

Article

Chemical Composition and Sources of Marine Aerosol over the Western North Pacific Ocean in Winter

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Abstract: Atmospheric deposition of long-range transported continental substances from natural and anthropogenic sources affects biogeochemical processes in marine systems. Emissions of sea spray contribute aerosol particles to the marine atmosphere. Despite the importance of continental dispersion and atmospheric processes involving aerosol particles within remote marine atmosphere, knowledge of the sources of various water-soluble ions is limited because of insufficient observations. Concentrations of Total suspended particulates (TSPs) and major inorganic ions (Cl^- , Na^+ , SO_4^{2-} , Mg^{2+} , Ca^{2+} , K^+ , NO_3^- , NH_4^+), as well as organic nitrogen (ON-N) values, were measured in marine aerosol collected over the western north Pacific (WNP) during a cruise from 3 December 2014 to 13 March 2015. Aerosol samples were analyzed to determine their chemical characteristics and a source apportionment for this region and the continental influence on the open ocean when air masses are from continent in winter. TSP mass concentrations ranged from 14.1 to 136.0 $\mu\text{g}/\text{m}^3$ with an average of $44.8 \pm 28.1 \mu\text{g}/\text{m}^3$. Concentrations of TSPs and major ions were higher near the coast (close to Qingdao and Xiamen) and lower over the open ocean. The total mass of inorganic ions and organic nitrogen accounted for 51.1% of the total TSP. Cl^- had highest concentrations among the major inorganic ions, followed by SO_4^{2-} , NO_3^- , Mg^{2+} , Ca^{2+} , K^+ , and NH_4^+ , respectively. However, Cl^- showed a deficit relative to Na^+ in most samples, likely related to heterogeneous reactions within the marine atmosphere. Most SO_4^{2-} , Mg^{2+} , Ca^{2+} , and K^+ were from sea salt, while other major ions were from continental sources. The non-sea-salt (nss) fractions of Ca^{2+} , Mg^{2+} and K^+ were derived from continental crust, while nss- SO_4^{2-} and NO_3^- were derived from anthropogenic sources. ON had several sources, including reactions of NO_x with volatile organic compounds (anthropogenic sources) or NH_3 with gaseous hydrocarbons, as well as crustal and marine biogenic sources.

Keywords: marine aerosol; source apportionment; Western North Pacific Ocean; total suspended particulates; ionic composition

1. Introduction

The long-distance transport of atmospheric aerosol particles from natural and anthropogenic sources affects ocean ecosystems, climate, and atmospheric chemical processes [1–3]. Atmospheric deposition is an important pathway in which nutrients fertilize the open ocean, stimulating

phytoplankton growth and influencing the marine biogeochemical cycles [3]. Marine aerosols are comprised of both primary and secondary inorganic and organic components [4]. Primary aerosol particles are emitted directly from natural and anthropogenic sources, while secondary aerosol particles form from condensable atmospheric gases [2]. In the remote marine atmosphere, ions from sea salt make significant contributions to the aerosol mass of Na^+ , Cl^- , SO_4^{2-} , Mg^{2+} , K^+ and Ca^{2+} [1–3]. Both natural and anthropogenic components from continental sources also may affect NO_3^- , SO_4^{2-} , Ca^{2+} , K^+ , and trace metal concentrations over open ocean [1,5–7]. Duce et al. [5] suggested that reactive nitrogen from anthropogenic sources accounts for a third of the ocean's external nitrogen supply. Nitrogen deposition to the open ocean is assumed to be 80–100% from anthropogenic sources [8]. In addition, the atmospheric inputs of sulfur and nitrogen oxides from anthropogenic or natural sources to the ocean may accelerate rates of ocean acidification [9]. Other atmospheric substances may also influence the ocean, such as iron, phosphorus, cadmium, and persistent organic pollutants. However, studies suggest that the effect of anthropogenic substances on oceans may not be as great as previously expected [10–12].

During the East Asian winter monsoon, strongly flowing cold air masses bring large amounts of mineral dust and anthropogenic pollutants from the Asian continent out over the western Pacific Ocean [13–15]. As a result of ongoing rapid economic and industrial development the western Pacific Ocean has received more anthropogenic substances from the Asian continent via long-range transport [2,13,14,16]. It has been reported that higher concentrations of NO_3^- and NH_4^+ were observed within air masses from the Asian continent in winter, compared with air masses from the Pacific Ocean in summer at subtropical Okinawa Island [16]. Both NO_x and NH_y can absorb onto dust or crust particles from the Asian continent to transport to the western North Pacific Ocean (WNP), producing higher concentrations of non-sea-salt (nss) SO_4^{2-} , nss- K^+ , and nss- Ca^{2+} in winter and spring. Such levels contrast the low levels in summer, when air masses are from the Pacific Ocean [2,13,17]. Thus, it is necessary to study the continental influence on the open ocean when air masses are from continent in winter.

In this study, we measured concentrations of major ions (Cl^- , Na^+ , SO_4^{2-} , Mg^{2+} , Ca^{2+} , K^+ , NO_3^- , NH_4^+) and organic nitrogen (ON) in total suspended particulates (TSP) sampled over the WNP in winter from 3 December 2014 to 13 March 2015 (winter) to identify the spatial variation and sources of major ions and ON. We used a back-trajectory analysis model to assign source regions to these components, with a particular focus on identifying continental (or nss) sources in this region.

2. Methodology

2.1. Aerosol Collection and Chemical Analysis

The TSP samples were collected using a special high-volume KC-1000 sampler ($1.05 \pm 0.03 \text{ m}^3/\text{min}$; Laoshan Institute for Electronic Equipment, China), which was fixed on the front uppermost observation deck of the Research Vessel Kexue to avoid contamination from the ship exhaust, as described in our previous study [18]. The cruise covered a period of 101 days from 3 December 2014 to 13 March 2015. The cruise occurred during the winter period of the East Asian monsoon. The aerosol optical thickness (AOT), dust, SO_2 and NO_2 show marked spatial variation during this period (Figure 1), with higher values nearshore to China and lower ones over the Pacific Ocean. The average wind streamlines of 1000 hPa over the western North Pacific Ocean (WNP) indicate a wind speed varying from 4 to 12 m/s. Wind speed increased from about 7 to 12 m/s, as the vessel traveled away from the shore. Sampling sites are shown in Figure 2.

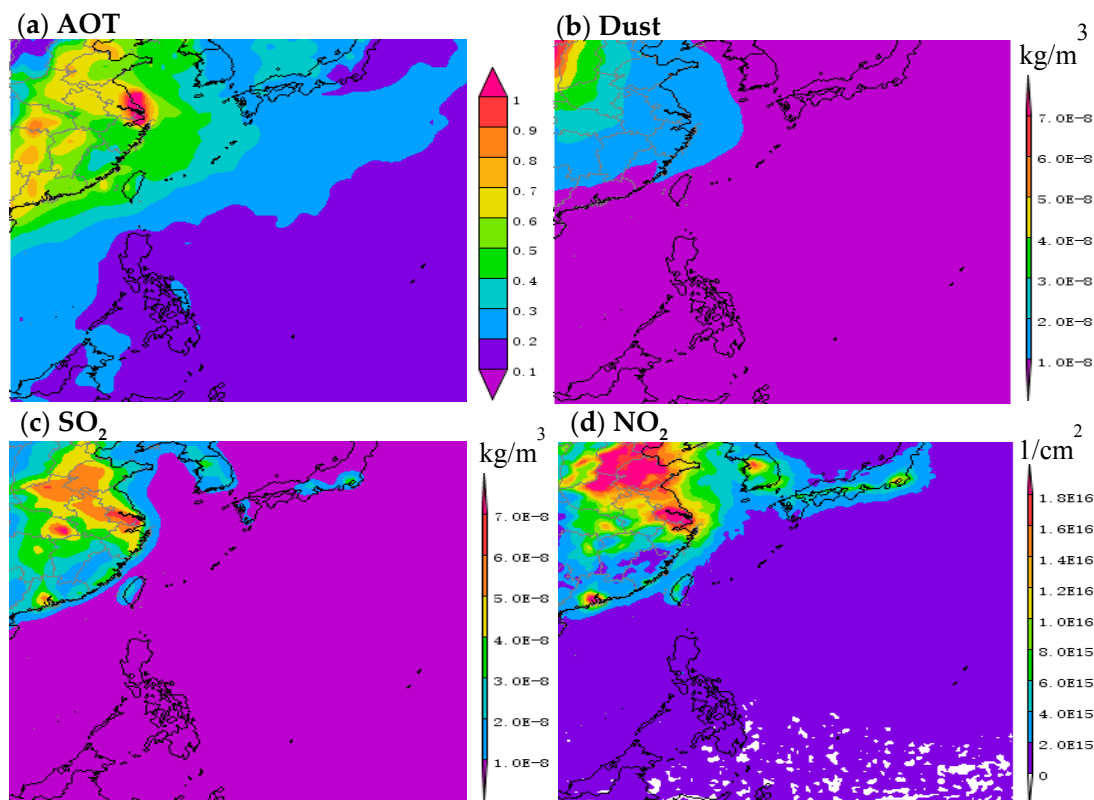


Figure 1. Distribution of seasonal average aerosol optical thickness (AOT), dust, SO_2 and NO_2 over the western North Pacific during the sampling period. Monthly AOT, dust, SO_2 products with $1 \times 1^\circ$ (MODIS-Terra MOD08_M3 v6), $0.5 \times 0.625^\circ$ (MERRA-2 Model M2TMNXAER v5.12.4), $0.5 \times 0.625^\circ$ (MERRA-2 Model M2TMNXAER v5.12.4) resolution, daily NO_2 products with $0.25 \times 0.25^\circ$ (OMI OMNO2d v003) resolution were downloaded from Giovanni (<https://giovanni.gsfc.nasa.gov>).

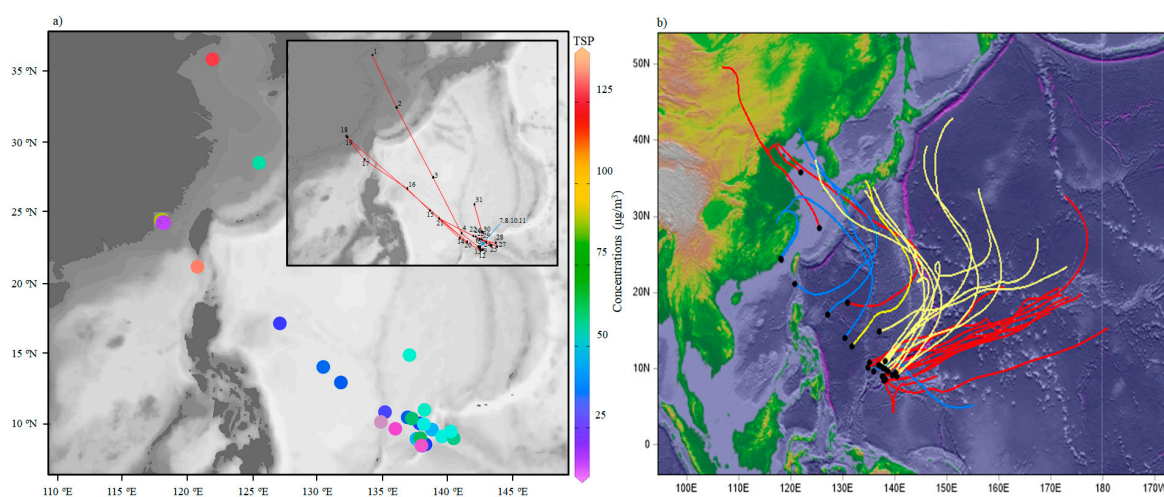


Figure 2. (a) Geographic distribution of TSP concentrations and start sampling sites over the western North Pacific. (b) Five-day back trajectories of every starting sampling site over the western North Pacific. Red, blue and yellow lines represent the start sampling site in December 2014, January and February 2015, respectively.

All samples ($n = 31$) were collected on precombusted Pallflex membrane filters (Type: Tissuquartz™ Filters, 2500 QAT-UP, reorder No. 7204; size: 8×10 in; Pall Corp.;

Port Washington, NY, USA). Collection time was 48 h (or 2 d/sample). Four blank filters were taken in this cruise. Once a sampling interval was completed, the sample holder and filter were placed in a clean zip-locked plastic bag, folded in two with exposed sides face to face. All samples and blank filters were stored in a refrigerator at $-20\text{ }^{\circ}\text{C}$ until analysis in the laboratory. Detailed information about sampling procedures and analysis of all major components are described in Xiao et al. [2].

Briefly, concentrations of major anions (Cl^- , SO_4^{2-} , NO_3^-), as well as water-soluble metal and nonmetal elements (Ca, K, Mg, Na) were determined using a DionexTM ICS-90 ion chromatography system (Thermo Fisher Scientific Inc.; Waltham, MA, USA) and a Vista-MPXTM simultaneous inductively coupled plasma optical emission spectrometer (ICP-OES; Varian Inc.; Palo Alto, CA, USA), respectively. The NH_4^+ concentrations were analyzed using spectrophotometry, following treatment with Nessler's reagent. The detection limits of Cl^- , NO_3^- , and SO_4^{2-} were 0.03, 0.08, and 0.075 mg/L, respectively, and the relative standard deviation of these ions of standard samples were 2.55%, 1.16%, and 1.36%, respectively [19,20]. The detection limits of Ca, K, Mg, and Na were 0.003, 0.06, 0.0005, and 0.02 mg/L, respectively, and the relative standard deviation of these ions of standard samples were less than 1.5% [19,20]. The detection limit of NH_4^+ was 0.1 mg/L and its relative standard deviation was less than 5.0% [19,20]. A good ionic balance between cations and anions in μeq ($y = 1.03x$, $R^2 = 0.92$) in our samples suggested that no major ionic components were missed [21]. Detailed information about ON analysis is described in Jickells et al. [22]. The ON-N fraction in the aerosol samples was determined by subtracting the amounts of NO_3^- and NH_4^+ from total nitrogen (TN), i.e., $\text{ON-N} = \text{TN} - \text{NH}_4^+ - \text{NO}_3^-$. The total nitrogen and carbon (TN and TC) concentrations in aerosols were determined by Elemental Analyzer (EA-IRMS, Flash 2000). Briefly, six circular filters with a diameter of 4.5 mm were weighted and placed in a tin cup, and then combusted subsequently at $980\text{ }^{\circ}\text{C}$ in combustion tube within the EA. During combustion, all the nitrogen- and carbon-containing species were converted to N_2 and CO_2 , respectively. These gaseous components were separated on a packed column of a gas chromatograph installed within the EA. Analytical precision determined from multiple measurements of various standards was and 5% for TN and TC concentrations.

2.2. Concentrations of the Non-Sea-Salt Fraction

To determine the source contributions of major ions within aerosol particles, the concentrations of the nss fractions were calculated, using Na^+ as a reference element for a marine origin, as follows [2,19]:

$$\text{nss} - X_{\text{aerosol}} = X_{\text{aerosol}} \times (1 - (X/\text{Na}^+)_{\text{sea}} / (X/\text{Na}^+)_{\text{aerosol}}) \quad (1)$$

where X represents various ionic components, and (X/Na^+) values for seawater are from Keene et al. [23].

2.3. Back Trajectories and Aerosol Optical Thickness

To determine the long-distance transport of atmospheric pollutants, 5-d (120 h) back trajectories of air masses arriving at the location at the start of each sampling period were computed using the open source software TrajStat (Version 1.2.2.6; Meteoinfo), with the top of the model set to 1000 m above sea level, as described in Xiao et al. [2,18,19,24]. The back trajectories of all starting locations are shown in Figure 2. Modeled AOT, dust, SO_2 and NO_2 data are shown in Figure 1 (data from <https://giovanni.gsfc.nasa.gov>). Meanwhile, to provide a regional context to this study, Navy Aerosol Analysis and Prediction System Global Aerosol Model data (<https://www.nrlmry.navy.mil/aerosol/#aerosolobservations>) for SO_2 , dust, and smoke surface concentrations are shown in Figure S1.

3. Results and Discussion

3.1. Total Suspended Particulate and Major Ion Concentrations of Marine Aerosol Samples

The average TSP concentration in samples was $44.8 \pm 28.1 \mu\text{g}/\text{m}^3$, with a range from 14.1 to $136.0 \mu\text{g}/\text{m}^3$. The highest TSP concentration was measured in the Bashi Channel near Xiamen, followed by coastal marine areas near Qingdao (Figure 2). The average TSP concentration over the WNP was much lower than those values in winter at Yongxing Island in the South China Sea ($158.9 \mu\text{g}/\text{m}^3$) [2] and over the Northern Yellow Sea ($145.2 \mu\text{g}/\text{m}^3$) [1], consistent with their respective AOT distributions (Figure 1), both which are marginal seas of northwestern Pacific Ocean. The average TSP concentration over the WNP was similar to the value obtained in previous study in this region ($36.7 \mu\text{g}/\text{m}^3$) [3]. The percentage of total major inorganic ions contributing to the TSP over the WNP was similar to values for the Pacific Oceans [3], but higher than over the South China Sea [2].

The total mass concentration of major inorganic ions (Cl^- , Na^+ , SO_4^{2-} , Mg^{2+} , Ca^{2+} , K^+ , NO_3^- , NH_4^+) and ON-N was $22.9 \mu\text{g}/\text{m}^3$; it had a similar spatial variation to TSP (Figure 1). These components accounted for 51.1% of the TSP. Cl^- had the highest concentration among these ions (Figures 3 and 4), ranging from $2.1 \mu\text{g}/\text{m}^3$ to $21.3 \mu\text{g}/\text{m}^3$, with an average value of $9.9 \pm 4.6 \mu\text{g}/\text{m}^3$. It was followed by Na^+ (2.0 – $10.9 \mu\text{g}/\text{m}^3$; $5.8 \pm 2.2 \mu\text{g}/\text{m}^3$), SO_4^{2-} (0.8 – $12.5 \mu\text{g}/\text{m}^3$; $3.3 \pm 3.0 \mu\text{g}/\text{m}^3$), NO_3^- (0.1 – $10.3 \mu\text{g}/\text{m}^3$; $1.2 \pm 2.3 \mu\text{g}/\text{m}^3$), Mg^{2+} (0.24 – $1.71 \mu\text{g}/\text{m}^3$; $0.78 \pm 0.34 \mu\text{g}/\text{m}^3$), Ca^{2+} ($0.15 \pm 2.44 \mu\text{g}/\text{m}^3$; $0.49 \pm 0.50 \mu\text{g}/\text{m}^3$), K^+ (0.08 – $0.94 \mu\text{g}/\text{m}^3$; $0.36 \pm 0.24 \mu\text{g}/\text{m}^3$), and NH_4^+ (0.02 – $1.81 \mu\text{g}/\text{m}^3$; $0.21 \pm 0.45 \mu\text{g}/\text{m}^3$), as shown in Figure 3. The average value of ON-N was $0.94 \pm 1.02 \mu\text{g}/\text{m}^3$, with a range from $0.05 \mu\text{g}/\text{m}^3$ to $4.64 \mu\text{g}/\text{m}^3$.

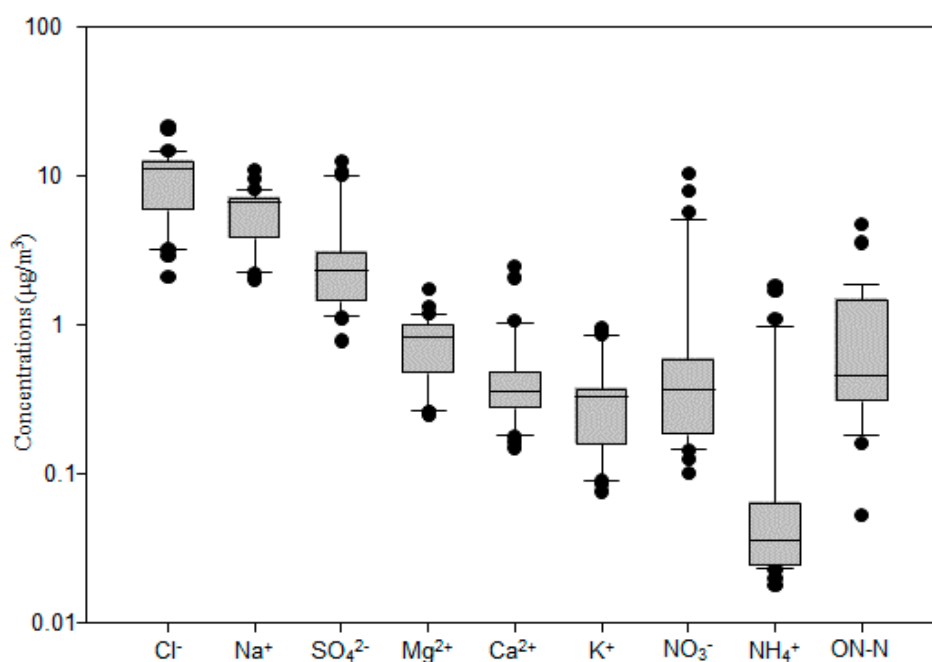


Figure 3. Mass concentrations of aerosol chemical species over the western North Pacific. The boundary of the box indicates the 25th and 75th percentile. The lines within the box mark the mean. Whiskers above and below the box indicate the 90th and 10th percentiles.

Sea salt (ss) and nss contributions to Cl^- , SO_4^{2-} , Mg^{2+} , K^+ , and Ca^{2+} were determined using Na^+ as the reference element for sea salt [2,19]. Generally, most of Cl^- values were deficit relative to Na^+ , with an average deficit of $14.8 \pm 16.4\%$, and a range of 0.9–85.9% (Figure 5). On average, fractions of $39.7 \pm 12.3\%$, $41.8 \pm 17.0\%$, $30.0 \pm 19.0\%$ and $13.1 \pm 6.9\%$ of the total SO_4^{2-} , Ca^{2+} , K^+ , and Mg^{2+} were derived from a non-sea-salt origin (Figure 6), indicating that most of the SO_4^{2-} , Ca^{2+} , K^+ , and Mg^{2+}

was derived from sea salt. The concentrations of nss-SO_4^{2-} , nss-Ca^{2+} , nss-K^+ and nss-Mg^{2+} were 1.89 ± 2.92 , 0.27 ± 0.47 , 0.15 ± 0.19 and $0.12 \pm 0.10 \mu\text{g}/\text{m}^3$, respectively (Figure 6).

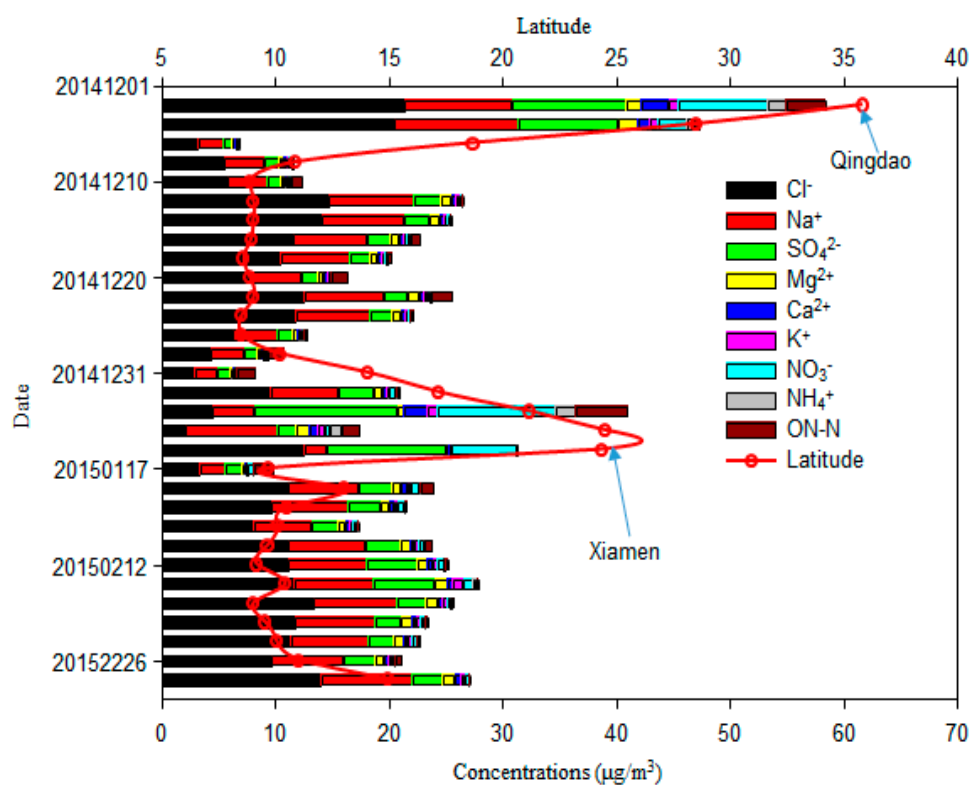


Figure 4. Concentrations of major ions in aerosols over the western North Pacific.

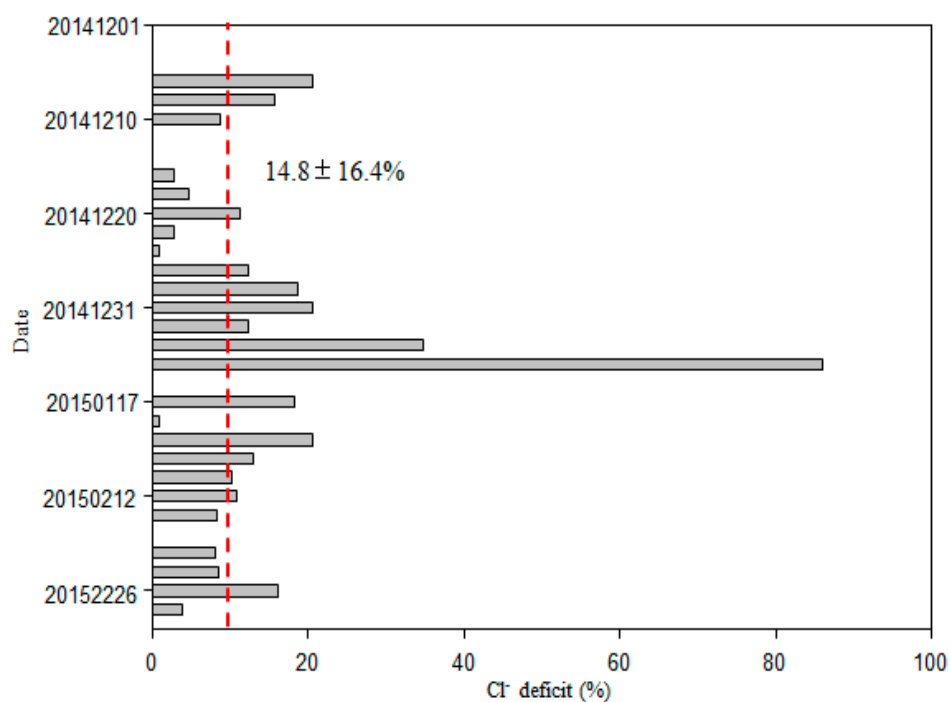


Figure 5. Percent Cl^- deficit at different sites, with the dashed line for the averaged value.

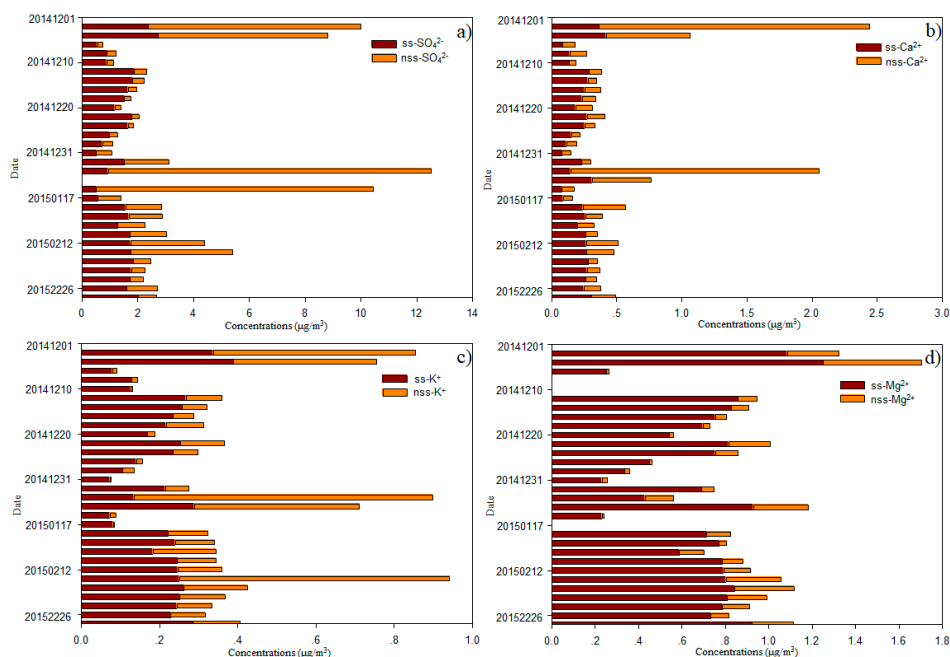


Figure 6. (a) nss- and ss- SO_4^{2-} , (b) nss- and ss- Ca^{2+} , (c) nss- and ss- K^+ , and (d) nss- and ss- Mg^{2+} at different sites over the western North Pacific.

3.2. Major Ion Characteristics and Sources of Aerosol over the Western North Pacific

Typical marine ions (Na^+ and Cl^-) accounted for 71.5% of total inorganic ions over the WNP, which was higher than levels reported over the South China Sea in winter [2]. The concentrations of Na^+ and Cl^- were close to those of a previous study of the Pacific Ocean (5.8 and $9.0 \mu\text{g}/\text{m}^3$, respectively) [3], but higher than those of other marine areas in winter, such as Northern South China Sea, and Northern Yellow Sea [1,25]. The average molar ratio of Cl^- to Na^+ was 1.14, which was slightly smaller than the value of 1.17 for seawater, indicating that Cl^- was in deficit. This may reflect Cl^- evaporation to the atmosphere. In the marine atmosphere, sea salt aerosol can react with inorganic acids, such as HNO_3 and H_2SO_4 , as well as NO_2 and SO_2 in heterogeneous reactions to release gaseous reactive chlorine compounds, including HCl , Cl_2 , ClNO_2 , and HOCl to the atmosphere [2,3,13,18,23,26,27]. Such reactions result in a deficit of Cl^- relative to Na^+ , when compared with fresh sea salt aerosol. The value of the Cl^- deficit measured on this cruise (Figure 5) was slightly lower than that reported for the Atlantic and Pacific Oceans by Zhang et al. [3], suggesting that the aerosol samples in this study were also influenced by anthropogenic emissions. In this study, Cl^- concentrations were correlated with SO_4^{2-} when air masses were from the open ocean ($R = 0.51$, $p < 0.01$; Table 1a) because they both originate from sea salt. However, it was not correlated with nss- SO_4^{2-} ($R = 0.59$, $p > 0.05$; $R = 0.18$, $p > 0.05$, respectively, Table 1) or NO_3^- ($R = 0.36$, $p > 0.05$; $R = 0.05$, $p > 0.05$, respectively, Table 1), suggesting that the Cl^- deficit cannot be fully explained by the formation of inorganic salts, but organic acids, such as oxalic acid, fulvic acid and malonic acid, and atmospheric processes must also contribute to the Cl^- deficit [13,18,26,28]. Generally, inorganic acids play a major role in chloride depletion, but in some cases this role may be played by organic acids, especially in the absence of continental anthropogenic emissions [13,28]. Over the open ocean, water-soluble organic acids can be transported from polluted continental regions or formed in situ by photochemical oxidation of precursor biogenic volatile organic compounds (VOCs) and unsaturated fatty acids [28]. We found that in some samples (e.g., 1, 2, 6, 7, 19 and 27) there was enrichment of Cl^- , suggesting that a small fraction of Cl^- is derived from anthropogenic sources [19]. Another study observed that dry particles were devoid of acetate but enriched in Cl^- related to evaporation of acetic acid [26].

Table 1. Correlation coefficients among major ions in aerosol over the western North Pacific with air masses from continent ((a), $n = 8$) and open ocean ((b), $n = 23$).

	Na ⁺	Cl [−]	SO ₄ ^{2−}	nss-SO ₄ ^{2−}	K ⁺	nss-K ⁺	Mg ²⁺	nss-Mg ²⁺	Ca ²⁺	nss-Ca ²⁺	NH ₄ ⁺	NO ₃ [−]	ON
(a)													
Na ⁺	1	0.65	0.21	0.17	0.73 *	0.49	0.99 **	0.92 **	0.52	0.40	0.44	0.05	−0.26
Cl [−]		1	0.59	0.37	0.38	0.19	0.64	0.56	0.45	0.38	0.18	0.36	−0.08
SO ₄ ^{2−}			1	0.98 **	0.55	0.61	0.24	0.29	0.67	0.69	0.57	0.92 **	0.66
nss-SO ₄ ^{2−}				1	0.62	0.71	0.20	0.29	0.63	0.65	0.68	0.93 **	0.78
K ⁺					1	0.95 **	0.75 *	0.77 *	0.89 **	0.83 *	0.91 **	0.57	0.56
nss-K ⁺						1	0.52	0.57	0.90 **	0.89 **	0.96 **	0.71 *	0.77
Mg ²⁺							1	0.96 **	0.52	0.40	0.45	0.07	−0.24
nss-Mg ²⁺								1	0.50	0.39	0.45	0.10	−0.16
Ca ²⁺									1	0.99 **	0.94 **	0.79 *	0.80
nss-Ca ²⁺										1	0.94 **	0.84 **	0.87 *
NH ₄ ⁺											1	0.74 *	0.86 *
NO ₃ [−]												1	0.89 *
ON													1
(b)													
Na ⁺	1	0.97 **	0.61 **	0.25	0.67 **	0.38	0.98 **	0.60 **	0.80 **	0.30	0.15	0.13	−0.44 *
Cl [−]		1	0.51 **	0.18	0.61 **	0.33	0.95 **	0.56 *	0.74 **	0.23	0.09	0.05	−0.41 *
SO ₄ ^{2−}			1	0.92 **	0.86 **	0.80 **	0.66 **	0.56 *	0.76 **	0.61 **	0.75 **	0.71 **	−0.35
nss-SO ₄ ^{2−}				1	0.72 **	0.79 **	0.33	0.42	0.53 **	0.60 **	0.84 **	0.80 **	−0.21
K ⁺					1	0.95 **	0.75 **	0.72 **	0.66 **	0.41	0.78 **	0.50 *	−0.39
nss-K ⁺						1	0.50 *	0.62 **	0.48 *	0.38	0.91 **	0.57 **	−0.30
Mg ²⁺							1	0.81 **	0.80 **	0.31	0.25	0.21	−0.41
nss-Mg ²⁺								1	0.49 *	0.22	0.45 *	0.46 *	−0.22
Ca ²⁺									1	0.81 **	0.34	0.50 *	−0.26
nss-Ca ²⁺										1	0.39	0.68 **	0.02
NH ₄ ⁺											1	0.61 **	−0.13
NO ₃ [−]												1	−0.10
ON													1

*: $p < 0.05$; **: $p < 0.01$.

Mg^{2+} , another typical ion in the marine atmosphere [2], was well-correlated with Na^+ ($R = 0.99$ and 0.98 , $p < 0.01$, respectively; Table 1). The ratios of Mg^{2+} to Na^+ were close to that of seawater, indicating that Mg^{2+} was mainly derived from this source. The concentrations of ss-Mg^{2+} versus nss-Mg^{2+} are shown in Figure 6d; these ratios further support this view. In this study, ss-Mg^{2+} and nss-Mg^{2+} accounted for 86.9% and 13.1% of the total Mg^{2+} in the TSP, respectively. Thus, the ratio of nss-Mg^{2+} to total Mg^{2+} was higher than ratios measured over the Pacific Ocean [3]. The nss-Mg^{2+} may originate from crustal sources, given that nss-Mg^{2+} is often treated as a crustal-derived ion and element in continental studies [19,29]. Ca^{2+} ions are also treated as tracers of crustal material [2,19]. No relationship between Ca^{2+} and Mg^{2+} was found when air masses were from continent ($p > 0.05$; Table 1a) and a strong relationship was found when air masses were from the open ocean ($R = 0.80$, $p < 0.01$; Table 1b), although a much larger portion of nss-Ca^{2+} ($41.8 \pm 17.0\%$) than nss-Mg^{2+} ($13.1 \pm 6.9\%$) was observed (Figure 6). Larger contributions of nss-Ca^{2+} were found over the South China Sea [2]. The average molar ratio of nss-Mg^{2+} to nss-Ca^{2+} (0.760) in samples of this cruise was higher than the average ratio for continental crust (0.569) [30], indicating that nss-Mg^{2+} is enriched relative to nss-Ca^{2+} within the marine atmosphere. K^+ is commonly used as a tracer for biomass and biofuel burning [31]; however, we found that K^+ showed little influence from biomass and biofuel burning sources over the remote WNP (Figures 2 and S1). Typically, in remote marine environments, sea salt provides 70% of K^+ , e.g., in the Pacific marine atmosphere, while nss-K^+ accounts for 30% of the total water-soluble K^+ (Figure 6). In this study, strong correlations between K^+ and Ca^{2+} (both $R = 0.89$, $p < 0.01$; Table 1), as well as nss-K^+ and nss-Ca^{2+} ($R = 0.66$, $p < 0.01$; Table 1a) when air masses from continent were observed. These relationships suggest that nss-K^+ over the WNP is primarily derived from the crust. In Figures 1 and S1, we show that Asian dust influences the remote WNP, which may provide crustal sources of Ca^{2+} , Mg^{2+} and K^+ . We found higher contributions of nss-Ca^{2+} to total Ca^{2+} , nss-Mg^{2+} to total Mg^{2+} and nss-K^+ to total K^+ in coastal regions (near Qingdao and Xiamen), compared with open ocean environments, suggesting that coastal regions were more affected by dust (Figures 1 and S1). Higher contributions of these nss ions, especially nss-Mg^{2+} and nss-K^+ , were also observed when air masses originated from near continent rather than from open ocean regions (Figure 2). Our results indicate that it is more difficult for Ca^{2+} (compared with Mg^{2+} and K^+) to transport to open ocean environments.

The fraction of nss-SO_4^{2-} accounted for 39.7% of total SO_4^{2-} in samples of this cruise. This percentage is close to that observed over the Northern and Southern Atlantic Oceans, but higher than that in a previous study of the Pacific Ocean [3]. However, the average concentration of nss-SO_4^{2-} was much higher than over the Atlantic Ocean, Pacific Ocean, and Mediterranean Sea [3,15], but lower than over the Yellow Sea, East China Sea, Yongxing Island in the South China Sea [1,2,32], Sea of Okhotsk, and Sea of Japan [15]. The concentrations of nss-SO_4^{2-} recorded near the coast on this cruise were close to values over the Yellow Sea and East China Sea [32]. Like our results for nss-Mg^{2+} and nss-K^+ , we found a higher percentage of nss-SO_4^{2-} to total SO_4^{2-} in coastal regions and in air masses originating from near the continent, compared with those from open ocean regions. However, Figure S1 clearly shows that continental SO_4^{2-} is transported to and affects these open ocean environments. The average concentration of NO_3^- was $1.2 \pm 2.3 \mu\text{g}/\text{m}^3$ in samples from our cruise; these are higher than those observed over the Northern Atlantic and Pacific Oceans [3], but lower than those from the China Sea [3,18,25,32]. Very high concentrations of NO_3^- were observed near the coast close to Qingdao and Xiamen (Figure 4), whereas much lower values were recorded over the open WNP in December 2014, especially in samples of air masses from the open ocean (Figure 2). Zhang et al. [3] surmised that the background level of NO_3^- was about $0.12 \mu\text{g}/\text{m}^3$ in this region. The average concentration of NO_3^- was $0.20 \mu\text{g}/\text{m}^3$ in December between 8.5°N and 14.0°N , which is close to this background level, suggesting that NO_3^- was not markedly influenced by continental surprising. When air masses were from the continent (in February 2015), the average concentration of NO_3^- reached $0.49 \mu\text{g}/\text{m}^3$. In this case, continental sources appear to have contributed majorly (76%) to the total NO_3^- . NO_3^- was also well correlated with SO_4^{2-} ($R = 0.92$ and 0.71 , respectively, $p < 0.01$;

Table 1) and nss-SO_4^{2-} ($R = 0.93$ and 0.80 , respectively, $p < 0.01$; Table 1), indicating that the relatively high concentrations of NO_3^- and nss-SO_4^{2-} were related to continental sources. In contrast, we found low concentrations of NH_4^+ over the open ocean in samples from our cruise, with average values of $0.04 \mu\text{g}/\text{m}^3$ in December and $0.07 \mu\text{g}/\text{m}^3$ in February. Other studies also reported low values of NH_4^+ over the WNP [4,33,34], and over the South China Sea [2]. These results indicate that gaseous ammonia or particulate ammonium are difficult to transport long distance, and that NH_4^+ over the WNP may be from local emissions. Zhao et al. [35] suggested that atmospheric nitrogen deposition to the northwestern Pacific decreased rapidly downwind of the Asian continent. In addition, ON-N ($0.94 \mu\text{g}/\text{m}^3$) had much higher concentrations than NO_3^- -N ($0.26 \mu\text{g}/\text{m}^3$) and NH_4^+ -N ($0.16 \mu\text{g}/\text{m}^3$), accounting for 74.0% of the total nitrogen (NO_3^- -N, NH_4^+ -N and ON-N), with a range of 28.0–97.4%. The average concentration of ON-N over the WNP was markedly lower than at other sites, including inland cities (Baoji), coastal cities (Chennai and Rondonia), and marine sites (Gosan, Okinawa Island, Bermuda, Northern South China Sea, Yellow Sea, and North Pacific Ocean) [4,10,36–41]. The ON content was correlated with nss-Ca^{2+} ($R = 0.87$, $p < 0.01$; Table 1a) when air masses from continent, indicating that part of this ON was derived from primary crustal material. A good correlation was also found between ON and both NO_3^- and NH_4^+ (Table 1a), suggesting that secondary formation was a major source of ON. The most common secondary ONs are organic nitrates, such as hydroxyalkyl nitrates produced by the gas phase reactions of NO_x and VOCs or water-insoluble ON produced by the reactions of gaseous hydrocarbons with NH_3 [22,38]. Altieri et al. [10] and Pavuluri et al. [38] also suggested that ON may be derived from marine biogenic emissions. A high concentration of protein in western North Pacific Ocean in the winter was found [42]. So, a weak relationship with Na^+ when air masses were from the open ocean (Table 1b) was found.

4. Conclusions

Cl^- and Na^+ made the greatest contributions to total inorganic ions, followed by SO_4^{2-} , NO_3^- , Mg^{2+} , Ca^{2+} , K^+ , and NH_4^+ . ON-N had higher average concentrations than NO_3^- and NH_4^+ in all samples. A Cl^- deficit in samples was likely caused by formation of acids and/or other atmospheric processes. Relate to most SO_4^{2-} , Mg^{2+} , Ca^{2+} , and K^+ were derived from sea salt, with the remainder from continental sources, including crustal and anthropogenic sources. When air masses were from continental regions, NO_3^- concentrations were significantly higher, indicating that the WNP was influenced by anthropogenic sources. In contrast, NH_4^+ may reflect local oceanic emissions. ON had several sources, including reactions of NO_x with VOCs, gaseous hydrocarbons with NH_3 , and both crustal and marine biogenic sources.

Supplementary Materials: The following are available online at <http://www.mdpi.com/2073-4433/9/8/298/s1>.

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Conflicts of Interest: The authors declare that they have no conflict of interest.

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