



Article Pore Structure and Migration Ability of Deep Shale Reservoirs in the Southern Sichuan Basin

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Abstract: The migration phenomenon of deep shale gas is a subject that has yet to be fully comprehended, specifically regarding the migration ability of deep shale gas. This study focuses on the Longmaxi Formation in the southern Sichuan Basin, utilizing it as an example. Various experimental techniques, such as temperature-driven nitrogen and carbon dioxide adsorption, high-pressure mercury intrusion, XRD, and TOC analysis, are employed. The goal is to analyze the pore structure and fractal characteristics of the Longmaxi Formation shale. Additionally, the study aims to calculate its Knudsen number based on parameters like temperature gradient and pressure coefficient. The migration ability of the Longmaxi Formation shale in southern Sichuan Basin is also discussed. The results show the following: (1) The pore volume distribution of the Longmaxi Formation shale in the study area ranges from 0.0131 to 0.0364 cm^3/g . Mesopores contribute approximately 56% of the pore volume, followed by micropores with a contribution rate of about 26%, and macropores contributing approximately 18%. Additionally, the Longmaxi Formation shale exhibits strong heterogeneity, with the fractal dimension (D1) of mesopores ranging from 2.452 to 2.8548, with an average of 2.6833, and the fractal dimension (D2) of macropores ranging from 2.9626 to 2.9786, averaging 2.9707. (2) The fractal dimensions of shale were significantly influenced by organic matter, inorganic minerals, and pore structure parameters. D1 and D2 were positively correlated with TOC, clay mineral content, and specific surface area, while exhibiting negative correlation with quartz. However, the correlations with calcite content, pore volume, and average pore size were not significant. (3) The proportion of pores satisfying Darcy flow in the Longmaxi Formation shale was approximately 3.7%-11.8%, with an average of 6.6%. Consequently, the migration capability of shale gas can be calculated using Darcy's law. Furthermore, the migration capability of shale gas is controlled by D2, specifically the surface area, and the connectivity of macropores.

Keywords: southern Sichuan Basin; Longmaxi Formation shale; deep shale gas; migration ability; fractal dimension

1. Introduction

Shale gas, an essential unconventional natural gas resource, has gained global prominence following the shale gas revolution in North America [1–3]. Continuous exploration and evaluation, spanning over a decade, have led to a steady increase in the production of marine shale gas in southern China. In 2022, China achieved a shale gas production milestone of 24 billion cubic meters [4]. Notably, the majority of this output is produced from the Upper Ordovician Wufeng Formation and the Lower Silurian Longmaxi Formation shale, situated at depths ranging from 2000 to 3500 m in the Sichuan Basin. This indicates the enormous exploration and development potential of shale gas in the Sichuan Basin [5,6]. With the development of exploration technology and theory, attention in China



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Copyright: © 2024 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). is gradually shifting towards deep shale gas, characterized by distinct occurrences and enrichment mechanisms compared to shallow shale gas [7–9]. Understanding the enrichment mechanism of deep shale gas is pivotal for effective exploration and development.

Regarding the deep shale gas enrichment mechanisms, previous researchers have proposed various models such as the binary enrichment theory [10], ternary enrichment theory [11], source-cap matching theory [12,13], and multi-factor coupled enrichment theory [14]. However, these models predominantly adopt static theories, with comparatively limited exploration into the migration and diffusion phenomena associated with shale gas.

In the context of the dynamic shale gas enrichment theory, previous studies have indicated that the migration and flow behavior of shale gas are significantly different from the conventional gas reservoirs. Notably, Darcy's law inadequately explains the flow dynamics [15]. Additionally, the migration process involves the gradual flow of nanoscale pores into larger spaces, such as microfractures [16]. To understand this migration mechanism, previous researchers have proposed microscale transport models for shale gas based on flow-diffusion-desorption mechanisms [17]. Additionally, equations addressing gas flow in nanoscale pores based on Knudsen diffusion and slip flow have also been developed [18]. However, these models primarily consider small pore sizes, neglecting the contributions of macropores and fractures to shale gas migration. Recent studies have discovered five distinct migration types for shale gas in bedding fractures and matrix pores, namely Darcy flow, slip flow, Fickian diffusion, Knudsen diffusion, and adsorption–desorption diffusion [19–21]. These migration types significantly influence the enrichment level and determine favorable areas within shale gas reservoirs, exerting control over the direction and intensity of shale gas migration [20–22].

Previous studies have elucidated the direct influence of pore structure and fractal dimension on gas occurrence and permeability characteristics [23–25]. Pore structure parameters, including pore volume, specific surface area, pore size, and their distribution patterns, are of significant importance in studying the occurrence state, desorption diffusion, and permeability of shale gas [25–29]. Among these parameters, pore volume determines the migration space for shale gas, while pore size governs its migration mode and capacity [20–22]. In addition, the complexity of the pore system also affects the migration efficiency of shale gas. Fractal dimension is currently one of the most effective parameters for evaluating the complexity of shale pores. Extensive research by previous scholars has led to the development of various mesoporous fractal dimension model such as the BET model [30], the FFH model [31], and thermodynamic methods [32]. Additionally, the J-function model [33], the capillary bundle model [34], the thermodynamic model [35], and the Menger model [36] have been employed to calculate the fractal dimension of macropores. These models facilitate the exploration of the intricate relationship between pore complexity and shale gas permeability.

Therefore, it is of great significance to achieve a quantitative and heterogeneous characterization of pore structure in shale reservoirs to understand the migration capacity of shale gas. In this study, TOC testing, XRD analysis, low-temperature nitrogen adsorption, low-temperature CO_2 adsorption, and high-pressure mercury experiments are employed to analyze the pore structure and fractal dimension characteristics of deep shale gas in southern Sichuan Basin. The objective is to gain insights into the migration capacity of deep shale gas reservoirs.

2. Geologic Background

The Sichuan Basin, located in Southwestern China, stands as one of the largest basins in the region, sprawling across an expansive 2.6 million square kilometers. Nestled in the western part of the Yangtze Plate, this basin exhibits an east–west elongated configuration, encircled by the Qinghai-Tibet Plateau and the Western Sichuan Plateau to the north and northeast, while bordering the Yunnan-Guizhou Plateau to the south and southwest [9,37]. Over time, the Sichuan Basin has evolved through multiple tectonic phases and sedimentary processes [38–41]. It comprises two principal components: the Sichuan Basin Foreland Structural Belt and the Sichuan Basin Inland Structural Belt. The Sichuan Basin Foreland Structural Belt, located in the northern and northeastern parts of the basin, serves as the marginal tectonic belt of the Qinghai-Tibet Plateau and the Western Sichuan Plateau. This tectonic belt has undergone complex structural deformations, resulting in the formation of numerous mountain ranges and fault zones [38,39]. In contrast, the Sichuan Basin Inland Structural Belt, located in the central and southern parts of the basin, is mainly composed of fault zones and fold belts, making it the primary region for crustal deformation in the central basin [40,41].

The southern Sichuan region, located on the northwest side of the Yangtze Platform, encompasses the steeply folded Chuanan area of the Huaying Mountain fold belt and the gently structured southwestern Sichuan area. Encompassing an area of approximately 22,000 square kilometers [42,43], this region is abundant in shale gas resources and hosts five prominent shale gas production bases in Fuling, Changning, Weiyuan, Zhaotong, and Weirong. The Longmaxi Formation shale gas reservoir in the study area is one of the main gas reservoirs and a key contributor to deep shale gas production in China. The study area is composed of four blocks: the Weiyuan Block, the Yuxi Block, the Luzhou Block, and the Changning Block (Figure 1). The sediment thickness of the Longmaxi Formation in the study area ranges from 200 to 500 m and is further categorized into the Longyi Member and the Long'er Member from top to bottom. The Longyi Member 2 sub-member, with a total of four sub-layers in the Longyi Member 1 sub-member 1 sub-member, which serves as the primary production layer.



Figure 1. Geological background of the study area.

3. Samples and Methods

3.1. Samples

To investigate the pore structure and migration ability of the deep Longmaxi Formation shale reservoir, we collected 16 shale core samples from the southern part of Sichuan. This collection comprised 3 samples each from the Weiyuan and Yuxi blocks, 4 samples from the Changning block, and 6 samples from the Luzhou block. The depths of these core samples ranged from 2349.5 m to 4926.8 m. They were systematically numbered and labeled according to the blocks they belong to, with corresponding depths provided in the Table A1.

Additionally, we divided these samples into two parts. The first part underwent grinding and sieving to obtain 80–mesh and 200–mesh samples. The 80–mesh samples were used for TOC determination, while the 200–mesh samples were used for XRD, low-temperature N₂ adsorption, and low-temperature CO₂ adsorption. The second part of the samples were transformed into cylinders with a diameter of 2.5 cm and a height of 5 cm. These cylindrical samples were used for high-pressure mercury injection experiments and permeability measurements, as shown in Figure 2.



Figure 2. Experimental and testing flowchart.

3.2. Methods

3.2.1. Experimental Testing

TOC tests were conducted at the laboratory affiliated with the Southwest Oil and Gas Field Company of PetroChina (Chengdu, China). The testing instrument used was the LECO-CS230 Carbon Sulfur Analyzer. Prior to the experiments, the samples were soaked in dilute hydrochloric acid (HCl) to remove inorganic carbon. Then, the samples were washed with distilled water and dried. The testing was carried out using the dried samples at a temperature of 25 °C and a humidity level of 57%.

XRD experiments were conducted at the Unconventional Petroleum Research Institute, China University of Petroleum (Beijing). The testing instrument used was the X'Pert Powder X-ray diffractometer. Prior to the experiments, the samples were crushed to a 200–mesh size and dried. The dried samples were then pressed onto glass slides, and mineral composition was determined through sample scanning. Additionally, the instrument used Cu radiation, with a voltage of 40 kV and a current of 40 mA. The scanning speed of the instrument was $2^{\circ}/min$, with a scanning step size of 0.02° .

High-pressure mercury intrusion, low-temperature nitrogen adsorption, and CO_2 adsorption experiments were conducted at the Key Laboratory of Unconventional Oil and Gas, China Petroleum Group. Low-temperature N₂ adsorption and CO_2 adsorption were performed using the Autosorb-IQ3 surface area and pore size analyzer produced by Quantachrome, Boynton Beach, FL, USA. Prior to the tests, the samples were crushed, dried at 105 °C for 8 h, and vacuum degassed for 18 h. Low-temperature N₂ adsorption was

conducted at 77 K, and low-pressure CO_2 adsorption was performed at 273 K. The specific surface area was calculated using the BET model, and the pore volume was calculated using the BJH model for low-temperature N₂ adsorption. High-pressure mercury intrusion was performed using the Pore Master 60 GT fully automated mercury intrusion porosimeter from Quantachrome, USA. The maximum working pressure was 60,000 PSI, and the pore size measurement range was 0.03 to 1000 μ m, with an accuracy of mercury intrusion volume measurement at 0.1 cm³.

Permeability measurements were carried out at Beijing Yanbo Times Co., Ltd. (Beijing, China), using the PoroPDP-200 pressure decay porosity permeability measurement instrument from Petrolab Company (Albany, NY, USA). The instrument's testing range extended from 0.00001 to 10 mD. In this study, the permeability of shale samples was measured using the non-steady-state method (pressure pulse decay method). It is worth noting that the sample permeability measurements were fully automated, with a testing error below 0.5%.

3.2.2. Fractal Theory

The fractal dimension of shale pores can be calculated for micropores, mesopores, and macropores. As shale predominantly occurs in micropores in the form of free gas, our study focused on calculating the fractal dimensions for mesopores and macropores. Various models exist for calculating the fractal dimension of mesopores, including the BET model [30], the FFH model [31], and thermodynamic methods [32]. Among these, the FFH model is the most widely adopted, offering two algorithms—one is based on the van der Waals force mechanism and the other on the capillary condensation mechanism [44]. Previous research [44] suggests that the capillary condensation mechanism is more suitable for studying the heterogeneity of porous media. Therefore, in this study, this method was adopted to calculate the fractal dimension of the gas adsorption part, and the calculation formula is as follows:

$$\ln V = C + (D - 3) \ln[\ln(P/P_0)]$$
(1)

In this equation, *V* represents the gas adsorption amount (cm^3/g) , *P* represents the equilibrium pressure of the system (MPa), P_0 represents the saturation vapor pressure of the gas (MPa), *C* is a constant (dimensionless), and *D* is the fractal dimension (dimensionless). In the scatter plot of $\ln V - \ln(P/P_0)$, the slope *K* of the double logarithmic curve corresponds to D - 3 in Equation (1), allowing the calculation of the fractal dimension of mesopores based on D = K + 3.

The current models for calculating the fractal dimension of macropores are the Jfunction model [33], the capillary bundle model [34], the thermodynamic model [35] and the Menger model [36]. After comparing previous research, it was found that the fractal dimension of the capillary bundle model has a better correlation with the physical properties of the reservoir and pore structure parameters, making it widely used by scholars [34,45–47]. The methods of calculating fractal dimension according to the capillary bundle model are based on water saturation and mercury saturation [47], but the capillary bundle calculation model based on water saturation is most widely used, expressed as follows:

$$lg(1 - S_{Hg}) = (D - 3)lgP_c - (D - 3)P_{min}$$
(2)

In the equation, S_{Hg} represents the volume fraction of mercury entering the pore (%), P_c represents the capillary pressure (MPa), and P_{min} represents the capillary pressure corresponding to the maximum throat radius (MPa). In the scatter plot of $lg(1 - S_{Hg}) - lgP_c$, the slope K of the double logarithmic curve corresponds to D - 3 in Equation (2), facilitating the calculation of the fractal dimension of macropores based on D = 3 + K.

3.2.3. Shale Gas Transport Capacity

The ability of shale gas migration can be determined by calculating the Knudsen number (Kn). Gas migration types are categorized as follows: when Kn < 0.001, the gas migration type is Darcy flow; when 0.001 < Kn < 0.01, the gas migration type is slippage

flow; when 0.01 < Kn < 1, the gas migration type is Fick diffusion; when Kn > 1, the gas migration type is Knudsen diffusion; and when the pore size is smaller than or equal to twice the thickness of the adsorption layer plus the exclusion layer thickness, the gas migration form is adsorbed gas diffusion. The formula for calculating Kn can be obtained by referring to reference [20]. Since the migration ability of Darcy flow is 100 times that of slip flow and 10,000 to 1,000,000 times that of Fickian diffusion [20,21], we consider that if more than 1% of the pore volume in the samples satisfies Darcy flow, it can be approximately considered as satisfying Darcy's law.

4. Results

4.1. Total Organic Carbon

This study analyzed the TOC across 16 samples in the study area, as shown in Figure 2. Based on Figure 3, the TOC content of Longmaxi Formation shale in the study area ranges from 2.9% to 6%, with an average value of 4.75%. The highest TOC content was found in the Yuxi block of the Yuxi Formation shale, with an average value of 5.4%, followed by the Changning block (average value of 4.78%) and the Luzhou block (average value of 4.63%). The Weiyuan block has the lowest TOC content in the Longmaxi Formation shale, with an average value of 4.3%.



Figure 3. Box diagram of TOC content distribution in the shale of Longmaxi Formation in southern Sichuan.

4.2. Mineralogy

The mineral composition of the Longmaxi Formation shale in the study area was analyzed, and the results are presented in Figure 4. According to Figure 4a, the diagenetic minerals of Longmaxi Formation shale include clay, quartz, calcite, plagioclase, K-feldspar, and pyrite. Calcite has the highest content, with an average value of 37.21%, followed by quartz (average content of 29.69%), clay minerals (average content of 18.09%), plagioclase (average content of 7.98%), K-feldspar (average content of 3.93%), and pyrite (average content of 3.07%). However, there are significant differences in the mineral composition of Longmaxi Formation shale among different blocks. The quartz content in the Changning block shale is significantly higher than that of clay and calcite, while calcite content is highest in the Weiyuan and Yuxi blocks. In the Luzhou block, shale with the highest quartz content and shale with the highest calcite content are equally distributed. Additionally, the mineral composition of shale plays a crucial role in determining the reservoir alteration difficulty, often represented by the brittleness index. This index is determined by the proportion of brittle minerals in the shale [48]. The brittleness index of Longmaxi Formation shale in the study area ranges from 21.23 to 72.7, with an average value of 44.53. The highest brittleness index was recorded in the Changning block Longmaxi Formation shale (59.29), while the



lowest brittleness index was observed in the Yuxi block Longmaxi Formation shale, with a value of 28.49.

Figure 4. Mineralogical characteristics of shale in the Longmaxi Formation of the research area. (a) Mineral composition of Longmaxi Formation shale; (b) classification of shale facies in the Longmaxi Formation.

In addition, mineral composition is also an important means to determine the shale lithofacies, which holds significant implications for analyzing sedimentary environments, among other aspects. The lithofacies composition of Longmaxi Formation shale was analyzed using a mineral ternary diagram (Figure 4b) [49]. The results show that the main lithofacies of Longmaxi Formation shale in the study area are siliceous shale, calcareous shale, and mixed shale. Specifically, the Changning block is mainly composed of siliceous shale and mixed shale. In the Luzhou block, the Longmaxi Formation shale exhibits a combination of siliceous shale, calcareous shale, and mixed shale.

4.3. Full Aperture Characterization

4.3.1. High-Pressure Mercury Injection

High-pressure mercury intrusion is an important method for analyzing the macroscopic pores in shale. In this study, the distribution characteristics of macroscopic pores in the Longmaxi Formation shale were analyzed using high-pressure mercury intrusion experiments. Figure 5 shows the mercury intrusion and extrusion curves of the Longmaxi Formation shale in the study area. According to Figure 5, the mercury intrusion volume of the Longmaxi Formation shale ranges from 0 to 0.016 cm³/g. Notably, the mercury intrusion volume in the Changning block (Figure 5a) and the Weiyuan block (Figure 5c) is higher than that in the Yuxi (Figure 5b) and Luzhou blocks (Figure 5d), indicating a higher macroscopic pore content in the Changning and Weiyuan blocks compared to the Yuxi and Luzhou blocks. Additionally, the mercury intrusion and extrusion curves of CN-1 and LZ-5 are similar, indicating poor connectivity of macroscopic pores in these samples.

Based on the Washburn equation, the pore diameters corresponding to each pressure point and their corresponding mercury intrusion volumes were calculated. This intrusion volume represents the pore volume corresponding to the pore diameter, providing the pore size distribution of the Longmaxi Formation shale in the study area (Figure 6). According to Figure 5, the pore volume of the Longmaxi Formation shale in the study area is mainly distributed between 1–100 nm. For pores larger than 100 nm, the pore volume in the Changning block (Figure 6a) is slightly higher than that in the Yuxi block (Figure 6b). However, the Weiyuan block has the highest proportion of pore volume larger than 100 nm in the Longmaxi Formation shale (Figure 6c), while the Luzhou block has the lowest pore volume larger than 100 nm (Figure 6d). This observation may be related to the deeper burial of the Longmaxi Formation shale in the Luzhou block.



Figure 5. Mercury intrusion and retreat curves of the Longmaxi Formation shale in the research area. (a) Changning block; (b) Yuxi block; (c) Weiyuan block; (d) Luzhou block.



Figure 6. Pore size distribution of high-pressure mercury injection in Longmaxi Formation shale in the research area. (a) Changning block; (b) Yuxi block; (c) Weiyuan block; and (d) Luzhou block.

4.3.2. Low-Temperature N₂ Adsorption

Low-temperature nitrogen adsorption is an important method for characterizing mesopores. In this study, the low-temperature nitrogen adsorption–desorption isotherms of 16 shale samples in the study area were analyzed (Figure 7). According to Figure 6, the adsorption–desorption isotherm types of the shale samples in the study area vary, but they all exhibit a steep drop in the desorption curve at relative pressures of 0.4–0.6, similar to

the H2 and H3 type hysteresis loops proposed by IUPAC. This indicates that the pores in the samples are mainly composed of micropores and mesopores. Additionally, the shapes of the hysteresis loops in the adsorption curves can be used to identify the shape of the micropores [50]. The mesopores in the Longmaxi Formation shale in the study area are mainly composed of slit-like pores and ink bottle-like pores. Among them, the mesopores in the Longmaxi Formation shale in the mesopores in the Changning block are predominantly slit-like pores (Figure 7a), while the mesopores in the Yuxi block and the Weiyuan block shale are mainly ink bottle-like pores (Figure 7b,c), and the mesopores in the Luzhou block shale are mainly slit-like pores (Figure 7d).



Figure 7. Low temperature nitrogen adsorption–desorption curve of Longmaxi Formation shale. (a) Changning block; (b) Yuxi block; (c) Weiyuan block; and (d) Luzhou block.

In this study, the pore size distribution of the Longmaxi Formation shale was analyzed using the BJH model. The range of pore size distribution characterized by low-temperature N_2 adsorption under this model is 2 to 180 nm. For this study, the pore size distribution in the range of 2 to 100 nm was extracted (Figure 8). The results show that the pore volume of the Longmaxi Formation shale in the study area is mainly contributed by pores with a diameter of 4 to 6 nm. This result differs from the pore volume obtained by high-pressure mercury intrusion, attributed to the difference in experimental methods. High-pressure mercury intrusion measures pore volume by the amount of mercury intrusion and extrusion, while low-temperature nitrogen adsorption relies on the gas adsorption capacity. Moreover, high-pressure mercury intrusion is more effective in characterizing macropore volume, while low-temperature nitrogen adsorption is more effective in characterizing mesopore volume.

4.3.3. Low-Temperature CO₂ Adsorption

 CO_2 adsorption is one of the important methods for characterizing the micropore structure of shale. Based on the CO_2 adsorption curves (Figure 9), the CN-3 sample in the Changning block has the lowest adsorption capacity, measuring 0.394 cm³/g (Figure 9a), while the WY-2 sample has the highest adsorption capacity, at 1.03 cm³/g (Figure 9c).



Additionally, the overall CO_2 adsorption capacity of the samples in the Changning block is lower than the other three blocks, with the Weiyuan block exhibiting the highest adsorption capacity, followed by the Luzhou (Figure 9d) and Yuxi blocks (Figure 9b).

Figure 8. Pore size distribution of Longmaxi Formation shale based on low-temperature nitrogen adsorption. (a) Changning block; (b) Yuxi block; (c) Weiyuan block; and (d) Luzhou block.



Figure 9. CO₂ adsorption curve of Longmaxi Formation shale. (**a**) Changning block; (**b**) Yuxi block; (**c**) Weiyuan block; and (**d**) Luzhou block.

The NL-DFT model was used to analyze the CO_2 adsorption capacity characterization in shale micropores. The results revealed a three-peak distribution in most micropores of the Longmaxi Formation shale in the study area, with peaks at 0.33–0.385 nm, 0.501–0.603 nm, and 0.805–0.85 nm. Only a few samples showed four peaks (Figure 10), such as CN-4 (Figure 10a), YX-3 (Figure 10b), WY-1 (Figure 10c), and LZ-2 (Figure 10d). According to



Figure 9, the micropore volume of the Longmaxi Formation shale samples in the Weiyuan block is the largest.

Figure 10. Micropore distribution of Longmaxi Formation shale based on CO₂ adsorption. (a) Changning block; (b) Yuxi block; (c) Weiyuan block; and (d) Luzhou block.

4.3.4. Full Aperture Distribution

The present study utilized a method proposed in previous studies [50,51] to comprehensively characterize the pore size distribution of the Longmaxi Formation shale in the study area using high-pressure mercury intrusion, low-temperature nitrogen adsorption, and low-pressure carbon dioxide adsorption. The results are shown in Figure 11. According to Figure 11, the pore volume distribution of the Longmaxi Formation shale ranges from 0.0131 to 0.0364 cm³/g. Micropores, mesopores, and macropores are all well-developed, with mesopores contributing the most to the pore volume of the shale, accounting for approximately 56%. Micropores are the second largest contributor, accounting for approximately 26% of the pore volume, while macropores contribute approximately 18% to the pore volume. Additionally, the average pore size distribution of the Longmaxi Formation shale in the study area ranges from 4.55 to 17.92 nm, with an average value of 7.69 nm. The specific surface area ranges from 12.01 to 21.99 m²/g.

4.4. Fractal Characteristics

In this study, the fractal dimensions of the meso-pores and micro-pores of the Longmaxi Formation shale in the study area were calculated using Equations (1) and (2) respectively. The results are shown in Figure 12 and Table A1. According to Figure 12 and Table A1, the fractal dimension (D₁) of the meso-pores in the Longmaxi Formation shale ranges from 2.452 to 2.8548, with an average of 2.6833. The D₁ values for the Changning block, Weiyuan block, Yuxi block, and Luzhou block are 2.6925, 2.6789, 2.7726, and 2.6346 respectively. The fractal dimension (D₂) of the macro-pores in the Longmaxi Formation shale ranges from 2.9626 to 2.9786, with an average of 2.9707. The D₂ values for the Changning block, Weiyuan block, Yuxi block, and Luzhou block are 2.9683, 2.9648, 2.9765, and 2.9726 respectively.



Figure 11. Full pore size distribution of shale in Longmaxi Formation. (**a**) Changning block; (**b**) Yuxi block; (**c**) Weiyuan block; and (**d**) Luzhou block.

In addition, there are differences in the fractal dimensions of shale with different lithologies. For example, the D_1 values for siliceous shale range from 2.6064 to 2.7312, with an average of 2.6619, while the D_2 values range from 2.9636 to 2.9746, with an average of 2.9693. For mixed shale, the D_1 values range from 2.452 to 2.8315, with an average of 2.6528, while the D_2 values range from 2.9626 to 2.9786, with an average of 2.9721. For calcareous shale, the D_1 values range from 2.6316 to 2.8548, with an average of 2.7394, while the D_2 values range from 2.9641 to 2.9779, with an average of 2.9713.

4.5. Permeability and Pores Satisfying Darcy Flow

The permeability of the Longmaxi Formation shale in the study area was measured using the pressure pulse decay method, and the results are shown in Figure 13. According to Figure 13, the permeability of the Longmaxi Formation shale in the study area ranges from 2.13 to 35.22 μ D, with an average of 14.53 μ D. There are notable differences in shale permeability among different blocks. Specifically, the average permeability of the Longmaxi Formation shale is 18.74 μ D in the Changning block, 25.31 μ D in the Weiyuan block, 6.53 μ D in the Yuxi block, and 10.35 μ D in the Luzhou block (Figure 13a). This indicates that as the uplift magnitude decreases, the shale reservoir exhibits a rebound in petrophysical properties, leading to a gradual increase in the permeability. Additionally, shale permeability exhibits variations with different lithologies, where the average permeability is 15.75 μ D for siliceous shale, 15.98 μ D for mixed shale, and 11.63 μ D for calcareous shale (Figure 13b).

The migration forms of shale gas can be determined by Kn [20,21]. In this study, Kn and the critical migration aperture of different blocks were calculated using the geothermal gradient and pressure coefficient. By combining Kn with the relationship between various migration forms, the aperture of different migration types in the Longmaxi Formation shale in the study area was determined. The aperture satisfying Darcy flow in the Changning block is approximately 195 nm, in the Weiyuan block is 335 nm, in the Yuxi block is 155 nm, and in the Luzhou block is 147 nm.

Figure 12. Fractal dimension characteristics of typical shale samples from Longmaxi Formation. (a) CN-3, 3166.9 m, D₁; (b) WY-1, 2569 m, D₁; (c) YX-3, 3345 m, D₁; (d) LZ-6, 4031.3 m, D₁; (e) CN-3, 3166.9 m, D₂; (f) WY-1, 2569 m, D₂; (g) YX-3, 3345 m, D₂; and (h) LZ-6, 4031.3 m, D₂.

Figure 13. Distribution characteristics of permeability in the Longmaxi Formation shale in the research area. (**a**) Permeability of Longmaxi Formation Shale in different blocks; (**b**) Permeability of Longmaxi Formation Shale in different lithofacies.

5. Discussion

5.1. Factors Controlling Pore Complexity

5.1.1. Influence of TOC on Pore Complexity

Organic matter serves as the material basis for hydrocarbon gas generation and plays a crucial role in the development of organic matter pores [52]. This study explores the correlation between the fractal dimensions (D_1 and D_2) of mesopores and macropores and TOC content. The results show that both D_1 and D_2 of the Longmaxi Formation shale in the study area increases with rising TOC content. Notably, the correlation between D_1 and TOC is significantly stronger than that between D_2 and TOC (Figure 14). This indicates that organic matter pores have a significant impact on shale porosity, and this impact is more pronounced on mesopores.

Figure 14. The relationship between TOC pairs and fractal dimensions (D₁, D₂).

5.1.2. The Influence of Mineral Composition on the Pore Complexity

The minerals present in the Longmaxi Formation shale in the study area include carbonate minerals (calcite), feldspar (K-feldspar + plagioclase), quartz, and clay minerals. This study investigates the correlation between the fractal dimension of the Longmaxi Formation shale and these mineral components. The results are shown in Figure 15.

Figure 15. The relationship between mineral composition and fractal dimensions (D₁, D₂). (a) Clay vs. D₁, D₂; (b) quartz vs. D₁, D₂; (c) calcite vs. D₁, D₂; and (d) feldspar vs. D₁, D₂.

Based on Figure 15a, clay minerals show a positive correlation with both D_1 and D_2 , although not significantly. This may be due to the fact that clay minerals in the diagenetic stage are prone to structural changes, resulting in various pore and fracture types. These changes increase the surface area and volume of pores, contributing to a more complex pore structure. On the other hand, quartz does not show a clear correlation with D_1 and D₂ (Figure 15b), but D₂ shows a weak negative correlation with quartz. Smooth internal surfaces of quartz particles result in poor development of secondary and primary pores. An increase in quartz content reduces the pore surface heterogeneity, pore development, and pore structure complexity, leading to a decrease in D_2 . Carbonate minerals do not show a clear correlation with D_1 and D_2 (Figure 15c). This is because carbonate minerals mostly exist in a cement form, filling and obstructing pores, thus reducing pore connectivity. Additionally, carbonate minerals are soluble and can create dissolution pores. The dual impact of carbonate minerals on pore development contributes to the poor correlation between carbonate minerals and D_1 and D_2 . The influence mechanism of feldspar on D_1 and D_2 is different. D_1 shows an increasing trend followed by a decreasing trend with rising feldspar content, while D₂ shows a decreasing trend followed by an increasing trend (Figure 15d). This is because feldspar is a soluble mineral, and at low feldspar content, feldspar dissolution occurs. Under compaction, large pores decrease while medium and micro pores increase, resulting in an increase in D_1 and a decrease in D_2 . Furthermore, feldspar dissolution pores are generally large pores, so at high feldspar content, D_2 increases while D₁ decreases.

5.1.3. The Influence of Pore Structure on the Pore Complexity

The pore structure significantly influences the pore fractal dimension. The study explores the relationship between the Longmaxi Formation shale in the study area and its fractal dimension, as illustrated in Figure 16. The analysis reveals no clear correlation between D1 and D2 and the average pore size or pore volume (Figure 16a,b). However, a strong correlation exists between D₁ and D₂ and the specific surface area, which increases

proportionally with the specific surface area (Figure 16c). This indicates that a larger specific surface area corresponds to a more complex pore structure. Additionally, the correlation between D_2 and specific surface area is stronger than that between D_1 and the specific surface area, suggesting that the specific surface area has a more substantial influence on the pore structure of macropores. This phenomenon stems from the composition of pores in the Longmaxi Formation shale, primarily consisting of micropores and mesopores (Figure 11), which also explains the lack of a clear correlation between D2 and average pore size or pore volume.

D1 • Luzhou • Weiyuan • Yuxi • Changning D2 • Luzhou • Weiyuan • Yuxi • Changning

Additionally, this study analyzed the correlation between TOC and pore volume, as well as specific surface area. The results show no clear correlation between TOC and pore volume, but a higher correlation between TOC and specific surface area (Figure 16d). Furthermore, a significant correlation exists between D₁ and D₂ and TOC content (Figure 15), indicating the presence of organic-matter-rich pores in the pore system of the Longmaxi Formation shale in the study area, with the specific surface area mainly contributed by organic matter pores.

5.2. Shale Gas Migration Capacity and Its Influencing Factors

Shale reservoirs contain various types of pores and fractures, acting as important pathways for shale gas migration. These pathways facilitate shale gas migration differently, but the permeability of shale is collectively contributed by these various types of migration. Therefore, permeability can be obtained by calculating the parallel permeability of each type of migration [53]. Additionally, previous studies have found that the mobility of Darcy flow is 100–1000 times greater than that of slip flow [21,53]. Therefore, if 1% of the pores satisfy Darcy flow, it can be approximated that the permeability is predominantly contributed by Darcy flow. By examining the pore structure of shale in each block, the proportion of pores satisfying Darcy flow in the study area was determined, as shown in Figure 17. According to Figure 17, the proportion of pores in the Longmaxi Formation shale in the study area that satisfies Darcy flow, ranges from approximately 3.7% to 11.8%,

with an average value of 6.6%. The Changning block is approximately 7.1%, the Weiyuan block is approximately 5.8%, the Yuxi block is approximately 6.2%, and the Luzhou block is approximately 7%. This indicates that the permeability of the Longmaxi Formation shale can be attributed to Darcy flow to a significant extent.

Figure 17. The proportion of pore sizes in the Longmaxi Formation shale in the research area that meets Darcy flow.

In addition, the pore volume, complexity, and specific surface area of the pores are also important factors affecting the shale gas migration capability. Permeability, as an important parameter for assessing shale gas migration capability, was analyzed for its correlation with the proportion of pores satisfying Darcy flow, specific surface area, D1, and D2, as shown in Figure 18. According to Figure 18, there is no significant correlation between permeability and the proportion of pores satisfying Darcy flow (Figure 18a). This lack of correlation is related to pore connectivity and is negatively correlated with the specific surface area. A larger specific surface area is more favorable for gas adsorption but hinders gas migration. Additionally, a larger specific surface area indicates a higher complexity of the pores, leading to poorer gas migration capability, consistent with the correlation shown in Figure 18b. However, there is no significant correlation between shale permeability and D₁, but there is a clear negative correlation with D₂. This indicates that the primary mechanism for shale gas migration is through macropores, while mesopores play a minor role.

Figure 18. The correlation between the permeability and the proportion of pores satisfying Darcy flow, specific surface area, D₁, and D₂. (a) Permeability vs. proportion of pores satisfying Darcy flow and specific surface area; (b) permeability vs. D₁ and D₂.

6. Conclusions

This investigation into the mineralogy, organic geochemistry, pore structure, and permeability of the Longmaxi Formation shale in the study area has yielded the following key findings:

- (1) The principal lithogenic minerals in the Longmaxi Formation shale in the study area consist of quartz and calcite, resulting in the classification of siliceous shale, calcareous shale, and mixed shale.
- (2) The pore structure of the Longmaxi Formation shale in the study area is complex, with an average pore size ranging from 4.55 to 17.92 nm, pore volume ranging from 0.0131 to 0.0364 cm³/g, and specific surface area ranging from 12.01 to 21.99 m²/g. The fractal dimensions (D₁) of the mesopores range from 2.452 to 2.8548, and the fractal dimensions (D₂) of the macropores range from 2.9626 to 2.9786.
- (3) Fractal dimensions of shale are influenced by organic matter, inorganic minerals, and pore structure parameters. Both D₁ and D₂ exhibit positive correlations with TOC, clay mineral content, and specific surface area, while showing negative correlations with quartz. However, correlations with calcite content, pore volume, and average pore size are not statistically significant.
- (4) Macropores emerge as the primary contributors to the migration capability of shale reservoirs, with the permeability of shale being influenced by D₂, specific surface area, and the connectivity of macropores.

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Appendix A

Table A1. TOC: mineralogy, pore structure, and permeability of the Longmaxi Formation shale in the study area.

Samples	Depth/m	Blocks	TOC/%	Mineral Content/%						Fractal Dimension			Down on hill:	Pore	BET	Average	
				Clay	Quartz	Calcite	Plagioclase	K- Feldspar	Pyrite	D1	D ₂	Lithofacies	μD	Volume /(cm ³ /g)	Surface Area/(m²/g)	Pore Size/nm	Proportion/%
CN-1	2520.4	– Changning –	4.62	14.73	45.49	21.29	11	4.32	3.17	2.6064	2.9697	Siliceous shale	16.85	0.0241	15.53	7.35	9.56
CN-2	3148.4		4.38	19.91	48.42	19.4	5.06	2.68	4.53	2.6902	2.9636	Siliceous shale	25.63	0.0253	13.36	5.37	6.86
CN-3	3166.9		5.2	13.58	64.9	19.19	0	1.27	1.06	2.66	2.9699	Siliceous shale	15.64	0.02321	16.78	8.17	6.83
CN-4	2349.5		4.91	26.22	42.85	28.53	0.8	0.9	0.7	2.8135	2.9701	Mixed shale	16.85	0.02431	16.31	12.76	5.00
WY-1	2569	_ Weiyuan _	5.2	15.13	12.71	55.75	8.9	4.22	3.29	2.7152	2.9641	Calcareous shale	15.08	0.0336	16.3	17.92	4.72
WY-2	3116.74		4.8	15.25	10.22	60.39	8.26	3.86	2.02	2.7523	2.9678	Calcareous shale	25.63	0.0293	14.84	6.14	6.73
WY-3	3784.96		2.9	9.38	20.04	49.26	12.36	2.36	3.7	2.5693	2.9626	Mixed shale	35.22	0.0312	12.01	6.98	5.84
YX-1	4077.1	– Yuxi	5.9	17.36	4.41	54.61	14.96	4.56	4.1	2.8315	2.9786	Mixed shale	4.26	0.0296	21.06	4.89	7.25
YX-2	3754		6	17.8	20.4	45.9	9.7	4	2.2	2.8548	2.9779	Calcareous shale	11.79	0.0364	21.99	5.16	5.40
YX-3	3345		4.3	18.32	1.25	60.65	12.93	4.65	2.3	2.6316	2.9731	Calcareous shale	3.52	0.0298	18.31	4.89	6.05
LZ-1	4039	 Luzhou 	5.6	18.8	13.8	52.9	6.7	2.7	5.1	2.7432	2.9736	Calcareous shale	2.13	0.0186	19.87	12.86	11.79
LZ-2	3839.7		5.1	25.5	16.3	42.3	7.99	4.31	3.6	2.5976	2.9718	Mixed shale	15.03	0.0216	18.06	4.55	4.00
LZ-3	4926.8		4.4	21.4	49.2	18.7	5.5	1.9	3.3	2.6458	2.9746	Siliceous shale	7.05	0.0196	20.01	5.73	7.58
LZ-4	4086.4		4.5	21.5	52.8	7.9	5.1	10.4	2.3	2.6375	2.9739	Siliceous shale	6.89	0.0131	15.84	4.98	6.76
LZ-5	4918		4	15.4	63.2	11.9	3.7	2.6	3.2	2.7312	2.9638	Siliceous shale	22.45	0.0216	18.06	5.43	7.99
LZ-6	4031.3		4.2	19.1	9.1	46.7	14.64	5.96	4.5	2.452	2.9776	Mixed shale	8.54	0.0342	14.53	9.93	3.73

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