



# Article A New Multi-Scale Method to Evaluate the Porosity and MICP Curve for Digital Rock of Complex Reservoir

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Abstract: The evaluation of rock porosity and the mercury injection capillary pressure (MICP) curve is fundamental for oil and gas exploration and production. Digital rock (DR) technology, incorporating 3D micro-CT imaging and numerical methods, has been widely employed to predict these properties. However, analyzing the pore structure of heterogeneous rocks, such as fractured rocks or glutenite, solely through single-scale DR analysis poses challenges. Existing upscaling methods have limitations in fully representing the complete range of pore structures at different scales, with limited comparison to experimental data. To address this, we propose a novel method that upscales porosity and simulates the MICP curve from nano-scale to core scale by merging results from micro-CT (at resolutions of 35 µm and  $2 \mu m$ ) and SEM (at resolutions of 6.5 nm and 65 nm). We validate the developed DR model by applying it to sandstones, glutenite, and igneous rocks, and achieve excellent agreement between the experimental data and the multi-scale DR model across 67 samples. The results demonstrate that the multi-scale model effectively captures the porosity and pore structures across the entire range. In contrast, the single digital rock (DR) model underestimates the porosity measurements for both homogeneous sandstones and heterogeneous cores. While the MICP model based on a single DR proves suitable for homogeneous rock samples, it introduces noticeable discrepancies when applied to heterogeneous rock samples. The developed multi-scale method significantly enhances the confidence in using DR to assess the pore structure of complex rocks.

**Keywords:** porosity; pore distribution; upscale modeling; mercury injection curve; digital rock; heterogeneous

## 1. Introduction

The characterization of porosity and pore throat distribution is crucial for reservoir classification and reserve evaluation [1–4]. Various techniques have been developed to assess the pore structure, including qualitative and quantitative methods [5–10]. Techniques such as casting thin section and scanning electron microscopy (SEM) have been used for qualitative analysis of pore throat size and structure, while high-pressure mercury intrusion porosimetry (MICP) and nuclear magnetic resonance (NMR) have been used for quantitative analysis. However, traditional methods for the characterization of porosity and pore throat distribution have certain limitations, including destructive nature, lack of spatial resolution, sampling bias, and limited resolution range [11,12].



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**Copyright:** © 2023 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). With advances in micro-CT technology and numerical methods, digital rock (DR) analysis based on 3D images has become an important tool for predicting petrophysical properties of rock [13–19]. Compared to traditional laboratory experiments, micro-CT combined with numerical simulations can provide quicker and cost-effective results [20–23]. However, traditional single-scale DR analysis is limited in its ability to capture the unresolved nano-sized pores, as well as the centimeter-sized intergranular pores, which limits its accuracy in assessing pore structure in heterogeneous reservoirs [24]. There are several studies on the effect of unresolved pores and how to predict petrophysical properties through upscaling methods. Inadequate digital rock upscaling can result in substantial consequences in field applications. For example, it can lead to the inaccurate depiction of pore structures, directly affecting subsequent tasks such as reserve evaluation and reservoir simulation. Additionally, improper prediction of physical properties, which is essential for evaluating reservoir quality and optimizing strategies for oil recovery, can also be influenced by poor digital rock upscaling [25–31].

To overcome the trade-off, one strategy is to combine imaging data from different sources. For instance, the integration of 3D micro-CT images at the micron scale and 2D scanning electron microscopy (SEM) images at the nano-scale can be employed. Traditionally, stochastic methods that rely on spatial statistical information, such as two-point correlation functions and multiple-point statistics, have been used to tackle this challenging problem. Okabe and Blunt [32] reconstructed the 3D pore space structure by integrating micro-CT images (capturing large pores) with statistically simulated high-resolution images from 2D thin sections (offering finer-scale features). Mohebi et al. [33] proposed a statistical approach to fuse low-resolution measurements with a high-resolution prior model, aiming to improve the overall resolution of the reconstructed images. Furthermore, Tahmasebi et al. [34] introduced a multi-scale and multi-resolution reconstruction method for generating 3D models of shales based on 2D images. However, these methods can be computationally expensive, and the generated results may not always capture the full range of possible pore structures at the core scale [35–39].

Existing upscaling methods exhibit limitations in adequately depicting the full spectrum of pore structures at various scales, while also having limited comparison with experimental data. These challenges hinder the effective utilization of digital rock techniques for assessing the intricate pore structure of complex rock formations. In this study, we used a multi-scale approach to evaluate porosity and the simulated MICP curve using two different micro-CT resolutions ( $35 \mu m$  and  $2 \mu m$ ) and two SEM resolutions (65 nm and 6.5 nm) to generate the DR model. Sixty-seven cores including sandstone, glutenite, and igneous rock from the eastern South China Sea were used, and the results obtained from the DR analysis were compared against laboratory measurements. We analyzed the differences between the two and corrected the simulated MICP curve obtained from the DR analysis to be consistent with laboratory results for practical application. This multi-scale approach provides a more accurate and comprehensive assessment of the pore structure of heterogeneous reservoirs.

#### 2. Methods

#### 2.1. Image Acquisition Procedures

The procedure for evaluating the porosity and MICP curve at the multi-scale is as follows: (1) Real 3D geometric structures are generated by micro-CT scanning for the core plug, with a resolution of 35  $\mu$ m. (2) Micro-samples with a size of 2~3 mm are drilled from the core plug and scanned by micro-CT with a resolution of 2  $\mu$ m. (3) The micro-samples are then scanned by SEM at two different magnifications (500× and 5000×), corresponding to the resolution of 65 nm and 6.5 nm, respectively. (4) The porosity and the pore throat distribution are calculated at each scale and then are combined and upscaled to the core plug scale. (5) The laboratory measurements for porosity and high-pressure mercury injection are performed to compare with the DR method. (6) The differences are analyzed, and then the corrected DR model for porosity and simulated MICP curve is obtained.

In order to better understand the pore structure characteristics of the core plugs, the traditional single-scale DR is not sufficient to capture the nano-pores due to the resolution of  $1 \sim 3 \mu m$ , as well as the large intergranular pores due to the physical size of  $1 \sim 3 mm$ . Thus, scanning of the whole core plug at a resolution of  $35 \mu m$  is conducted first to capture the intergranular pores, and then three micro-samples are drilled from the core plug and scanned by micro-CT with a resolution of  $2 \mu m$ . In order to capture the pores smaller than  $2 \mu m$ , SEM is used to describe the porosity and the pore structure characteristics at a finer scale. The sample size is  $5 mm \times 5 mm$ , and two magnifications,  $500 \times$  and  $5000 \times$ , are selected at three different locations, respectively, in order to capture the heterogeneity. The SEM results are then used as a supplement to the micro-CT results.

Micro-computed tomography (CT) refers to a non-destructive imaging technique that allows researchers to obtain high-resolution 3D images of objects. The micro-CT instrument used in this study is nanoVoxel-3502E, manufactured by QingNengSanYing in Tianjin, China, one of the leading manufacturers of micro-CT systems in China. The micro-CT system is equipped with a high-resolution detector and a high-precision X-ray tube that allows it to achieve high-quality imaging of objects. Its advanced data acquisition system enables the acquisition of images at a high speed with high spatial resolution, making it ideal for studying the micro-structure of objects. The view field of the equipment is a 1920  $\times$  1920 grid, and the scanner is equipped with a micro-focus ray source up to 300 KV. The micro-CT instrument consists of a radiation source, a sample table, and a signal receiver. The signal attenuated by the tested sample after the radiation source emits X-rays is received by the receiver to obtain the projected image of the sample. Images are taken from different angles, and reconstruction is carried out to obtain the 3D structures, with a resolution up to 0.5  $\mu$ m.

The TESCAN MIRA3 XMH (TESCAN, Brno, Czech Republic) is a latest generation scanning electron microscope (SEM) designed for high-performance imaging and analysis of various types of samples, which is used in this study to obtain the porosity and pore distribution at the nano-scale.

## 2.2. Image Segmentation Method

X-ray CT scanning technology uses a synchrotron to generate high-energy X-rays, and different components have different absorption of X-rays. The gray value of each pixel on the CT image is represented by the CT number, which is related to the linear attenuation coefficient of the corresponding minerals [40,41]. Image-processing methods are performed, including three-dimensional reconstruction, image denoising, and intelligent image segmentation, and a series of petrophysical properties are then calculated based on the segmented 3D model.

In this study, we use Weka for two-phase image segmentation. Weka is a highly beneficial tool when it comes to two-phase image segmentation, as it significantly reduces the need for human intervention. By leveraging its powerful algorithms and machine-learning capabilities, Weka automates the process of dividing images into two distinct phases, eliminating the manual effort required for this task. This automation not only saves time and resources but also enhances accuracy and consistency in image segmentation, making Weka an invaluable asset in various image analysis applications [42,43]. The Weka procedure for image segmentation involves preparing image data, extracting relevant features, generating labeled training data, training the segmentation model, evaluating the model's performance, performing image segmentation on new images, and post-processing the segmented images. This automated process uses machine-learning algorithms to accurately segment images and provides consistent results with reduced manual intervention. For example, the 2D gray image of the core obtained by CT scanning is shown in Figure 1a, where gray represents rock matrix and dark areas represent pores. The CT image segmented by the machine-learning algorithm is shown in Figure 1b, where the pores are represented by blue. The machine-learning algorithm is also used for two-phase SEM segmentations to obtain the corresponding porosity.



**Figure 1.** Images before and after segmentation by machine-learning method. (**a**) Original image obtained by micro-CT, (**b**) image segmented by the machine-learning algorithm (blue represents pores and gray represents rock matrix).

#### 2.3. Calculation of the Porosity at Multi-Scale

Due to the limitation of the resolution and its physical size, it is difficult to capture the pore information of the core plug using the traditional single-scale DR analysis, so a new upscale method is established to capture the unresolved pores by integrating the porosity at four scales with resolutions of 6.5 nm, 65 nm using SEM, and 2  $\mu$ m and 35  $\mu$ m using micro-CT, respectively. We first scan the whole core with a diameter of 1 inch and a length of 2 inches, using micro-CT at a resolution of 35  $\mu$ m in order to capture its entire information. After that, representative sub-samples with a physical size of 2~3 mm are taken from the core for further analysis using micro-CT at a resolution of 2  $\mu$ m. Subsequently, micro-samples with a size of 0.13~0.15 mm and nano-samples with a size of 13~15  $\mu$ m from the sub-samples are taken for SEM analysis at 65 nm and 6.5 nm, respectively. At each scale, we obtain the porosity by summation of the pore pixels through the segmentation and assume the rock pixels contain the unresolved porosity which can be obtained from the finer-scale analysis. We can then obtain the total core porosity by combining the four scales:

$$\begin{aligned} \varphi_{nano} &= \varphi_{65 \text{ nm}} + (1 - \varphi_{65 \text{ nm}}) \times \varphi_{6.5 \text{ nm}} \\ \varphi_{micro} &= \varphi_{2 \mu m} + (1 - \varphi_{2 \mu m}) \times \varphi_{nano} \\ \varphi_{total} &= \varphi_{35 \mu m} + (1 - \varphi_{35 \mu m}) \times \varphi_{micro} \end{aligned}$$
(1)

where  $\varphi_{nano}$  is defined as the nano-porosity where the pore size is below 2 µm, by summation of the porosity for SEM at 65 nm and 6.5 nm resolutions,  $\varphi_{micro}$  is defined as the micro-porosity where the pore size is below 35 µm, by summation of the porosity for SEM and micro-CT with 2 µm resolution, and  $\varphi_{total}$  is defined as the total DR porosity for the core plug. To prevent the redundancy of pore volume calculations, for each scale, if the pore size is larger than the resolution of the coarser scale, the pore size is less than or equal to the resolution of the coarser scale will it be considered and contribute to the overall pore volume calculation. This approach avoids multiple calculations of the same pore, and we then update Equation (1) to:

$$\varphi_{nano} = f_{2} \mu_{m} \varphi_{65 nm} + (1 - \varphi_{65 nm}) \times \varphi_{6.5 nm}$$

$$\varphi_{micro} = f_{35} \mu_{m} \varphi_{2} \mu_{m} + (1 - \varphi_{2} \mu_{m}) \times \varphi_{nano}$$

$$\varphi_{total} = \varphi_{35} \mu_{m} + (1 - \varphi_{35} \mu_{m}) \times \varphi_{micro}$$
(2)

where  $f_{2\mu m}$  is the ratio of the pore volume greater than 2 µm to the total pore volume at the resolution of 65 nm, and  $f_{35\mu m}$  is the ratio of the pore volume greater than 35 µm to the total pore volume at the resolution of 2 µm.

We performed a study aimed at evaluating the influence of the representative elemental volume (REV) on porosity. The objective was to determine the appropriate number of

sub-samples, micro-samples, and nano-samples to obtain a comprehensive representation of pore information during micro-CT and SEM analysis of a core sample. Figure 2 depicts a comparative analysis between the porosity values obtained using varying numbers of samples at each scale and the experimental porosity. The relative error quantifies the average deviation for the respective formations (consisting of 29 glutenite, 33 sandstone, and 5 igneous rocks).



(c)

**Figure 2.** Comparison of DR porosity to experimental porosity at various scales. (**a**) Impact of number of sub-samples on porosity, CT at 35  $\mu$ m, (**b**) impact of number of micro-samples on porosity, SEM at 65 nm, (**c**) impact of number of nano-samples on porosity, SEM at 6.5 nm.

Our findings indicate that increasing the number of samples for each corresponding scale leads to a decrease in relative error. This can be attributed to the REV effect, which ensures a more accurate representation of the overall pore characteristics. In the case

of homogeneous cores like sandstone, the impact of the number of samples is relatively small across all scales. However, for glutenite cores, the number of sub-samples has a significant influence due to the heterogeneity at the pore scale. Similarly, for igneous cores, the numbers of both micro-samples and nano-samples play a significant role due to the heterogeneity at the nano-scale.

Based on our study, we selected three REV samples for each corresponding scale, ensuring that the relative error remains below 15%.

#### 2.4. Calculation of the Pore Structure and the Simulated MICP Curve

After obtaining 2D and 3D rock images, the pore morphology model is used to extract the pore network topology and pore throat distribution for CT and SEM images [37]. The pore morphology model is advantageous for extracting pore network topology and pore throat distribution from CT and SEM images due to its ability to accurately represent complex pore structures, handle various types of pores, and provide quantitative analysis of pore throats. It is computationally efficient, enabling analysis of large image datasets, and aligns with the non-destructive nature of CT and SEM imaging techniques. More complete descriptions of the basic concepts and techniques can be found elsewhere [44-46]. The model is established by two steps: Building a maximal inscribed sphere list and assembling the spheres into chains. In the first step, each void point is taken as the starting point and the center of an expanding sphere, where the sphere grows until it touches a rock or a boundary point, becoming an inscribed sphere. Then, the list of all the inscribed spheres is sorted in descending order of their radii. For each of the inscribed spheres, one locally searches for smaller or equal-sized inscribed spheres overlapped with it, namely, its neighbors, and removes them completely, including ones from the list. The remaining spheres consist of the maximal inscribed sphere list. In the second step, an auxiliary index pair (node, rank) is built for each maximal inscribed sphere. The maximal inscribed sphere list in descending order of their radii is traversed again to find all the chains between neighboring maximal inscribed spheres. Hence, traverse process is able to assemble all maximal inscribed spheres into different chains, with end node spheres as pores and middle spheres as throats. Finally, according to the morphological information of the original pore space, the pores and throats are simplified as spherical and pipe objects. After the division of pore throats, parameters such as pore size distribution, average pore throat ratio, and average pore radius can be obtained by calculating the link length and inscribed sphere radius.

The relationship between the capillary pressure of reservoir rocks and the saturation of the wetting phase is called the MICP curve [47]. After the pore distribution is obtained using the pore morphology model at various scales, the simulated MICP curve can then be derived using the Young–Laplace equation:

$$p_c = \frac{2\sigma \cos\theta}{r_c} \tag{3}$$

where  $p_c$  is the capillary pressure,  $\sigma$  is the surface tension,  $\theta$  is the wetting contact angle, and  $r_c$  is the pore radius.  $\sigma = 0.48$  N/m and  $\theta = 140^\circ$  are obtained from the laboratory under ambient conditions:

$$p_c = \frac{0.735}{r_c} \tag{4}$$

Then, the MICP curve can be obtained as follows:

- 1. Pore size distribution: Based on the segmented pore spaces for micro-CT and SEM images, the frequencies or percentages of pores at different size ranges present in the rock sample can be combined together.
- 2. Pore radius assignment: Each identified pore is assigned a specific radius to provide the necessary input for Equation (4).
- 3. Young–Laplace equation: Equation (4) is then employed to relate the capillary pressure  $(\Delta P)$  to the radius of curvature (R).

- 4. Capillary pressure calculation: Using the assigned pore radii and Equation (4), the capillary pressure for each pore is calculated to quantify the pressure difference for mercury injection.
- Mercury saturation determination: The calculated capillary pressure values are then used to determine the corresponding mercury saturation to represent the ratio of the mercury-filled pore volume to the total pore volume.

#### 2.5. Laboratory Measurement of the Porosity and MICP Curve

A comprehensive laboratory measurement is conducted for determining the porosity of core plugs using the water saturation technique. This technique is widely employed in the fields of rock physics and lithology analysis. The 67 plugs used in this experiment are cylindrical, with a diameter of 1 inch and a length of 2 inches. To ensure accuracy and repeatability, a detailed description of the experimental procedure is as follows:

- 1. Core selection and preparation: Representative core plugs are chosen from extracted rock samples, ensuring removal of impurities and uneven sections.
- Sample drying: Selected core plugs are placed in a temperature-controlled oven at 60 °C until a constant weight is achieved.
- 3. Core size measurement: Digital calipers are used to accurately measure the diameter and length of the plugs, with the results recorded.
- 4. Core saturation: Dry plugs are placed in a sealed container and saturated with deionized water to remove any residual air or gas.
- 5. Mass measurement in saturated state: The saturated plugs are removed and accurately weighed to determine their mass, with the results recorded.
- 6. Porosity calculation: The percentage of porosity is calculated based on the dry and saturated masses of the plugs.

By following the aforementioned experimental procedure and maintaining strict control over laboratory conditions, the porosity of core plugs can be accurately measured, providing reliable data for comparison to the DR porosity at the core scale.

The laboratory measurement of the mercury injection capillary pressure (MICP) curve is conducted to characterize the pore structure. The core plugs used in the experiment were cylindrical in shape with a diameter of 1 inch and a length of 2 inches as well. Prior to the measurements, the core plugs were cleaned thoroughly using a series of organic solvents to remove any contaminants that could potentially interfere with the measurements. The experiment was performed under ambient temperature and pressure conditions to mimic the subsurface reservoir conditions. The core plugs were placed in a high-pressure vessel, and mercury intrusion was performed using a state-of-the-art MICP apparatus, which allows precise control of the pressure and monitoring of the intrusion of mercury into the rock's pore space. The MICP curve was obtained by measuring the pressure at which mercury intruded into the sample versus the volume of mercury that intruded. The entire process was conducted meticulously to ensure accurate and reliable results. The data obtained from the MICP measurements were further analyzed to quantify the pore size distribution using Equation (4).

#### 3. Results and Discussion

### 3.1. Evaluation of the Porosity at Multi-Scale

Sixty-seven core plugs with a diameter of 1 inch and a length of 2 inches were used to validate the multi-scale DR method, including sandstone (quartz sandstone and feldspathic sandstone), glutenite, and igneous rock from the eastern South China Sea. The depositional environment in this region is primarily influenced by various factors including sea level changes, sediment supply, tectonic activity, and paleoclimate conditions. Extensive research has revealed that the reservoirs in the South China Sea consist of a variety of lithologies such as sandstone, mudstone, glutenite, and igneous rocks. The reservoirs in the South China Sea are commonly associated with clastic sedimentation, indicating a predominantly terrestrial input of sediments. However, the influence of marine processes, such as currents

and waves, is also evident in certain areas. Depositional environments vary across the region, ranging from fluvial-deltaic systems near river mouths to deepwater turbidite systems in the basin floor. These variations in depositional environments contribute to the heterogeneity of the reservoirs, affecting their petrophysical properties and hydrocarbon distribution. Figure 3 shows the photos of the samples for the representative core plugs of sandstone, glutenite, and igneous rocks, respectively.



Figure 3. Appearance of the samples: (a) Sandstone; (b) Glutenite; (c) Igneous.

The experimental porosity is measured and compared with the DR porosity at the same size. Figure 4 shows the CT and SEM images for sandstone, glutenite, and igneous rocks at four scales, respectively.



**Figure 4.** Comparison of CT and SEM images at four scales. (**1a**) Sandstone, CT at 35 μm; (**1b**) Sandstone, CT at 2 μm; (**1c**) Sandstone, SEM at 65 nm; (**1d**) Sandstone, SEM at 6.5 nm; (**2a**) Glutenite, CT at 35 μm; (**2b**) Glutenite, CT at 2 μm; (**2c**) Glutenite, SEM at 65 nm; (**2d**) Glutenite, SEM at 6.5 nm; (**3a**) Igneous, CT at 35 μm; (**3b**) Igneous, CT at 2 μm; (**3c**) Igneous, SEM at 65 nm; (**3d**) Igneous, SEM at 6.5 nm.

Figure 5a shows the comparison between the laboratory porosity and traditional DR porosity at a resolution of 2  $\mu$ m ( $\varphi_{micro}$ , the third column in Table 1), for three different reservoir rocks, sandstone, glutenite, and igneous. The red line represents where the DR porosity is equal to the experimental porosity, and the black dashed line represents a margin of error of  $\pm 2\%$  in porosity. We can clearly observe that the traditional single-scale DR estimates the porosity lower than the laboratory, especially for glutenite. This is because the unresolved pores with a resolution less than 2 µm cannot be captured, and the physical dimensions of the samples are only 1.6 mm  $\times$  1.6 mm  $\times$  1.6 mm (800  $\times$  800  $\times$  800 voxels) which fails to capture the large pores greater than 1.6 mm. The  $\varphi_{total}$  calculated by the multi-scale method agrees well with the experimental result with an R<sup>2</sup> of 0.92 as shown in Figure 5b, which verifies the reliability of the upscale method ( $\varphi_{total}$ , the fourth column in Table 1). Table 1 summarizes the porosity from the laboratory experiments and the multi-scale DR. There are very few instances where the measured porosity deviates significantly from the experimental values. This is primarily due to the fact that the REV does not accurately represent sub-samples, micro-samples, and nano-samples. To capture the heterogeneity of the REV, additional samples are required.



**Figure 5.** Comparison of porosity between DR and laboratory measurements: Red line means the DR porosity is equal to the experimental results, and black line represents a margin of error of  $\pm 2\%$  in porosity. (a) Porosity by traditional DR method, (b) porosity by multi-scale method.

Туре	$k_{exp}$ (mD)	$\varphi_{exp}(\%)$	$arphi_{total}$ (%)	$\varphi_{micro}(\%)$	$\varphi_{nano}(\%)$	Index
Glutenite	25.29	19.25	19.34	12.81	2.17	1
Glutenite	12.91	18.34	15.83	7.84	0.90	2
Glutenite	38.09	21.44	19.68	13.44	0.65	3
Glutenite	54.44	17.47	15.12	5.71	2.24	4
Glutenite	2.35	13.56	14.18	5.47	2.29	5
Glutenite	13.29	17.72	15.93	7.23	2.46	6
Glutenite	1.24	15.43	15.25	3.75	3.69	7
Glutenite	0.62	12.83	13.24	4.93	2.88	8
Glutenite	54.13	15.74	11.41	8.88	1.90	9
Glutenite	206.19	14.51	12.78	8.85	3.08	10
Glutenite	15.64	9.76	10.28	6.10	2.74	11
Glutenite	3.62	12.75	13.07	8.92	2.28	12
Glutenite	0.33	7.41	7.22	3.08	2.52	13
Glutenite	0.24	8.85	7.77	5.82	1.09	14

Table 1. Summary of the porosities from the laboratory and multi-scale DR.

Index	$\varphi_{nano}(\%)$	$\varphi_{micro}(\%)$	$\varphi_{total}(\%)$	$\varphi_{exp}(\%)$	$k_{exp}$ (mD)	Туре
15	2.13	2.35	6.49	11.36	2.04	Glutenite
16	2.32	4.74	9.62	8.10	0.49	Glutenite
17	1.19	1.73	3.73	1.92	0.00	Glutenite
18	2.46	1.79	5.80	6.86	0.10	Glutenite
19	1.59	2.41	4.65	6.39	0.05	Glutenite
20	1.20	0.62	4.81	6.78	0.15	Glutenite
21	0.84	1.10	4.12	7.36	0.21	Glutenite
22	1.60	1.62	6.01	7.13	0.09	Glutenite
23	2.14	2.22	5.42	4.16	0.02	Glutenite
24	2.05	0.14	3.25	4.27	0.11	Glutenite
25	2.05	1.88	6.03	7.01	0.11	Glutenite
26	1.83	0.40	5.37	2.96	0.02	Glutenite
27	1.42	2.61	8.25	5.03	0.03	Glutenite
28	1.64	1.37	6.39	5.92	0.01	Glutenite
29	2.95	3.72	8.74	11.30	4.92	Glutenite
30	3.72	1.56	9.57	11.51	1.95	Sandstone
31	4.96	3.16	10.20	12.86	5.93	Sandstone
32	5.50	10.47	19.50	21.09	1.21	Sandstone
33	2.25	6.74	15.14	14.80	0.29	Sandstone
34	9.20	0.86	13.29	12.87	0.08	Sandstone
35	4.51	20.48	26.15	25.46	280.09	Sandstone
36	5.28	6.34	12.08	10.81	0.16	Sandstone
37	3.07	16.08	19.77	21.14	25.72	Sandstone
38	3.35	8.31	19.95	18.67	216.63	Sandstone
39	8.27	15.08	25.08	28.96	666.30	Sandstone
40	6.60	16.79	27.05	31.05	1563.74	Sandstone
41	3.85	11.02	18.58	21.16	80.12	Sandstone
42	3.65	20.66	25.17	24.08	39.27	Sandstone
43	3.12	23.51	27.33	24.87	758.01	Sandstone
44	4.81	12.25	18.36	24.02	41.06	Sandstone
45	1.58	9.54	13.80	10.40	5.07	Sandstone
46	2.32	14.01	18.90	17.06	969.66	Sandstone
47	2.25	8.43	11.88	11.02	7.69	Sandstone
48	2.56	7.86	12.30	14.98	259.27	Sandstone
49	2.90	11.93	16.84	14.37	58.94	Sandstone
50	3.11	6.15	11.22	11.30	53.04	Sandstone
51	2.98	5.27	12.23	14.54	2.27	Sandstone
52	2.57	4.66	10.02	10.52	0.29	Sandstone
53	1.75	4.71	8.57	9.44	0.91	Sandstone
54	2.09	6.55	11.39	12.73	96.31	Sandstone
55	2.74	16.71	27.39	27.44	2994.98	Sandstone
56	2.81	19.40	27.04	28.60	5770.15	Sandstone
57	4.12	12.95	20.37	17.32	20.95	Sandstone
58	2.16	4.75	9.63	10.31	2.75	Sandstone
59	3.40	2.13	7.46	9.70	1.41	Sandstone
60	2.33	3.98	12.18	13.83	1.37	Sandstone
61	1.48	4.13	10.52	7.87	0.06	Sandstone
62	1.60	1.92	4.79	5.50	0.01	Sandstone
63	2.35	1.94	4.71	3.95	0.17	Igneous
64	2.35	2.34	5.41	3.88	0.05	Igneous
65	1.80	1.37	4.52	3.58	0.05	Igneous
66	1.74	1.11	3.64	1.54	0.03	Igneous
67	2.23	0.89	4.70	3.58	0.01	Igneous

3.2. Evaluation of the Simulated MICP Curve at Multi-Scale

The pore throat distribution can be simply combined at four different scales into one, and the simulated MICP curve can be calculated following the procedure in Section 2.4. Figure 6 compares the MICP curve and the pore throat distribution between the multi-scale



DR and the laboratory results for the sandstone with a good homogeneity (sample #49 in Table 1). The good agreements prove that the method is suitable for homogeneous rock.

**Figure 6.** Comparison of the MICP curve and the pore distribution for sandstone (sample #49). (a) MICP curve; (b) Pore throat distribution.

Figure 7 compares the MICP curve for heterogeneous glutenite between the multi-scale DR model and the laboratory results. The dashed line for each corresponding pore size is calculated following the procedure in Section 2.4. A good agreement for the saturation at the entry point and ending point can be observed. However, a large discrepancy exists in between.



**Figure 7.** Comparison of the MICP curve between multi-scale DR and laboratory results. (**a**) Sample 1; (**b**) Sample 2.

A further analysis is performed to explain this discrepancy. DR combines the pore throat distribution at four different scales, and then calculates the simulated MICP curve from the pore distribution through Equation (4). Meanwhile, for experiments the MICP curve is measured first, then the pore throat distribution is calculated. When the simulated MICP curve is derived from DR based on the pore throat distribution, the default sequence of packings is from large pores to small pores as shown in Figure 8a, while in reality the packings or the pore throat distribution more accurately, yet it does not account for the random distribution of pore throats within the sample. Thus, a homogenous rock leads to little discrepancy, while heterogenous rock and poor pore connectivity lead to a large discrepancy between DR and the laboratory results. A better simulated MICP model is necessary for DR in practical application.



**Figure 8.** Difference in mercury injection sequence. (**a**) Digitized sequence from multi-scale MICP; (**b**) Laboratory mercury injection sequence.

## 3.3. A Corrected DR Model for the Simulated MICP Curve

A corrected model is proposed in this study in order to improve the simulated MICP curve from DR for glutenite. The model involves adjusting the saturation points of the simulated MICP curves at 35  $\mu$ m, 2  $\mu$ m, and 65 nm to align with those obtained from experimental MICP data. Subsequently, the saturations between these points from the simulated MICP curves are stretched while keeping pressure unchanged. We examine the correlation between the saturation points of the simulated MICP and experiments at 35  $\mu$ m, 2  $\mu$ m, and 65 nm, respectively, for the six glutenite cores in order to obtain a constant correction coefficient. First, the mercury inlet pressure for three pore throat sizes of 0.065  $\mu$ m, 2  $\mu$ m and 35  $\mu$ m are calculated by Equation (4), then the corresponding saturation points can be obtained from Figure 7. The relationship for the three saturation points between the laboratory and DR is compared and shown in Figure 9.





A good linear relationship can be observed, thus the DR saturation at the three corresponding nodes can be corrected:

$$s_{1,2,3}^* = c_{1,2,3} \times s_{1,2,3} \tag{5}$$

where subscripts 1,2,3 represent three nodes of 0.065  $\mu$ m, 2  $\mu$ m, and 35  $\mu$ m, respectively,  $s_{1,2,3}^*$  is the corrected saturation, and  $s_{1,2,3}$  is the original saturation calculated from DR. The correlation coefficient  $c_{1,2,3}$ , obtained from Figure 9, has values of 0.78, 0.22, and 0.25,

respectively. Keeping the mercury saturation at the inlet and outlet unchanged, we stretch the saturation at each corresponding scale where:

$$\begin{cases} s_{i}^{35\mu m} = s_{inlet} + \frac{(S_{1}^{*} - S_{inlet})}{(S_{1} - S_{inlet})} \times (s_{i} - s_{inlet}) \\ s_{i}^{2\mu m} = s_{1}^{*} + \frac{(S_{2}^{*} - S_{1}^{*})}{(S_{2} - S_{1})} \times (s_{i} - s_{1}) \\ s_{i}^{65nm} = s_{2}^{*} + \frac{(S_{3}^{*} - S_{2}^{*})}{(S_{3} - S_{2})} \times (s_{i} - s_{2}) \\ s_{i}^{6.5nm} = s_{3}^{*} + \frac{(S_{outlet} - S_{3}^{*})}{(S_{outlet} - S_{3})} \times (s_{outlet} - s_{i}) \end{cases}$$
(6)

where  $s_i^{35\mu m}$ ,  $s_i^{2\mu m}$ ,  $s_i^{65nm}$ , and  $s_i^{6.5nm}$  are the mercury saturation at the corresponding scales,  $s_{inlet}$  and  $s_{outlet}$  are the mercury saturation of inlet and outlet, respectively, and  $s_i$  is the original mercury saturation from DR.

Figure 10 shows the comparison for the MICP curve between the corrected DR model and the laboratory results for the six glutenite samples. An improved agreement can be clearly observed by comparing to Figure 7.



**Figure 10.** Comparison of the MICP curve between the corrected DR and laboratory results. (**a**) Sample 1; (**b**) Sample 2; (**c**) Sample 3; (**d**) Sample 4; (**e**) Sample 5; (**f**) Sample 6.

## 4. Conclusions

This study developed a new DR method to upscale the porosity and the simulated MICP curve from nano-scale to core scale. The porosity agrees well with the laboratory experiments for 67 cores, including sandstone, glutenite, and igneous rock. The MICP curve, however, has a large discrepancy between the DR and laboratory for the heterogeneous rock. A corrected simulated MICP method is developed in order to match the experiments. The corrected DR method will provide a basis to guide reservoir evaluation, for example, to support the interpretation for nuclear magnetic resonance logging where a pseudo capillary pressure curve is usually used which is indirect and inaccurate (Xiao et al., 2016) [46]. Based on the results, we can conclude that:

- (1) The traditional DR method by the single-resolution micro-CT scanning (usually  $1 \sim 3 \mu m$ ) cannot capture all the pore information, including the nano-pores less than 1  $\mu m$  and macro-pores greater than 1 mm. Thus, the traditional single-scale DR often leads to lower porosity for both homogeneous and heterogeneous rock. The upscale DR model combining four scales represents the porosity and the pore throat distribution better by comparing to the laboratory experiments.
- (2) The pore distribution and the simulated MICP curve from the multi-scale DR agree well with the experiments for sandstones with good homogeneity, but a large discrepancy has been observed for heterogenous rock such as glutenite. A further analysis is performed to investigate the discrepancy, where DR obtains the pore throat distribution from images first, and then calculates the simulated MICP curve from the pore throat distribution through Equation (4), while the laboratory experiment measures the MICP curve first, and then calculates the pore throat distribution. Also, the multi-scale DR accounts for all the pore information at four scales, while the mercury injection experiment only accounts for the connected pores. Thus, the pore throat distribution by multi-scale DR is more accurate, while the MICP curve by the laboratory experiments is more representative. The simulated MICP curve by the multi-scale DR model assumes the mercury enters the pore throats from large to small, while in the experiment mercury enters large and small pores at the same time, as shown in Figure 8.
- (3) The multi-scale DR method involves the integration of high-resolution digital imaging and advanced modeling techniques to assess the porous structure of rocks at multiple scales. In comparison to previous upscaling methods, the new method demonstrates a higher level of accuracy in depicting the full range of pore structures across different length scales. This work also allows for more comprehensive comparison with experimental data, which addresses the limitations of previous methods and improves the confidence for the DR application. The implications for petroleum exploration and development are crucial, as rock pore structure greatly affects the flow of fluids, such as oil and gas, through rock formations. Consequently, a better understanding of pore structures via DR techniques can significantly enhance the accuracy of reservoir characterization, forecasting production rates, and optimizing petroleum recovery strategies. The multi-scale DR method offers a potentially valuable tool for these purposes.
- (4) The present study utilized the DR model at four different scales to upscale porosity values from the nano-scale to core scale. However, it is worth noting that the characterization of pore connectivity is essential to gain a comprehensive understanding of the pore structure within rock samples. Unfortunately, the pore connectivity can only be determined using the single-scale DR at the 2  $\mu$ m resolution, which restricts the accuracy in the evaluation of pore connectivity. Thus, the total pore connectivity cannot be accurately captured by the multi-scale DR method and can lead to discrepancies for some rocks. Further research is recommended to investigate this limitation and explore methods that can accurately capture pore connectivity down to the nano-scale.

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