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Measuring of Thermal Conductivities of Soils and Rocks to Be Used in the Calculation of A Geothermal Installation

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Abstract: The thermal conductivity of soils and rocks constitutes an important property for the design of geothermal energy foundations and borehole heat exchange systems. Therefore, it is interesting to find new alternatives to define this parameter involved in the calculation of very low enthalpy geothermal installations. This work presents the development of an experimental set-up for measurements of thermal conductivity of soils and rocks. The device was designed based on the principle of the Guarded Hot Plate method using as heat source a laboratory heater. The thermal conductivity of thirteen rocky and soil samples was experimentally measured. Results are finally compared with the most common thermal conductivity values for each material. In summary, the aim of the present research is suggesting a procedure to determine the thermal conductivity parameter by a simple and economic way. Thus, increases of the final price of these systems that techniques such as the “Thermal Response Test” (TRT) involves, could be avoided. Calculations with software “Earth Energy Designer” (EED) highlighted the importance of knowing the thermal conductivity of the surrounding ground of these geothermal systems.

Keywords: thermal conductivity; very low enthalpy geothermal installation; Guarded Hot Plate method; Thermal Response Test (TRT); Earth Energy Designer (EED)

1. Introduction

With respect to the lithosphere, heat transfer is produced by thermal conduction; heat diffuses without transfer of matter. Conduction is the principal mechanism of thermal propagation that takes part in the process of thermal exchange in a very low temperature geothermal installation [1].

Thus, the parameter of thermal conductivity (W/mK) plays a fundamental role in these systems. When this value increases, the capacity of the ground to transmit the heat to the components of the installation is also bigger, increasing its efficiency. Therefore, thermal conductivity constitutes a reference to evaluate the speed of the energetic extraction through the geothermal pipes or the dissipation of heat through the ground. For these reasons, it is recommendable to define this parameter to carry out a suitable calculation of a low enthalpy geothermal installation [2–4].

There are many different ways to measure this thermal property. Generally, experimental methods can be grouped into two categories: (i) stable methods, which provide more precise results despite requiring long measurement periods; and (ii) non-stable methods, which stand out for their rapidity, although they offer a lower precision.

Regarding the geothermal field, in practice, tables providing reference values of thermal conductivity for a set of materials are commonly used. In such cases, only approximate values are used, so the calculation of the installation may not be completely correct. This fact usually causes over-measurements that, in some cases, involve considerable increases of the final cost [5,6].

In large projects, another less frequent practice is the execution of a “Thermal Response Test” (TRT) that allows obtaining “in situ” the thermal conductivity parameter. The test allows studying the behavior of the ground when a constant thermal power is transmitted through the geothermal pipe. It constitutes a suitable solution in spite of the high cost that its realization implies, especially in small installations where a TRT could mean an important increase of the global budget.

Laboratory studies of the thermal conductivity of soils and rocks are usually carried out on samples collected from the ground.

The “Guarded Hot Plate” (GHP) method is the standard technique for measuring thermal conductivity of solid materials in the range from about 0.01 to 15 W/mK [7]. Two main versions of the GHP method can be differentiated: the double-sided (2S-GHP) which implies the use of two identical specimens, and the single-sided (1S-GHP) that only uses one specimen. Different techniques based on this method are used in the laboratory to measure the thermal conductivity of rocky and soil samples [8].

Ramstad et al. [9] designed equipment to measure the thermal diffusivity of rock samples. Thermal conductivity was calculated as a product of density, specific heat capacity and thermal diffusivity. Lira-Cortés et al. [10] implemented a system of thermal conductivity measurement for solid conductive materials. The system measured the thermal conductivity of an aluminum bar using a reference material.

Liou and Tien [11] estimated the thermal conductivity of granite using a combination of techniques, the “Transient Plane Source” (TPS) method [12], the thermal probe method and heat transfer test.

Krishnaiah et al. [13] designed a thermal probe to estimate different thermal properties such as thermal resistivity and diffusivity, and specific heat of rocks.

Kukkonen and Lindberg [14] measured the thermal conductivity of rocks making use of the steady-state divided bar method.

Jorand et al. [15] used the TCS method based on contact-free thermal conductivity scanning of a plane or cylindrical surface [16]. This instrument uses a focused, mobile and continuously operating heat source, together with two infrared temperature sensors at small distances behind and in front of the source, for measuring the thermal conductivity along scanning lines.

Table 1 presents a comparison among the works previously cited and the method proposed in the present study.

As examples of more recent studies, Xiao et al. [17] proposed an analytical model for effective thermal conductivity of nanofluids, while Cai et al. [18] provide a complete review about the recent investigations on the fractal models and fractal-based approaches applied for effective thermal conductivity.

Devices commercially produced to measure this property are numerous at present. As example, the equipment commercially known as KD2-PRO is commonly used to determine the thermal conductivity of different materials including rocks and soils [19,20]. However, most of the equipment is not cheap and the price of the whole geothermal system immediately grows.

The present research offers a description of a thermal conductivity measuring apparatus based on the 1S-GHP (a single-sided guarded hot plate) principle. The aim is to present a new alternative to estimate this parameter making use of usual equipment in a soil science laboratory: a laboratory heater. Throughout this work, we describe a new experimental application of the 1S-GHP method; we apply the method to a set of heterogeneous rocks and soil samples; and we compare these values with the theoretical ones.

The novelty of this study is the combination of both rocks and soils thermal conductivity measurements in the same device. The main strength of the method is the simplicity in the calculation

of the thermal conductivity parameter by measuring temperatures in four horizons. The measurement range reaches the thermal conductivity of aluminum, used as reference sample in the current work.

Table 1. Contributions and limitations of past thermal conductivity works.

Authors	Contributions	Limitations
Terzic et al., 2016	The model offers an estimation of the both separate and total parasitic heat fluxes, improving the accuracy and uncertainty of final results.	The method is limited to solid materials with thermal conductivity values from 0.1 to 2 W/mK, in the temperature range 10–50 °C.
Ramstad et al., 2009	The measurement time is around 200 s. It allows thermal conductivity measurements up to 4.5 W/mK.	The improved version of this method requires a heat source with constant temperature of 300 °C.
Lira-Cortés et al., 2008	The system is suitable to measure the thermal conductivity of conductive materials with a design error of 2% order.	The method cannot be used in poor conductivity materials given the high design errors
Liou and Tien, 2016	The three techniques provide fairly similar values for the same granite sample. TPS method presents high accuracy and simplicity to sample preparation.	For the thermal probe test, it is difficult to ensure that the air in filled holes is completely removed.
Krishnaiah et al., 2004	The method considers the thermal conductivity variations with the samples porosity.	Trends variations of thermal conductivity with porosity were not established and validated in a standard way.
Kukkonen and Lindberg, 1995	Thermal conductivity is calculated from using the arithmetic, harmonic and geometric mean values and compared with the measured one.	The grain size and the textural variation of some samples affect the thermal conductivity measurements.
Jorand et al., 2013	It combines high-accuracy laboratory measurements and numerical petro physics. They use thermal conductivity scanning for obtaining 2-D thermal conductivity maps reflecting the structural heterogeneity in two samples. Measurements are made in two profiles along the core axis and perpendicular to it. Anisotropy of thermal conductivity is also estimated.	Heterogeneous proportions of lateral heat flow within the sample affect both heat transport in general and the determination of effective thermal conductivity. The method might not solve the entire upscaling problem.
Proposed method	The method allows both rocks and soils thermal conductivity measurements. It also allows thermal conductivity measurements up to the aluminum thermal conductivity. Thermal conductivity results agree with the common values officially accepted for each material.	Long term measurements. It requires a constant ambient temperature to avoid influencing the measurements. Anisotropy is not considered.

2. Materials and Methods

2.1. Description of the Suggested Method

In the laboratory of Rock Mechanics of the Higher Polytechnic School of Avila (Spain), a procedure to determine the thermal conductivity of different samples of rocks and soils was developed.

The method suggests, from two pattern samples (S_{R1} – S_{R2}) with well-known value of thermal conductivity, quantifying this property in any other sample (S) whose value of thermal conductivity wants to be known. A heat source (laboratory heater) that generates a constant heat flow Q_x , was used. This flow goes through the three samples (S_{R1} – S – S_{R2}) placed contiguous as shown in Figure 2. Temperatures are controlled in four horizons (T_1 , T_2 , T_3 , and T_4) by thermocouples (Figure 1). Once these temperatures are stabilized and known, they can be used in the corresponding calculation of the thermal conductivity parameter [21,22].

The cold source equally schematized in Figure 1 represents the temperature of the room where the measurement equipment was placed. It is important to highlight that this cold source was kept

constant during the whole process of measurement, given that any variation could involve important changes in the results. Thus, the temperature of the cold source was controlled and set in the value of 296.65 K.

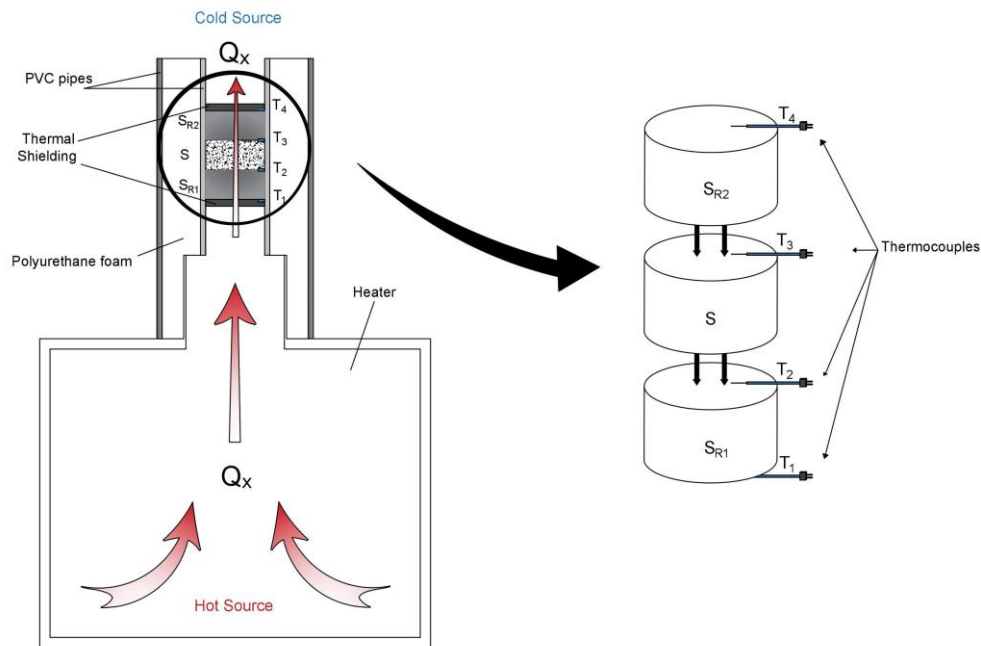


Figure 1. Schema of the suggested method.

2.2. Theoretical Basis

In the system described in Figure 1, the transfer of heat only occurs by conduction with one-dimensional flux and permanent state. The possible convection phenomena were prevented placing two aluminum sheets that behaved as insulations. The action of these sheets made the thermal transfer among the different samples purely conductive. Thermal conductivity is the physical property that controls the conduction of heat in a solid. It is materialized by Fourier's law (Equation (1)) which relates the specific heat flow and the gradient of temperature [23–25]:

$$\dot{Q}_x = -k A \frac{dT}{dx} \quad (1)$$

where \dot{Q}_x = Heat flux in the direction x ; k = Thermal conductivity; A = Area of the transverse section of the conductive object; and $\frac{dT}{dx}$ = Gradient of temperature in the direction x .

If we apply Equation (1) to each one of the samples that are part of the system, the resultant equations are:

$$\begin{aligned} \dot{Q}_1 &= -k_1 A_1 \frac{dT_1}{dx_1} \\ \dot{Q}_2 &= -k_2 A_2 \frac{dT_2}{dx_2} \\ \dot{Q}_3 &= -k_3 A_3 \frac{dT_3}{dx_3} \end{aligned} \quad (2)$$

Heat flux \dot{Q}_x was calculated using temperatures T_1 , T_2 , T_3 and T_4 (when they become steady) and Equation (1). This calculation was possible because of two main facts: the heat flux is constant and thus the same for each of the samples, and we use two reference samples with known thermal conductivity values (S_{R1} and S_{R2}).

From \dot{Q}_x value and considering the steady temperatures, the dimensions of sample S and Equation (1), thermal conductivity of the sample S can be easily calculated [26,27].

Therefore, by the implementation of the procedure detailed in the present paper, it is viable to obtain the value of thermal conductivity of a particular material S .

Equations (1) and (2) can be obtained based on Fourier's law. Heat transport in geo-materials is known of dual-phase-lagging type. It is important to mention that Fourier's law of heat conduction is valid only for some limiting cases, so the method will be limited to those situations.

2.3. Heat Flux Analysis

Until now, the heat flux has been considered as one-dimensional; it only flows in the longitudinal direction. However, after several tests and temperatures analysis, it was experimentally verified that, in spite of the insulation used, there was an additional heat flux in radial direction.

Analyzing temperatures T_1 , T_2 , T_3 and T_4 , it was observed that, through the first pattern sample S_{R1} , a substantial quantity of heat flux is lost as radial flux. As a result, sample S_{R1} was not used in the corresponding thermal conductivity calculations (it was only used as a stabilizing element of the system). The location of this sample minimizes the loss of heat as radial flux in the remaining samples (S and S_{R2}).

Nonetheless, axial heat flux in samples S and S_{R2} must also be considered, because, although smaller, it alters the final results too. To quantify this radial flux, the method was previously used on a sample with known thermal conductivity value. Once this flux was quantified, final thermal conductivity results were exempt from this kind of error.

The distribution of the heat flow is represented in Figure 2.

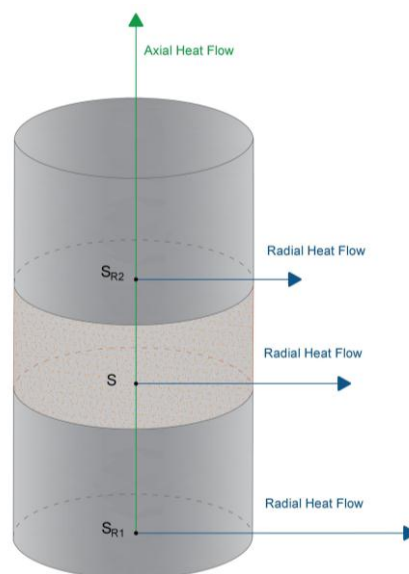


Figure 2. Heat flux distribution.

2.4. Equipment Description

The device designed to measure the thermal conductivity parameter consists of the following components:

- Sterilization and drying heater “Dry-Big” (Figure 3):

Heater with air force circulation mechanism, regulated by a microprocessor and with temperature and time digital reading. It constitutes the heat source used for the calculation of the thermal conductivity parameter.

A set of working temperatures was tested to analyze the evolution of the thermal conductivity with the temperature. Thus, heat source was regulated according to the most suitable temperature.

- PVC pipes (Figure 1):

Two PVC hollow cylinders were used in the construction of the equipment.

1. Hollow cylinder of diameter slightly higher to the air outlet placed on the top of the heater. This PVC pipe of diameter (0.10 m) coupled to the air outlet was adiabatically insulated in the whole contour by polyurethane foam.
2. Hollow cylinder 0.052 m of diameter, placed inside the previous pipe. It behaves as fastener of the samples. The space between both pipes was adiabatically insulated by polyurethane foam.

- Polyurethane foam (Figure 1):

Polyurethane foam was used to insulate the system from any external influence getting at the same time one-dimensional circulation of the heat flux through the samples.

It is a porous plastic material made up of a bubble aggregation. It consists of the chemical reaction of two polyol and isocyanate, although it accepts multiple additives. Its insulation capacity comes from the low thermal conductivity of the gas that its closed cells send.

- Pattern samples (Figure 2):

Two reference samples with known thermal conductivity values were used. These samples are made of pure aluminum. Given the high thermal conductivity of this element, it facilitates the heat flux transmission through the system. The dimensions of both patterns are: 0.10 m of thickness and 0.05 m of diameter.

- Thermocouples (Figure 1):

Four sounding lines (constituted by chrome and aluminum alloys) connected to a digital thermometer made possible the measurement of temperatures in four areas. Before its use, thermocouples were duly calibrated [28].

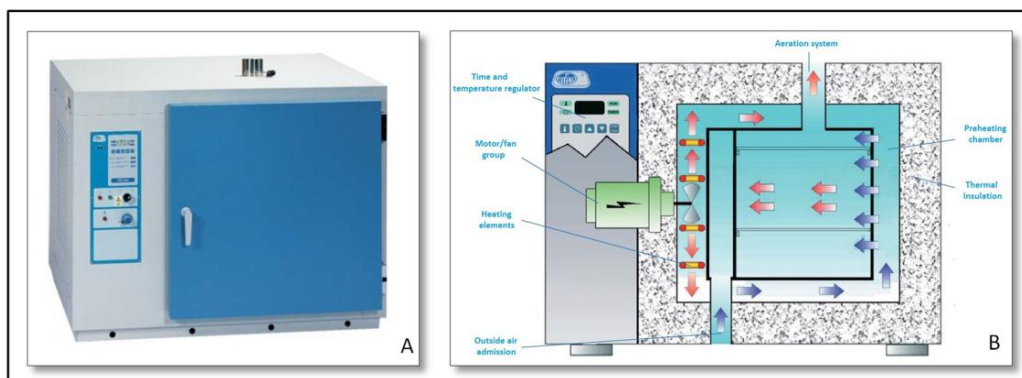


Figure 3. (A) Heater outside view; and (B) heater internal schema.

As explained in Section 2.2, calculations of the thermal conductivity parameter can only be carried out when the four temperatures recorded by the thermocouples keep a constant value over time. Figure 4 shows a graphical example of readings of these sounding lines. After a certain period of time, thermocouples record constant temperature values; these data are the ones used in the calculation of the thermal conductivity of sample S .

First, the PVC pipe of 0.1 m was coupled to the heater air outlet. A central opening was left for the subsequent placement of the second PVC pipe. Space between the first pipe and the opening was filled with polyurethane foam. A pattern sample (S_{R1}) followed the sample whose conductivity wants to be measured (S) and the other pattern sample (S_{R2}) was introduced in the second pipe. After placing thermocouples in the corresponding horizons, the second pipe was placed in the central opening to begin the test.

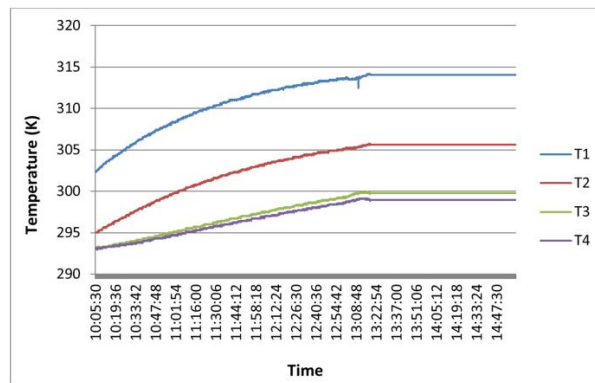


Figure 4. Measuring of temperatures with thermocouples.

3. Methodology of the Thermal Conductivity Test

The methodology of the proposed thermal conductivity test includes the next stages.

3.1. Materials Selection

A series of materials (rocks and soil) were selected to carry out the thermal conductivity measurements. These materials have different composition and nature so the study covers a varied geological range, defining with more precision the reliability of the methodology in question.

Table 2 shows the materials used in the study.

Table 2. Materials selected for the test.

Sample	Description
1	<i>Common Granite</i> Constituted by quartz, feldspar and micas and very varied group of secondary minerals in percentages under 5% like: apatite, esfena, oxides, allanite, zircon, etc.
2	<i>Adamellite</i> Plutonic igneous rock, with more than 65% of silica and more than 20% of quartz.
3	<i>Granodiorite</i> Plutonic rock of quartz, plagioclases, potassium feldspar, biotite and amphibole.
4	<i>Red Granite</i> Igneous rock known as “Bleeding Granite” with a high silica corn (more than 80%).
5	<i>Common Slate</i> Thin grain metamorphic rock with sericite, muscovite, chlorite and quartz.
6	<i>Quartzite</i> Hard metamorphic rock composed by quartz (more than 90%), it can also contain muscovite, orthoses or albite. Its structure presents soldered quartz crystals.
7	<i>Sandstone</i> Sedimentary rock with clasts about the size of the sand. The grains have quartz, feldspars or rock fragments.
8	<i>Gypsum</i> Mineral of hydrated calcium sulphate giving mono mineral sedimentary rocks.
9	<i>Pumice</i> Vitreous volcanic igneous rock. It is grey with silica dioxide, aluminum oxide and other oxides.
10	<i>Orto gneiss</i> Gneiss generated by dynamic metamorphism of eruptive rocks of silica. It is composed by quartz, feldspar and mica.
11	<i>White marble</i> Metamorphic compact rock with calcium carbonates (more than 90%). It is predominately white.
12	<i>Basaltic sandstone</i> Sandstones (<15% of matrix) whose content in rock fragments is superior to 25% and higher to the feldspar content. Its origin is basaltic, coming from a volcanic igneous rock characteristic for its dark color and mafic structure.
13	<i>Tertiary Soils</i> Tertiary materials (clays, sands, sandstones and conglomerates) without compaction among the grains.

3.2. Samples Preparation

Samples used to test the thermal conductivity equipment, required a specific preparation whether they are rocks or soils.

3.2.1. Rocky Samples

Rocky samples are cylindrical blocks of 0.05 m in diameter and variable thickness. The preparation of these samples was carried out as follows:

Sample extraction (Figure 5): Using a rotating drilling machine equipped with diamond circular crown, it was possible to obtain cylinder blocks of each of the rocky materials. The diameter of these blocks was of 0.05 m and variable thickness depending on the size of the origin rock. During the process of extraction, the crown was cooled by water.

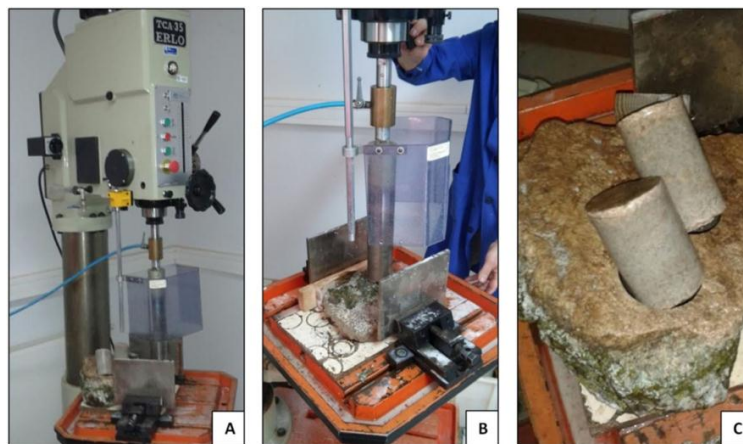


Figure 5. Sample extraction: (A) rock placing; (B) drilling; and (C) final samples.

Carving of samples (Figure 6): Cylinder samples were cut using a cutting-machine supplied with diamond disk to give the samples a certain thickness. Samples of different thickness were prepared, with the aim of analyzing the influence of this factor in the calculation of the thermal conductivity. Thickness of each one of the samples was measured by electronic caliber.

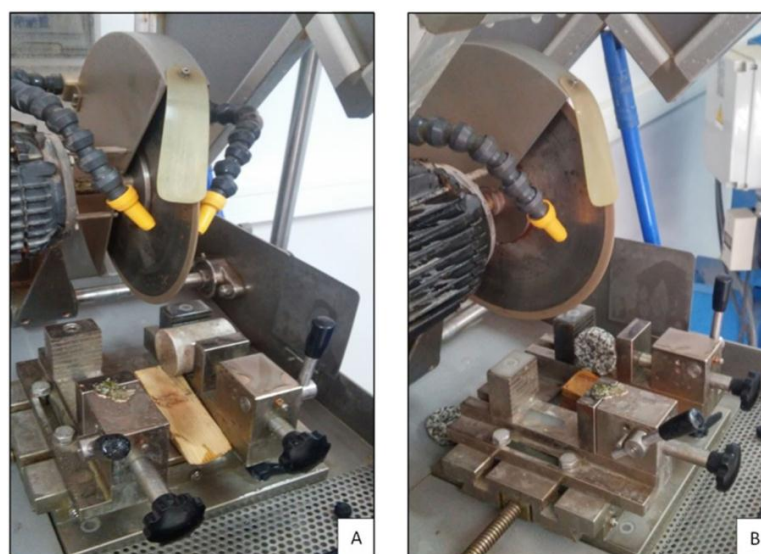


Figure 6. (A) Cutting of a quartzite sample; (B) Cutting of a granitic sample.

Samples cleaning: Samples surfaces were thoroughly cleaned to minimize any possibility of error at the heat transmission. It facilitates the contact with the temperature sounding lines or thermocouples. Figure 7 shows the final appearance of some of these samples.



Figure 7. Samples ready to be used in the suggested equipment.

3.2.2. Soil Samples

Soils materials followed a different procedure to equally get cylinder blocks of 0.05 m in diameter and variable thickness. The preparation of these samples was made according to the next steps.

Determination of humidity by drying in heater [29]: Thermal conductivity depends on the water content that a certain material has, thus, it is important to know the humidity conditions when measuring this parameter. Natural humidity of the soil at its origin was increased by adding a particular percentage of water. The addition of water facilitates the soil compaction in a mold to obtain cylinder samples that will be introduced into the measuring equipment.

Humidity was set to 11.55%.

Soil compaction: As already explained, the proposed system works with cylinder blocks of certain dimensions. Given that, in the case of soils, the material cannot be cut as rocks, it was compacted in a suitable mold. This compaction made easier the obtaining of cylinder samples ready for use in the thermal conductivity device

Soil compaction was made according to the *Proctor Test* conditions, in the point of the optimal humidity defined in the mentioned law [30].

3.3. Placing of Samples in the Measuring Equipment and Determination of Thermal Conductivities

Firstly, one of the aluminum reference samples S_{R1} was introduced in the carrier pipe, then sample S (whose thermal conductivity value wants to be measured) and finally the second aluminum reference sample S_{R2} . It is important to highlight that, before the first reference sample and after the second one, two thin aluminum sheets were placed. The function of these sheets is to get a shielding that avoids convection phenomena. In this way, all the heat transfer just happens by thermal conduction.

Once placed the respective samples and thermocouples in the thermal conductivity equipment, it starts working, sending a constant heat flux that goes through the samples. After letting enough time to make the stabilization of temperatures T_1 , T_2 , T_3 and T_4 , possible, the last step was making the correspondent calculations (as explained by the Section 2.2). Finally, thermal conductivities values of each of the samples were obtained.

Figure 8 shows the measuring equipment expounded over this work and schematized in Figure 1.



Figure 8. Equipment designed to measure thermal conductivities.

4. Analysis of the Measuring Process

Before the measuring of the thermal conductivity parameter in different materials, a series of tests were carried out on the same rocky sample (granite). They were used to analyze how the thermal conductivity changes with the heater temperature and the thickness of the sample in question (S).

These tests allowed the establishment of the appropriate working conditions (temperature of the heat source and sample thickness) to be used in the subsequent measuring of thermal conductivities of the samples presented in Table 2.

4.1. Evolution of the Thermal Conductivity with Temperature

In a crystalline solid, thermal conductivity depends on temperature; however, this dependence is not homogeneous. This dependence can be divided in four regions, so the variation of conductivity will be different based on the region where it is. In region I, of low temperature ($T \leq 20$ K), thermal conductivity quickly increases with temperature, being proportional to T^3 . In region II, it achieves a maximum value, usually at a temperature close to $T \approx \theta_D/20$ (where θ_D is the Debye's temperature). At higher temperatures, in region III, thermal conductivity decreases proportionally to T^{-1} . Finally, at very high temperatures ($T \geq \theta_D$) in region IV, it stops being dependent on temperature [31,32].

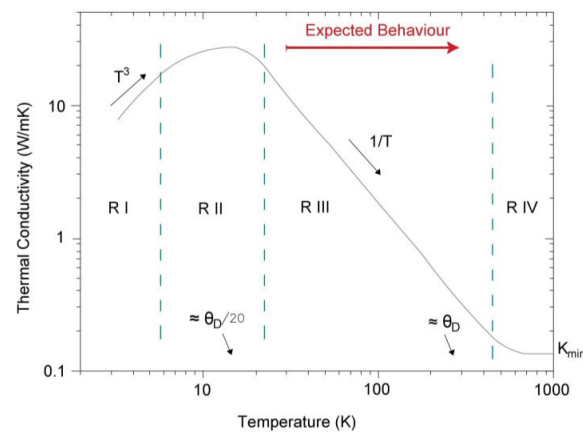
In this particular case, to analyze the behavior of the thermal conductivity with the temperature, the region where the present study is must be defined. To that end, in the first place, Debye's temperature was determined. Table 3 shows the values of Debye's temperatures for a series of substances.

The substance in Table 3 with the most similar composition to the studied material (granite) is silica. It has a Debye's temperature of 645 K, so that, if $T \approx 645 \text{ K}/20 = 32.25 \text{ K}$, in this value, thermal conductivity will get its maximum value and will decrease at higher temperatures until the point of $T \geq 645 \text{ K}$ where it starts being independent of temperature. Therefore, the assumption studied is in region III, which establishes an inversely proportional relation between temperature and thermal conductivity.

Figure 9 presents the distribution of the regions and the behavior that the system should have in the area where it is, region III.

Table 3. Debye's temperatures for a series of substances.

Substance	Debye Temperature (K)
Aluminum	428
Cadmium	209
Chromium	630
Copper	343.5
Gold	165
Iron	470
Lead	105
Manganese	410
Nickel	450
Platinum	240
Silicon	645
Silver	225
Tantalum	240
Tin	200
Titanium	420
Wolfram	400
Zinc	327
Carbon	2230
Ice	192

**Figure 9.** Dependence of thermal conductivity with temperature for a crystalline solid.

Graphically, taking from the graphic in Figure 9 two temperatures and the corresponding values of thermal conductivity according to the curve represented in region III, it is possible to establish the reduction of thermal conductivity with the temperature in that region. Thus, if we select the temperatures of 95 K and 100 K, we can verify the decrease of thermal conductivity with an increase of temperature of 5 K. In this way, and according to Debye's graphical, for the temperature of 95 K, the corresponding value of thermal conductivity is 2.01 W/mK, while for 100 K the value of thermal conductivity is 1.87 W/mK. It means that an increase of 5 K of temperature involves a reduction of thermal conductivity of 0.14 W/mK. That is to say, in a crystalline solid, per each grade of temperature increased, thermal conductivity decreases 0.028 W/mK.

Nevertheless, the granitic sample used to examine the variation of the thermal conductivity with the temperature, contains about 30% of silica. Thus, the reduction for this sample would not be of 0.028 W/mK (for a material constituted by 100% of silica), but 0.0084 W/mK per each grade of temperature increased in this material.

The above is the expected behavior of thermal conductivity based on theoretical knowledge. However, to know what really happens in the practice of this procedure, tests with this method were carried out at different working temperatures (313.15 K, 338.15 K and 358.15 K) and always with the same sample (granite whose content in silica is around 30%). It will allow determining the evolution of the thermal conductivity with these temperatures (Table 4).

4.2. Variation of the Thermal Conductivity with the Sample Thickness

Another of the tests consisted in analyzing the variation of the thermal conductivity parameter with different thicknesses of the sample S . In this way, the range of sample thickness for which the equipment properly worked was established, discarding those ones where the results obtained moved away from the reference values. Thus, through these tests, the limits of the system regarding the sample thickness were set.

As in the previous case, a series of measurements were made with the same granitic sample modifying in this case its thickness.

The results of these tests (modifying the working temperatures and the sample thickness) are described in Table 4.

Table 4. Thermal conductivity for different values of temperature and sample thickness.

Heat Source Temperature (K)	Sample Thickness (m)	Sample Medium Temperature (K)	Heat Flux "Q" (W)	Thermal Conductivity "k" (W/mK)
313.15	0.0061	295.95	8.13	2.11
	0.0090	295.92	5.42	2.16
	0.0131	295.80	5.42	2.13
	0.0162	299.55	8.13	5.59
338.15	0.0061	296.20	19.22	2.06
	0.0090	302.30	13.31	2.10
	0.0131	302.30	9.24	2.13
	0.0162	305.55	18.49	4.77
358.15	0.0061	305.90	23.58	1.98
	0.0090	309.90	22.43	2.09
	0.0131	308.15	17.83	2.12
	0.0162	306.10	21.28	4.75

Analyzing the results presented in Table 4 and focusing on the variation of the working temperature, thermal conductivity was measured for three values of working temperature and for each of these cases, four thickness of the same granitic sample. Tables 5 and 6 show the variation of the thermal conductivity for each thickness when the working temperature, increases from 313.15 K to 338.15 K and from 313.15 K to 338.15 K. Additionally, Tables 5 and 6 present the decrease of the thermal conductivity parameter for each grade that the sample temperature increases.

Table 5. Evolution of thermal conductivity when temperature increases from 313.15 K to 338.15 K.

Increase of Temperature from 313.15 K to 338.15 K			
Thickness (m)	Increase of "T" among Samples (°)	Difference of "k" (W/mK)	Decrease by Grade (W/mK)
0.0061	0.25	0.05	0.2000
0.0090	6.38	0.06	0.0094
0.0131	6.50	0.00	0.0000
0.0162	6.00	0.82	0.1366

Table 6. Evolution of thermal conductivity when temperature increases from 338.15 K to 358.15 K.

Increase of Temperature from 338.15 K to 358.15 K			
Thickness (m)	Increase of "T" among Samples (K)	Difference of "k" (W/mK)	Decrease by Grade (W/mK)
0.006	9.70	0.08	0.0082
0.009	7.60	0.01	0.0013
0.013	5.85	0.01	0.0017
0.016	0.55	0.02	0.0360

Some deductions can be drawn:

- Most of the thermal conductivity values are around 2 W/mK. Increasing the temperature, these values decrease as it was expected for a crystalline solid. However, the reduction of the thermal conductivity parameter is not constant in the different sample thicknesses, and is not the expected 0.0084 W/mK calculated in Section 4.1. Evolution of thermal conductivity with temperature. Therefore, in a crystalline solid, thermal conductivity decreases when temperature grows. However, it was not possible to set a model of behavior of this reduction because it does not follow any constant pattern.
- Regarding the different sample thickness, Table 4 shows the measurements carried out at the laboratory equipment (four different thicknesses for each one of the three work temperatures). The optimal dimensions of the sample *S* could be established based on the results of these measurements. Thus, analyzing Table 4, it can be observed that, for the three temperatures, the values of thermal conductivity for each of the thickness are around the same value (~2 W/mK). These data agree with the expected thermal conductivity value for a granitic material. However, for the case of the highest thickness, the result of thermal conductivity moves away from the rest of results for lower thickness. All this made it possible to set the sample thicknesses for which the present method works properly. On the basis of these results, with sample thicknesses greater than 0.0131 m, the procedure described in this paper does not provide reliable values. In these cases, results are highly anomalous due to a high dissipation of the heat flux through the sample *S*.
- The following working conditions were established in the thermal conductivity apparatus:
 - Temperature of the heat source was set in 313.15 K. Although results were acceptable in the three temperatures (313.15 K, 338.15 K and 358.15 K), this value is closer to the ground temperature in a very low enthalpy geothermal installation.
 - Thickness of the sample *S* could not exceed in any case the mentioned 0.0131 m for the reasons previously justified.

5. Thermal Conductivity Results

The results of thermal conductivity measurements are presented in Table 7. Three measurements of this parameter were carried out on each of the samples considered.

Table 7. Thermal conductivities, standard deviation and derivate error for each material studied.

Sample	Thickness (m)	K ₁ (W/mK)	K ₂ (W/mK)	K ₃ (W/mK)	Medium K <x> (W/mK)	Standard Deviation σ_x	Derivate Error
1	0.0090	2.16	2.12	2.09	2.12	0.035	±0.10
2	0.0102	3.58	3.42	3.51	3.50	0.080	±0.10
3	0.0131	2.57	2.61	2.39	2.52	0.117	±0.10
4	0.0103	2.15	2.11	2.13	2.13	0.020	±0.10
5	0.0130	2.24	2.31	2.25	2.27	0.038	±0.10
6	0.0124	3.11	3.10	3.18	3.13	0.043	±0.10
7	0.0132	2.99	2.97	3.02	2.99	0.025	±0.10
8	0.0085	0.53	0.53	0.51	0.52	0.012	±0.10
9	0.0051	0.21	0.18	0.19	0.19	0.016	±0.10
10	0.0082	3.15	3.24	3.17	3.19	0.047	±0.10
11	0.0091	3.39	3.38	3.44	3.40	0.032	±0.10
12	0.0063	2.63	2.59	2.61	2.61	0.020	±0.10
13	0.0090	1.77	1.58	1.63	1.66	0.098	±0.10

When a certain measurement is repeated several times, medium values group around a central value. This distribution can be described by statistical the mean $\langle x \rangle$ (Equation (3)) and the standard deviation σx (Equation (4)) [33,34]:

$$\langle x \rangle = \frac{1}{N} \sum_{i=1}^N x_i = \frac{1}{N} (x_1 + x_2 + x_3 + \dots + x_{N-1} + x_N) \quad (3)$$

$$\sigma x = \sqrt{\frac{1}{N-1} \sum_{i=1}^N (x_i - \langle x \rangle)^2} \quad (4)$$

Errors derived from the precision of the tools used to measure the different parameters (thickness and temperatures) must also be considered. For this reason, the total differential of our equation of calculation of conductivity was calculated (Equation (5)):

$$dk = \frac{\partial K}{\partial e} de + \frac{\partial K}{\partial T_{2-3}} dT_{2-3} + \frac{\partial K}{\partial T_{3-4}} dT_{3-4} \quad (5)$$

where $k = k(e, T_{2-3}, T_{3-4})$

where

k = Thermal conductivity (W/mK);

e = Sample thickness (m);

T_{2-3} = Increase of temperature between thermocouples 2 and 3 (Figure 1); and

T_{3-4} = Increase of temperature between thermocouples 3 and 4 (Figure 1).

Equation (5) was transformed into increases, and absolute values were taken to each partial derivative to estimate the derivative error (Equation (6)):

$$\Delta k = \left| \frac{\Delta k}{\Delta e} \right| \Delta e + \left| \frac{\Delta k}{\Delta T_{2-3}} \right| \Delta T_{2-3} + \left| \frac{\Delta k}{\Delta T_{3-4}} \right| \Delta T_{3-4} \quad (6)$$

Increases represent the absolute errors of the measuring dispositive and the growth of k symbolizes the derivative error.

From each one of the three thermal conductivities measurements of each material, derivative error was calculated. It was found that the reduction in precision did not exceed in any case one order of magnitude with respect to the precision of the least precise dispositive (thermocouples with ± 0.01 K). Derivative error was estimated as ± 0.10 for all samples.

6. Validity of the Method

The validity of the suggested thermal conductivity device was analyzed by comparing the results presented in Table 7 with the ones commonly accepted at the “*Technical Code of Building*” (CTE). From this comparison, the difference, with respect to the officially accepted value for that sample, was calculated.

CTE provides a certain thermal conductivity value for a wide variety of materials, including rocks and soils. Given the heterogeneity of samples 1, 2, 3 and 4 (granitic rocks) and the high presence of these rocks in numerous European geothermal installations, a different reference value was assigned to each of these samples starting from the CTE value. This assignment was made based on studies that relate the thermal conductivity of a rock with its quartz content [35,36].

Thus, instead of the thermal conductivity value of 3 W/mK provided by the CTE for a granite rock, an interval of 2.0 W/mK–3.8 W/mK was taken for quartz contents between 3% and 50%.

Table 8 shows the reference thermal conductivity values set for each of the samples according to CTE. It also presents the thermal conductivity values measured with the equipment developed in this research and the difference between both values (common and measured values).

It is important to highlight that thermal conductivity parameter easily changes depending on different factors (temperature, anisotropy, humidity, etc.) and could be quite different in materials

of similar geological origin. Despite these facts, the differences between the measured values and the reference ones are considerably low. The most unfavorable case was for sample 12 (0.37 W/mK of difference) and the most favorable one for sample 7 (0.01 W/mK of difference). Figure 10 shows a graphic of deviations presented in Table 8.

Table 8. Comparison between values of thermal conductivity measured and the reference ones.

Sample	K Measured (W/mK)	K Reference (W/mK)	Deviation-Reference Value (W/mK)
1	2.12	2.4 (15% of quartz)	0.28
2	3.50	3.2 (40% of quartz)	0.3
3	2.52	2.7 (25% of quartz)	0.18
4	2.13	2.1 (5% of quartz)	0.03
5	2.27	2.2	0.07
6	3.13	3	0.13
7	2.99	3	0.01
8	0.52	0.56	0.04
9	0.19	0.12	0.07
10	3.19	3.5	0.31
11	3.40	3.5	0.1
12	2.61	3	0.37
13	1.66	2	0.34

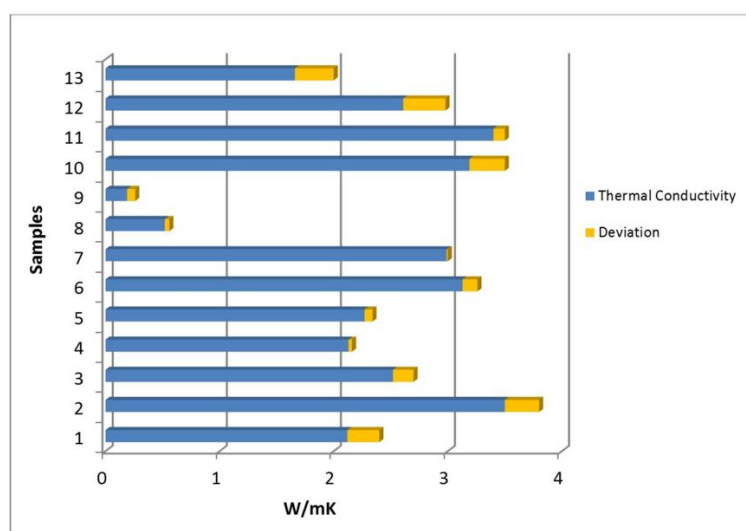


Figure 10. Deviations between the common and the measured thermal conductivity values.

7. Influence of the Thermal Conductivity Parameter in the Geothermal Measuring

“Earth Energy Designer” (EED) is software developed by “Blocon Software” that allows knowing the total drilling depth of a vertical closed-loop system. The calculation process of EED is based on a series of initial data (provided by the user) of the ground where the installation is going to be placed. One of these initial data is the thermal conductivity of the surrounding ground.

In order to understand the importance that the thermal conductivity parameter has in the dimensioning of a geothermal installation, some calculations were made with this software. The total drilling depth was calculated with the same conditions but changing the thermal conductivity of the ground for each of the samples of this study. Thus, for each sample, calculations were made with the measured thermal conductivity value and with interval of $\pm 15\%$ of that value (Table 9).

Depending on the material, with a variation of only $\pm 15\%$ in the thermal conductivity value, the total drilling length significantly changes. The most notable case is sample 9, where increasing the

thermal conductivity from 0.19 W/mK to 0.22 W/mK, the drilling length decreases 27 m, and reducing the thermal conductivity to 0.16 W/mK, the drilling length increases 34 m.

As Table 9 shows, for the rest of samples, large variations are also experimented. Therefore, a proper knowledge of the thermal conductivity of the ground means important variations in the total drilling depth of a very low enthalpy geothermal installation.

Table 9. Measuring with EED software.

EED Measuring			
Sample 1			
k (W/mK)	2.12 *	1.80 **	2.44 ***
Drilling length (m)	185	194	174
Sample 2			
k (W/mK)	3.50 *	2.97 **	4.02 ***
Drilling length (m)	153	161	143
Sample 3			
k (W/mK)	2.52 *	2.14 **	2.90 ***
Drilling length (m)	174	182	163
Sample 4			
k (W/mK)	2.13 *	1.81 **	2.45 ***
Drilling length (m)	185	193	173
Sample 5			
k (W/mK)	2.27 *	1.93 **	2.61 ***
Drilling length (m)	181	192	172
Sample 6			
k (W/mK)	3.13 *	2.66 **	3.60 ***
Drilling length (m)	160	170	151
Sample 7			
k (W/mK)	2.99 *	2.54 **	3.44 ***
Drilling length (m)	163	174	154
Sample 8			
k (W/mK)	0.52 *	0.44 **	0.60 ***
Drilling length (m)	324	333	316
Sample 9			
k (W/mK)	0.19 *	0.16 **	0.22 ***
Drilling length (m)	449	483	422
Sample 10			
k (W/mK)	3.19 *	2.71 **	3.67 ***
Drilling length (m)	158	169	150
Sample 11			
k (W/mK)	3.40 *	2.89 **	3.91 ***
Drilling length (m)	155	166	147
Sample 12			
k (W/mK)	2.61 *	2.22 **	3.00 ***
Drilling length (m)	172	183	163
Sample 13			
k (W/mK)	1.66 *	1.41 **	1.91 ***
Drilling length (m)	206	216	196

* Measured thermal conductivity value; ** −15% of the measured thermal conductivity value; *** +15% of the measured thermal conductivity value.

8. Conclusions

Thermal conductivity is a fundamental property in the process of measuring of a geothermal installation but, at the same time, it is a parameter of difficult quantification.

A series of difficulties of diverse nature appeared throughout this research:

- Temperature of the cold source (ambient temperature) was controlled at all times and set in a constant value to avoid external thermal influences on the thermal conductivity device.
- The insulation placed around the system minimized the radial heat flux but did not eliminate it. As a result, corrections of this heat flux were made for each of the temperatures set at the heater.
- A high precision was needed when measuring the sample thickness due to its excessive influence in the calculation of the thermal conductivity parameter.
- Thermocouples were carefully placed to ensure a complete contact with the faces of the samples.
- Long-term measurements to guarantee the stabilization of the heat flux.

Despite these facts, this method means an excellent solution when measuring the thermal conductivity parameter in an economical and simple way. It provides accurate results taking advantage of equipment present in most laboratories such as the heater. These results are applicable to the calculation of very low enthalpy geothermal installations, avoiding over measuring that raise the price of these renewable installations.

Finally, EED software allowed highlighting the importance of knowing the thermal conductivity of the surrounding ground in a geothermal system. This knowledge could mean important savings regarding the drilling length.

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