



# Article Impacts of Temporal-Spatial Variant Background Ionosphere on Repeat-Track GEO D-InSAR System

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Academic Editors: Zhong Lu, Xiaofeng Li and Prasad S. Thenkabail Received: 25 July 2016; Accepted: 28 October 2016; Published: 4 November 2016

Abstract: An L band geosynchronous synthetic aperture radar (GEO SAR) differential interferometry system (D-InSAR) will be obviously impacted by the background ionosphere, which will give rise to relative image shifts and decorrelations of the SAR interferometry (InSAR) pair, and induce the interferometric phase screen errors in interferograms. However, the background ionosphere varies within the long integration time (hundreds to thousands of seconds) and the extensive imaging scene (1000 km levels) of GEO SAR. As a result, the conventional temporal-spatial invariant background ionosphere model (i.e., frozen model) used in Low Earth Orbit (LEO) SAR is no longer valid. To address the issue, we firstly construct a temporal-spatial background ionosphere variation model, and then theoretically analyze its impacts, including relative image shifts and the decorrelation of the GEO InSAR pair, and the interferometric phase screen errors, on the repeat-track GEO D-InSAR processing. The related impacts highly depend on the background ionosphere parameters (constant total electron content (TEC) component, and the temporal first-order and the temporal second-order derivatives of TEC with respect to the azimuth time), signal bandwidth, and integration time. Finally, the background ionosphere data at Isla Guadalupe Island (29.02°N, 118.27°W) on 7–8 October 2013 is employed for validating the aforementioned analysis. Under the selected background ionosphere dataset, the temporal-spatial background ionosphere variation can give rise to a relative azimuth shift of dozens of meters at most, and even the complete decorrelation in the InSAR pair. Moreover, the produced interferometric phase screen error corresponds to a deformation measurement error of more than 0.2 m at most, even in a not severely impacted area.

**Keywords:** Geosynchronous SAR (GEO SAR); temporal-spatial variation; background ionosphere; D-InSAR

# 1. Introduction

Geosynchronous synthetic aperture radar (GEO SAR) [1–3] runs at geosynchronous orbit which has the height of about 36,000 km. It has a short revisit time of less than 24 h and extensive imaging coverage of more than 1000 km [4,5], which helps to realize the fast revisit and imaging of scenes of interest [6,7]. Therefore, the combination of GEO SAR and differential SAR interferometry (D-InSAR) can realize timely surface deformation detection, which has great advantages in the evaluation and forecast of natural disasters (earthquakes, landslides, volcano eruptions, etc.) [8–10].

The study of D-InSAR technology has a long history of more than 20 years. It can help to obtain the deformation information by virtue of accurate interferometric phases, which are deeply demonstrated in [11–15]. Nevertheless, the related study of D-InSAR processing in GEO SAR is rarely conducted. In 2002, Madsen et al. proposed the basic conception of using GEO D-InSAR to

monitor the worldwide earthquake activities and analyzed the feasibility of the system [16]. Later, Monti-Guarneri and Bruno [2,17,18] studied the temporal advantages of a GEO D-InSAR system in surface deformation monitoring and the decorrelation impacts on GEO SAR data processing. Hu et al. [19] proposed methods of obtaining the optimal SAR interferometry (InSAR) pair for height and deformation retrieval in the repeat-track GEO InSAR system.

Although a GEO D-InSAR system has great potential in fast deformation detection, unfortunately, an L band GEO D-InSAR system will be seriously impacted by the ionosphere due to its low working frequency [20–22]. Meyer et al. [23] pointed out that the ionosphere could not only give rise to the relative image shift and the decorrelation of the InSAR pair, which will degrade the quality of the generated interferogram, but also the interferometric phase screen error could also be induced as a result of the degraded deformation retrieval accuracy. In Low Earth Orbit (LEO) SAR, because of the short integration time (less than 1 s) and small imaging scene (about dozens of kilometers width), a significant amount of research about the ionosphere impacts on D-InSAR was conducted only based on the assumption of the ionosphere temporal-spatial frozen model, which considers the ionosphere to be invariant within the integration time and the scene [23,24]. However, GEO SAR has an integration time of more than 1 min (Figure 1a) and more than 1000 km  $\times$  1000 km coverage, and the ionospheric total electron content (TEC) is obviously changing within the long integration time (Figure 1b) and the large spatial scope (Figure 1c). Therefore, the ionosphere temporal-spatial frozen model is invalid for the analysis of the impacts of background ionosphere variation on D-InSAR processing in a GEO SAR case. In addition, although some researches have applied the temporal-spatial background ionosphere variation model in previous studies, they mainly focus on the analysis of the impacts of the ionosphere on GEO SAR imaging [1,25,26]. Considering the above reasons, impacts of the temporal-spatial variant ionosphere on the deformation retrieval in GEO D-InSAR need to be specially considered. Generally, the ionospheric variation can be classified as two parts: the slow background ionosphere variation caused by the large-scale irregularities and the fast ionospheric scintillation caused by the small-scale irregularities [1,27,28]. Since occurrence of ionospheric scintillation has a special diurnal and geographical pattern (from early evening after sunset to midnight and at low and high latitude), ionospheric scintillation often occurs in a low probability. Even in the tropical zone, only 14% of the SAR images could be impacted by ionospheric scintillation [29]. More importantly, ionospheric scintillation could be avoided by applying the proper orbit optimization and the autofocusing algorithms [30,31]. Therefore, we focus on the impacts of the background ionosphere variation on GEO D-InSAR in this paper.



**Figure 1.** (a) Integration time of L band geosynchronous synthetic aperture radar (GEO SAR) in different orbit positions (0.24 m wavelength, 20 m azimuth resolution, a "figure-8" inclined orbit (0.07 eccentricity, 42,164 km semi-major axis and 53° inclination)); (b) An example of the vertical *TEC* variation at the site ( $10^{\circ}$ N,  $150^{\circ}$ W) within 250 s (UTC Time 23:06:00–23:10:10, 16 September 2013); (c) An example of the vertical *TEC* variation within a spatial scope (UTC Time 23:00, 16 September 2013). The data in (**b**,**c**) derive from the United States Total Electron Content (USTEC) data released by the National Oceanic and Atmospheric Administration (NOAA) [32], which is near real-time assessment of the *TEC* by GPS signals. The USTEC uncertain in the vertical *TEC* (*VTEC*) is about 2 TECU (TECU is the unit of *TEC*, 1 TECU = 1 × 10<sup>16</sup> electrons/m<sup>2</sup>) during quiet geomagnetic conditions.

In this paper, aimed at the issue that the frozen background ionosphere model fails under the long integration time and large imaging scene of GEO SAR, we firstly deduce the GEO D-InSAR signal model in the presence of the temporal-spatial background ionosphere variation in Section 2. Then, based on the proposed model, expressions of the relative image shift and the decorrelation of the GEO InSAR pair, and the interferometric phase screen error model are given in Section 3. In addition, we illustrate the boundary parameters of the background ionosphere when these influences could be ignored in the D-InSAR processing. In Section 4, by utilizing the National Oceanic and Atmospheric Administration (NOAA) [32] TEC data at Isla Guadalupe Island, the relative image shift and the decorrelation in the interferometry processing brought by the temporal-spatial background ionosphere variation are simulated and analyzed. Moreover, the corresponding interferometric phase screen error is also evaluated and the results show that the obvious deformation measurement errors will be brought under the impacts of the temporal-spatial background ionosphere variation. Our studies are conducted based on the classical "figure-8" inclined orbit with an observation of America by an L band (0.24 m) GEO SAR system. Because of different orbits, required integration times and working frequencies, the performances of different GEO SAR systems under the impacts of the temporal-spatial background ionosphere variation are distinct. Finally, Section 5 concludes this paper.

# 2. GEO D-InSAR Signal Model in the Presence of the Temporal-Spatial Background Ionosphere Variation

The sketch map of a GEO D-InSAR system in the presence of the background ionosphere is shown in Figure 2. Point **P** is the target on the ground.  $R_G$  is the slant range of GEO SAR.  $L_G$  is the synthetic aperture of GEO SAR, and  $l_P$  is the equivalent projections of the corresponding synthetic apertures on the ionosphere.  $H_{iono}$  is the representative height of the ionosphere. The color variation of the ionosphere represents the *TEC* variation. Because of the long integration time of GEO SAR, *TEC* is obviously changing, not only between the revisiting interval, but within the integration time.



Figure 2. Sketch map of a GEO D-InSAR system in the presence of the background ionosphere.

Considering that the impacts of the background ionosphere variation on GEO SAR signals vary along with the slow time and the spatial position, the *TEC* variation in one track of GEO SAR, based on the series expansion with respect to the slow time, is written as:

$$TEC(t_a; \mathbf{P}) = TEC_0(\mathbf{P}) + k_1(\mathbf{P})t_a + k_2(\mathbf{P})t_a^2 + \dots$$
(1)

where **P** is the position vector of a target in the scene,  $t_a$  is the azimuth time,  $TEC(t_a; \mathbf{P})$  is the *TEC* of point **P** changing with the slow time along the signal propagation path (*STEC*), and *TEC*<sub>0</sub> is the constant component of  $TEC(t_a; \mathbf{P})$ , which corresponds to the *TEC* value in the aperture center moments.  $k_i(\mathbf{P})$  (i = 1, 2, ...) is the temporal *i*th-order derivative of *TEC* with respect to the azimuth time.

Generally, under a design of moderate resolution (20 m), the ionospheric pierce trajectory of the GEO SAR with a "figure-8" nadir-point trajectory is less than 3 km (smaller than that in the LEO SAR case (5–10 km)). The *TEC* variation caused by the spatial ionospheric pierce points (IPPs) variation (about  $10^{-2} \sim 10^{-1}$  TECU) is far smaller than the *TEC* variation of a spatial invariant IPP within the integration time of GEO SAR (about  $10^{-1} \sim 1$  TECU). Therefore, we ignore the impacts of *TEC* variation caused by the spatial IPPs variation. Based on Equation (1), the GEO SAR signal *s* (*f<sub>a</sub>*, *t<sub>a</sub>*; **P**) is given as (in the range frequency domain and the azimuth time domain):

$$s(f_a, t_a; \mathbf{P}) = W(f_a, t_a; \mathbf{P}) \cdot \exp\left[j\left(\frac{\pi f_a^2}{k_r} - \underbrace{\frac{4\pi K \cdot TEC(t_a; \mathbf{P})}{cf_a}}_{\phi_i(f_a)} - 4\pi \frac{f_0 R(t_a; \mathbf{P})}{c}\right)\right]$$
(2)

where  $k_r$  is the frequency modulation rate,  $f_a$  is the range frequency,  $f_0$  is the carrier center frequency, c is light velocity,  $R(t_a; \mathbf{P})$  is the accurate slant range,  $K = 40.28 \text{ m}^3/\text{s}^2$  is a constant,  $W(f_a, t_a; \mathbf{P})$  is the signal envelope, and  $\phi_i(f_a)$  represents the phase induced by the background ionospheric impacts.

Considering that the designed frequency bandwidth of GEO SAR is about 20 MHz under the moderate resolution of 20 m and the variation of *TEC* is generally smaller than 100 TECU in natural conditions, dispersion in range could be ignored [25]. Thus, after imaging processing in the time domain, we have

$$s(t - \tau, t_a, \mathbf{P}) = \overline{W}(t - \tau, t_a; \mathbf{P}) \cdot \exp\left[-j\left(4\pi \frac{f_0 R_c(\mathbf{P})}{c} + \phi_{ion}(\mathbf{P})\right)\right]$$
(3)

where *t* is the fast time,  $\tau$  is the imaged position of the target in range,  $\overline{W}(t - \tau, t_a; \mathbf{P})$  is the envelope of the signal,  $R_c(\mathbf{P})$  is the slant from the target to the aperture center, and  $\phi_{ion}(\mathbf{P})$  is the integrated phase error brought by the background ionosphere variation after imaging, which is expressed as

$$\phi_{ion}\left(\mathbf{P}\right) = \arg\left\{\int_{-T_a/2}^{T_a/2} \exp\left[-j\frac{4\pi K \cdot TEC(t_a;\mathbf{P})}{cf_0}\right] dt_a\right\}$$
(4)

where  $T_a$  represents the integration time.

Considering the D-InSAR processing, after complex multiplication of the GEO InSAR pair, the interferometric signal is expressed as

$$s_{M}\left(t-\tau,t_{a,T_{0}};\mathbf{P}\right)s_{S}^{*}\left(t-\tau,t_{a,T_{0}+nT_{r}};\mathbf{P}\right)=\widetilde{W}\left(t-\tau,t_{a};\mathbf{P}\right)\exp\left[-j\left(4\pi\frac{f_{0}\Delta R_{c}\left(\mathbf{P}\right)}{c}+\Delta\phi_{ion}\left(\mathbf{P}\right)\right)\right]$$
(5)

where  $s_M$  and  $s_S$  are the master image and the slave image of the InSAR pair, respectively.  $\widetilde{W}(t - \tau, t_a; \mathbf{P})$  is the envelope of the interferometric signal,  $\Delta R_c(\mathbf{P})$  is the slant difference from the target to the aperture center.  $T_0$  is the aperture center moments of the first track and  $T_r$  is the revisit period. *n* is an integer to represent the number of revisit periods which have been passed.  $\Delta \phi_{ion}$  (**P**) is the interferometric phase brought by the background ionosphere variation, which is shown as:

$$\Delta\phi_{ion}\left(\mathbf{P}\right) = \arg\left\{\int_{-T_a/2}^{T_a/2} \exp\left[-j\frac{4\pi K \cdot TEC_M(t_a;\mathbf{P})}{cf_0}\right] dt_a\right\} - \arg\left\{\int_{-T_a/2}^{T_a/2} \exp\left[-j\frac{4\pi K \cdot TEC_S(t_a;\mathbf{P})}{cf_0}\right] dt_a\right\}$$
(6)

where  $TEC_M(t_a; \mathbf{P})$  and  $TEC_S(t_a; \mathbf{P})$  are the *TECs* of point **P**, changing with the slow time during two acquisitions of the GEO InSAR pair.

#### 3. Impacts of the Temporal-Spatial Background Ionosphere Variation on GEO D-InSAR

#### 3.1. Relative Image Shift and the Decorrelation of the InSAR Pair

The signal envelope delay brought by the temporal-spatial background ionosphere variation during two acquisitions of the GEO InSAR pair will give rise to a relative range shift of the InSAR pair  $\Delta R$ . The range shift of each image in the InSAR pair is determined by the first-order derivative of  $\phi_i(f_a)$  in Equation (2) with respect to  $f_a$  at  $f_0$ . Thus,  $\Delta R$  is shown as:

$$\Delta R(\mathbf{P}) \approx \frac{K}{f_0^2} \{ TEC_{0,M}(\mathbf{P}) - TEC_{0,S}(\mathbf{P}) \}$$
(7)

where  $TEC_{0,M}(\mathbf{P})$  and  $TEC_{0,S}(\mathbf{P})$  are the constant components of  $TEC_M(t_a; \mathbf{P})$  and  $TEC_S(t_a; \mathbf{P})$ , respectively.

As the azimuth shift of each SAR image is proportional to the temporal first-order derivative of *TEC* with respect to the azimuth time [25], the relative azimuth shift of the InSAR pair  $\Delta a$  can be expressed as:

$$\Delta \mathbf{a}(\mathbf{P}) \approx \frac{2K}{cf_0} \left\{ \frac{v_{M,0} k_{1,M}(\mathbf{P})}{f_{dr0,M}} - \frac{v_{S,0} k_{1,S}(\mathbf{P})}{f_{dr0,S}} \right\}$$
(8)

where  $v_{M,0}$  and  $v_{S,0}$  are the velocities of the two tracks of the InSAR pair at the aperture center moments, respectively.  $f_{dr0,M}$  and  $f_{dr0,S}$  are the Doppler frequency modulation rates of the two tracks of the InSAR pair at the aperture center moments, respectively.  $k_{1,M}$  and  $k_{1,S}$  are the temporal first-order derivatives of *TECs* during two acquisitions of the GEO InSAR pair, respectively.

According to Equations (7) and (8), the constant *TEC* components during two acquisitions of the GEO InSAR pair and the carrier center frequency determine the relative range shift, while the relative azimuth shift depends on the temporal first-order derivatives of *TECs* during two acquisitions of the GEO InSAR pair, the velocities of the two tracks of the InSAR pair at the aperture center moments and the Doppler frequency modulation rates of the two tracks of the InSAR pair at the aperture center moments. Generally, the relative range shift and the azimuth shift will give rise to the mismatch of the InSAR pair, and finally cause the decorrelation, which is shown as:

$$\gamma_i = \gamma_1 \gamma_2 \tag{9}$$

where  $\gamma_i$  is the coefficient that corresponds to the mismatch decorrelation,

$$\gamma_1 = \operatorname{sinc}\left\{\frac{2KB\operatorname{sin}\theta_r}{cf_0^2} \left[TEC_{0,M}(\mathbf{P}) - TEC_{0,S}(\mathbf{P})\right]\right\}$$
(10)

where  $\gamma_1$  represents the mismatch decorrelation in range,  $\theta_r$  is the incident angle, B is the bandwidth,

$$\gamma_2 = \operatorname{sinc}\left\{\frac{2KT_a}{cf_0} \left(k_{1,M}(\mathbf{P}) - k_{1,S}(\mathbf{P})\right)\right\}$$
(11)

where  $\gamma_2$  represents the mismatch decorrelation in azimuth.

On one hand, large differences in the constant *TEC* components and the temporal first-order derivatives of *TECs* during two acquisitions of the GEO InSAR pair will aggravate the decorrelation. On the other hand, a larger bandwidth and incident angle (higher range resolution), and a longer integration time (higher azimuth resolution) can also make the coherence decrease. In addition, as the constant components and the temporal first-order derivatives of *TECs* are spatial-varying, the produced relative shifts and the decorrelation of the InSAR pair are not uniform in the scene. To obtain a high coherence in the generated interferogram, we need co-registration to eliminate the relative shifts of the InSAR pair. Nevertheless, the regularity of the relative shifts of the InSAR pair will decrease when the constant components and the temporal first-order derivatives of *TECs* change dramatically within the scene. A high-order polynomial model could be utilized under the condition to realize a more accurate co-registration. Likewise, we can ignore the relative shifts of the InSAR pair when the differences between the constant components and the temporal first-order derivatives of *TECs* during two acquisitions are small. Generally, the accuracy requirement for co-registration is the mismatch of the pixels that should be smaller than 1/10 of a pixel [12], which can only bring a decorrelation of less than 4%. Hereby, without considering the changes of the velocities and the Doppler frequency

modulation rates of the two tracks, the boundary parameters of the background ionosphere, when the impacts of the relative image shift and the decorrelation of the GEO SAR interferometry pair could be ignored, can be expressed as:

$$\begin{cases} |\Delta TEC_0(\mathbf{P})| \le \frac{cf_0^2}{20B\sin\theta_r K} \\ |\Delta k_1| \le \frac{cf_0}{20T_a K} \end{cases}$$
(12)

# 3.2. Interferometric Phase Screen Error

Based on the previous signal model,  $\Delta \phi_{ion}$  will introduce errors in the interferometric phase screen, resulting in deformation errors. Thus, it is necessary to discuss its impacts. Combining Equation (1) in Equation (4), we obtained:

$$\phi_{ion}\left(\mathbf{P}\right) = \arg\left\{\int_{-T_a/2}^{T_a/2} \exp\left[-j\frac{4\pi K}{cf_0} \cdot \left(TEC_0(\mathbf{P}) + k_1(\mathbf{P})t_a + k_2(\mathbf{P})t_a^2 + \ldots\right)\right] dt_a\right\}.$$
 (13)

Generally, impacts of the temporal third-order derivative of *TEC* with respect to the azimuth time are about  $10^{-8}$ – $10^{-10}$  TECU and it is smaller than 0.01 TECU after the integration within the integration time. Thus, we ignore the impacts of the *TEC* variation components above the temporal second-order derivative. By deduction (See Appendix A for details), the temporal-spatial variation interferometric phase screen error  $\Delta \phi_{ion}$  can be written as

$$\Delta\phi_{ion}\left(\mathbf{P}\right) = \frac{4\pi K}{cf_0} \left\{ \underbrace{\underbrace{\widehat{\Phi}_0(\mathbf{P})}_{\text{Space Variation}} - \underbrace{\underbrace{\widehat{\Phi}_1(\mathbf{P}) + \widehat{\Phi}_2(\mathbf{P})}_{\text{Time-spatial Variation}} \right\}$$
(14)

where  $\Phi_0(\mathbf{P})$  is the component of spatial background ionosphere variation,  $\Phi_1(\mathbf{P})$  and  $\Phi_2(\mathbf{P})$  are the components of temporal-spatial background ionosphere variation, which are shown as:

$$\begin{aligned}
\widehat{\Phi}_{0}(\mathbf{P}) &= TEC_{0,M}(\mathbf{P}) - TEC_{0,S}(\mathbf{P}) \\
\widehat{\Phi}_{1}(\mathbf{P}) &= \left(\frac{k_{1,M}(\mathbf{P})^{2}}{4k_{2,M}(\mathbf{P})} - \frac{k_{1,S}(\mathbf{P})^{2}}{4k_{2,S}(\mathbf{P})}\right) \\
\widehat{\Phi}_{2}(\mathbf{P}) &= \frac{cf_{0}}{4\pi K} \left( \arctan\left[\frac{4\pi K}{3cf_{0}} \left(\frac{3k_{1,M}(\mathbf{P})^{2}}{4k_{2,M}(\mathbf{P})} + \frac{T_{a}^{2}k_{2,M}(\mathbf{P})}{4}\right)\right] \\
&- \arctan\left[\frac{4\pi K}{3cf_{0}} \left(\frac{3k_{1,S}(\mathbf{P})^{2}}{4k_{2,S}(\mathbf{P})} + \frac{T_{a}^{2}k_{2,S}(\mathbf{P})}{4}\right)\right] \right)
\end{aligned}$$
(15)

where  $k_{2,M}$  and  $k_{2,S}$  are the temporal second-order derivative of *TECs* during two acquisitions of the GEO InSAR pair, respectively.

According to Equation (14), it not only contains the component of spatial background ionosphere variation, but it includes the components of the temporal-spatial background ionosphere variation.  $\widehat{\Phi}_0(\mathbf{P})$  only relates to the constant components of *TECs* during two acquisitions of the GEO InSAR pair, while  $\widehat{\Phi}_1(\mathbf{P})$  and  $\widehat{\Phi}_2(\mathbf{P})$  highly depend on the temporal derivatives of *TECs* and the integration time. When  $k_{2,M}$  and  $k_{2,S}$  are small, as the impacts of the  $k_{1,M}$  and  $k_{1,S}$  are symmetrical within the whole aperture, the phase error after integration is dominated by  $\widehat{\Phi}_0(\mathbf{P})$ . In contrast, when  $k_{2,M}$  and  $k_{2,S}$  are large, only Equation (14) could be utilized to describe the interferometric phase screen error brought by the temporal-spatial background ionosphere variation accurately.

 $\Delta \phi_{ion}$  (**P**) will disturb the correct interferometric phase and introduce the deformation detection error  $\Delta d_{ion}$ , which is given as:

$$\Delta d_{ion} \left( \mathbf{P} \right) = -\frac{c}{4\pi f_0} \cdot \Delta \phi_{ion} \left( \mathbf{P} \right).$$
(16)

#### 4. Simulation and Discussion

# 4.1. Experiment Data and the Preprocessing

To verify the theoretical analysis, the USTEC data released by the NOAA is utilized. USTEC data has a temporal sampling rate of 15 min and the spatial sampling interval of 1°. The constant component, the temporal first-order derivative component and the temporal second-order derivative component of *TEC*, with respect to the azimuth time, are obtained by a second or higher order polynomial fitting of the multi-temporal sampling *TEC* data.

Shown in Figure 3a, the inclined GEO SAR orbit with a "figure-8" nadir-point trajectory is utilized. Its orbit parameters are given in Table 1. Our scene of interest in the simulation is an area of more than 1000 km  $\times$  1000 km, which has a center at Isla Guadalupe Island (29.02°N, 118.27°W) in the Pacific Ocean. As shown in the sketch map in Figure 3b, the GEO SAR signal obliquely traversed in the space to illuminate target A. Thus, the calculated IPPs of target A should be located at point B. Meanwhile, the *VTEC* should be transformed into *STEC* as well. Moreover, as discussed in Section 2, since the *TEC* variation caused by the spatial IPPs variation is negligible, we ignore the impacts of *TEC* variation caused by the spatial IPPs variation in the following data processing.



**Figure 3.** (a) GEO SAR "figure-8" nadir-point trajectory; (b) Sketch map of the oblique path of the GEO SAR signal in space.

Items	Value	Items	Value
Semi-major axis (km)	42,164.170	Wavelength (m)	0.24
Inclination (degrees)	53	Eccentricity	0.07
Argument of Perigee (degrees)	270	Right Ascension of Ascending Node (degrees)	210

Table 1. GEO SAR orbit parameters.

The GEO SAR system has a signal bandwidth of 80 MHz and an integration time of 250 s. The SAR data are obtained at orbit positions near apogee. Considering the short revisit time of the GEO SAR system, we select the temporal baseline of 1 day in the following simulation. The interferometric baseline is about 1.3 km. The *TEC* data used in the InSAR pair were acquired at UTC 22:00–23:00, 7 October 2013 and UTC 22:00–23:00, 8 October 2013, respectively. The acquisition time interval corresponds to the local time interval from noon to the sunset, which makes the utilized *TEC* data characterized by larger absolute *TEC* values and obvious *TEC* variation, and helps to study the impacts of temporal-spatial background ionosphere variation on GEO D-InSAR. After the IPPs calculations and the transformation from the *VTEC* to the *STEC*, based on the selected orbit, the *TEC* data during two acquisitions of the InSAR pair is shown in Figure 4 (0.1° spatial interval after triangle-based linear interpolation processing).



**Figure 4.** *TEC* data during two acquisitions of the InSAR pair. The first track: (**a**) *TEC*<sub>0</sub>; (**b**)  $k_1$ ; (**c**)  $k_2$ ; the second track: (**d**) *TEC*<sub>0</sub>; (**e**)  $k_1$ ; (**f**)  $k_2$ .

#### 4.2. Impacts of the Relative Image Shift and the Decorrelation of the InSAR Pair

Based on the *TEC* data (as shown in Figure 4) and the GEO SAR configuration, the relative image shift and the decorrelation of the InSAR pair are obtained and shown in Figure 5. It can be concluded that the largest range relative shift is less than 4 m, which gives rise to a loss of correlation coefficient of no more than 0.13. As the difference of the constant components of the *TECs* during the InSAR pair acquisitions are only about 10 TECU, the corresponding range relative shift is limited. The obvious range relative shift will only be generated in the case of a large difference in the constant components of the *TECs* during the InSAR pair acquisitions. For instance, in a condition of a 50 TECU difference for the constant components of the *TECs* during the InSAR pair acquisitions, the generated range relative shift is larger than 13 m and a total decorrelation will occur. Nevertheless, it is generally uncharacteristic that there will be a large bias between the *TEC* data during the two acquisitions of the InSAR pair in natural conditions.



**Figure 5.** Impacts of the relative image shift and the decorrelation of the InSAR pair. (**a**) Range relative shift; (**b**) Azimuth relative shift; (**c**) Correlation coefficient only under the impacts of the range relative shift; (**d**) Correlation coefficient only under the impacts of the azimuth relative shift.

In contrast, the azimuth relative shift and the corresponding decorrelation are obvious under the selected temporal first-order derivatives of *TECs* during two acquisitions of the GEO InSAR pair and the integration time. There are many parts with the azimuth relative shifts of more than 20 m in Figure 5b, which result in a total decorrelation in the interferogram. Only less than half of the interferogram keeps the correlation coefficients of more than 0.6.

Based on Equation (12) and the parameters in Table 1, the boundary parameters of the background ionosphere, when the impacts of the relative image shift and the decorrelation of the GEO InSAR pair can be ignored at apogee, is given in Table 2. It could be concluded that the impacts are negligible when the resolution of the SAR image is coarse and the variation of the parameters of the *TECs* between the two acquisitions of the GEO InSAR pair is limited. Otherwise, a high accurate co-registration is needed in the processing.

Resolution (m)	$ \Delta TEC_0 $ (Total Electron Content Unit (TECU))	$ \Delta k_1 $ (TECU/s)
100	<38.8	$< 9.65 \times 10^{-4}$
20	<7.8	$< 1.93 \times 10^{-4}$
5	<1.9	$< 4.83 \times 10^{-5}$

**Table 2.** Boundary parameters of the background ionosphere when the impacts of the relative image shift and the decorrelation of the GEO InSAR pair could be ignored under different resolutions (apogee).

#### 4.3. Interferometric Phase Screen Error

Based on Equation (14), the relationships between the integration time and the interferometric phase under the different cases of the constant components, and the temporal first-order and the second-order derivatives of TEC with respect to the azimuth time, are shown in Figure 6. Firstly, under a specified *TEC* condition, the interferometric phase brought by the impacts of the temporal-spatial background ionosphere variation on GEO SAR increases when the integration time becomes longer. In additional, the tendencies of variations of the blue line and the red line with respect to the integration time in the figure are the same. Only the initial interferometric phase is different. Thus, the difference in the constant components of the *TECs* only gives rise to a fixed bias in the interferometric phase. In contrast, the variations of  $\Delta k_1$  and  $\Delta k_2$  have obvious impacts on the interferometric phase. Based on the changing tendencies of the blue line, the light blue line and the green line, it could be concluded that the interferometric phase decreases with the increase of  $\Delta k_1$  under a specified integration time and  $\Delta k_2$ . According to Equation (15), when  $\Delta k_1$  is small, the product between  $\Delta k_2$  and the integration time approximately dominates the phase error. Conversely, a larger  $\Delta k_1$  will decrease the effect of the product between  $\Delta k_2$  and the integration time on the phase error. As a result, the interferometric phase becomes insensitive to the variation of the integration time for a larger  $\Delta k_1$ . As for  $\Delta k_2$ , when  $\Delta k_2$  is increasing, the interferometric phase becomes larger under a specified integration time and  $\Delta k_1$  (the blue line, the yellow line and the purplish red line). In this case, the interferometric phase becomes more sensitive to the variation of the integration time.



**Figure 6.** Relationships between the integration time and the interferometric phase under the different cases of the constant components, and the temporal first-order and the temporal second-order derivatives of *TEC* with respect to the azimuth time (*TEC* parameters in the first track: *TEC*<sub>0</sub> = 10 TECU,  $k_1 = 1 \times 10^{-5}$  TECU/s,  $k_2 = 1 \times 10^{-5}$  TECU/s<sup>2</sup>,  $\Delta TEC_0$  is the differences of the constant components of *TECs* during two acquisitions of the GEO InSAR pair,  $\Delta k_1$  and  $\Delta k_2$  are the differences of the GEO InSAR pair, respectively).

Utilizing the USTEC data in Figure 4 and Equation (14), the interferometric phase screen errors generated by  $\Phi_1(\mathbf{P})$  and  $\Phi_2(\mathbf{P})$  are shown in Figure 7a. The spatial distributed fringe frequency is brought by the interferometric phase screen ranges from less than one circle to multiple circles per one million square kilometers. Thus, the interferometric phase screen errors could give rise to deformation errors of up to several meters in the areas with dense fringes. We select the interferometric phase screen in the area marked by the white rectangle of Figure 7a, which has a relatively lower fringe frequency for the detailed deformation error analysis. The corresponding deformation retrieval error is given in Figure 7b. Even in the area with a lower fringe frequency, the largest deformation retrieval error in the scene is higher than 0.2 m and the mean square root error of the deformation is 0.13 m, which cannot satisfy the requirements for the deformation retrieval accuracy in any engineering applications. Therefore, some compensation algorithms based on Persistent Scatterer technology (PS) [4,33–37] or some similar methods based on the high quality coherent points are really needed to eliminate the serious impacts of the interferometric phase screen errors brought by the temporal-spatial background ionosphere variation in GEO D-InSAR processing in the future.



**Figure 7.** (a) Interferometric phase screen error generated by  $\Phi_1(\mathbf{P})$  and  $\Phi_2(\mathbf{P})$ ; (b) Deformation retrieval error in the white rectangle area.

# 5. Conclusions

Based on the temporal-spatial background ionosphere variation model, this paper focused on the analysis of impacts of the temporal-spatial background ionosphere variation on GEO D-InSAR processing, including the relative image shift and the decorrelation of the GEO InSAR pair, and the interferometric phase screen errors. It addresses the issue that the conventional frozen background ionosphere model used in LEO SAR is no longer valid for GEO SAR because of its long integration time and extensive imaging scene. The quantitative analysis is conducted by simulations based on the USTEC data. We draw two meaningful conclusions from our research.

Firstly, the differences of the constant *TEC* components and the temporal first-order derivatives of *TEC* with respect to the azimuth time during two acquisitions of the GEO InSAR pair, signal bandwidth and the integration time will give rise to an obvious relative image shift and the decorrelation of the GEO InSAR pair. Under the condition of GEO SAR parameters and the selected *TEC* data in our paper, a serious relative image shift in azimuth (dozens of meters) and the decorrelation of the GEO InSAR pair (almost total decorrelation) occur. When the parameters of the *TEC* variation during the acquisition of the GEO InSAR pair are small, the impacts of the relative image shift and the decorrelation of the GEO SAR interferometry pair could be ignored. For instance, the requirement of the *TEC* parameters are  $|\Delta TEC_0| < 7.8$  TECU and  $|\Delta k_1| < 1.93 \times 10^{-4}$  TECU/s under the moderate resolution of 20 m.

Secondly, the temporal-spatial part of the interferometric phase screen error  $\Phi_1(\mathbf{P})$  and  $\Phi_2(\mathbf{P})$  highly depend on the temporal derivatives of *TECs* and the integration time. During two acquisitions of the GEO InSAR pair, a small variation of the temporal first-order derivatives of *TEC* with respect to

the azimuth time and a larger variation of the temporal second-order derivatives of *TEC* with respect to the azimuth time, will give rise to the obvious temporal-spatial part of the interferometric phase screen error. Under the selected *TEC* parameters, the temporal-spatial part of the interferometric phase screen error will cause a deformation retrieval error of more than 0.2 m, even in the area with relatively small impacts of the ionosphere. Some compensation algorithms based on PS technology could have great potential in suppressing the deformation retrieval error, which will be our study focus in the future.

Acknowledgments: This work is supported by National Natural Science Foundation of China (Grant No. 61225005, 61471038, 61501032, 61427802, 61120106004), Beijing Natural Science Foundation (Grant No. 4162052) and Chang Jiang Scholars Program (T2012122) and 111 project of China under Grant B14010.

Author Contributions: Y.L. and C.H. conceived and designed the methods; X.D. and Y.L. performed the simulation; C.C. and T.L. analyzed the data; X.D. and Y.L. wrote the paper.

Conflicts of Interest: The authors declare no conflict of interest.

# Appendix A

Simplifying Equation (13), it is expressed as (take  $k_2 > 0$  as an example):

$$\phi_{ion}\left(\mathbf{P}\right) = \arg\left\{ \exp\left[-j\frac{4\pi K}{cf_{0}}\left(TEC_{0}(\mathbf{P}) - \frac{k_{1}(\mathbf{P})^{2}}{4k_{2}(\mathbf{P})}\right)\right] \cdot \int_{-\frac{T_{a}\sqrt{k_{2}(\mathbf{P})}}{2} + \frac{k_{1}(\mathbf{P})}{2\sqrt{k_{2}(\mathbf{P})}}} \exp\left[-j\frac{4\pi K}{cf_{0}} \cdot u^{2}\right] du \right\}$$
(A1)

Using  $u = \sqrt{k_2(\mathbf{P})} t_a + \frac{k_1(\mathbf{P})}{2\sqrt{k_2(\mathbf{P})}}$  and  $y = \sqrt{\frac{8K}{cf_0}} u$  to realize the variable substitution in Equation (10) and expanding it by Euler's formula, we have:

$$\phi_{ion}\left(\mathbf{P}\right) = \arg\left\{\exp\left[-j\frac{4\pi K}{cf_0}\left(TEC_0(\mathbf{P}) - \frac{k_1(\mathbf{P})^2}{4k_2(\mathbf{P})}\right)\right] \cdot \sqrt{\frac{cf_0}{8K}} \int_{\varepsilon_2}^{\varepsilon_1} \left[\cos\left(\frac{\pi}{2}y^2\right) - j\sin\left(\frac{\pi}{2}y^2\right)\right] dy\right\}$$
(A2)

where

$$\begin{cases} \varepsilon_1 = \sqrt{\frac{8K}{cf_0}} \left( \frac{T_a \sqrt{k_2(\mathbf{P})}}{2} + \frac{k_1(\mathbf{P})}{2\sqrt{k_2(\mathbf{P})}} \right) \\ \varepsilon_2 = \sqrt{\frac{8K}{cf_0}} \left( -\frac{T_a \sqrt{k_2(\mathbf{P})}}{2} + \frac{k_1(\mathbf{P})}{2\sqrt{k_2(\mathbf{P})}} \right) \end{cases}$$
(A3)

Since there are two Fresnel integrals inside Equation (A2), we apply the series expansions to them. Then, we obtain:

$$\phi_{ion}\left(\mathbf{P}\right) = \frac{4\pi K}{cf_0} \left( TEC_0(\mathbf{P}) - \frac{k_1(\mathbf{P})^2}{4k_2(\mathbf{P})} + \Phi_h\left(k_1(\mathbf{P}), k_2(\mathbf{P})\right) \right)$$
(A4)

where  $\Phi_h(k_1(\mathbf{P}), k_2(\mathbf{P}))$  is expressed as

$$\Phi_{h}\left(k_{1}(\mathbf{P}),k_{2}(\mathbf{P})\right) = \left( \left[ \frac{\left(\sum_{k=0}^{\infty} \frac{(-1)^{k}}{(2k)!} \left(\frac{\pi}{2}\right)^{2k} \frac{z^{4k+1}}{4k+1}\right)\right|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{a}\sqrt{k_{2}(\mathbf{P})}}{2} + \frac{k_{1}(\mathbf{P})}{2\sqrt{k_{2}(\mathbf{P})}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{(-1)^{k}}{(2k)!} \left(\frac{\pi}{2}\right)^{2k} \frac{z^{4k+1}}{4k+1}\right)\right|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{a}\sqrt{k_{2}(\mathbf{P})}}{2} + \frac{k_{1}(\mathbf{P})}{2\sqrt{k_{2}(\mathbf{P})}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{(-1)^{k}}{(2k+1)!} \left(\frac{\pi}{2}\right)^{2k+1} \frac{z^{4k+3}}{4k+3}\right)\right|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{a}\sqrt{k_{2}(\mathbf{P})}}{2} + \frac{k_{1}(\mathbf{P})}{2\sqrt{k_{2}(\mathbf{P})}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{(-1)^{k}}{(2k+1)!} \left(\frac{\pi}{2}\right)^{2k+1} \frac{z^{4k+3}}{2\sqrt{k_{2}(\mathbf{P})}}\right)|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{a}\sqrt{k_{2}(\mathbf{P})}}{2} + \frac{K_{1}(\mathbf{P})}{2\sqrt{k_{2}(\mathbf{P})}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{(-1)^{k}}{(2k+1)!} \left(\frac{\pi}{2}\right)^{2k+1} \frac{z^{4k+3}}{2\sqrt{k_{2}(\mathbf{P})}}\right)|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{a}\sqrt{k_{2}(\mathbf{P})}}{2} + \frac{K_{1}(\mathbf{P})}{2\sqrt{k_{2}(\mathbf{P})}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)^{-\left(\sum_{k=0}^{\infty} \frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{k}}(\mathbf{P})}{2\sqrt{k_{k}}}\right)|_{z=\sqrt{\frac{8K}{cf_{0}}} \left(\frac{T_{k}}(\mathbf{P$$

where *z* and *k* are the variables related to the series.

Especially when (A3) is small, we have approximately:

$$\phi_{ion}\left(\mathbf{P}\right) = \frac{4\pi K}{cf_0} \left( TEC_0(\mathbf{P}) - \frac{k_1(\mathbf{P})^2}{4k_2(\mathbf{P})} + \frac{cf_0}{4\pi K} \arctan\left[\frac{4\pi K}{3cf_0} \left(\frac{3k_1(\mathbf{P})^2}{4k_2(\mathbf{P})} + \frac{T_a^2k_2(\mathbf{P})}{4}\right)\right] \right)$$
(A6)

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