

# Spatial and Temporal Investigation of Dew Potential based on Long-Term Model Simulations in Iran

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## S.1. Detailed Model description

The original model setup included phase change between water vapor and liquid (i.e. vapor-to-liquid; condensation) as well as solid (i.e. vapor-to-solid; desublimation). The model assumes condensation on a condenser surface, which is horizontally aligned sheet of a suitable material such as low-density polyethylene (LDPE) or polymethylmethacrylate (PMMA). The condenser sheet is also assumed to be at 2 meters from the ground and thermally insulated from the ground. In the model, only dew is considered whereas precipitation and fog were not accounted.

The model describes the dew formation on a flat surface (i.e. condenser) and it is based on the mass and heat balance equation

$$\frac{dT_c}{dt}(C_c m_c + C_w m_w + C_i m_i) = P_{rad} + P_{cond} + P_{conv} + P_{lat}, \quad (1)$$

where  $dT_c/dt$  is the change rate in the condenser temperature.  $C_c$ ,  $C_w$ , and  $C_i$  are the specific heat capacities of the condenser, water, and ice; respectively. Here,  $m_c$ ,  $m_w$ , and  $m_i$  are the masses of the condenser, water, and ice; respectively. The right-hand side includes the heat exchange involved in the process:  $P_{rad}$  is the incoming and outgoing radiation,  $P_{cond}$  is the conductive heat exchange between the condenser surface and the ground,  $P_{conv}$  is the convective heat exchange, and  $P_{lat}$  is the latent heat released by the condensation or desublimation of water. This equation was solved by using a fourth-order Runge-Kutta with 10 s time-step. All terms and nomenclature are described in more details in [Tables S1](#) and [S2](#).

The model was setup so that it assumes similar conditions for the phase-change of pre-existing water or ice on the condenser sheet. For instance, if the water on the condenser is in liquid phase (i.e.  $m_w > 0$ ) and the condenser temperature  $T_c < 0$  °C, then the sheet is losing energy (i.e. the right-hand side of [Equation \(1\)](#) is negative). In that case, instead of solving [Equation \(1\)](#),  $T_c$  is assumed to be constant and the lost mass from the liquid phase of water is transferred to the cumulated mass of ice; i.e. the water is transformed from liquid phase to solid phase. Consequently, [Equation \(1\)](#) is replaced by

$$L_{wi} \frac{dm_w}{dt} = P_{rad} + P_{conv} + P_{lat}, \quad (2)$$

where  $L_{wi}$  is the latent heat of fusion. If the water on the condenser is in solid phase (i.e.  $m_i > 0$ ) and the condenser temperature  $T_c > 0$  °C, a similar equation is assumed for the change rate of ice mass ( $m_i$ ).

Note that [Equation \(2\)](#) is not related to the condensation of water; it only describes the phase change of the already condensed water or ice on the condenser. For the water condensation rate, which is assumed independent of [Equation \(2\)](#), the mass-balance equation is then assumed as

$$\frac{dm}{dt} = \max[0, S_c k (P_{sat}(T_d) - P_c(T_c))], \quad (3)$$

where  $m$  represents either the mass of ice ( $m_i$ ) or water ( $m_w$ ) depending on whether  $T_c$  is below or above 0 °C.  $P_{sat}(T_d)$  is the saturation pressure at the dew point temperature and  $P_c(T_c)$  is the vapor pressure over the condenser sheet.  $k = h / L_{vw} \gamma = 0.622h / C_a p$  is the mass transfer coefficient, where  $L_{vw}$  is the specific latent heat of water vaporization,  $\gamma$  is the psychrometric constant,  $C_a$  is the specific heat capacity of air, and  $p$  is the atmospheric air pressure. Here,  $h = 5.9 + 4.1 u (511 + 294) / (511 + T_a)$  is the heat transfer coefficient, where  $u$  and  $T_a$  are the prevailing horizontal wind speed and the ambient temperature at 2 meters from the ground.

In practice, the wettability of the surface affects the vapor pressure  $P_c$  directly above it. In other words,  $P_c$  is lower over a wet surface; and thus, condensation may take place even if  $T_c > T_a$ . It is also assumed that the processes included in [Equation \(3\)](#) undergo irreversible condensation; i.e. there is no evaporation or sublimation during daytime even if  $T_c > T_a$ . Furthermore, the model simulation resets the cumulative values for water and ice condensation at noon and takes the preceding maximum value of  $m_w + m_i$  as the representative daily yield. This way, the model simulation replicates the daily manual dew water collection of the condensed water around sunrise; i.e. after which  $T_c$  is often above the dew point temperature.

**Table S1:** Description of the dew formation model by listing the terms in Eq. (1).

<b>Term</b>	<b>Unit</b>	<b>Description</b>
$dT_c/dt$	$K s^{-1}$	Change rate of the condenser temperature
$T_c$	K	Temperature of the condenser
$t$	s	Time. Here the time step in the model was 10 s
$C_c$	$J kg^{-1} K^{-1}$	Specific heat capacity of the condenser. For low-density polyethylene (LDPE) and polymethylmethacrylate (PMMA) it is $2300 J kg^{-1} K^{-1}$
$C_i$	$J kg^{-1} K^{-1}$	Specific heat capacity of ice ( $2110 J kg^{-1} K^{-1}$ )
$C_w$	$J kg^{-1} K^{-1}$	Specific heat capacity of water ( $4181.3 J kg^{-1} K^{-1}$ )
$m_c$	kg	Mass of the condenser given by $m_c = \rho_c S_c \delta_c$ where $\rho_c$ , $S_c$ , and $\delta_c$ are the density (here it is $920 kg m^{-3}$ ), surface area (here it is $1 m^2$ ), and thickness of the condenser (here it is $0.39 mm$ )
$m_i$	kg	Mass of ice
$m_v$	kg	Mass of water, representing the cumulative mass of water that has
$P_{rad}$	W	Heat exchange due to incoming and outgoing radiation  $P_{rad} = (1 - a)S_c R_{sw} + \epsilon_c S_c R_{lw} - S_c \epsilon_c \sigma T_c^4$ where $a$ is the condenser short-wave albedo (here it is $0.84$ ), $S_c$ is the condenser surface area (here it is $1 m^2$ ), $\epsilon_c$ is the emissivity of the condenser (here it is $0.94$ ), $\sigma$ is Stephan-Boltzmann constant ( $5.67 \times 10^{-8} W m^{-2} K^{-4}$ ), $T_c$ [K] is the temperature of the condenser, and $R_{sw}$ and $R_{lw}$ [ $W m^{-2}$ ] are the incoming short-wave radiation (i.e. surface solar radiation downwards) and incoming long-wave radiation (i.e. surface thermal radiation downwards)
$P_{cond}$	W	Conductive heat exchange between the condenser surface and the ground. For simplicity, we assumed that the condenser is perfectly insulated from the ground; i.e. $P_{cond} = 0$
$P_{conv}$	W	Convective heat exchange  $P_{conv} = S_c (T_a - T_c) h$ where $S_c$ is the condenser surface area (here it is $1 m^2$ ), $T_a$ [K] is the ambient temperature at 2 meters from the ground, $T_c$ [K] is the temperature of the condenser, and $h$ [ $W m^{-2} K^{-1}$ ] is the heat transfer coefficient that is estimated based on a semi-empirical equation (Richards, 2009)  $h = 5.9 + 4.1 WS (511 + 294) / (511 + T_a)$ and here $WS$ [ $m s^{-1}$ ] is the prevailing horizontal wind speed at 2 meters from the ground.
$P_{lat}$	W	Latent heat released by the condensation or desublimation of water  $P_{lat} = \begin{cases} L_{vw} \frac{dm_w}{dt} & T_c > 0 \text{ } ^\circ C \\ L_{vi} \frac{dm_i}{dt} & T_c < 0 \text{ } ^\circ C \end{cases}$ where $L_{vw}$ [ $J kg^{-1}$ ] is the specific latent heat of water vaporization and $L_{vi}$ [ $J kg^{-1}$ ] is specific latent heat of water desublimation. Here, $dm_w/dt$ is the change rate of water whereas $dm_i/dt$ is the change rate of ice

**Table S2:** A list of nomenclature.

<i>Parameter</i>	<i>Unit</i>	<i>Description</i>
$\alpha$	--	Albedo of condenser sheet
$C_a$	$\text{J kg}^{-1} \text{K}^{-1}$	Specific heat capacity of air
$C_c$	$\text{J kg}^{-1} \text{K}^{-1}$	Specific heat capacity of the condenser
$C_i$	$\text{J kg}^{-1} \text{K}^{-1}$	Specific heat capacity of ice
$C_w$	$\text{J kg}^{-1} \text{K}^{-1}$	Specific heat capacity of water
$DP$	K	Dew point temperature
$h$	$\text{W K}^{-1} \text{m}^{-2}$	Heat transfer coefficient
$k$	$\text{Per s}^{-1}$	Mass transfer coefficient
$L_{vi}$	$\text{J kg}^{-1}$	Specific latent heat of desublimation for water
$L_{vw}$	$\text{J kg}^{-1}$	Specific latent heat of vaporization for water
$L_{wi}$	$\text{J kg}^{-1}$	Latent heat of fusion
$m_c$	kg	Mass of the condenser
$m_i$	kg	Mass of ice
$m_w$	kg	Mass of water
$p$	Pa	Atmospheric air pressure
$p_c$	Pa	Vapour pressure over condenser
$p_{sat}$	Pa	Saturation pressure of water
$P_{cond}$	W	Conductive heat exchange between the condenser surface and the ground
$P_{conv}$	W	Convective heat exchange
$P_{lat}$	W	Latent heat released by the condensation or desublimation of water
$P_{rad}$	W	Heat exchange due to incoming and outgoing radiation
$R_{lw}$	$\text{W m}^{-2}$	Surface thermal radiation downwards
$R_{sw}$	$\text{W m}^{-2}$	Surface solar radiation downwards
$S_c$	$\text{m}^2$	Surface area of condenser
$T_a$	K	Ambient temperature at 2 meters
$T_c$	K	Temperature of the condenser
$U_{10}$	$\text{m s}^{-1}$	Horizontal wind speed component at 10 meters
$V_{10}$	$\text{m s}^{-1}$	Horizontal wind speed component at 10 meters
$WS$	$\text{m s}^{-1}$	Prevailing horizontal wind speed at 2 meters
$z_0$	m	Surface roughness
$\delta_c$	mm	Condenser sheet thickness
$\epsilon_c$	--	Emissivity of condenser sheet
$\gamma$	$\text{Pa K}^{-1}$	Psychrometric constant
$\sigma$	$\text{W m}^{-2} \text{K}^{-4}$	Stefan-Boltzmann constant

## S.2. Meteorological data as input for the model

The model input parameters includes:

- horizontal and vertical wind components ( $U_{10}$  and  $V_{10}$ ) at 2 meters
- surface roughness ( $z_0$ )
- ambient temperature and dew point ( $T_a$  and  $D_P$ ) at 2 meters
- short-wave and long-wave surface solar radiation ( $R_{sw}$  and  $R_{lw}$ )

These were obtained from the European Centre for Medium Range Weather Forecast (ECMWF) Interim Reanalysis. ERA-Interim is a reanalysis of the global atmosphere covering the data-rich period since 1979 and continuing in real time [37,38], which has a horizontal resolution of  $0.75^\circ$  (approximately 80 km) and 60 vertical levels. The reanalysis combines available data sources (satellite, radiosondes, air craft, buoy data, stations etc.) into a coherent and balanced form of the atmospheric dynamic and thermodynamic state [39].

The ECMWF data-base is differentiated into three main categories: reanalysis, instantaneous forecast, and accumulated forecast. In our case,  $U_{10}$ ,  $V_{10}$ ,  $T_a$ , and  $DP$  were obtained from the reanalysis fields at 00:00, 06:00, 12:00 and 18:00 UTC whereas at 03:00, 09:00, 15:00 and 21:00 UTC were instantaneous forecasted fields. Both the forecasted and the reanalysis fields provided every 3 hours. According to the ECMWF data-base, the horizontal wind components ( $U_{10}$  and  $V_{10}$ ) are provided at 10 meters. Therefore, the wind speed at 2 meters was calculated by using a logarithmic wind profile

$$WS = \frac{\log\left(\frac{z+z_0}{z_0}\right)}{\log\left(\frac{10+z_0}{z_0}\right)} \sqrt{U_{10}^2 + V_{10}^2}, \quad (4)$$

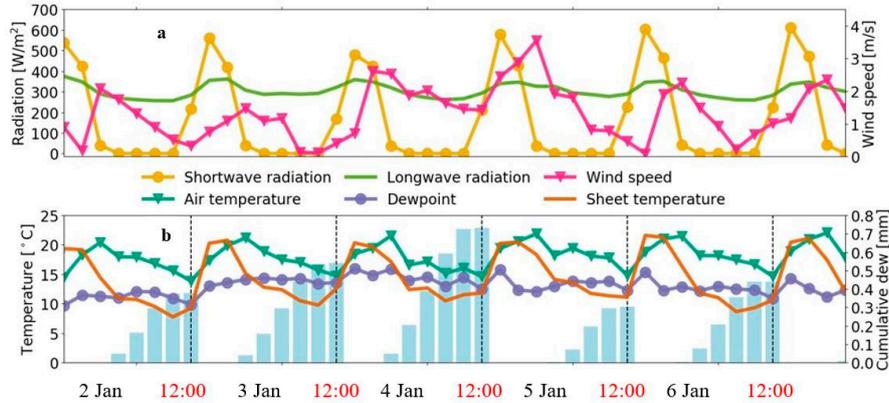
where  $z_0$  is the surface roughness and  $U_{10}$  and  $V_{10}$  are the horizontal wind speed components at 10 meters. It is important to understand that Equation (4) is a simple equation that is valid during certain conditions. For instance, in stable conditions (such as at night) Equation (4) overestimates wind speed at 2 meters whereas in unstable conditions equation (4) underestimates the wind speed at 2 meters.

According to the ECMWF data-base,  $z_0$  was obtained as an instantons forecast parameter whereas  $R_{sw}$  and  $R_{lw}$  were accumulated forecasted fields. The mean  $R_{sw}$  and  $R_{lw}$  in a time interval is obtained by taking the difference of the accumulated values between the corresponding time steps divided by the time difference in seconds. The result is a mean value for that time interval 00:00 or 12:00. All input parameters had a horizontal resolution 0.25 degree (approximately 30 km).

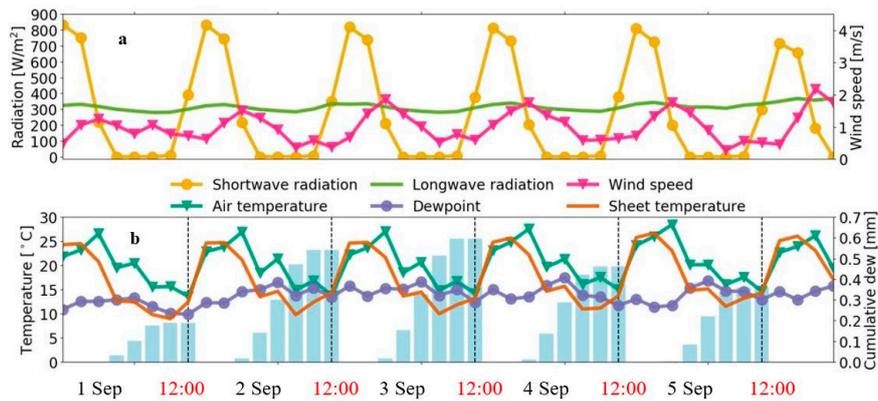
The model output for dew occurrence is given in liters of water collected on a  $1 \text{ m}^2$  condenser sheet (i.e.  $\text{L}/\text{m}^2$ ). However, this can be converted to units of mm (i.e. equivalent to rainfall).

### S.3. Illustrations for detailed model input and output

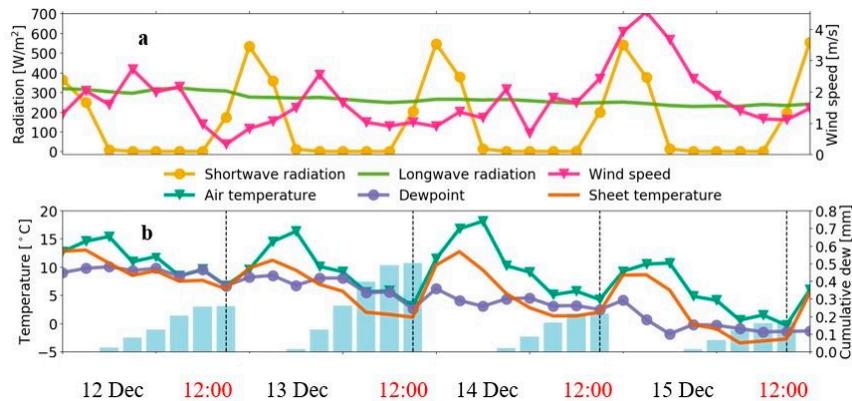
Here we selected three examples to illustrate the model input and output parameters. The selected examples were: Bandarabas during 2–6 January 2005 (Figure S1), Ramsar during 1–5 September 2011 (Figure S2), and Tabas during 12–15 December 2002 (Figure S3). As mentioned in the main part of the manuscript, the selection of these examples was to represent three regions in Iran. As for the time periods, they were based on the model simulation result for maximum yield in that year and season.



**Figure S1:** An example model simulation for dew formation during 2-6 January 2005 at Bandarabas, Iran. (a) the short-wave and long-wave radiation (left y-axis) and wind speed (right y-axis) whereas (b) the cumulative dew formation on the condenser (bars linked to the y-axis right axis) and the ambient temperature, dew point, and condenser sheet temperature (left y-axis).



**Figure S2:** An example model simulation for dew formation during 1-5 September 2011 at Ramsar, Iran. (a) the short-wave and long-wave radiation (left y-axis) and wind speed (right y-axis) whereas (b) the cumulative dew formation on the condenser (bars linked to the y-axis right axis) and the ambient temperature, dew point, and condenser sheet temperature (left y-axis).



**Figure S3:** An example model simulation for dew formation during 12-15 December 2002 at Tabas, Iran. (a) the short-wave and long-wave radiation (left y-axis) and wind speed (right y-axis) whereas (b) the cumulative dew formation on the condenser (bars linked to the y-axis right axis) and the ambient temperature, dew point, and condenser sheet temperature (left y-axis).