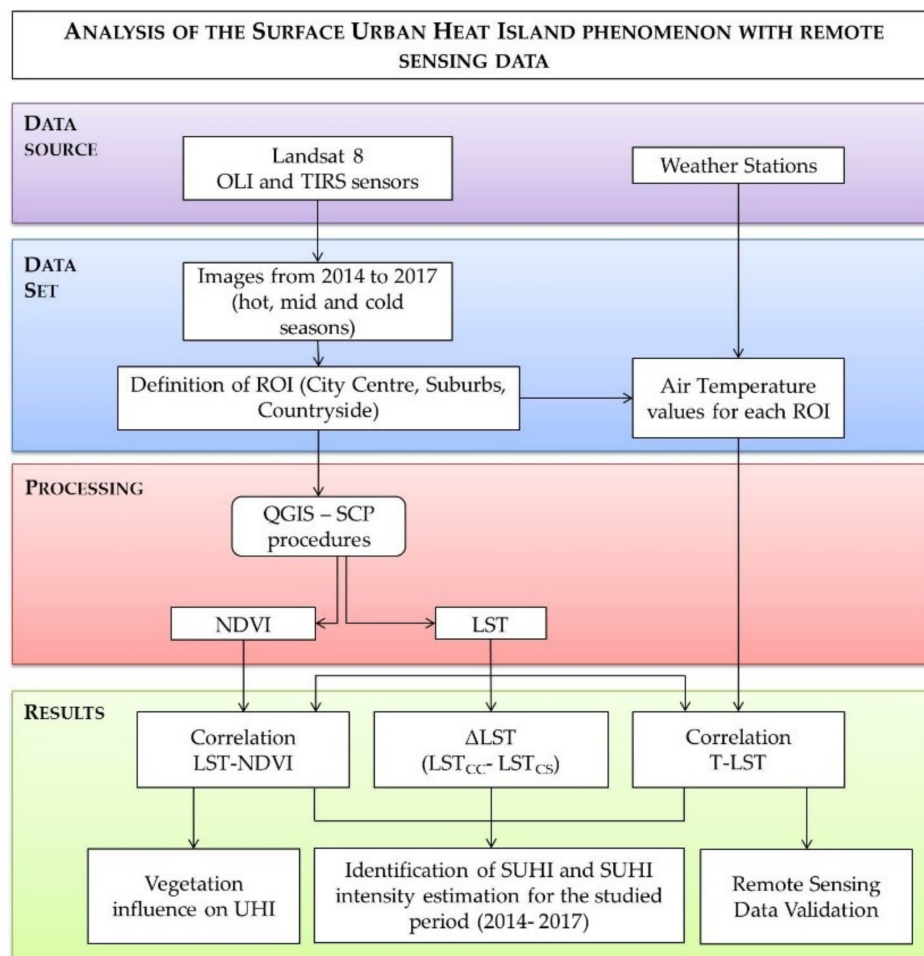
Images were divided into three seasons depending on mean LST values retrieved by satellite data: hot season ( $\text{LST} > 30\text{ }^{\circ}\text{C}$ ); mid season ( $10\text{ }^{\circ}\text{C} < \text{LST} < 30\text{ }^{\circ}\text{C}$ ); cold season ( $\text{LST} < 10\text{ }^{\circ}\text{C}$ ).

As reported in Table 3, there are 15 daily images covering about the hot, mid, and cold seasons of four years. The scene center time is about 10:00 UTC. Only images with a clear-sky condition were selected.

There are only two images from 2015 because of the cloud cover, but winter 2015 can be represented by the image taken on 12 December 2014. Autumn 2015 can be represented by the image taken on 10 September 2015, which presents typical autumnal climatic conditions. Therefore, the year 2015 can be considered completely covered, except for spring.

For the entire procedure, the Semi-Automatic Classification Plugin (SCP) on the open source software QGIS was used. The SCP plugin allows one to use remote sensing functions in QGIS [22,24,67]. The methodology, from input data to obtained results, is shown in the conceptual diagram in Figure 3. Each step is described in detail in the following subsections.





**Figure 3.** Conceptual diagram of the proposed methodology.

### 2.3. Data Processing

#### a. Conversion to Top of Atmosphere (TOA) Radiance

OLI and TIRS band data can be converted to top of atmosphere (TOA) spectral radiance using radiance rescaling factors provided in the metadata file [80]:

$$L_{\lambda} = M_L \times Q_{cal} + A_L \quad (1)$$

where:

- $L_{\lambda}$  = top of atmosphere spectral radiance (TOA) ( $W/(m^2 \text{ srad } \mu m)$ ).
- $M_L$  = band-specific multiplicative rescaling factor from the image metadata ( $W/(m^2 \text{ srad } \mu m)$ ).
- $A_L$  = band-specific additive rescaling factor from the image metadata ( $W/(m^2 \text{ srad } \mu m)$ ).
- $Q_{cal}$  = digital number (DN).

#### b. Conversion to TOA Reflectance

OLI band data were converted to TOA planetary reflectance using reflectance rescaling coefficients provided in the image metadata files [80]. The following equation is used to convert DN values to TOA reflectance for OLI data:

$$\rho_{\lambda} = M_{\rho} \times Q_{cal} + A_{\rho} \quad (2)$$



where:

- $\rho\lambda' =$  TOA reflectance.
- $M_p =$  band-specific multiplicative rescaling factor from the metadata.
- $A_p =$  band-specific additive rescaling factor from the metadata.
- $Q_{ca} =$  digital number.

c. Conversion to At-Satellite Brightness Temperature

TIRS band data were converted from spectral radiance to brightness temperature using the thermal constants provided in the image the metadata file [80]:

$$T = K_2 / \ln ((K_1 / L_\lambda) + 1) \quad (3)$$

where:

- $T =$  at-satellite brightness temperature (K).
- $L_\lambda =$  TOA spectral radiance ( $W/(m^2 \text{ srad } \mu m)$ ).
- $K_1 =$  band-specific thermal conversion constant from the metadata ( $W/(m^2 \text{ srad } \mu m)$ ).
- $K_2 =$  band-specific thermal conversion constant from the metadata (K).

d. Clip of the image using the shapefile of Modena municipality, provided for free by the cartographic archive of the Emilia Romagna region (official website “Geoportale dell’Emilia Romagna”)

e. Calculation of Normalized Different Vegetation Index (NDVI)

This index is computed using spectral reflectance in the near infrared band and in the red band. It provides a rapid estimation of the presence of vegetation. NDVI values ranges from  $-1$  to  $1$ . Higher NDVI values indicate dense vegetation while lower values (typically from  $0$  to  $0.2$ ) identify light or dark soils [24]. NDVI is computed using the following equation:

$$NDVI = (\rho_{NIR} - \rho_{Red}) / (\rho_{NIR} + \rho_{Red}) \quad (4)$$

where:

- $\rho_{NIR} =$  near infrared band reflectance.
- $\rho_{RED} =$  red band reflectance.

f. Fractional Vegetation Cover (FVC)

NDVI was used to calculate FVC, an index that estimates the proportion of an area covered by a set of predefined type of vegetation or soil cover [24].

The following equation is used to calculate FVC:

$$FVC = (NDVI - NDVI_s) / (NDVI_v - NDVI_s) \quad (5)$$

where:

- $NDVI_s =$  NDVI value for bare soils.
- $NDVI_v =$  NDVI value for fully vegetated soils.

The  $NDVI_s$  value was set equal to  $0.1$ , while the  $NDVI_v$  value was set equal to  $0.65$  [22].

g. Calculation of Land Surface Emissivity (LSE), which is a measurement of the capacity of a material to radiate energy [24]

$$\varepsilon = (\varepsilon_s \times (1 - FVC)) + (\varepsilon_v \times FVC) \quad (6)$$

where:

- $\varepsilon_s$  = typical soil emissivity (0.93).
- $\varepsilon_v$  = typical vegetation emissivity (0.98).

#### h. Estimation of LST

LST can be calculated from the at-satellite brightness temperature  $T_B$  [23]. LST was first computed in K, then converted to °C.

$$LST = T_B / (1 + (\lambda \times s \times T_B / c^2) \times \ln(\varepsilon)) \quad (7)$$

where:

- $\lambda$  = wavelength of the emitted radiance (11.5  $\mu\text{m}$  for band 10 in Landsat-8 OLI).
- $c^2 = h \times c / s = 1.4388 \times 10^{-2} \text{ m K}$ .
- $s$  = Boltzmann constant =  $1.38 \times 10^{-23} \text{ J/K}$ .
- $h$  = Planck's constant =  $6.626 \times 10^{-34} \text{ Js}$ .
- $c$  = velocity of light =  $2.998 \times 10^8 \text{ m/s}$ .
- $\varepsilon$  = emissivity.

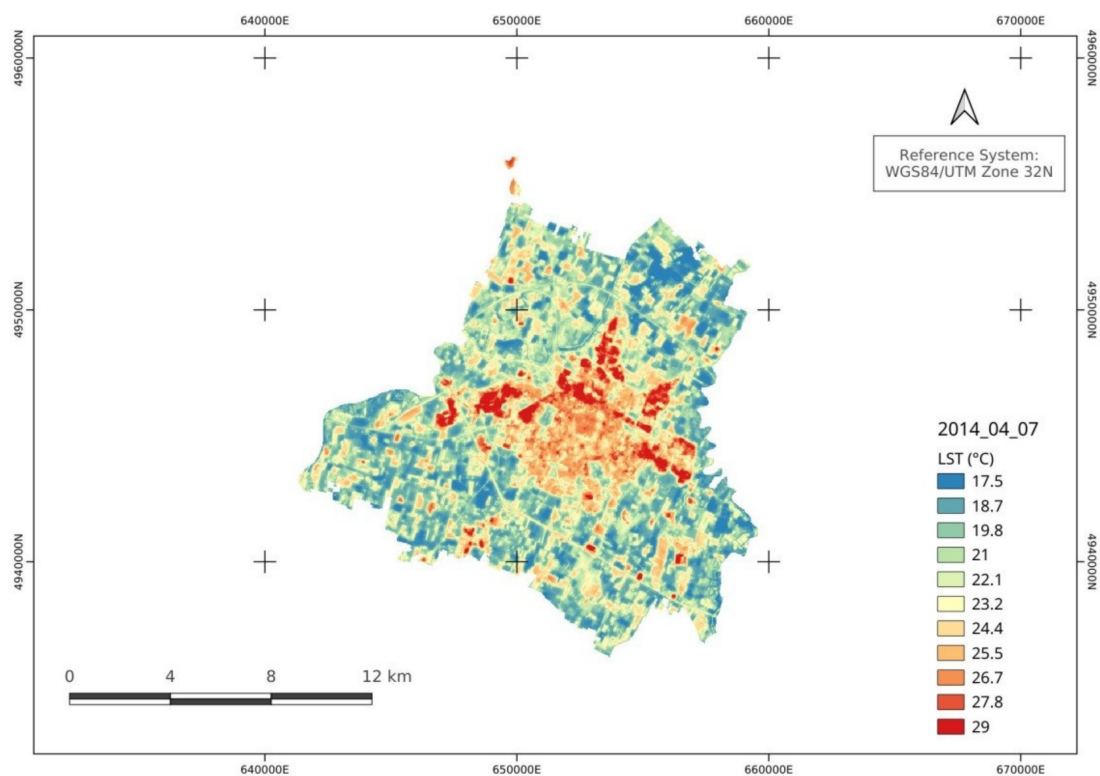
### 3. Results and Discussion

#### 3.1. LST Maps

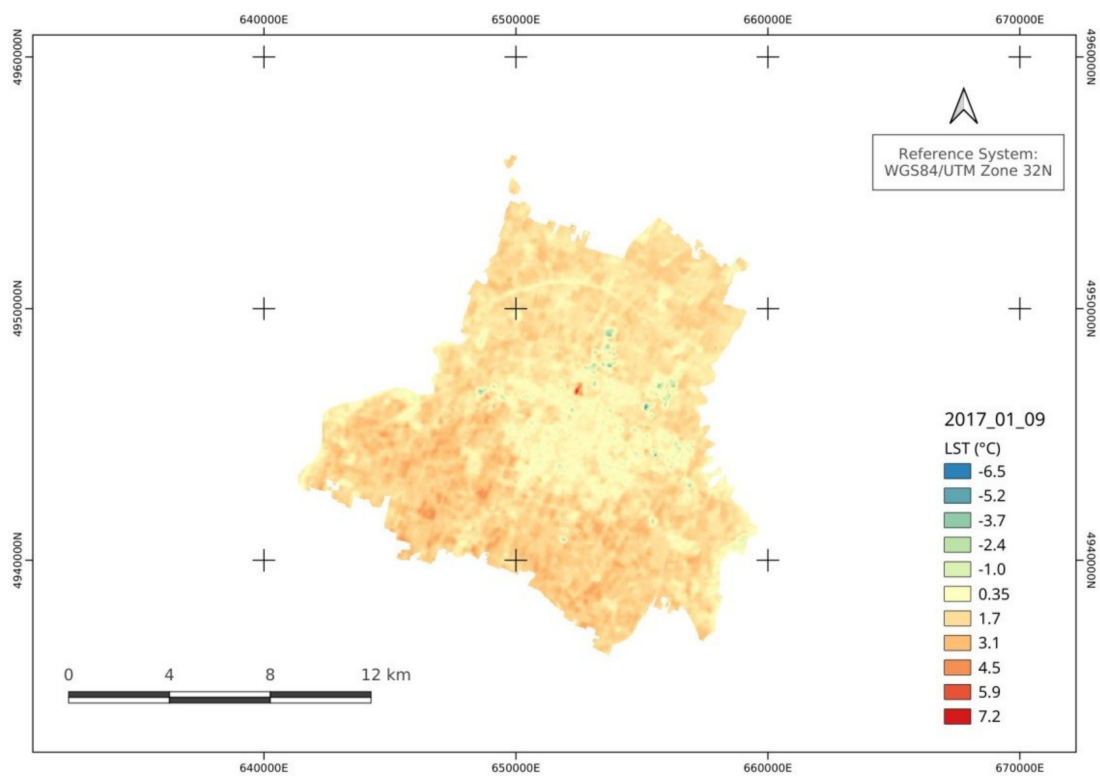
The methodology was applied to the dataset in order to obtain LST maps of Modena for different climate seasons. In this section only a few significant images are shown of the 15 processed. The statistical results on the next pages were, however, retrieved using all of the images listed in Table 3.

Figure 4 shows the LST map of the study area on 7 April 2014. The minimum value is 17.5 °C, the maximum value is 29 °C. Thanks to the spatial resolution of the TIRS sensor, inside the city in the industrial areas (LST values between 28 °C and 29 °C) it is possible to identify dense areas (most with LST values between 25.5 °C and 27.8 °C) and vegetation areas (LST values between 17.5 °C to 22 °C). Also, small residential areas and bare soils (LST values from 23 °C to 29 °C) could be identified. Furthermore, main roads are clearly visible (LST values of about 23 °C). In the main urban area, LSTs are on average higher than in the countryside (the maximum temperature difference between the urban area and rural area is about 11 °C), confirming the presence of the SUHI phenomenon. In the following section, the difference between the urban LST and the rural LST will be further analyzed.

In winter, the LST maps show a more homogeneous distribution of temperatures and sometimes the urban LST values (from 0 °C to 2 °C) are slightly lower than the rural LST values (highly variable values that average between 0 °C and 5 °C). The map in Figure 5, for example, shows the study area on 9 January 2017. The situation is significantly different from that of Figure 4. LST values are very similar both in urban and in rural zones, thus in this period the SUHI phenomenon is not visible from this kind of data. In accordance with the scientific literature [61,81], during the winter period an inversion of temperature trends could occur (urban areas have lower temperatures than rural areas), especially in medium-sized municipalities like Modena. In order to explain this map, an investigation of the presence of water in the study area is required. When humidity levels are very high, the water evaporation and the plant transpiration saturate the atmosphere, thus countryside and city temperatures tend to flatten, out becoming much more similar to each other [61,81]. For this reason, in humid periods the SUHI is barely identifiable using remote sensing techniques, which however does not mean that the SUHI does not exist. For example, a temperature increase of just one or two degrees with humidity levels exceeding 90% does not allow the definition of the UHI phenomenon on an urban scale, but the population feels health discomforts nonetheless [82].



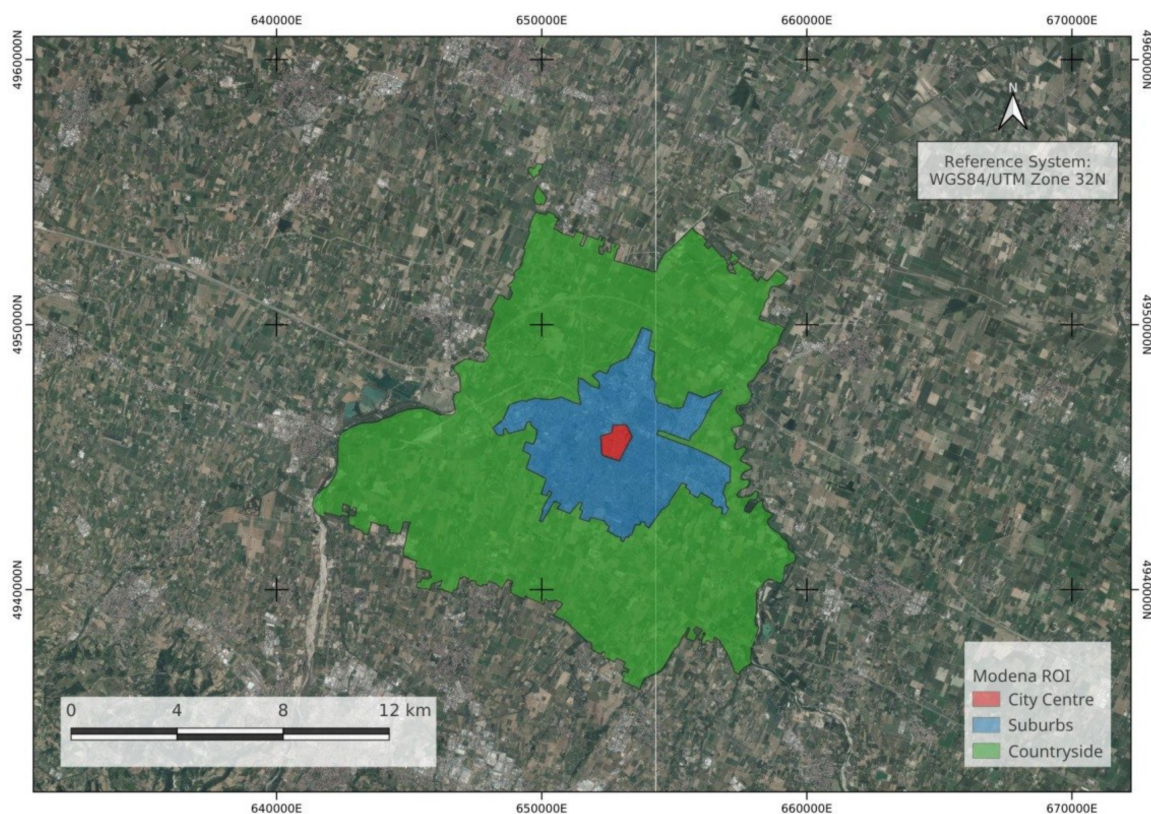
**Figure 4.** Land Surface Temperature (LST) map of Modena, 7 April 2014. (Source: Quantum Geographic Information System (QGIS) elaboration of LST map provided by the authors).



**Figure 5.** LST map of Modena, 9 January 2017. (Source: QGIS elaboration of LST map provided by the authors).

### 3.2. Region of Interest

To analyze temperature variations between urban areas and rural areas, three ROIs were created. A shapefile built thanks to municipality information was overlaid on each image in order to divide the municipality of Modena in three areas: City Center, Suburbs, and Countryside (Figure 6). These ROIs correspond to three specific Local Climate Zones (LCZs) defined by Stewart and Oke as “regions of uniform surfaces cover, structure, material, and human activity that span hundreds of meters to several kilometers in horizontal scale” [83]. City Center corresponds to LCZ1 (Compact Midrise) characterized by a building surface fraction between 40% and 70%. Suburbs corresponds to LCZ5 (Open Midrise) characterized by a building surface fraction between 20% and 40%. Countryside corresponds to LCZD (Low Plants) characterized by a building surface fraction less than 10% [83]. The building surface fraction represents the ratio of building plan area to total plan area (%). For Modena’s ROIs, the building surface fraction is equal to 46% for City Center, 20% for Suburbs, and 2% for Countryside. LCZs are defined in a univocal way based on specific parameters. The link between ROIs and LCZs allows the repeatability of the present methodology as well as comparison with other study areas.



**Figure 6.** Region of Interest (ROI) map of Modena. ROIs shapefile source: municipality of Modena. (Source: Geodatabase Emilia Romagna region—<http://geoportale.regione.emilia-romagna.it>, image elaborated by the authors).

Table 4 shows the average LST value of each ROI for every image and the difference ( $\Delta$ LST) between the City Center LST values ( $LST_{CC}$ ) and the Countryside LST values ( $LST_{CS}$ ). These values allow to have a first evaluation of the intensity of the SUHI phenomenon. Suburbs LST values are not considered for this estimation because they are often very similar to City Center LST values.



**Table 4.** Average LST value of each ROI for the processed images.

Date	Season	LST Mean (°C)			$\Delta\text{LST (LST}_{\text{CC}} - \text{LST}_{\text{CS}})$
		City Center	Suburbs	Countryside	
26 January 2014	Cold	7.6	7.0	5.9	1.7
7 April 2014	Mid	26.1	25.5	21.0	5.1
6 August 2014	Hot	34.6	33.1	28.2	6.4
25 October 2014	Mid	19.9	19.0	17.9	2.0
12 December 2014	Cold	8.1	7.4	6.2	1.9
15 January 2015	Hot	36.3	35.4	31.4	5.2
10 September 2015	Mid	25.4	25.6	25.2	0.2
16 January 2016	Cold	4.8	4.8	4.8	0.0
21 April 2016	Mid	21.7	21.7	19.2	2.5
26 July 2016	Hot	30.5	30.3	27.4	3.1
5 October 2016	Mid	21.2	21.5	21.5	−0.3
9 January 2017	Cold	0.7	0.7	1.8	−1.1
17 May 2017	Mid	30.3	30.2	26.6	3.7
4 July 2017	Hot	32.0	32.6	30.2	1.8
22 September 2017	Mid	22.4	23.4	22.0	0.4

Looking at Table 4, it can be observed that the images of cold seasons are usually associated with low LST variations ( $\leq 2$  °C). Accordingly, with Figure 4 the SUHI phenomenon is not visible in these seasons with Landsat-8 data and thus it is not possible to correlate SUHIs with the UHI phenomenon.

Otherwise, images of hot seasons (spring and summer) are mostly associated with high LST variations ( $> 2$  °C). In this case it is possible to clearly identify the SUHI phenomenon and estimate its intensity. In Table 4 it is possible to observe the maximum and the minimum LST variation in the three seasons of interest. For the hot season, this value reached 6.4 °C, confirming the presence of the SUHI phenomenon and thus of the UHI phenomenon.

An exception of this analysis is represented by the image acquired on 9 January 2017. LST values for different regions are very close and the  $\Delta\text{LST}$  between City Center and Countryside is lower than 2 °C. This anomaly is due to particular weather conditions of the days prior to the image acquisition; 2 and 3 July 2017 were in fact cloudy days with temperatures below the usual seasonal means.

Figure 7 shows a graphical representation of mean LST values for each acquired image and for each ROI—during hot seasons, the SUHI phenomenon is clearly visible.

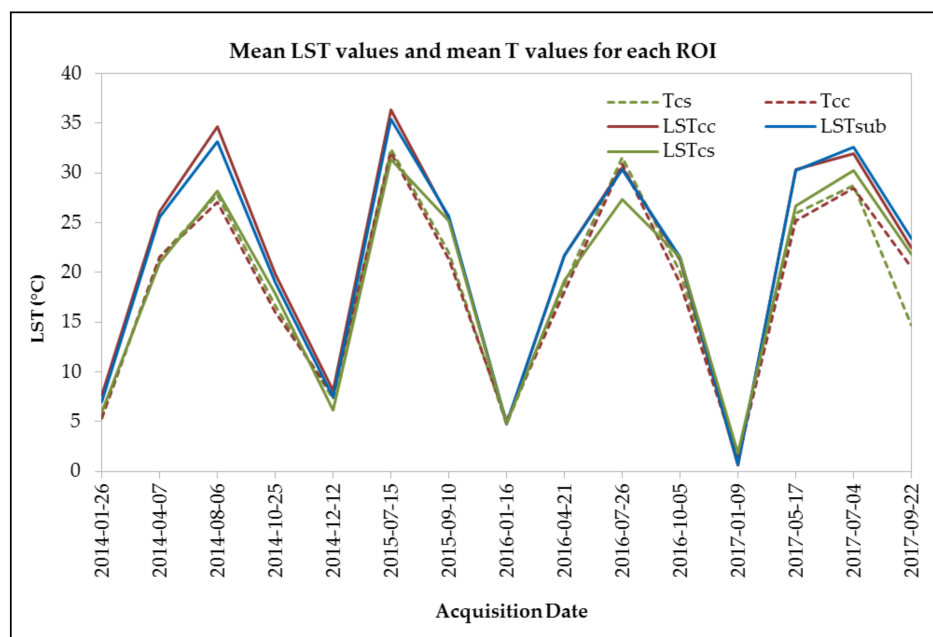
For example, between April 2014 and October 2014, the difference between City Center LST values and Countryside LST values is clear, ranging from 3 °C and 5 °C. Similarly, from April 2015 to September 2015, the LST variation ranges from 1 °C to 5 °C (July 2015). From April 2016 to October 2016 and also from April 2017 to September 2017, this difference in LST values is not so remarkable, with  $\Delta\text{LST}$  values ranging from 1 °C to 3 °C.

During cold seasons, the temperature lines in Figure 7 are almost completely overlapped, showing no SUHI presence. Thus, from October 2014 to March 2015, from October 2015 to April 2016, and from October 2016 to April 2017, LST values are very close for all three regions of interest.

In general, the LST peaks are higher in 2014–2015 and lower in 2016–2017. Winter 2017 had very low LST values compared to the other winters.

The results highlighted in Table 5 and visually shown in Figure 7 could be compared to scientific studies on UHI even if the present study deals with SUHI. In Asian large cities, the UHI intensity ranges from 0.4 to 11 °C. These values are largely influenced by the monsoon circulation [40]. In Europe, in particular in the Mediterranean area, it is possible to find UHI intensity values from 3 °C to 7 °C for cities like Athens and Salonicco [44,45]. Focusing on small cities, two interesting studies show the presence of UHI even in cities with less than 100,000 inhabitants [84,85]. These studies revealed a mean summer UHI intensity from 2.6 to 8 °C.





**Figure 7.** Mean LST values for each ROI from 2014 to 2017 compared with air temperature values for City Center ( $T_{CC}$ ) and Countryside ( $T_{CS}$ ). (Source: Data elaboration provided by the authors).

**Table 5.** Minimum  $\Delta LST$  and maximum  $\Delta LST$  divided by seasons.

Season	$\Delta LST_{\min} - \Delta LST_{\max}$ (°C)
Hot	(6.4)–(1.8)
Mid	(−0.3)–(5.1)
Cold	(−1.1)–(1.9)

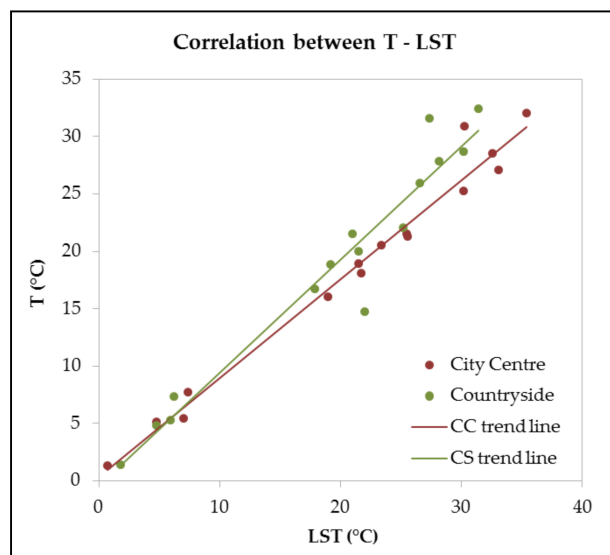
Table 6 shows the UHI intensity retrieved for Italian cities in the scientific literature during the hot seasons. Even if these values are related to UHI and not to SUHI and these cities exhibit some differences from Modena (size, position, and inhabitants), it is possible to notice that the results obtained from this study are comparable with the values in Table 6. In particular, Zauli Sajani et al. [51] analyzed a large area that includes Bologna as well as Modena: the UHI intensity values measured by the meteorological station during the summer period are coherent with the results obtained from the methodology presented in this paper.

**Table 6.** Summer Urban Heat Island (UHI) intensity for major Italian cities.

City	Methodology for UHI Intensity Estimation	Summer UHI Intensity (°C)	References
Milan	Remote sensing data (Moderate Resolution Imaging Spectroradiometer—MODIS sensor)	Maximum of 8–10 °C	Anniballe et al., 2014 [49]
Rome	Temperature data collected by meteorological stations/remote sensing data	2.5–7.5 °C	Guattari et al., 2018 [50] Fabrizi et al., 2010 [86]
Bologna-Modena	Temperature data collected by meteorological stations	4–8 °C	Zauli Sajani et al., 2016 [51]
Padua	Temperature data collected by meteorological stations	2–6 °C	Busato et al., 2014 [52]
Venice	Temperature data collected by meteorological stations	4–7 °C	Peron et al., 2015 [53]

In Figure 7, the air temperature values are also reported. In particular, air temperature data for City Center and Countryside were retrieved from two weather stations provided by ARPAE Emilia Romagna, TCC and TCS, respectively.

These air temperature values are highly correlated to satellite LST values as it is possible to observe both in Figures 7 and 8. The Pearson coefficients ( $r$ ) shown in Figure 8 and in Table 7 have values closer to 1, which means there is an excellent correlation between LST values and air temperature values.



**Figure 8.** Scatterplot showing the correlation between air temperature (T) and LST considering the whole dataset. (Source: Data elaboration provided by the authors).

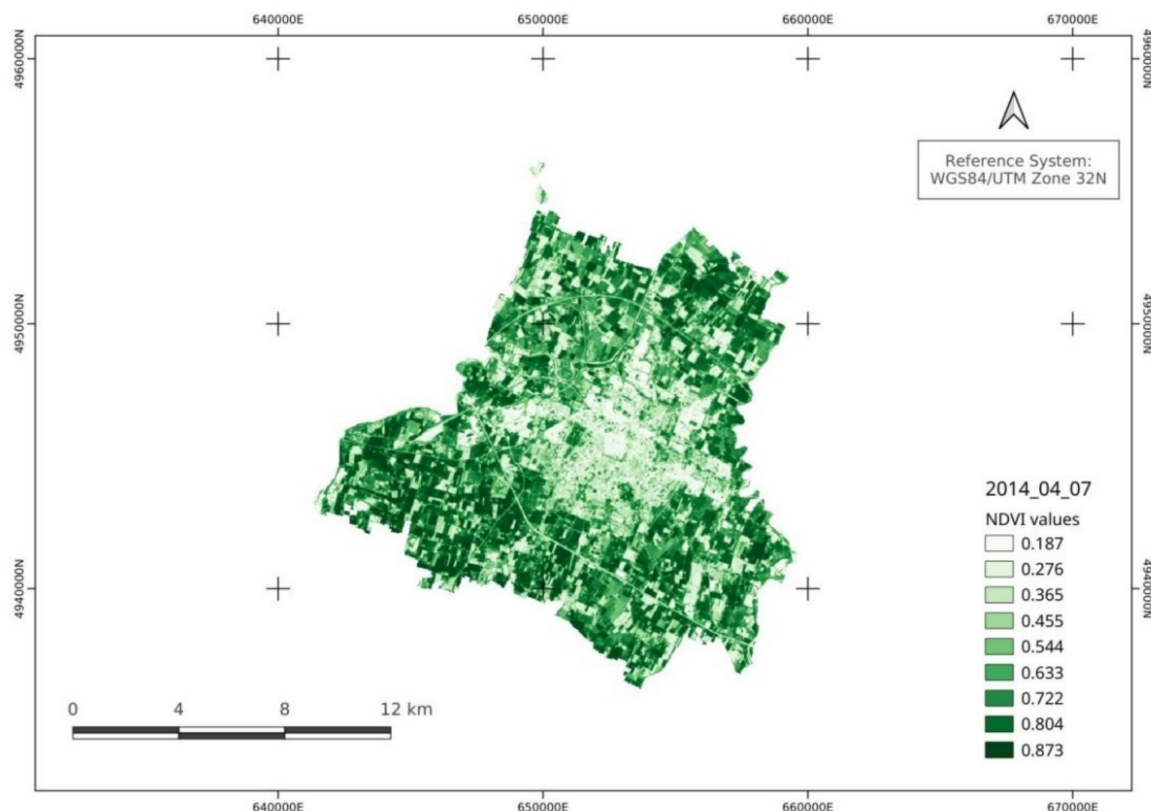
**Table 7.** Pearson coefficient of LST vs. T for each ROI.

ROI	Pearson Coefficient $r$
City Center	0.98
Countryside	0.97

The results shown in Figure 8 are very important because they prove the effectiveness and reliability of satellite measurements for the study of SUHIs and consequently of UHIs. The high correlation with air temperatures measured by weather stations demonstrates that the methodology is repeatable and can be transferred to other areas similar to the city of Modena (medium-sized municipalities). In urban areas with different dimensions (large cities or small municipalities), the methodology can still be applied, only requiring the foresight to choose carefully the ROIs. In particular, the number and size of these regions must be appropriate to the study area. In this framework, the reference to LCZs [83] allows, for cities of different dimensions, the identification of the correct ROI number and size to study the SUHI phenomenon. Weather stations must also be chosen accurately as they are representative of the investigated ROI. Possible influences from traffic or emissions from industrial activities must be avoided if possible.

### 3.3. NDVI Maps

NDVI maps were retrieved in order to partially identify surface land cover and correlate it with LST values [87]. Figure 9 represents the NDVI map of the “Modena 7 April 2014” Landsat image. High values of NDVI (0.6 to 0.9) suggest a high density of vegetation, and low values of NDVI (from 0 to 0.3) suggest the presence of bare soil or artificial surfaces.



**Figure 9.** Normalized Different Vegetation Index (NDVI) map of Modena, 7 April 2014. (Source: QGIS elaboration of NDVI map provided by the authors).

Obviously, urban areas present low NDVI values and rural areas have high NDVI values. From the NDVI map presented in Figure 9, it is possible to recognize built-up areas (or bare soil areas) and vegetation areas. NDVI maps are consistent with scientific literature data [28,88]. As reported by Yuan et al. [15], lower LSTs usually are found in areas with high NDVI. This negative correlation between NDVI and LST is valuable for urban climate studies. However, NDVI measurements are subject to seasonal variations, which may influence the results of SUHI studies.

### 3.4. LST-NDVI Correlation

For every pixel of the images from 7 April 2014, 15 July 2015, and 9 January 2017, the LST and NDVI values were extracted to make a correlation analysis.

Focusing on the image from 7 April 2014 and the image from 15 July 2015, Figure 10 shows that high values of NDVI correspond to low values of LST and vice versa. Thus, the presence of vegetation corresponds to lower LST values.

Regarding the 9 January 2017 image, it is difficult to make observations by just looking at the maps; it is necessary to analyze Table 8. This table reports the Pearson correlation coefficient between LST and NDVI computed for each ROI in the three images analyzed in Figure 10 (“Modena 7 April 2014”, “Modena 15 July 2015”, and “Modena 9 January 2017”).

For the images of April 2014 and July 2015,  $r$  values are close to 1, suggesting a strong correlation between NDVI and LST. Instead, for the image of January 2017, the  $r$  value is close to 0, thus a poor correlation between NDVI and LST is highlighted.

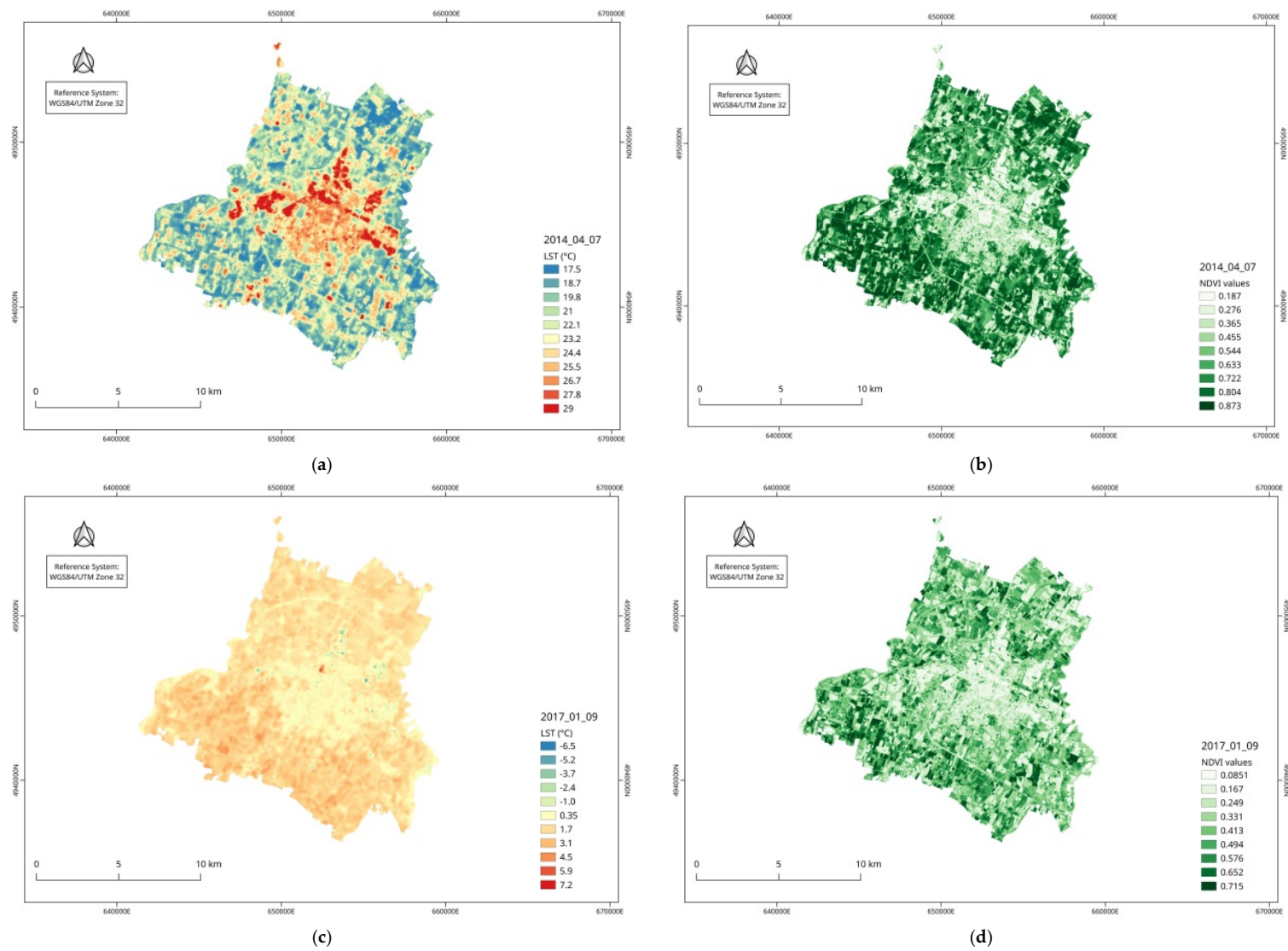
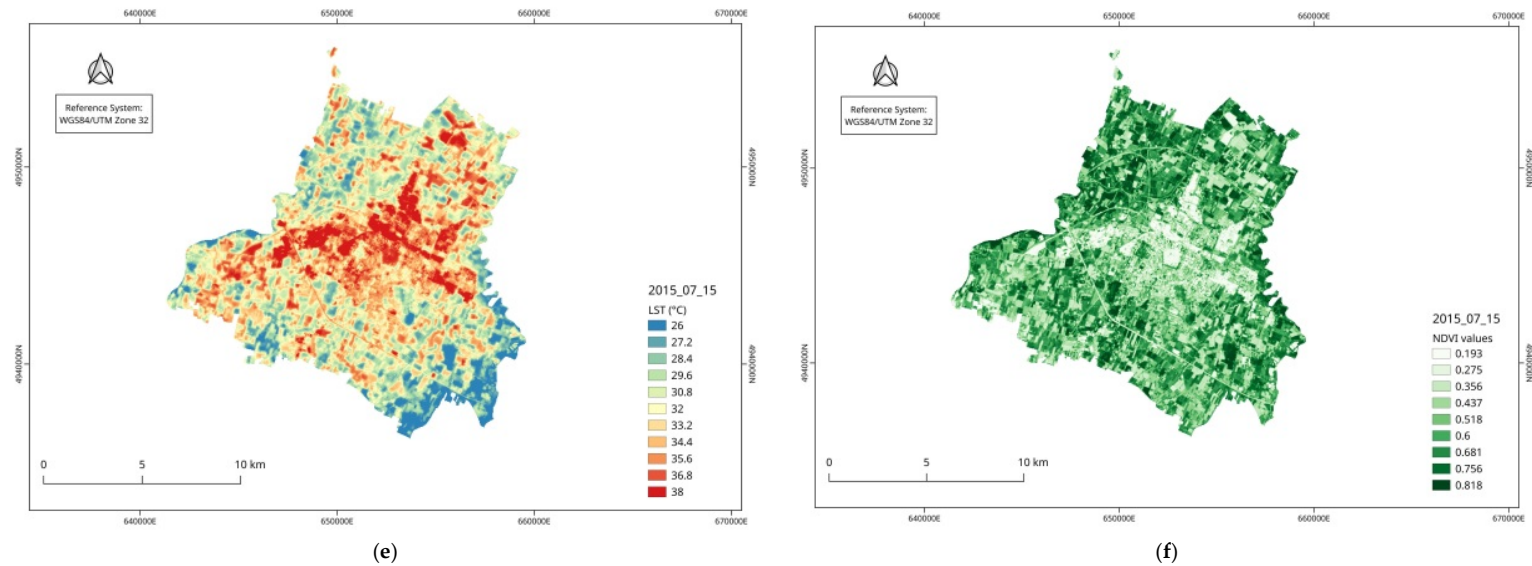


Figure 10. Cont.



**Figure 10.** (a) LST map of Modena, 7 April 2014. (b) NDVI map of Modena, 7 April 2014. (c) LST map of Modena, 9 January 2017. (d) NDVI map of Modena, 9 January 2017. (e) LST map of Modena, 15 July 2015. (f) NDVI map of Modena, 15 July 2015. (Source: QGIS elaboration of LST and NDVI maps provided by the authors).



Compared to “Modena 7 April 2014” and to “Modena 15 July 2015”, the January 2017 image has two main differences: the low LST-NDVI correlation and lower LST values in the city Center. This is probably due to the difference in thermal inertia of the materials mainly present in each ROI. When, the night before the image acquisition day, the air temperature is really low and there is a clear sky condition, the surface of the city becomes very cold. Thus, on the image acquisition day, more time is needed by the urban surfaces (mainly concrete roofs and asphalt) to increase their temperature compared to the countryside [59–61]. The countryside has more vegetation and bare soils compared to the city, therefore the temperature increase is faster here than in urban areas.

**Table 8.** Pearson correlation coefficient  $r$  between LST and NDVI for each ROI of Modena within different images (“Modena 7 April 2014”, “Modena 15 July 2015” and “Modena 9 January 2017”).

Date	ROI	Pearson Coefficient $r$
7 April 2014	City Center	0.92
	Suburbs	0.85
	Countryside	0.88
15 July 2015	City Center	0.91
	Suburbs	0.79
	Countryside	0.77
9 January 2017	City Center	0.14
	Suburbs	0.20
	Countryside	0.06

#### 4. Conclusions

In this work the presence, extension, and intensity of the SUHI in the municipality of Modena were studied using Landsat-8 data.

The results showed that with satellite data, the SUHI is clearly visible during hot and mid seasons, while an opposite phenomenon is observed during cold seasons.

In particular, during hot seasons, Modena records a significant SUHI phenomenon, with a difference between City Center LST values and Countryside LST values of up to 6.4 °C.

Observing the entire studied period (from 2014 to 2017), the following assessments can be made:

- Mean LST values are on average higher from 2014 to 2015 than from 2016 to 2017;
- The SUHI phenomenon is more evident during hot seasons, especially in the years 2014 and 2015.

These assessments are certainly consistent with the policies implemented by the municipality of Modena in recent years concerning European Green Infrastructures Strategies. Thus, the application of the methodology presented in this paper highlights the mitigation effect on the SUHI (and therefore on the UHI) of the Modena SEAP agreement.

NDVI maps were useful to identify the vegetation cover and to correlate it with LST maps. This study confirmed that, for hot and mid seasons, high values of LST correspond to low values of NDVI (bare soil or artificial covers). For example, for the image of 15 July 2015, the Pearson coefficient ( $r$ ) shows a correlation between LST values and NDVI values equal to 0.91 for City Center, 0.79 for Suburbs, and 0.77 for Countryside. For the mid season, the correlation between NDVI and LST is still strong. For example, for the image from 7 April 2014, the Pearson coefficient values are between 0.85 and 0.92. Otherwise, during cold seasons, this study revealed a null or opposite trend of the phenomenon and a low correlation between NDVI values and LST values.

Remote sensing LST values were also compared with air temperature values measured by two weather stations (one for the City Center region and one for the Countryside region). These values showed a high correlation with the Pearson coefficient equal to 0.98 for City Center and to 0.97 for Countryside. Remote sensing data are therefore representative of real surface temperatures and thus they can be used for the identification of the SUHI and the estimation of its intensity.

In conclusion, this study showed that the SUHI is a phenomenon that can also be observed in medium-sized municipalities like Modena and that can be investigated with a free and accessible methodology. The proposed methodology could be easily used not only for the identification of SUHI but also for its intensity estimation thanks to the good correlation between LST values and air temperature values. This study could thus represent a powerful tool for public administration for sustainable planning policies based on UHI mitigation strategies. Using the proposed methodology for several images acquired in different years allows one to monitor the UHI mitigation actions and verify their effect on the territory.

Furthermore, the proposed methodology could be used not only for cities similar to Modena but also for different kind of cities, from large metropolises to small municipalities. For the transferability of the methodology, the choice of ROIs is important: each ROI has to be compared with LCZs definitions in order to set universal benchmarks to compare obtained results.

This study is not concluded. More images will be analyzed, in particular during hot seasons, and, for each image, the Albedo parameter will be calculated to correlate the SUHI to surface reflectance. Moreover, complete meteorological data will be collected in order to better understand the SUHI phenomenon, taking the local climate into consideration. Additionally, data such as the daily precipitation of the three days before the image acquisition time, the average daily wind speed, and the main wind direction will provide a complete climatic and environmental characterization of the acquisition time for each image.

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## Abbreviations

ARPAE	Agenzia Regionale Prevenzione Ambiente ed Energia
DN	Digital Number
FVC	Fractional Vegetation Cover
LCZ	Local Climate Zone
LSE	Land Surface Emissivity
LST	Land Surface Temperature
LWIR	Long-Wave Infrared
MODIS	Moderate Resolution Imaging Spectroradiometer
NDVI	Normalized Difference Vegetation Index
NIR	Near Infrared
OLI	Operational Land Imager
QGIS	Quantum Geographic Information System
ROI	Region of Interest
SCP	Semi-Automatic Classification Plugin
SEAP	Plan of Action for Sustainable Energy
SUHI	Surface Urban Heat Island
SWIR	Short-Wave Infrared
TIRS	Thermal Infrared Sensor
TOA	Top of Atmosphere
UHI	Urban Heat Island
USGS	United States Geological Survey
UTC	Coordinated Universal Time
VIS	Visible

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