

Article

Calcite Twinning in the Ordovician Martinsburg Formation, Delaware Water Gap, New Jersey, USA: Implications for Cleavage Formation and Tectonic Shortening in the Appalachian Piedmont Province

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Abstract: A traverse across the Stone Church syncline in the Ordovician Martinsburg turbidites reveals an axial planar cleavage (N40°E, SE dips) in regional thrust-related folds (N40°E, shallow plunges) and five phases of sparry calcite. Calcite fillings are bedding-parallel, cleavage-parallel, and one vein set cross-cuts both earlier phases; the youngest calcite filling is a bedding-parallel fault gouge that crosscuts the cleavage and preserves top-down-to-the-southeast normal fault kinematics. Calcite veins unique to disharmonically-folded calcareous siltstones (Maxwell, 1962) were also analyzed. Stable isotopic analysis (O, C) of all of the calcite phases indicates a uniform fluid source (δ^{13} C -2.0, $\delta^{18}O - 13.3$ VPDB) and, potentially, a similar precipitation and mechanical twinning age. The twinning strains (n = 1341; average $\Delta \sigma = -32$ MPa; average $\varepsilon_1 = -2.9\%$) in the calcite suite are consistent with SE-NW thrust shortening, and sub-horizontal shortening perpendicular to evolving axial planar cleavage planes in the Stone Church syncline. Calcareous siltstone layers within the Martinsburg Fm. turbidites share concordant bedding planes and are unique, chemically (XRF), but folded and cleaved differently than the surrounding clay-rich Martinsburg turbidites. Neither sediment type yielded detrital zircons. Electron backscatter X-ray diffraction (EBSD) and calcite twinning results in a folded calcareous siltstone layer preserving a layer-normal SE-NW shortening strain and Lattice Preferred Orientation (LPO). Shortening axes for the five-phase calcite suite trends ~N40°W, consistent with tectonic transport associated with crystalline nappe emplacement of the Reading Prong within the Piedmont province.

Keywords: structural geology; Appalachian Piedmont; strain analysis

1. Introduction

Since Chapple [1] first unraveled the order and dynamics of shortening in a thin-skinned fold-and-thrust belt, using the Appalachian Mountains as a model, later workers have made enormous headway in understanding the deformational style in the Valley and Ridge (thin-skinned thrust belt) and foreland provinces. Recent studies in the Valley and Ridge and foreland provinces, which are separated by an oroclinal boundary, has generated controversy about the role of thrust belt translation and rotations relative to the adjacent, stable foreland [2,3]. There is no calcite twinning strain data from the internal, thick-skinned Piedmont province. The numerous calcite fillings (phases) in the folded



and cleaved Ordovician Martinsburg Fm. in the Stone Church syncline provide a field site to evaluate the stress-strain relations that existed during the evolution of this fold and its axial planar cleavage and add the twinned calcite results to the existing literature on deformation and strain partitioning at this classic Piedmont province locale [4–7]. How does this strain fabric correlate with other indicators of tectonic shortening across the Appalachian orogenic belt and did the collisional orogeny involve significant thrust sheet rotation during shortening [8–10]?

Field relations suggest that the five-phases of calcite precipitated, and deformed plastically, while the host Ordovician Martinsburg Fm. was being folded and cleaved as the crystalline Taconic Reading Prong nappe moved to the northwest on the Martic thrust [11–16]. Sparry calcite is present as bedding-parallel veins, cleavage-parallel veins, and veins that cross-cut both. Also present are calcite veins unique to disharmonically-folded and cleaved calcareous siltstone layers (previously thought to be clastic dikes as the Martinsburg Fm. "dewatered"; see [4-7]) and we discovered a bedding-parallel, layered calcite fault zone with normal fault kinematics. We have compared the bulk geochemistry of the Martinsburg Fm. host and a calcareous siltstone (XRF), the stable isotope (O, C), and calcite twinning strain characteristics of each calcite phase in a section of folded Martinsburg Fm. turbidites exposed parallel to the Taconic orogeny thrust transport direction (SE–NW) along the Delaware River in New Jersey. Our goal was to determine if, (1) the many calcite vein phases hosted in the Martinsburg Fm. precipitated from a common fluid source and preserve twinning strains that are consistent with other observations of Taconic (or younger) shortening; (2) the relationship of the stress-strain field, as preserved by calcite precipitated parallel to the Martinsburg Fm. axial planar cleavage, is consistent with the results of Craddock et al. ([17]; horizontal shortening normal to the cleavage strike, horizontal extension parallel to strike) in the Gondwanide belt, and (3) the shortening strain fabrics in the Cambro-Ordovician sediments deformed in the thick-skinned Piedmont province of the Taconic orogeny remain parallel to the inferred thrust transport directions of younger (Acadian, Alleghenian) shortening events to the northwest in the Valley and Ridge and foreland provinces.

2. Previous Work

Structural studies in the Appalachian Mountains have revealed a variety of observations about deformation mechanisms and thrust belt mechanics, including finite strain studies by Cloos [18] and the role of pressure solution in the removal of rock materials during cleavage formation [19] in the Valley and Ridge province. Strain studies in the frontal Pine Mountain block record a layer-parallel shortening strain oriented parallel (NW–SE) to the inferred thrust transport direction and this fabric is parallel in the adjacent autochthonous footwall carbonates suggesting the thrust sheets moved without rotation [20–22]. Ong *et al.* [8] and Hnat and van der Pluijm [22] have documented the shortening strain parallelism between the thrust belt and foreland in additional Appalachian locations. Appalachian thrust belt deformation is also recorded in the foreland as salt-cored anticlines (200 km inboard from the thrust front; [23]), finite strains (200 km; [24,25]) and calcite twinning strains (2000 km; [26,27]) in the Paleozoic sediments of cratonic North America. The Delaware Water Gap exposures represent a rare traverse across the poorly exposed Piedmont province.

The internal (hinterland) Piedmont portion of the Appalachian Mountains is characterized by Cambrian (limestone)-Ordovician marine (Martinsburg Fm. flysch; 12,000 feet) sediments that were deformed by the late Ordovician Taconic orogeny. Geologic mapping by Drake *et al.* ([28,29] is the basis for the most recent local fieldguide prepared by Epstein ([30]; see geologic map and stratigraphic column) along the Delaware Water Gap. Folds trend ~N40°E with shallow plunges and are often underlain by SE-dipping thrusts. Minor folds, of many geometries, are observed and multiple, local cleavages are reported in the Martinsburg Fm. except along the Delaware Water gap section where only the axial-planar cleavge is observed. Grenville basement rocks, including marble inliers, were mobilized as thrust-cored nappes (Wilmington complex, Reading Prong) at ~450 Ma deforming the Cambrian-Ordovician marine section along thrust faults that dip to the SE and transported these nappes to the NW [14,31,32]. These structures, including various Taconic klippe [33], were

immediately buried by Taconic (450–443 Ma) molasse deposits, followed by deposition of the Silurian Shawangunk (Tuscarora sands)–Bloomsberg Fms. (redbeds) and Devonian Marcellus Shale [14] and constitute the hinterland Piedmont province of the Appalachians (Figure 1). The Delaware River valley exposes the Stone Church syncline (with minor folds), largely in the Martinsburg Fm. turbidites (Figure 1), which has revealed the interplay of folding [4], pressure solution volume loss due to cleavage formation [5–7,34], framboid-pressure shadow evolution in cleavage lithons [5,34], finite strain measurements [4,6,7], and AMS studies [35–37] toward developing the concept of strain partitioning in deformed, supracrustal sediments. Mechanical twinning in calcite is another piece of the strain history puzzle and there are numerous phases of sparry calcite veins to measure and interpret in this Piedmont structure relative to younger Acadian (Devonian) and Alleghenian (Permian) convergence and younger-still Triassic extension.



Figure 1. LIDAR base map of the Delaware Water Gap area (**B**) in New Jersey and Pennsylvania relative to the Appalachian Mountains (inset **A** with box for **B**). Sample sites are numbered. Cross section line (A-B) is from Maxwell [4] and in Figure 3; see also Herman *et al.* ([32], Figure 14 for regional subsurface structural relations).

3. Methods

3.1. Geochemistry

Stable isotopes were analyzed at the University of Michigan stable isotope laboratory. Drilled carbonate powder samples weighing a minimum of 10 micrograms were placed in stainless steel boats. Samples were then placed in individual borosilicate reaction vessels and reacted at 77 \pm 1 °C with four drops of anhydrous phosphoric acid for 8 min (a total of 12 min for dolomites, 17 min for apatite, and 22 min for siderites) in a Finnigan MAT Kiel IV preparation device coupled directly to the inlet of a Finnigan MAT 253 triple collector isotope ratio mass spectrometer. O¹⁷ corrected data are corrected for acid fractionation and source mixing by calibration to a best-fit regression line defined by two NBS standards, NBS 18 and NBS 19. Data are reported in % notation relative to VPDB. Precision and accuracy of data are monitored through daily analysis of a variety of powdered carbonate standards. At least four standards are reacted and analyzed daily. Measured precision is maintained at better than 0.1‰ for both carbon and oxygen isotope compositions.

Martinsburg Fm. turbidites and calcareous siltstone materials were analyzed using X-ray flouresence methods at McGill University, Montreal. These materials were also powdered and

separated at Macalester to extract detrital zircons; none were recovered. We also attempted to use U-Pb laser ablation methods to date the syn-orogenic calcite but this was unsuccessful due to low U levels and high levels of ²⁰⁴Pb [38].

3.2. Calcite Twin Analysis

The Calcite Strain-Gage Technique (CSGT) of Groshong [39] allows investigation of intraplate stresses as constrained by intracrystalline twinning of rock-forming calcite grains. Although the result is actually a strain tensor, a similar orientation of the stress tensor appears likely in case of coaxial deformation [40,41]. The CSGT has been used to constrain strain tensor directions in veins [21,42,43], limestones [20,26,44–49], marble [50,51], amygdaloidal basalts [52–54], and lamprophryres [55].

Under temperatures of *ca.* 200 °C intracrystalline deformation of calcite results in the formation of e-twins. The formation of calcite e-twins requires a shear stress exceeding *ca.* 10 MPa [56–59]. Calcite offers three glide systems for e-twinning. From U-stage measurements of width, frequency, and orientation of twins, and the crystallographic orientation of the host crystals, a strain tensor can be calculated using a least-squares technique [39]. In order to remove "noise" from the dataset, a refinement of the calculated strain tensor can be achieved by stripping 20% twins with highest deviations [60]. This procedure has been used if the number of measured grains were large (n > 20). In cases where the data appear to be inhomogeneous, the separation of incompatible twins ("NEV" = negative expected values) from compatible twins ("POS" = positive expected values) of the initial dataset allows separate calculation of two or more least-squares deviatoric strain tensors. Thus, the CSGT can be used to obtain information on superimposed deformations [39,61] and differential stress magnitudes [62,63].

The validity of this stripping procedure was demonstrated in experimental tests where the reliability depends on the overall complexity of deformation and the number of grains with twins [61,64]. The stripping procedure was used in cases of high proportions of NEV and a large number of measured grains. An experimental re-evaluation of the CSGT has shown that measurements of about 50 grains on one thin-section or 25 grains on two mutually-perpendicular thin sections yield the best results [60,65,66]. The chance to extract the records of more than two deformations from one dataset is limited when dealing with natural rocks [57]. Individual analyses of veins, matrix, nodules, *etc.*, allows the acquisition of several strain tensors without applying statistical data stripping. The complexity of rotational strains in fault zones has limited the application of this method to the efforts of Gray *et al.* [67]. Application of the CSGT requires the following assumptions to be valid: (1) low temperatures (dominance of Type I and Type II twins); (2) random c-axis orientations of calcite; (3) homogenous strain; (4) coaxial deformation; (5) volume constancy; (6) low porosity materials, and (7) low strain (<15%). If these conditions are not fully met, the underlying dataset of the calculated strain tensor could be biased, modified, or random [66]. Strain tensors were calculated from calcite e-twin datasets using the software package of Evans and Groshong [65].

3.3. EBSD Methods

Thin sections were polished using a Buehler Mastermet colloidal silica and Buehler grinderpolisher for 1.5 h in Beijing. The Lattice Preferred Orientation (LPO) data acquisition was finished on a Hitachi S-3400N-II (Singapore, Singapore) scanning electron microscope mounted with a HKL Channel 5 EBSD detector with the thin section surface inclined at 70° to the incidental beam. Acceleration voltage was 15 kV and the working distance was 18.4 mm. Optic axis measurements were done on calcite grains using the interactive mode, *i.e.*, the data were collected grain by grain, manually with six to eight kikuchi bands detected. EBSD analysis is finished using the CHANNEL5 analysis software. All the LPO data are plotted on lower hemisphere equal-area Schmidt projections.

4. Results

4.1. Stable Isotopes

Various calcite elements were sampled and powdered along the traverse (Figure 1) and these results are presented in Table 1. The Cambrian Kittatiny limestones are consistent with other early Paleozoic marine carbonates and are consistent across the limb of the syncline (samples 1 and 18). All the various calcite veins and fault fillings are also remarkably uniform ($\delta^{13}C - 2.0$, $\delta^{18}O - 13.3$ VPDB) across the syncline, are not derived locally from the Kittatiny limestones and have an isotopic signature suggestive of calcite precipitation from warm metamorphic fluids ([68]).

Sample	δ13C (VPDB)	δ18O (VPDB)	Туре
1	-0.98	-6.60	Kittatiny Limestone
2	-1.63	-11.35	Martinsburg Fm. vein
4	-1.27	-13.31	Martinsburg Fm. vein
5	-1.88	-13.20	Martinsburg Fm. vein
6	-2.29	-13.68	Martinsburg Fm. vein
9	-1.69	-13.37	Martinsburg Fm. vein
10	-1.82	-13.58	Martinsburg Fm. vein
11	-1.81	-13.05	Martinsburg Fm. vein
12(A)	-1.59	-13.20	Martinsburg Fm. vein
13	-1.72	-13.10	Martinsburg Fm. vein
14	-1.75	-13.11	Martinsburg Fm. vein
15	-2.06	-13.24	Martinsburg Fm. vein
16	-2.00	-13.70	Martinsburg Fm. vein
17(B)	-2.11	-12.80	Martinsburg Fm. vein
18	-0.56	-6.54	Kittatiny Limestone

Table 1. Stone Church Syncline Stable Isotope Results.

4.2. Martinsburg Fm. Traverse

We collected oriented calcite phases, bracketed by the underlying Cambrian Kittatiny limestones along a SE-NW traverse through the Stone Church syncline (Figures 1–3; Table 2). The Kittatiny limestones preserve a layer-parallel shortening strain in the direction of tectonic transport with modest strain values and no strain overprint. Synorogenic calcite is found as bedding-parallel veins and gouge zones, cleavage-parallel veins and veins that crosscut bedding, and cleavage-parallel veins in the Martinsburg Fm. turbidites (Figure 2). All of these calcite elements preserve a NW–SE layer-parallel shortening strain (n = 14 strain results) with extension normal to Taconic transport, or parallel to the strike of the axial planar cleavage (N45°E). Specifically, for calcite precipitated in a cleavage plane, the shortening axis is normal to the cleavage plane with horizontal extension parallel to the strike (Figure 4). Strain values are modest, there is no strain overprint (low NEVs), and the differential stresses responsible for twinning was -323 bars (-32 MPa).

Across from the powerplant, along the Delaware River and Highway 46 (GPS: $40^{\circ}54'43.95''$ N, $75^{\circ}4'37.04''$ W), we observed a series of en echelon, calcite-filled veins (see [43]) that merge with a bedding-parallel calcite-filled fault zone (Figure 3a,b). The en echelon array is filled with sparry calcite which merges into the layers of the fault zone above. Both the en echelon array and fault zone kinematics indicate a normal fault offset (N40°E, 40°SE; Top-down-to-the SE). Six twinning samples record a NW–SE sub-horizontal shortening that is within the plane of the fault (which is parallel to bedding; Figure 4). We combined the six normal fault data sets into one (n = 161) and separated positive (n = 135) and negative expected values (n = 27). The 27 NEV grains represent 14% of the total and would normally be purged with routine data cleaning [39,60,61]. The contrived NEV result records a subtle twinning strain overprint of sub-vertical shortening (Figure 5). The axial planar cleavage does not crosscut the fault plane or en echelon veins.



Figure 2. Photo collage of field photos along Maxwell's [4] cross-section (see Figure 1). (**A**,**B**) SE-dipping, bedding-parallel normal fault on the northwest limb of the Stone Church syncline; (**C**,**D**) bedding and cleavage-parallel calcite fillings; (**D**–**G**) the folded, vertical calcareous siltstone layer with horizontal, bedding-normal calcite veins and the locations of the XRF samples (Table 3). The calcareous siltstone layer is part of the Martinsburg turbidite sequence and the cleavage is sheared along the sedimentological interface (F,G). Dark lines are stylolites and cleavage selvages. Red lines are the orientations of the ϵ 1 shortening axis for each site.



Figure 3. Cont.



Figure 3. (a) Normal fault gouge zone with folded fault drag kinematics. Shear zone is grooved (**A**; top view) and, in cross-section, reveals normal faulted folds (**B**) shown in thin section (**C**; full section in plane light) with close-up of fault drag and calcite twins (**D**; plane light). Coin for scale (3.8 mm diameter); (**b**) top photo of striated, bedding-parallel normal fault zone with steps down-to-the-SE (**A**), and a cross-section of the same sample (**B**) showing normal fault offsets, listric to the local shear zone base.





Figure 4. Lower hemisphere projections of calcite twin shortening axes for the various calcite phases, Delaware Water Gap, New Jersey. See Table 2 for details.

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Sample	Rock Unit	Orientation (Bedding)	Orientation (Cleavage)	Orientation (vein)	Grains (n =)	e1	e2	e3	e1 (%)	NEV (%)	-Δσ (bars)	Fabric Interp. (vein)	Fabric Interp. (Bedding)	Comment
							Southeast							
1	Kittatiny LS	N45°E, 38°SE	None	None	36	171°, 40°	345°, 62°	88°, 6°	-3.8	19	377.76021	_	LPS	_
2	Om, qtz vein	N45°E, 84°SE	N45°E, 42°SE	_	-	-	_	-	-	-	_	-	_	Fold, no twins
3-1A	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	30	341°, 7°	163°, 87°	53°, 12°	-2.9	10	373.300662	VPS	LNS	Fold (Calcareous siltstone)
3-1B	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	23	101°, 3°	281°, 87°	42°, 6°	-1.4	21	366.464023	VPS	LNS	Fold (Calcareous siltstone)
3-2A	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	8	322°, 4°	143°, 88°	253°, 6°	-1.3	0	370.339565	VPS	LNS	Fold (Calcareous siltstone)
3-3A	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	28	155°, 4°	318°, 88°	$\frac{48^{\circ}}{7^{\circ}}$	-1.7	10	373.7829	VPS	LNS	Fold (Calcareous siltstone)
3-4A	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	42	138°, 6°	317°, 86°	47°, 16°	-1.8	21	360.67426	VPS	LNS	Fold (Calcareous siltstone)
3-4B	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	42	148°, 5°	321°, 88°	$54^{\circ},$ 6°	-1.5	12	373.7829	VPS	LNS	Fold (Calcareous siltstone)
3-5A	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	34	318°, 11°	141°, 87°	35°, 9°	-1.1	29	367.25548	VPS	LNS	Fold (Calcareous siltstone)
3-5B	Om, vein	N45°E, 84°SE	N45°E, 45°SE	Horizontal	16	99°, 8°	33°, 4°	272°, 88°	-1.1	0	402.519949	VPS	LNS	Fold (Calcareous siltstone)
3: All-PEV	Om, veins	N45°E, 84°SE	N45°E, 45°SE	Horizontal	223	158°, 4°	221°, 12°	342°, 87°	-7.2	0	383.247517	VPS	LNS	Fold (Calcareous siltstone)
3: All-NEV	Om, veins	N45°E, 84°SE	N45°E, 45°SE	Horizontal	26	36°, 6°	326°, 2°	147°, 88°	-3.6	100	358.639492	VPS	LPS	Fold (Calcareous siltstone)
4	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	29	330°, 12°	225°, 32°	62°, 3°	-3.7	17	405.4012551	VPS	LPS	Bedding-parallel normal fault
5	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	33	155°, 30°	326°, 47°	$\begin{array}{c} 65^{\circ}, \\ 4^{\circ} \end{array}$	-1.4	18	259.142013	VPS	LPS	Bedding-parallel normal fault
6	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	36	162°, 21°	252°, 12°	6°, 83°	-1.7	5	276.1623538	VPS	LPS	Bedding-parallel normal fault
7	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	36	349°, 5°	249°, 83°	71°, 6°	-2.1	1	302.2918232	VPS	LPS	Bedding-parallel normal fault
8	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	29	342°, 22°	223°, 61°	68°, 5°	-1.1	10	297.7502294	VPS	LPS	Bedding-parallel normal fault
ALL-PEV	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	163	163°, 2°	253°, 43°	43°, 5°	-2.7	0	304.256842	VPS	LPS	Bedding-parallel normal fault

 Table 2. Martinsburg Fm. Cleavage-Clastic Dike Calcite Twinning Strain Results.

Table 2. Cont.

Sample	Rock Unit	Orientation (Bedding)	Orientation (Cleavage)	Orientation (vein)	Grains (n =)	e1	e2	e3	e1 (%)	NEV (%)	$-\Delta\sigma$ (bars)	Fabric Interp. (vein)	Fabric Interp. (Bedding)	Comment
							Southeast							
ALL-NEV	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 20°SE	27	168°, 52°	31°, 17°	311°, 18°	-9.9	100	343.852106	VNS	LNS	Bedding-parallel normal fault
9	Om, vein	N45°E, 20°SE	N45°E, 40°SE	N45°E, 80°SE	28	152°, 32°	338°, 28°	71°, 3°	-1.3	0	283.807718	VNS	LPS	S1-normal vein
10a	Om, vein	N45°E, 18°SE	N45°E, 41°SE	N45°E, 18°SE	36	331°, 28°	149°, 82°	73°, 2°	-5.1	16	279.657792	VPS	LPS	S1-parallel vein
10b	Om, vein	N45°E, 18°SE	N45°E, 41°SE	N45°E, 18°SE	36	324°, 11°	128°, 82°	72°, 6°	-1.9	16	282.175469	VPS	LPS	S1-parallel vein
11	Om, vein	N45°E, 18°SE	N45°E, 41°SE	N45°E, 61°SE	36	162°, 55°	26°, 5°	338°, 12°	-1.8	11	319.114753	VPS	LNS	S1-normal vein
12a	Om, vein	N45°E, 15°SE	N45°E, 38°SE	N45°E, 15°SE	34	155°, 28°	252°, 78°	68°, 3°	-3.3	5	282.996078	VPS	LPS	S1-parallel vein
12b	Om, vein	N45°E, 15°SE	N45°E, 38°SE	N45°E, 15°SE	29	152°, 12°	302°, 82°	193°, 6°	-1.2	3	276.162354	VPS	LPS	S1-parallel vein
13a	Om, vein	N45°E, 15°SE	N45°E, 38°SE	N45°E, 15°SE	34	161°, 13°	342°, 52°	76°, 2°	-1.3	3	299.73073	VPS	LPS	S1-parallel vein
13b	Om, vein	N45°E, 15°SE	N45°E, 42°SE	N45°E, 42°SE	34	$168^{\circ},$ 8°	311°, 68°	75°, 4°	-1.3	19	319.114753	VPS	LPS	S1-normal vein
14	Om, vein	N45°E, 36°SE	N45°E, 58°SE	N45°E, 36°SE	35	323°, 12°	171°, 62°	63°, 8°	-4.8	2	292.195165	VPS	LPS	S1-parallel vein
15	Om, vein	N45°E, 36°SE	N45°E, 58°SE	N45°E, 36°SE	36	331°, 27°	133°, 71°	51°, 11°	-6.1	22	302.291823	VPS	LPS	S1-parallel vein
16a	Om, vein	N45°E, 36°SE	N45°E, 58°SE	N45°E, 36°SE	37	$160^{\circ},$ 8°	321°, 62°	70°, 3°	-1.5	16	305.975108	VPS	LPS	S1-parallel vein
16b	Om, vein	N45°E, 36°SE	N45°E, 58°SE	N45°E, 36°SE	35	326°, 11°	168°, 82°	61°, 5°	-4.9	0	281.345691	VPS	LPS	S1-parallel vein
17a	Om, vein	N45°E, 21°SE	N45°E, 65°SE	N45°E, 21°SE	28	318°, 28°	147°, 56°	69°, 8°	-5.3	17	277.930636	VPS	LPS	S1-parallel vein
17b	Om, vein	N45°E, 21°SE	N45°E, 65°SE	N45°E, 21°SE	36	141°, 22°	251°, 72°	72°, 5°	-4.9	8	269.619124	VPS	LPS	S1-parallel vein
17c	Om, vein	N45°E, 21°SE	N45°E, 65°SE	N45°E, 21°SE	28	157°, 22°	328°, 51°	66°, 8°	-3.3	13	293.62337	VPS	LPS	S1-parallel vein
18	Kittatiny LS	N45°E, 14°SE	None	None	31	160°, 42°	352°, 8°	282°, 28°	-1.8	25	339.5959185	-	LPS	_
Northwest														
_	_	_	_	_	n = 1341	_	-	_	Avg.: -2.9%	_	Avg: –323 bars		-	_

Om: Ordovician Martinsburg Fm. Fabric Interpretation: LPS (layer-parallel shortening); VPS (vein-parallel shortening); VNS (vein-normal shortening).



Figure 5. Lower hemisphere stereoplots of "contrived" strain overprint data (NEVs) for (**A**) the calcite veins in the calcareous siltstone and (**B**) the bedding-parallel normal fault. Strain axes (E1 = shortening, E2 = intermediate axis, E3 = extension) are plotted with Turner [40] compression axes contoured. See Table 2.

4.3. Calcareous Siltstone Layers

We sampled the Martinsburg Fm. turbidites and an interbedded calcareous siltstone across a concordant depositional contact (Figure 2). Both major (e.g., SiO₂, Al₂O₃, CaO, LOI) and trace element (e.g., Rb, Sr) counts are remarkably different (Table 3). Heavy mineral separations, with the intent of using detrital zircons for provenance and depositional age data (see [69]) resulted in the recovery of pyrite and no zircons from either sample.

One large oriented hand sample of a calcareous siltstone was collected north of the classic roadcut (Figure 2; see Maxwell, [4], Figures 9 and 10). This sample was cut into five vertical slabs from which two oriented thin sections per slab were recovered for a total of 10 thin sections. Bedding is sub-vertical and the horizontal veins are bed-normal. Petrographic observations reveal a sheared boundary between the turbidites and siltstone, local vein and stylolite development, and poor cleavage development in the siltstone (Figure 2e–g). Calcite twin analysis revealed a uniform sub-horizontal shortening strain that is normal to the local bedding (layer-normal shortening) and fold axis, yet within the locally-horizontal cleavage (Figure 2, Table 2). For the combined data set, the strain overprint was negligible (low NEVs; 14%) and removed with routine data cleaning, but we did combine 26 NEV grains for a separate strain analysis. This contrived strain overprint records a sub-horizontal shortening strain approximately parallel to the Stone Church syncline fold axis and Taconic margin (SW-NE; Figure 5).

EBSD fabric results for the siltstone layer show a robust Lattice Preferred Orientation (LPO), even though the calcite grains (10–50 microns) are widely spaced (~200 microns) in the matrix between stylolite seams implying that the deformation was homogeneous on the scale of the thin section. The calcite (<001>) axes are sub-horizontal and parallel to the Taconic shortening direction (NW-SE), similar to the a-axes (<11-20>), both of which are similar to the fabrics of calcite polycrystals (both experimentally and naturally deformed) at low temperature (~100 °C) reported by Wenk *et al.* [56,70] (Figure 6). The inferred horizontal shortening is also perpendicular to bedding, which is sub-vertical.

_	Samples										
	Martinsburg Fm. (Calc. siltstone)	Martinsburg Fm.	Detection Limits (ppm)								
	Major Elements										
SiO ₂	18.32	56.64	60								
TiO ₂	0.2989	0.8694	25								
Al_2O_3	6.2	18.16	120								
Fe ₂ O ₃	9.359	7.2402	25								
MnO	0.3959	0.0494	25								
MgO	11.3	3.26	95								
CaO	20.61	2.14	15								
Na ₂ O	0.2208	0.9132	35								
K ₂ O	1.29	4.45	25								
P_2O_5	0.376	0.158	35								
BaO	1113.2	750.3	12								
Ce	42	104	15								
Co	<d 1<="" td=""><td>15</td><td>10</td></d>	15	10								
Cr_2O_3	51.3	138.3	10								
Ċu	31	50	2								
Ni	12	45	3								
Sc	15	13	7								
V	71.4	129.9	7								
Zn	43	81	2								
LOI	31.34	5.84	100								
Total	99.85	99.85	-								
Trace Elements											
Ga	7.1	23.9	1								
Nb	5.8	16.6	0.3								
Pb	2.8	19	1								
Rb	53.3	171.8	1								
Sr	723.6	100	1								
Th	3.7	8.5	1								
U	1.9	3.6	1								
Y	24.6	19.9	1								
Zr	60.7	161	1								

Table 3. Whole-rock geochemistry (XRF).



Figure 6. Cont.



Figure 6. Photomicrograph of calcareous siltstone in the Martinsburg Fm. (**A**: polarized light) with yellow rectangles indicating areas of EBSD analysis (**B**,**C**) and lower hemisphere pole figure calcite results below (n = 202 for each area, A,B).

5. Discussion

Studies of deformation in the Martinsburg Fm., including along the Delaware Water Gap, is quite comprehensive, although few of the published results include oriented strain data. Fold shortening, by bed length curvature calculations, accounts for ~40% of the shortening in the Delaware Water Gap section [4] which is comparable to the ~50% regional shortening of a thin-skinned thrust belt. Volume loss by pressure solution during orogenic cleavage formation is well-documented in Appalachian Paleozoic rocks, ranging from 40%–75% [5–7,19] a value similar to the 10%–25% AMS proxy value reported by Pares and van der Pluijm [37]. Recrystallization, observed as pressure shadows parallel to cleavage planes and new chlorite growth, is a significant portion of the Martinsburg deformation [35] and involved a Permian remagnetization [36]. Beutner [7], using 80 measurements on chlorite, report no grain rotation, a result not in agreement with our EBSD measurements in the calcareous siltstones of the Martinsburg Fm. (see below).

Our contribution to understanding the partitioning of strain is the analysis of the five phases of secondary calcite in the Martinsburg Fm. The various calcite phases reported are of a uniform stable isotopic signature and presumably from the same hot, synorogenic Taconic source (Table 1). Calcite twinning, then, represents a minor portion of the total Taconic strain (-3% shortening) partitioning in these rocks, with a uniform SE–NW sub-horizontal shortening strain preserved by the secondary calcite across the Stone Church syncline. Strain magnitudes and differential stresses (-32 MPa) are typical of the internal portions of a thrust belt [71] and there is no twinning strain overprint (low NEVs). If the five calcite phases in the Martinsburg Fm. are the result of fluid migration during the Taconic orogeny then we might expect to see a twinning strain related to Acadian terrane accretion [72] and/or the collisional Alleghenian orogeny in the Permian [73,74]. Twinning strain overprints require a significantly different stress field orientation as calcite strain hardens [64] so the Acadian or Alleghenian stress fields were likely parallel (co-axial) to the Taconic stress-strain field. On a regional scale our Piedmont twinning shortening strain (N40°W) is parallel (co-axial) to other twinning and finite strains across the Valley and Ridge and foreland provinces (Figure 7, [8,10,24,25,44,45,75]). The parallelism of layer-parallel shortening fabrics from the Piedmont, across the Valley and Ridge province, and far

into the foreland (~2000 km; [26,27]) suggests that the pre-Appalachian sedimentary prism (foreland + restored orogenic belt) approaches ~3000 km and there is little rotation (no vertical axis rotations) between the LPS fabrics in the allochthonous rocks *versus* the same pre-thrusting LPS fabric in the autochthonous foreland [8,10,21]. This result suggests that the Pennsylvania orocline (salient) formed in response to a pre-orogenic margin shape that controlled the structural evolution of the Appalachians. Kinematic data across and around the Valley and Ridge province in the Pennsylvania salient supports a rotational fault-fold genesis in two stages [9,76] a model not supported by the calcite strain data herein. The intense cleavage spacing in the Martinsburg Fm. precludes acquisition of a reliable Ordovician paleomagnetic paleopole and Housen *et al.* [36] report a complex, four-component paleopole in the Martinsburg Fm. near Lehigh Gap, PA. The mid-late Paleozoic sediments of the Appalachian Valley and Ridge and foreland provinces preserve a late Paleozoic chemical remagnetization overprint making the role of pre-thrusting salient shapes difficult to interpret relative to thrust sheet rotations and migrations [77–85]. These complex data then preclude addressing the role of thrust sheet rotations in the Piedmont with younger magnetic signatures preserved in younger Paleozoic sediments inboard of the hinterland.

Finite strains in the Martinsburg Fm. are reported with aspect ratios up to 5:1 ([4], quartz) with the long axis of the ellipsoidal shape parallel to the cleavage plane dip direction, with shortening normal to the cleavage plane. Syn-cleavage calcite twins record a NW–SE shortening strain, normal to the cleavage strike and parallel to the inferred Taconic thrust transport direction. Extension axes are sub-horizontal and strike-parallel, a result identical to that reported for syn-cleavage twinning strains in the Permian Gondwanide orogeny in the Cape belt, S. Africa [17]. In both cases, the shaly rocks are folding as part of a thrust belt while the rocks are losing volume to pressure solution while syn-orogenic calcite is precipitating along bedding and cleavage planes. Most of the synorogenic calcite is sparry and twins immediately after precipitation, leaving a mechanical paradox: how does a cleavage plane open (dilate) while filling with calcite that shortens normal to the dilation direction? This strain-cleavage relationship is observed everywhere in the Stone Church syncline except on the southeast fold limb in the vertical calcareous siltstone layer within the turbidite sequence (see below).



Figure 7. Strain data for the Piedmont, Valley and Ridge, and Foreland provinces across the central Appalachian Mountains and the Pennsylvania Salient orocline. Sources are [8,10,24,25,44,45,85] and this study. Stereonet is a summary plot of calcite twinning shortening axes. Inset is a schematic of a flexural-slip interpretation of Figures 2a,b and 3a,b and the Solor Church folds.

Calcareous siltstone layers (<0.5 m) are infrequent within the Martinsburg Fm. and are part of the Ordovician turbidite sequence. These layers are geochemically different than the enclosing turbidites (Table 3) and there is literature about non-tabular clastic enclaves and dikes in the Martinsburg Fm. [4–6,86]. Isotopic values for calcite veins in this siltstone layer are similar to the other syn-orogenic calcite fluids (Table 1) and the twinning strains correlate well within the multi-layer sample and with twinning strain results across the syncline (Figure 4). Shear between the Martinsburg turbidite and calcareous siltstone demonstrates the rheological difference between these materials relative to the contemporaneous cleavage and the presence of minor folds in the Stone Church syncline. Internally, the fine-grained calcite in the silty layer has a preferred orientation parallel to the inferred Taconic shortening orientation (Figures 2 and 4) but normal to bedding. The calcareous siltstone layer LPO strain proxy and calcite vein twinning shortening strain are sub-horizontal and oriented SE-NW, and both preserve a layer-normal shortening suggesting the strain was imposed in these rocks after the bedding was folded into a sub-vertical orientation as part of the Stone Church syncline. The calcite shortening strain here is parallel to the cleavage dip direction and the extension axis is vertical. The contrived NEV calcite twinning overprint (n = 26) preserves a sub-horizontal shortening strain that is parallel to the Stone Church syncline axis, which is roughly parallel to the strike of normal faults on the margin of the Triassic Newark basin (Figure 5). Is this strain overprint related to Acadian [14,15,87] or Triassic [88] strike-slip motion within the Piedmont?

The youngest structural features observed along the Martinsburg Fm. traverse include a few bedding-parallel shear zones (N40°E, 25°SE) with layered calcite gouge zones that are not crosscut by the axial-planar cleavage. These SE-dipping fault zones are more common on the northwest end of the traverse where the northwest limb of the Stone Church syncline rolls into the adjacent anticline (Figure 2). Sparry, en echelon fillings are materially continuous into the layered fault zone and both preserve a top-down-to-the-southeast kinematic sense, indicating that this normal faulting (N40°E, 45°SE) post-dated cleavage formation. Beutner et al. [6] describe calcite boudins that look like normal fault steps (their Figure 6) and folded calcite veins (their Figure 3) that look like the normal fault offsets we describe in Figure 2e-g. This fault zone calcite is chemically similar to the other phases of syn-orogenic calcite (Table 1) and records a NW-SE horizontal shortening strain (Table 2, Figure 4). For a section of NW-vergent folds in the Martinsburg to have SE-dipping normal fault offsets (Figure 3), with a horizontal shortening strain, seems out of place in a Taconic thrust sheet. One explanation could be that the bedding-parallel slip surfaces are zones of flexural-slip within the Martinsburg Fm. and are slip surfaces at the inflection point along the common limb of the Stone Church syncline and the anticline to the northwest (Figure 2). So, for each SE-dipping slip surface, material motion would be down-to-the-SE into the hinge of the syncline (out of the anticline), and the lower surface would move material up-to-the-NW into the (poorly exposed) anticline. If the field observation is taken at face value, as an individual exposure, the stress-strain field of the syn-faulting calcite is not in agreement with the Anderson [89] expectation for principal stress planes and normal faults. We also have evidence of a (contrived) twinning strain overprint in this normal fault zone (Figure 5), a shortening strain that is sub-vertical. Vertical shortening strains are reported in tectonic settings with nappe burial [17,90] but the normal fault zone that crosscuts the Martinsburg Fm. cleavage is the youngest calcite phase, making a vertical shortening strain related to Taconic nappe (e.g., Reading Prong) burial unlikely. A sub-vertical twinning strain related to Triassic extension is a possibility as the youngest strain event, but this result is not in agreement with the NE-trending (5° - 60° plunges) shortening result reported by Lomando and Engelder [88] for calcite cements in the Newark Basin New Brunswick sandstones. Based on one detailed outcrop description from the hinge of a Taconic fold, the twinning strains (and contrived strain overprint) suggest a complex stress-strain field related to flexural-slip dynamics again (see inset, Figure 7), similar to the results reported in the Cape belt in Southern Africa [17] who reported a layer-normal twinning strain overprint where these secondary shortening axes rotated through the fold axis during folding.

6. Conclusions

As the Martinsburg Fm. folded, translated, and cleaved (fold axis: N40°E) from SE to NW during the Taconic orogeny five phases of syn-orogenic calcite, of a uniform isotopic signature, precipitated throughout the Stone Church syncline and was mechanically twinned. The calcite strains record a uniform sub-horizontal shortening parallel to the Taconic thrust translation direction (SE–NW), a shortening fabric that is generally parallel to Martinsburg Fm. bedding and at a steep angle to the axial planar cleavage, but normal to the cleavage strike. Extension axes were horizontal during cleavage formation. This result is identical to that reported by Craddock *et al.* [17] for syn-cleavage Gondwanide deformation calcite twinning strains in the frontal Cape belt. Twinning strains in cleavage planes within the vertical calcareous siltstones also record a SE–NW sub-horizontal shortening strain that is bedding-normal and cleavage-parallel. This relationship indicates a consistent stress-strain field during development of the axial-planar cleavage but rotation of the (minor) fold during formation of the syncline. Some SE-dipping bedding planes contain down-to-the-SE normal fault offsets that are post-cleavage and are likely related to flexural-slip dynamics at the inflection point between migrating Taconic anticlines and synclines and not the Triassic evolution of the Newark-Gettysburg basin.

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