



Article Wave Power as Solution for Off-Grid Water Desalination Systems: Resource Characterization for Kilifi-Kenya

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Abstract: Freshwater scarcity is one of humanity's reoccurring problems that hamper socio-economic development in many regions across the globe. In coastal areas, seawater can be desalinated through reverse osmosis (RO) and transformed into freshwater for human use. Desalination requires large amounts of energy, mostly in the form of a reliable electricity supply, which in many cases is supplied by diesel generators. The objective of this work is to analyze the wave power resource availability in Kilifi-Kenya and evaluate the possible use of wave power converter (WEC) to power desalination plants. A particular focus is given use of WECs developed by Uppsala University (UU-WEC). The results here presented were achieved using reanalysis—wave data revealed that the local wave climate has an approximate annual mean of 7 kW/m and mode of 5 kW/m. Significant wave height and wave mean period are within 0.8–2 m and 7–8 s respectively, with a predominant wave mean direction from southeast. The seasonal cycle appeared to be the most relevant for energy conversion, having the highest difference of 6 kW/m, in which April is the lowest (3.8 kW/m) and August is the peak (10.5 kW/m). In such mild wave climates, the UU–WEC and similar devices can be suitable for ocean energy harvesting for water desalination systems. Technically, with a capacity factor of 30% and energy consumption of 3 kWh/m³, a coastal community of about five thousand inhabitants can be provided of freshwater by only ten WECs with installed capacity of 20 kW.

Keywords: wave power resource; desalination; freshwater; wave energy converter; Kilifi; Kenya

1. Introduction

The worldwide deficiency of clean freshwater causes sanitation problems, food shortage and sometimes even conflicts [1–3]. UN's sustainable development goal (SDG), number 6 aim to ensure access to water and sanitation for all [4]—a goal currently far from met. For example, in 2015, 58% of the Kenyan population had access to safe drinking water, 30% had access to safe sanitation and only 14% had access to a proper handwashing facility [5,6]. The energy-food-water nexus describes an entangled relationship for the need of fulfilling basic human needs; water shortage in the agricultural sector inevitably leads to a decrease in food production [7]. Therefore, a holistic and interdisciplinary view on this problem is necessary. As an example, recent studies reflect upon effects of water usage in rural regions of Kenya [8,9]; discussing the farmers' willingness to pay for water for agricultural purposes; payments for water services in rural communities; the overall wealth; the use of water saving technologies; and the possibility of handling different types of future changes in the climate in this region [8,9].

In many coastal places around the world, the inhabitants still live without reliable source of clean freshwater even though there are an overflow of seawater and ocean waves. One such example is

small desalination plant located in an educational center in Kilifi-Kenya, which also was taken as an example for the present work. Instead of using conventional diesel generators [10], this desalination plant can be powered by renewable energy technologies such as wave energy converts (WECs).

Seawater desalination is used to produce freshwater around the world, using thermal processes or technologies involving membranes [11,12]. Three different desalination processes are reverse osmosis (RO), electrodialysis (ED) and multistage flash (MSF). To briefly describe these processes: With RO, the salt in the seawater is excluded by the use of a semipermeable membrane and an applied pressure, as described in [13]; with ED, an electric field is applied to the water, producing freshwater with the use of ion-selective membranes, with applications such as recently discussed in [14,15], and with MSF distillation, there are different evaporation stages in the steps towards production of freshwater [12,16]. RO technology is one of the most energy efficient for seawater desalination [17,18]. Several studies such as [18–22], have estimated that the energy consumption in RO desalination plants are mainly within 2–5 kWh/m³. This is mostly due to substantial advances in RO technologies that may include energy recovery techniques and efficient membranes [18]. Energy consumption will decrease even more with gains in membrane and energy recovery systems performances [18,23].

Dependable electric grid is not always available where the freshwater is needed. Therefore, solutions to implement renewable energy systems (RES) via smaller autonomous grids connected to desalination plants are suggested [24]. An issue regarding RES powered desalination, for technologies such as RO, is the intermittency of the RES. Whereas, a variable power output from the RES (if no energy storage system is included) causes a pressure variation on the RO membranes which may damage or lower their lifetime and can affect the freshwater output. As such, a RES with lower variability/intermittency is preferable.

Among the RES, wave power has high potential in terms of energy density, resource availability and predictability [4,25–28]. Especially in locations where seas and swell waves occur permanently, suggesting that wave power can be a better choice to combine with RO than solar or wind power. Moreover, as the energy source (ocean waves), and the water resource to desalinate (seawater) are found at the same location, there is a clear opportunity in combining wave power and desalination, taking care of the natural resources at a certain location [29]. However, the technology for wave power conversion is still in its early stages. This infers that there are still constrains on the implementation of large-scale wave power farms that are able to deliver large amounts of power (in the order of GWh) in a reliable manner. However, with advances in information technologies, computational modeling and prototyping, the actual WECs have gained substantial improvements in power take-off as well as decrease in production costs, making them ready-enough to be used in small scale. Information on WECs concepts can be found in for example [28,30].

Previous studies such as [31,32], have estimated that the wave climate in Kenya is similar to other regions situated within the equatorial belt, with offshore peak values within 20–30 kW/m. However, there is a need to access the nearshore wave power resource. Therefore, the overall objective of this research is to access the feasibility of utilizing wave power in desalination plants. In particular this study aims to investigate the local wave power resource in Kilifi as a starting step of evaluating the possibility of using wave power converter developed by Uppsala University (UU-WEC) technology for desalination systems. The present work lies within the research domain of the authors and an existent collaboration between researchers at Uppsala University in Sweden and Strathmore Energy Research Center in Kenya. The results of the present work will provide insight information of a potential solution for freshwater shortage in coastal areas.

The UU–WEC Technology

The UU-WEC (Figure 1) is a point absorber that comprises a submerged linear generator connected to a heaving buoy e.g., [33]. The WECs are connected through an offshore marine substation before the electricity is transmitted to an onshore connection point [34]. The voltage output of UU-WECs is about 450 V, and the estimated operational capacity is between of 10 kW and 100 kW but can be

scaled up. The power rating can be adjusted to match the local wave climate [33,35,36], and is mostly determined by the electrical configurations and by the size and shape of the buoy and generator. These three parameters can be adjusted to fit a local wave climate. This WEC technology is currently used for offshore experiments in Lysekil on the west coast of Sweden, and the experimental research has progressed for more than ten years [33], including studies on environmental aspects of wave energy conversion. These studies revealed very limited negative impacts of the UU-WECs to the marine environment e.g., [37,38]. Instead, positive effects, such as artificial reefs, were observed [37,38]. Other recent studies on the same WEC concept include control strategies, compensations for tides and discussions on the design of wave power farms with several WECs [39–41].



Figure 1. Comprehensive illustration of Uppsala University's main WEC concept, comprising of a linear generator on the ocean floor, with a movable part (the translator) connected via a line to a buoy.

2. Studied Region

The studied region covers the coastal area of Kilifi, located north of Mombasa in Kenya (Figure 2). The local bathymetry varies from 0 m to 30 m depth at close proximity to the beach. The seabed in this area is dominated by sedimentary rocks, loose clay, sand, mud flats, and coral reef that is normally distributed between 16 m and 45 m depth at distances from shore between 500 m to 2 km [42].

The marine environment is generally characterized by seagrass communities, mangrove forest and sandy subtracts [43]. The marine fauna is diverse and, includes sensitive species such as turtles and dugongs [42]. The wind pattern in this region is mostly dominated by monsoons, sea breeze, and occasionally cyclones. The tide in Kilifi is semi-diurnal with an amplitude of 3.5 m [27]. These tides contribute to intertidal platforms, mud flats and rather rocky communities that get exposed for several hours during low tides, and sub tidal platforms that are highly productive and populated by coral reefs, reef flats and susceptible to nearshore wave action.



Figure 2. Studied location in Kilifi (3.815° S, 39.846° E), ca. 60 km north of Mombasa, Kenya. The Yellow dashed circle represents the data extraction point (*Fugro*-No. 162230-1-R0). Local bathymetry and costal features are shown within the circle.

3. Wave Data

The wave data used was provided by *Fugro* (document No. 162030-1-R0) (see Supplementary Materials). The data covers a period from January 1997 to December 2015 with a temporal resolution of 6 h, and was based on analysis of the world waves data source of the *European Centre for Medium-range Weather Forecasts* (*ECWMF*) wave model. Re-analyses were conducted by *Fugro Oceanor* combined with a global buoy at location with 34 m bottom depth, and multi-satellite altimeter database (*Topex, Jason, Geosat, GFO* and *Envisat*) by transforming offshore grid points to a nearshore point (3.815° S, 39.846° E) using *Simulating Wave Nearshore* (*SWAN*) model.

The data set contains wave field variables: significant wave height (H_s), mean wave energy period (T_e), and mean wave direction (MDir). The variables were estimated from a wave energy spectrum (f, θ) with moment m_0 . The H_s is equivalent to the mean height of the highest one-third of the waves in a sea-state ($H_s = 4\sqrt{m_0}$). The T_e is equivalent to the spectral period and is given by $T_e = m_{-1}/m_0$. MDir is the direction of the most energetic spectral band.

The wave power resource (P_w) was estimated using Equation (1), which defines P_w as the average transport rate of energy per meter of wave front, it depends mostly on the wave height and wave period. By definition, P_w can be affected by wave dissipation, shoaling, reflection, refraction, trapping, diffraction, and other non-linear phenomena [44–46].

$$P_{w} = \rho g \int \int Cg(f, \theta) S(f, \theta) df d\theta$$
(1)

where ρ is the water density (1025 kg/m³ for sea water), *g* is the acceleration due to gravity (9.8 m/s²), *Cg* is the group velocity which depends of the water depth *h*, *f* is the frequency, and θ is the direction of propagation.

Assuming that the water depth is deep enough so that h/L > 0.5, then $Cg = g/4\pi f$; therefore the simplified form of Equation (1) (in (kW/m)) becomes:

$$P_w \cong 0.5 \ H_s^2 T_e \tag{2}$$

The validation of the possibility of utilizing wave energy for desalination through RO was conducted taking into account the following parameters: Available wave power resource (P_w), installed capacity per unit WEC (P_{wec}), energy per unit WEC per day at capacity factor of 30% ($W_{day.wec} = P_{wec} \times 24 \times 0.3$ h), energy needed per unit volume of treated water ($E_{ro} = 3$ kWh/m³), taken from an interval between 2 kWh/m³ and 5 kWh/m³ [18–22]. Taking into account that the basic daily need of volume of water for one person (V_i), is around 0.02 m³ [1]; however, this study uses $V_i \cong 0.1$ m³. A fundamental question can be: How many WECs (Equation (3)) are needed to meet the water demand of the local inhabitants at daily basis?

$$X_{wec} = X_p / X_{p.wec} \tag{3}$$

where X_{wec} is the number of WECs, X_p is the number of inhabitants, and $X_{p.wec} = \frac{W_{day.wec}}{E_{ro}}/1000$ L, is the number of inhabitants supplied by freshwater from a single WEC, recalling that a 1000 L is equivalent to 1 m³.

4. Results

4.1. Local Wave Climate

The mean value of wave power, P_w , for this data set is 7 kW/m, the median values is 5 kW/m, the mode is 5 kW/m, and the minimum and maximum are 0.2 and 53 kW/m respectively. This data reveals a wide difference of wave power between the periods of slack versus periods of high seas. It is important is to analyse the frequency of occurrence and magnitude of sea states (Figures 3–5) in order to understand how the energy is distributed over time. The significant wave height, H_s , (Figures 3 and 4) shows that the predominant values are within 0.75 m to 2 m, in which the southeast sea state, with a frequency of occurrence of 56%, has values of H_s in the interval between 1 m and 1.5 m. The second most dominant sea state, with 28% of occurrence, has easterly waves within 1.5–2 m. 12% of the waves compose the east-north-easterly seas state with H_s within 0.4–0.7 m. Higher sea states account for approximately 4% of occurrence and have values of H_s within 2–3 m. In what regards the wave mean period, results in Figure 3 show that the most prevalent sea state, 79% of occurrence, has values of within 7–8 s, followed by 10% of sea states with T_e of 5 s, 10% with 9–10 s, and 1% for sea states with values of T_e of 4 s and above 11 s, respectively.

The distribution of H_s and T_e (Figures 3 and 4) in time also show sea states with different wave power values starting from 2 kW/m up to 50 kW/m with a higher density between 10 kW/m and 20 kW/m (Figure 5). Approximately 52% of the waves have power values within 2 kW/m–5 kW/m, 33% have P_w within 7–10 kW/m, and the remaining 15% has values of P_w greater than 12 kW/m (Figure 6).



Figure 3. Total probability distribution of significant wave height H_s wave mean period T_e and wave mean direction versus mean wave direction MD_{ir} in Kilifi. The frequencies of occurrence of sea states are given in percentage.



Figure 4. Total probability distribution of sea sates at six-hourly occurrence in Kilifi.



Figure 5. Combined scatter plot showing the occurrence of H_s , T_e and P_w in Kilifi.



Figure 6. Total probability distribution of wave power P_w versus mean wave direction MD_{ir} . The frequencies of occurrence of sea states are given in percentage in Kilifi.

4.2. Diurnal Variability

The results show diurnal variations on the wave power resource in Kilifi. In Figure 7, there are the mean, mode and median graphs referent to observations at midnight, six, twelve and eighteen hours. From this date it was estimated that wave power values in evening and night are in average higher than in the day. However, the highest value of mode is observed in the morning. The difference between midday and mid night mean values is about 0.3 kW/m.



Figure 7. Diurnal variability of wave power in Kilifi based on 6–hourly wave data, showing lower values during day comparing to night.

4.3. Seasonal Variability

The seasonal cycle was obtained by extracting the total average for each month, over the entire time series (Figure 8: red line). There is a unimodal distribution of P_w over a period of 12 months, from January to December. Peak values are observed in August, when the mean is 10.5 kW/m followed by September with 10.2 kW/m. The box plot results shows that August and September had similar median values of approximately 10 kW/m. Although August has the highest mean, July has the highest occurrence of extreme values.

January–May had the lowest H_s mean values of approximately 1 m (Figure 9a). While July, August and September had higher mean values up to 1.5 m with a peak in August when H_s is 1.7 m. Considering T_e , there are three peaks occurring in April, September and December, being the highest mean of 7.5 s. The lowest T_e value, 6.8 s, was observed in July (Figure 9b). The wave mean direction MD_{ir} also had a seasonal variability signal. From February to April MD_{ir} is from east, then the waves gradually propagate from southeast from May to August, and again the mean direction reverts towards east (Figure 9c).



Figure 8. The seasonal variability of P_w in Kilifi. The red line represents monthly mean values of P_w , the box plot contains the median-middle mark, the 25th and 75th percentiles—the bottom and top edges of the box, respectively. The red crosses are outliers of P_w , representing for example storms and tropical cyclones.



Figure 9. The seasonal variability in Kilifi: (**A**) for H_s ; (**B**) for T_e ; and (**C**) for MD_{ir} . T_e (**B**) has a rather different pattern with three-modal seasonal variability compared with a unimodal pattern observed in both H_s (**A**) and MD_{ir} (**C**) respectively.

4.4. Interannual and Long-Term Variability

The interannual variations of P_w along the 19-year time-series are evident. The total mean is 6.5 kW/m, total standard deviation is 4.6 kW/m and the mean absolute deviation is 3.4 kW/m. For this period of time, the anomaly of P_w was calculated from the difference between the total mean and the mean of each month. Results show that there is an oscillatory pattern in the anomaly of P_w with periods of 3–5 years (Figure 10). Probably, with a longer data set, it would be possible to better detect the variability.



Figure 10. P_w from 1997 to 2016 in Kilifi, plotted as annual mean (black dots) and respective anomalies (red bars).

4.5. Wave Powered Desalination

Results showed that not many WECs are needed to supply water for small communities, assuming a capacity factor of 30% (Figure 11). For example, only 10 UU-WECs with installed capacity of 20 kW per unit would be required for daily production of freshwater for about 5000 inhabitants.



Figure 11. Number of WECs necessary for daily freshwater production as a function of number of inhabitants and installed capacity per unit WEC assuming a capacity factor of 30% and fixed energy consumption of 3 kWh/m³.

5. Discussion

5.1. Wave Climate

There is a relatively high frequency of occurrence of waves with simultaneously small H_s and longer T_e and vice-versa, see Figure 5. This may explain the reason why P_w occurs densely within 5 kW/m to 25 kW/m. It is also observed that the typical sea sates are determined by values of T_e rather than H_s (Figures 4–6). For instance, the frequency of occurrence of waves of H_s of 1–1.25 m are noticeable the highest and corresponds to values of T_e that widely range from 3–9 s.

Diurnal variability of wave power may be related to the coastal wind dynamics, such as offshore and sea breezes triggered by pressure gradients due to uneven heating and cooling of the land and sea masses. In this study, the wave power's time-variation signal share similarities in shape with the wind field's time-variations obtained by, for example Mahongo et al. [47]. Studies on coastal wind patterns such as [47], revealed the existence of similar diurnal variations on wind intensity using three-hourly wind data of four locations along the coast of Tanzania. The wind is strongest in the afternoon (14:00–15:00), and weakest in early morning (4:00). From the wind pattern described by [47], it is possible to infer that wave power values in the second quarter of the day occur when the wind direction is from north-northeast in the austral summer and from southwest during austral winter. Peak diurnal values of wave power do occur when the wind direction are dominantly form east-northeast (austral winter) and south-southeast (austral winter).

With regard to the seasonal variability, April has the lowest mean P_w of 3.8 kW/m. The minimum monthly mean coincides with calm north-easterly wind conditions in Tanzania and Kenya, which occur in March–April and November [48–50]. Moreover, low values of P_w observed during the same period of the northeast monsoon (November to March). The monthly mean values of P_w differ from what was found by other authors, e.g., Barstow 2008 [51], who estimated 5–10 kW/m for January and 20–30 kW/m for July. The reason for this discrepancy may be due to the fact that they used altimetry-model data from deep water points, while the present study used data from a shallower water point. The P_w and H_s have similar seasonal variability signals (Figure 11). The maximum H_s is in contra-phase with the maximum T_e may justify the reason to why there are considerably low values of P_w in this region. Other contributing factors can be latitude and weather systems including the inter-tropical convergence zone that oscillates within 15° S and 15° N and the monsoon regime [51,52]. The presence of a large land mass-Madagascar may also block larger and longer southerly waves from reaching the Kenya.

Interannual and long term variability of the wave power resources in Kilifi is influenced by local, regional and mesoscale weather systems. For example, previous studies have shown that the coastal wind in Kenya is associated with the monsoon wind system [53]. By looking at the coastal wind intensity, it is found that the shape of its seasonal cycle signal is similar to the one observed in the wave power's signal, with peaks in July and dips in March–April. The peaks are influenced by the austral seasonality of wind and wave fields generated in the southern ocean. With the absence of such influence, the minima of wave power should be registered around December–January. However, for this particular region of the Indian Ocean, the climate dynamics are strongly influenced by the Indian Ocean Dipole, that is an oscillatory mode that couples the variability of both atmosphere and the ocean. The lowest negative mode of the Indian Ocean Dipole occurs in March–April, that coincides with the minima in the magnitude of both wind and wave power.

Calculated anomalies of P_w correlate well with monthly mean sea level pressure at the surface linked to the Southern Ocean Mode (with correlation coefficient r < -0.6), Indian Ocean Dipole (r > 0.4) and slight influenced by the El Niño ($r \sim 0.4$). The 3 years–4 years variability can be observed with El Niño Southern Oscillation ($r \sim -0.5$). There is a strong correlation between P_w and sea surface temperature of tropical and southern ocean basins (Figure 9). Studies done by Reguero et al., 2015 [54] also found significant correlation of monthly mean wave power with the North Atlantic Oscillation ($r \sim 0.2$), El Niño Southern Oscillation ($r \sim 0.2$), Atlantic Multidecadal Oscillation ($r \sim -0.2$) and Dipole Mode Index ($r \sim -0.3$).

5.2. Wave Powered Desalination

The estimative results in Figure 10 were achieved assuming that wave power has an annual capacity factor near 30%. Several studies suggest that the capacity factor of wave power is ca. 20–40% at the most [55–58]. Being so, the number of WECs (20 kW) required for a daily supply of freshwater to 5000 inhabitants would increase or decrease in function of the capacity factor. Moreover, the capacity factor would vary seasonally, being lower during austral summer months (January to May) and higher during July to November. On the other hand, the number of required WECs would also vary according to the efficiency of the desalination system in terms of energy consumption. For example, similarly to above, assuming a capacity factor of 30%, but with an energy consumption of 2 kWh/m³, the required number of WECs (20 kW) for 5000 inhabitants would decrease to seven units.

The results indicate a mild wave climate in Kilifi. This is particularly suitable for the UU-WEC technology that is versatile to variety of wave steepness. Due to few extreme wave days and lower forces acting on the system—reducing the risk of technical failure, and lowering its predicted maintenance. The UU-WEC was designed to optimally operate in depths within 20 m and 50 m, in rather flat seabed, conditions which normally occur in Kilifi and on other coastal areas across the globe. Several layout designs of wave power farms can be used to maximize energy abortion, as described in [59–61]. In practice, these wave power farm layouts should also consider the distance between devices, substation and onshore cable connections which are key elements for cost reduction and feasibility, as described in [62,63]. In the case of Kilifi where a rather small number of WECs is required, a rather simple and cost effective wave power farm would be required. For example, it would include an array of WECs deployed along the same isobaths in order to maximize energy absorption.

Compared to the other wave climates, such as of the North Atlantic [54], the waves in Kenya are more predictable, which may enhance the reliability and predictability of the power or freshwater production from a wave powered desalination plant. Therefore, the use of UU-WEC for desalination purposes would be suitable to similar coastal areas where the water need is severe, requiring a reliable source of drinking water, irrigation among other purposes. Apart from the resource assessments aspects, the costs of electricity, water production and the entire life cycle of a project should be holistically analyzed.

6. Conclusions

The wave climate in Kilifi is mild, with mean values of 1.3 m, 7.1 s, and 6.5 kW/m for H_s , T_e , and P_w , respectively. The difference in monthly mean values between the lower and higher seasons was in the order of ± 0.7 m for H_s and of about ± 6.5 kW/m for P_w . This is due to the entire region is located within the tropical and equatorial zone where the weather patterns experience only gradual changes and the variations are rather small. The dynamics of the sea breezes and coastal winds influence the local wave power resource which is higher in the evening and night comparing to daytime values, with an estimated difference of 0.3 kW/m. The seasonal signal indicated that P_w and H_s share the same pattern in which both have peaks in August, while the dip occurred between January and May. In contrast, T_e have maxima in February, May and September, whereas the minima are in February, July and October. This seasonal variability of wave power is also similar in pattern with the coastal winds, and both are conditioned by the seasonality of the monsoon system and the inter-tropical convergence zone. The interannual variability has anomalies within 2 kW/m with periodicity of 3 years to 5 years, which can be influenced by mesoscale weather and climate systems such as Indian Ocean Dipole, Southern Ocean Mode, El Niño Southern Oscillation.

The wave power resource available at the Kenyan coast is estimated to be enough for wave power conversion throughout an entire annual cycle. The peak of wave power availability and conversion would be in the austral winter that match the dry season when freshwater is needed the most. This fact

highlights the importance of a desalination plant powered by WECs in arid and remote coastal areas. Assuming a capacity factor of 30% and energy consumption of 3 kWh/m³, a rather small coastal community of about five thousand inhabitants can be provided of freshwater by only ten WECs with installed capacity of 20 kW a unit, and certainly the number of WECs would decrease with the decrease of energy consumption and/or with an increase of installed capacity. There are still grey areas in what regards techno-economic and life cycle analysis of such projects involving renewable energy technologies. Uncertainties are even higher with non-stablished technologies such as wave power. Notwithstanding, near future projections appoints to reduction of energy consumption to below 3 kWh/m³ [23], and a possible decrease of water production costs [64]. On the other hand, there will be an increasing freshwater consumption which reflects on growing investments on desalination plants around the world. It is only left to the renewable energy sector to promote itself to competitive standards within the energy-water nexus. This brings us to a final conclusion that the electricity needed to run a desalination plant or any other type of electric water treatment system, can be technically supplied by renewable energy sources such as wave power.

7. Future Work

The present work is part of a larger study being undertaken by Uppsala University on use of wave power for freshwater production through desalination, with focus on the UU-WEC technology. Therefore, after investigating the wave climate in Kilifi and introducing the idea of utilizing WECs as potential power supplier, the following steps would proceed: Full description of the power system with several UU-WECs including voltage amplitude and frequency variations; Then, coupling of the power system with a RO desalination plant in Kenya that will include an energy storage system and freshwater storage seen as buffer for shifting power rates; Techno-economic and life cycle analysis would be investigated; assessment of local environmental parameters such as seabed conditions, residue management among other aspects required by law would precede a possible deployment of WECs that would culminate with an implementation of wave powered desalination system. The aforementioned steps may not follow the presented sequential order. Even so, research is needed before producing freshwater using the UU-WEC or similar technologies. The success of projects such as this, would revolutionize rural electrification in remote and coastal areas across the globe.

Supplementary Materials: The following are available online at http://www.mdpi.com/1996-1073/11/4/4/s1, the wave data used was provided by Fugro (document No. 162030-1-R0) and can be accessed upon a request to the authors.

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