Regional Water Balance Based on Remotely Sensed Evapotranspiration and Irrigation: An Assessment of the Haihe Plain, China

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Abstract: Optimal planning and management of the limited water resources for maximum productivity in agriculture requires quantifying the irrigation applied at a regional scale. However, most efforts involving remote sensing applications in assessing large-scale irrigation applied (IA) have focused on supplying spatial variables for crop models or studying evapotranspiration (ET) inversions, rather than directly building a remote sensing data-based model to estimate IA. In this study, based on remote sensing data, an IA estimation model together with an ET calculation model (ETWatch) is set up to simulate the spatial distribution of IA in the Haihe Plain of northern China. We have verified this as an effective approach for the simulation of regional IA, being more reflective of regional characteristics and of higher resolution compared to single site-specific results. The results show that annual ET varies from 527 mm to 679 mm and IA varies from 166 mm to
We confirm that the region along the Taihang Mountain in Hebei Plain has serious water resource sustainability problems, even while receiving water from the South-North Water Transfer (SNWT) project. This is due to the region’s intensive agricultural production and declining groundwater tables. Water-saving technologies, including more timely and accurate geo-specific IA assessments, may help reduce this threat.

**Keywords:** ETWatch; remote sensing technique; ET; irrigation requirement

1. Introduction

Competition for water resources in agriculture is increasing with water scarcity and becoming more serious in the 21st century. The agricultural sector is one of the biggest consumers of water resources, accounting for more than 70% of the world’s fresh water use from rivers and groundwater [1]. This figure is even greater in Asia and the Pacific region, where it accounts for as much as 90% [1]. For optimal planning and management of the limited water resources for maximum crop productivity in agriculture, especially in arid and semi-arid regions of the world, techniques are required to quantify the irrigation applied (IA) at regional scales.

Haihe Basin, located in North China, has especially severe water shortage problems. However, this basin undertakes the heavy task of grain production to support a large proportion of the Chinese population of this area. To sustain high-yield grain production, irrigation demand has been satisfied by excessive groundwater pumping, which has caused sharply declining groundwater tables. To alleviate water shortage problems in North China, the South-North Water Transfer (SNWT) project is under construction. The transferred water will be mainly used to meet high priority domestic and industrial demands. However, a large amount of surface water will then be released for agricultural use. The problem of how to redistribute water among different sectors and how much will be available for agricultural use will emerge with the completion of the project. Thus, the ability to accurately assess the spatial distribution of irrigation demand is necessary.

The science of assessing IA on larger regional scales has undergone development from a focus on statistical data to the present combination of techniques of crop growth simulation, a geographical information system (GIS) and remote sensing [2]. Early methods applied to assess irrigation requirements include water metering, questionnaires, water use coefficients, water rights, pumping hours or electricity consumption [2]. For example, Boken et al. [3] used irrigation requirements (mm) for cotton, peanut and maize measured at a limited number of sample sites combined with GIS and geospatial techniques to estimate regional water use at a county level in Georgia. These methods are direct and can be quite accurate, but are laborious and lacking in predictive capacity. Model-based estimation for IA has been reported in many documents, mainly including the Food and Agriculture Organization (FAO) Penman–Monteith method combining soil water balance models and crop growth simulation models.

The FAO Penman-Monteith method [4] has been widely used to estimate IA. The method is the internationally recommended standard for evapotranspiration (ET) estimation in both humid and arid
environments [5]. Ozdogan et al. [6] used the CROPWAT system (S) (a computer program for irrigation planning and management) to estimate water use in south-eastern Turkey, which computes Penman-Monteith-based reference (potential) evapotranspiration (ETp) and then adjusts this generalized variable for each crop using seasonal growth curves, or so-called Kc parameters. Leenhardt et al. [7] applied a simulation platform called ADEAUMIS (an irrigation water demand assessment tool), which mainly includes the FAO Penman-Monteith method and a spatial database to estimate IA. However, FAO guidelines are often used to estimate crop evapotranspiration from observed or standard values of crop development [7].

Crop models that dynamically simulate plant growth and the water demand of one or several crops can provide quantitative contributions to the environmental impact assessment and be very useful for water management [7]. One of the limitations of current crop simulation models is that they are site specific. By combining with spatial data, crop models have been used by many scientists to estimate IA at regional or global scales. Ines et al. [8] applied crop simulation models CERES-rice (Crop Environment Resource Synthesis for rice), CERES-maize and CROPGRO-peanut (a process-oriented model for crop growth of peanut) in DSSAT (Decision Support System for Agrotechnology Transfer) version 3.0 coupled with GIS to the estimate irrigation needs for rice, maize and peanut, and then water productivity in the Laoag River Basin in Ilocos Norte, Philippines. The cropping system model, CROPGRO-Cotton, combined with kriging interpolation, was used to evaluate the spatial distribution of monthly irrigation water use for cotton in the Coastal Plain region of Georgia [9]. Wriedt et al. [2] estimated the spatial distribution of irrigation applied at a 10 × 10-km grid resolution in Europe with the crop growth model, EPIC (Erosion Productivity Impact Calculator), combining available regional statistics on crop distribution, crop-specific irrigated area and spatial data sources on soils, land use and climate. The global scale water use of wheat was estimated by combining EPIC with GIS (geographic information system) [10].

The shortage of geolocated data on agricultural practices was considered to be a limiting factor for the operational use of crop models over large scales [11]. Fortunately, public domain Internet satellite data and scientific development make remote sensing an attractive option for assessing irrigation performance from individual fields to scheme or river basin scale [1]. One of the applications of remote sensing in assessing irrigation performance is in providing information on some key variables of crop production and, especially, variations in these over time and space. For example, Ozdogan et al. [6] successfully estimated irrigated area with the aid of remote sensing to compute water use in south-eastern Turkey, in the period 1993 to 2002. Ines et al. [12] estimated the distributed data, e.g., sowing dates, irrigation practices, soil properties, depth to groundwater and water quality, required as inputs to regional modelling by minimizing the residuals between the distributions of field-scale ET simulated by regional application of the SWAP (Soil, Water, Atmosphere and Plant) model and by the surface energy balance algorithm for land (SEBAL) using pairs of Landsat-7 ETM+ images. Another application of remote sensing in assessing irrigation performance was to estimate ET. The satellite-based model SEBAL has been applied in more than 30 countries worldwide and is now an operational instrument for targeting, monitoring and evaluating irrigation and drainage systems [13]. ETWatch is another ET monitoring system using remote sensing, which integrates SEBAL and SEBS and is commonly used in North China [14].
As stated above, most documents on RS applications for assessing irrigation water applied at larger scales focused on providing spatial variables for crop models or studies on ET inversion. Little literature is available on directly using RS to determine irrigation water requirements. The objectives of the present study are to: (1) set up a model based on RS-based ETWatch and a soil water balance model to evaluate irrigation water requirements; (2) analyse the spatial distribution of IA to evaluate regional water sustainability.

2. Study Area

The Haihe Plain (112–118°E, 35–42°N) is selected for this study. Located in the northern part of China, it is one of the three parts of the North China Plain, with an approximate area of 131,000 km² [15]. There are five administrative regions in this catchment, including Beijing, Tianjin, most of Hebei Province and the northern parts of Shangdong and Henan Provinces (Figure 1). This area has a moderate temperate, semi-arid and sub-humid climate. Long-term annual mean air temperature in the catchment ranges from 1.5 °C–14 °C and annual mean precipitation is 535 mm (84% occurring between June and September). Haihe Catchment is one of the most important grain production areas in China, accounting for about 10% of the nation’s total. The catchment contains 11% of the arable land to support 10% of China’s population. However, the Catchment only has 1.3% of the national water resources. Long-term average annual water resource volume is 37.2 billion m³ (37.2 km³), and water resources per capita are only 305 m³, 14% of the national average level and 3.7% of the world average [16]. Water resources in Haihe Catchment are the lowest among the seven biggest catchments in China.

Figure 1. Location of study area, the Haihe plain.
There is 8.0 million hectares of arable land in the Haihe Plain. Agriculture in this area is characterized by a diverse cropping pattern of winter wheat, corn, millet, sorghum, rice, soybean, cotton, potato, sweet potato and a variety of horticultural crops. Winter wheat-summer corn is the most important crop rotation system. As winter wheat grows in the dry season (sown in early October and harvested in early June in the next year), it is the biggest irrigation consumer. Some crops sown in the spring, such as spring corn, millet and cotton, are the second greatest irrigation users. These differ from the summer crops, which grow during the monsoon and only require supplementary irrigation to fill the gap when there is a shortage of precipitation. Planted areas of horticulture crops, such as vegetables and fruit trees, are increasing; becoming a larger potential irrigation user, because they commonly need more water per hectare than annual crops.

In Haihe Plain, 70% of irrigation water is sourced by groundwater pumping. Intensive use of groundwater has resulted in sharp declines in groundwater levels. The biggest groundwater drawdowns in the world are located in this area. This phenomenon has triggered severe environmental issues, such as land subsidence, aquifer compression and sea water intrusion [17]).

3. Model Description

3.1. ETWatch

ETWatch is an ET monitoring system using remote sensing. A brief description of ETWatch and an application for monitoring regional evapotranspiration were reported by Wu et al. [14]. ETWatch integrates SEBAL and SEBS to estimate surface fluxes under clear-sky conditions together with the Penman–Monteith approach to calculate daily ET, based on a surface resistance model, meteorological data and surface parameters from remote sensing. A brief description of ETWatch is given in the following text, and the main procedure is demonstrated in Figure 2.

The SEBAL model uses the following formula to convert global radiation into net radiation on a daily scale (Bastiaanssen et al. 2005 [13]),

\[
R_{n24} = \tau ((1 - \rho) R_{e24} - 110)
\]  

(1)

where \( \tau = R_{24}/R_{e24} \), which accounts for the fraction of daily incoming solar radiation (\( R_{24} \) (MJ·m\(^{-2}\)·day\(^{-1}\))) to the extra-terrestrial solar radiation (\( R_{e24} \) (MJ·m\(^{-2}\)·day\(^{-1}\))), and \( \tau_0 \) is the surface albedo.

The parameter, \( \rho \), measures the daily total atmospheric transmittance of extra-terrestrial solar radiation. The incoming solar radiation can be directly measured by pyranometers, but the availability of such data is limited, particularly in temporal and spatial contexts. ETWatch uses sunshine hours to compute \( \rho \) as:

\[
\rho_{sw} = a_s + b_s \frac{n}{N}
\]  

(2)

where \( n \) is the sunshine hours for the day, \( N \) is the maximum hours of sunshine on a typical cloud-free day, while \( a_s \) and \( b_s \) are empirical coefficients determined by regression analysis for daily data from weather stations in Haihe Catchment [18]. The parameters are determined separately for December–February (winter dry season) and March–November (wet season).
Figure 2. Operational flow-chart of ETWatch for the monitoring and estimation of evapotranspiration (ET) (adapted after Wu et al. [14,19]). SEBAL, surface energy balance algorithm for land.

Momentum roughness ($Z_{om}$) is an important parameter for calculating energy balance. ET models were developed based generally on smooth surfaces that only take into account the influence of vegetation height on roughness or use empirical values linked to land use classifications. However, this approach performs poorly in hilly land. ETWatch calculates the roughness length, based on land cover by vegetation or non-vegetation and hypsography. The roughness length of the land covered by vegetation, $Z_{om}^v$, is calculated as a function of the normalized difference vegetation index (NDVI) as follows,

$$Z_{om}^v = 0.001 + \left( \frac{NDVI}{NDVI_{max} - NDVI_{min}} \right)^{2.5}$$

(3)

where $NDVI = (\rho_{NIR} - \rho_{RED})/(\rho_{NIR} + \rho_{RED})$ and $\rho_{NIR}$ and $\rho_{RED}$ are the near-infrared and infrared band reflectivities. $NDVI_{min}$ and $NDVI_{max}$ are the minimum and maximum vegetation indices in the region. The effect of topography and slope on surface roughness is then estimated as:

$$Z_{om}^T = Z_{om}^v \left( 1 + \frac{slope - a}{b} \right)$$

(4)

where $Z_{om}^T$ is the surface roughness length for topography, $a$ and $b$ are empirical coefficients and $slope$ stands for topographic slope, which can be obtained from a digital terrain map. The surface roughness of the non-vegetation is computed from the surface geometric roughness derived from radar images:
\[
\log(Z_{om}') = -1.221 + 0.0906(\sigma_o)
\]

where \(Z_{om}'\) is the surface roughness of the non-vegetation fraction of the basin and \(\sigma_o\) is the backscatter coefficient acquired from the Radarsat satellite image. The total basin roughness is finally computed as:

\[
Z_{om} = w_1Z_{om}' + w_2Z_{om}^T + w_3Z_{om}'
\]

where \(w_1, w_2\) and \(w_3\) are the weighted factors for vegetation, topography and bare-land on surface roughness. The values of \(w_1, w_2\) and \(w_3\) are 1.0, 1.0 and 0.25, respectively for the Hai Basin, which were determined by using the empirical fitted ways [18].

Resistance for the days with clouds (\(RS_{unc}\)) is computed as:

\[
RS_{unc} = \frac{LAI_{clr}RS_{clr}}{LAI_{unc}m(T_{min})m(VPD)}
\]

where \(RS_{clr}\) is the resistance for the cloudless day, \(LAI_{unc}\) is the leaf area index for the day with clouds, \(LAI_{clr}\) is the leaf area index for cloud-free day and \(m(T_{min})\) and \(m(VPD)\) are the functions of minimum temperatures and vapour pressures, respectively, which account for the effect of extreme temperature and moisture on plant stomata.

### 3.2. Irrigation Applied (IA) Estimation Model

The sequence of calculations used to estimate IA is given in Figure 3. The model starts from the date when a crop is sown. Several practical assumptions are made. The initial soil water content is set as 60% of maximum plant available water (SW\(_{pa}\)), which is generally used by farmers for crops to be sown. The depth of soil profile is set as 80 cm. When soil water content is less than 40% SW\(_{pa}\), irrigation operation is to be triggered. The irrigation applied is equal to the difference between field capacity (SW\(_{fc}\)) and the soil water content on the previous day (SW\(_{i-1}\)). When there is precipitation on the day and the soil water content is greater than field capacity, the amount of water in excess of field capacity is assumed to go to deep drainage, and the current soil water content is set equal to field capacity. Since farmland in the Haihe Plain is very flat and soils are deep, there is a large water holding capacity, especially in relation to precipitation shortages; runoff (RO\(_i\)) is minimal and may be ignored [20].

Water content at field capacity and permanent wilting point (SW\(_{fc}\) and SW\(_{wp}\)) are calculated by the SPAW model [21]) using the basic soil texture parameters, including clay content (%), sand content (%) and organic matter content (%). These soil parameters come from 76 representative sites in Haihe Catchment. Precipitation data are collected from 46 meteorological stations. All these data are finally interpolated into a 1-km\(^2\) grid and prepared for use in the above-mentioned spatial calculation procedures.

### 4. Data Collection

#### 4.1. Remote Sensing Data

The MODIS data used in this study are the MOD021KM, MOD02QKM, MOD02HKM and MOD03 products provided by the NASA Goddard Space Flight Center (GSFC) Level 1 and Atmosphere Archive and Distribution System (LAADS) [22]. The MOD02QKM and MOD02HKM products, including the 250-m and 500-m resolution bands, are aggregated to 1-km resolution TOA
(top of atmosphere) radiances and reflectance. The MOD03 products are the geolocation field data calculated for each 1-km MODIS Instantaneous Field of View (IFOV) for all daytime orbits. The geolocation fields include geodetic latitude, longitude, surface height above geoid, solar zenith and azimuth angles, satellite zenith and azimuth angles and a land/sea mask for each 1-km sample. The solar zenith and azimuth angles, satellite zenith and azimuth angles are used to estimate net shortwave radiation in this study. The MODIS 2.1-μm band was used for the detection of dark targets, estimating their reflectance in the blue and red channels and using them for remote sensing of aerosol [23]. Using the aerosol optical thickness as input, a lookup table was established based on the 6S model to carry out the atmospheric correction [24]. A practical split-window approach was used to retrieve LST from MODIS data [25].

**Figure 3.** Flow chart of irrigation estimation from ET, where: $SW_{pa}$, the maximum extractable soil water content (mm); $SW_{fc}$, water content at field capacity ($m^3/m^3$); $SW_{wp}$, water content at permanent wilting point ($m^3/m^3$); DP, soil depth (mm), which is set to 800 mm; $SW_0$, initial water content (mm), which is set to 60% $SW_{pa}$; $SW_i$ and $SW_{i-1}$, water content at day $i$ and $i-1$ respectively (mm); ET, evapotranspiration at day $i$ (mm); $P_i$, precipitation at day $i-1$ (mm); $RO_i$, runoff at day $i$ (mm); $DR_i$, drainage at day $i$ (mm); $IA_i$, irrigation amount at day $i$ (mm).

\[
SW_{pa} = (SW_{fc} - SW_{wp}) \times DP \\
SW_0 = SW_{pa} \times 0.6, DR_0 = 0, i = 1 \\
\text{Input ET}_i, P_i \\
SW_i = SW_{i-1} - ET_i + P_i - RO_i \\
i = i + 1 \\
\text{Output } IA_i, DR_i \\
\text{If } SW_i > SW_{fc} \times DP \text{ then } \text{Y} \\
\text{Dr}_i = SW_i \times SW_{fc} \times DP \\
SW_i = SW_{fc} \times DP \\
\text{If } SW_i < SW_{pa} \times 0.4 \text{ then } \text{Y} \\
IA_i = SW_{fc} \times DP - SW_i \\
SW_i = SW_{fc} \times DP \\
DR_i = 0 \\
IA_i = 0 \\
\text{Else } \text{N} \\
DR_i = 0 \]

Output $IA_i$, $DR_i$.
Landsat TM images with Bands 2, 3, 4 and 5 at different periods from the years 2000–2001 at a spatial scale of 1:100,000 were used to interpret land use data over the Haihe Plain. The work was done by another group of “land cover dynamic monitoring” of the Key Innovation Project (KZCX1-YW-08-03-04) of the Chinese Academy of Sciences.

4.2. Meteorological Data

Daily meteorological data at 83 meteorological stations, including maximum and minimum air temperature, sunshine duration, relative humidity, wind speed and air pressure (2002–2007), are collected from the National Meteorological Bureau. The variables are interpolated into daily maps at 1-km resolution by the inverse distance squared method.

4.3. Lysimeter Data

Monthly ET and irrigation amount from large Lysimeter provided by Luangcheng Station (Chinese Ecosystem Research Network).

5. Results

5.1. Verification

Validation of ET estimated from ETwatch was described in the documents of [14] and [26]. According to these reports, the ET result exhibited less error than the results produced by the earlier method of constant-EF (evaporative factor). The mean absolute percent difference (MAPD) of gap-filling approach is reasonable (22.5%), compared to MAPD produced by constant-EF (46.3%). Alternatively, the gap-filling method used in ETWatch gives a 50% improvement in accuracy.

The grid (1 km²) of farmland where the large lysimeter was installed was selected to compare its ET and IR. The simulated ET and IA for 2002 to 2007 generally matched observed ET in most of months (Figure 4), but the simulated values in a number of months in the wet season were much larger than the observed values. The error came from the different resolutions of observed and simulated data. As noted, the simulated ET and IA by remote sensing represented 1-km² grid areas, while the observed ET and IA from the lysimeter was a site-specific point value. In the study region, a 1-km² area usually consists of different crops, roads and several villages.

To evaluate the model prediction of IA at the regional scale, water use by agricultural departments in 87 counties in the Hebei Plain (more than 80% of the Haihe Plain) from 2005 to 2006 (the data are from the Yearbook of Hebei Province Water Resources) was used to compare with the simulated result. Legates and McCabe [27] recommended that a complete assessment of model performance should include at least one “goodness-of-fit” or relative error measure (e.g., E1’ (modified coefficient of efficiency) or d1’ (modified index of agreement)) and at least one absolute error measure (e.g., root mean square error (RMSE) or mean absolute error (MAE), with additional supporting information (e.g., a comparison between the observed and simulated mean and standard deviations). E1’ or d1’ was recommended to replace the determination coefficient (R²) or Pearson’s correlation coefficient (P), because these are oversensitive to extreme values and are insensitive to additive and proportional differences between model predictions and observations. Therefore, in this paper, both correlation-related methods
(e.g., d1′) and absolute error measure (e.g., MAE) are used to evaluate the model (Table 1). The results show that the mean absolute differences between modelled and observed values in 2005, 2006 and 2007 are 17.9%, 1.36% and 17.8%, respectively, which are all lower than 20%. In 2006, the mean predicted value is very close to the observed value. Both MAE and RMSE in 2007 are higher than those in 2005 and 2006. Most of the values of the correlation related statistics, including P, R² and d1 for all years, are over 0.60, indicating that the modelled results were reasonable.

Figure 4. Comparison of simulated and observed monthly ET (mm) and irrigation applied (IA) (mm) from the years from 2002 to 2007.
Table 1. Means and standard deviations of observed and simulated yearly irrigation amount for county of Haihe Plain for the years of 2005, 2006 and 2007, and the statistics comparing the observed and simulated values.

<table>
<thead>
<tr>
<th>Year</th>
<th>Observed Mean, mm</th>
<th>Observed Standard deviation, mm</th>
<th>Modelled Mean, mm</th>
<th>Modelled Standard deviation, mm</th>
<th>Observed Mean absolute error, mm</th>
<th>Modelled Mean absolute error, mm</th>
<th>Observed Root mean square error, mm</th>
<th>Modelled Root mean square error, mm</th>
<th>Observed Pearson’s correlation</th>
<th>Modelled Pearson’s correlation</th>
<th>Observed Coefficient of determination</th>
<th>Modelled Coefficient of determination</th>
<th>Observed Modified index of agreement</th>
<th>Modelled Modified index of agreement</th>
</tr>
</thead>
<tbody>
<tr>
<td>2005</td>
<td>240</td>
<td>74</td>
<td>224</td>
<td>50</td>
<td>38.88</td>
<td>53</td>
<td>53</td>
<td>0.73</td>
<td>0.53</td>
<td>0.61</td>
<td>0.68</td>
<td>3.88</td>
<td>3.68</td>
<td></td>
</tr>
<tr>
<td>2006</td>
<td>295</td>
<td>91</td>
<td>299</td>
<td>72</td>
<td>39.75</td>
<td>55</td>
<td>55</td>
<td>0.80</td>
<td>0.64</td>
<td>0.65</td>
<td>0.65</td>
<td>52.05</td>
<td>52.25</td>
<td></td>
</tr>
<tr>
<td>2007</td>
<td>246</td>
<td>95</td>
<td>202</td>
<td>80</td>
<td>91</td>
<td>71</td>
<td>71</td>
<td>0.81</td>
<td>0.81</td>
<td>0.65</td>
<td>0.65</td>
<td>80</td>
<td>80</td>
<td></td>
</tr>
</tbody>
</table>

5.2. Yearly Spatial Distribution of ET and IA

The simulation of spatial distribution of ET in the Haihe Plain is illustrated in Figure 5. There is a general pattern for the modelled annual ET distributions. The regions along Taihang Mountain (western Hebei Province) and in Shandong Province have the highest ET, with a range of 650–1050 mm in most years, while the regions located in the low plain close to Bohai Sea have the lowest ET, with less than 500 mm. The north and middle regions indicate moderate ET, with a range of 500–700 mm.

Figure 5. Spatial distribution of yearly ET in Haihe Plain during 2002–2007.

The spatial pattern of IA (Figure 6) is similar to that of ET. The IA in the piedmont regions of Taihang Mountain and regions in Shandong Province are higher with the range of 200–600 mm; while the IA in the low plain and middle part of the south region is lower, with a range of 0–350 mm. The reason for the higher IA in the piedmont region and Shandong Province is that these regions have better water resources: relatively plentiful groundwater conditions, due to ample recharge from the mountain area for the piedmont region, and ample surface water for the regions in Shandong Province. With better water supplies and also fertile soil conditions, a winter wheat-summer maize planting...
system is used in these regions. With the intensive planting pattern and high water consuming crop (winter wheat), the ET and IA are higher in these regions. In contrast is the lower plain, where water conditions are worse for winter crops, due to a shortage of fresh water. However, even in a dry year, such as 2002, the IA is very low. The reason for the lower IA in the middle part of the southern region is its greater cotton cultivation areas and fewer winter crops. Cotton is a drought-resistant crop and grows through the rainy reason. Therefore, in the rainy or normal years, such as 2003, 2004, 2005 and 2007, IA is lower in the region. However, for very dry years, such as 2002 and 2006, some quantity of irrigation is still required.

**Figure 6.** Spatial distribution of yearly irrigation applied (IA) (mm) in Haihe Plain during 2002–2007.

![Spatial distribution of yearly irrigation applied (IA) (mm) in Haihe Plain during 2002–2007.](image)

**Figure 7.** Yearly precipitation (mm), ET (mm) and irrigation applied (mm) during 2002–2007.

![Yearly precipitation (mm), ET (mm) and irrigation applied (mm) during 2002–2007.](image)

Averaged across the whole region, yearly ET and IA vary from 527 mm to 679 mm and 166 mm to 289 mm, with mean values of 602 mm and 225 mm, respectively, from 2002 to 2007 (Figure 7).
Yearly variations in ET and IA are highly related to precipitation, which varies from 351 mm to 601 mm, with a mean value of 493 mm from 2002 to 2007. ET is positively correlated to precipitation and IA is negatively correlated to it. For example, in 2003, the precipitation is as high as 601 mm; accordingly, the ET is the highest (679 mm); however, the IA is lowest (225 mm) in that year.

5.3. Monthly Spatial Distribution of ET and IA

The monthly distributions of ET and IA are simulated for each year of the study (Figures 8 and 9). Figure 8 illustrates that the months of May, July and August have the largest ET. June has comparatively lower ET, because it is a period when the spring crop is harvested and the summer crop begins to be planted. All these months have higher temperature and precipitation. The highest ETs in May are located in the western region along the Taihang Mountain piedmont. Higher ET in July and August are also distributed in the middle part of the study area. Figure 9 illustrates the monthly spatial distribution of IA. The pattern of the spatial distribution of IA in months with less precipitation, such as April and May, is similar to that of ET. Contrarily, the spatial distribution of IA in months with more precipitation is different from that of ET, because the precipitation provided required water. Averaged monthly ET, IA and precipitation are indicated in Figure 10. The three months with the highest ETs are May, July and August; whereas June, July and August are the three months with the highest precipitation. The three months with the highest IA are April, May and September.

Figure 8. Spatial distribution of monthly irrigation applied in Haihe Plain during 2002–2007.
5.4. Reallocation of Water Resources after SNWT Project in Hebei Plain

The RS results for irrigation requirements are used to evaluate the future water resource redistribution in the Hebei Plain after the SNWT project. Since precipitation is insufficient to recharge...
groundwater in the North China Plain [28], SNWT has been under construction to stop groundwater table decline. Figure 11 indicates the planned amounts of water transfer by middle and east routes to different districts of Hebei Plain. It shows that except for the capital city Shijiazhuang, which is to receive a large relative share of transfer water, some other districts may be far from satisfied in meeting their irrigation requirements. Especially for Cangzhou and Hengshui, these two districts have the worst water conditions: less available surface water, rapid depletion of confined aquifers and brackish unconfined aquifer water qualities [29]. These two districts appear to require more irrigation water than officially indicated. Therefore, this study suggests more water could be distributed to these two regions to remedy their deficits.

**Figure 11.** The required irrigation amount (10^9 m^3 or km^3) estimated by remote sensing models, district agricultural water use (official data) and planned amount of water to be allocated from the South-North Water Transfer (SNWT) project in different districts of Hebei Plain.

6. Discussion

6.1. Comparison with Contemporary Studies

Both the ET and IA simulated results from our study are reasonable when compared to the results from contemporary studies. The average ET estimated by ETWatch for farmland of Haihe Plain during 2002–2007 is 602 mm, which is very similar to the result from Li [30], who demonstrated that the ET of farmland in Haihe Plain was 600–660 mm, by using a water balance method. They are also validated by the results from Yu [31], who concluded that ET in irrigation land is above 660 mm and that in rain-fed land, it is 400 mm. However, the latter result is low when compared to other site-specific experiments, most of which show ET under a winter-wheat and summer-maize rotation system is about 800–900 mm [20]. The differences were, at least partially, due to ETWatch representing land and water use on regional scales, which include roads, trees and villages.
The percentage differences between simulated and observed IA results are under 20% (Table 1). However, they show a lower trend, which is better explained by the difference between the simulated applied irrigation and actual applied irrigation by farmer. The irrigation applied, estimated here by modelling, is based on an optimum state of soil water. However, the statistical data (Table 1) is the gross irrigation, including the irrigation loss, including drainage infiltrating to the groundwater. In practice, farmers determine the amounts of irrigation water largely by experience and often apply more water than required by our modelling criteria, which only calls for refilling to field capacity. With the aid of crop growth modelling [32], an earlier result was that the irrigation applied to Hebei Plain (most parts of Haihe Plain) averaged 348 mm from 1986 to 2006, which is higher than the result in this study. This is inferred from 8.7 billion m$^3$ of agricultural water use in official data by the Water Resources Department of Hebei Province (WRDHP) in Yang et al. [32], that irrigation applications averaged 279 mm from 1986 to 2006. The irrigation applied of 225 mm (for Hebei Plain) in this study is lower than either the statistical data of 279 mm or the crop modelling result of 348 mm from Yang et al. [32]. However, the site result of 328 mm in Luancheng Station (experimentally observed result from the Lucheng Station average from 2002 to 2007) is closer to 348 mm of the crop modelling result. This indicates that the crop modelling is more site-specific, while the results of this study are more reflective of regional characteristics. Moreover, the present study has much higher spatial resolution (1 km$^2$) compared to former studies with crop growth models at resolutions averaging across county areas of about 1100 km$^2$.

6.2. Bias Analysis

Since the IA estimation is based on a simple water balance model, it still needs further improvement. Our simplifying assumption is that the irrigation triggering mechanism (when soil water content is less than 40% of field capacity) is the same throughout the study area. It does not take crop type, irrigation application practice, weather conditions, etc., into account. Some error is inevitable under this hypothesis. For example, the soil water deficiencies of different crops for triggering irrigation may not be the same, because different crops have different drought-tolerance potentials. Thus, errors in irrigation frequency will occur. Moreover, irrigation quotas should be different under different irrigation application practices, such as dripping, spraying and furrowing, other than simply to differences between soil water content when irrigation is triggered and field capacity. Though error will occur in irrigation frequency and quota, the total irrigation amount will keep balanced with ET, because the model is based on the soil water balance. However, the error in total irrigation estimates will occur under the condition when precipitation takes place on days closely following irrigation, which will be modelled as increased water drainage down to groundwater. In fact, farmers decide whether to irrigate depending on short-term rainfall forecasts and current soil moisture status. The present study provides a simple method to estimate spatial and temporal distributions of irrigation requirements. As for the exact irrigation frequency and volumes for farmer, there is a need for further research. Further work should be focused on more reasonable ways to calculate irrigation triggering time and quota by including consideration of the water use characteristics of different crop types and different irrigation application methods.
6.3. Suggestion for the Sustainability of Water Resources

Intensive agricultural production in Haihe Basin, especially Hebei Plain, makes the water shortage situation more and more pressing. As 70% of irrigation comes from groundwater, the groundwater table has declined drastically over several decades. According to the report from Yang et al. [33], an approximate 100-mm irrigation application in cultivated land leads to about a 640-mm groundwater decline. However, groundwater recharge is not sufficiently met by precipitation. Han et al. [28] reported that precipitation recharge is only expected when a single or successive rainfall are in excess of the minimum amount of 170 mm, a very rare condition in the piedmont region (Baoding, Shijiazhuang, Xingtai and Handan districts). While in the coastal and middle area, Cangzhou and Hengshui district, groundwater age has been found to exceed 10,000 years, this means there is only a remote possibility of confined aquifer recharge in these districts [34].

Though SNWT aims to alleviate the water shortage problem, it may not completely meet water demand in the study area (Figure 11). More strategies should be applied to stop ground water decline. One of the ideal ways is to change the current double cropping system, winter wheat and summer corn, to a single cropping system, corn. Winter wheat is the higher irrigation crop, because it is sown in the dry season. However, the Chinese population is large and needs more food to feed itself. Therefore, completely abandoning irrigated winter wheat production is impossible in the real world. The more reasonable and feasible approach is to develop water-saving agricultural technologies. This should be promoted by decision-makers, irrigation planners and agro-scientists. Hu et al. [35] suggested that a 29.2% or 135.7 mm reduction in irrigation could stop groundwater drawdown in the plain. An additional 10% reduction in irrigation pumping (i.e., a total of 39.2% or 182.1 mm) would induce groundwater recovery and restoration to the pre-development hydrologic conditions of 1956 in about 74 years. Total yield loss for wheat and maize under the appropriate irrigation timing and rate would be less than 10%. However, this should be supplemented with strong political will, water-saving incentives and water policy regulation and enforcement. The present study calculated IA based upon soil water deficits, without regard to the values of different crops received by farmers in the various districts, which translate into economic incentives for them. Where water quotas are limited, farmers will have incentives to avoid waste by putting available water to its best possible use in their farms. A market for water can provide further incentives for the most economically efficient use of this scarce resource across different production zones.

7. Conclusions

An irrigation applied (IA) estimation model together with an ET-calculating model (ETWatch) was set up to simulate the spatial distribution of IA in the Haihe Plain during the years from 2002–2007. Through comparison with the regional observed values, simulated IA shows a reasonable result with the correlation related statistics, including P, R2 and d1, all above 0.60. The present model is verified as an effective and feasible approach to simulate regional irrigation requirements and has the advantage of better reflecting local regional characteristics and at a higher resolution than the limited site-specific results upon which earlier catchment-wide estimates have been made. The simulated results show that annual ET and IA vary from 527 mm to 679 mm and 166 mm to 289 mm, with
average values of 602 mm and 225 mm, respectively, over the years from 2002 to 2007. Regions along Taihang Mountain and areas in Hebei Province have the highest ET and IA, with major problems in water resource sustainability, because of their intensive agricultural production and declining groundwater tables, even with new water expected from the SNWT project. Besides the need for more SNWT water for this area, agricultural water saving techniques are urgently required to stop groundwater decline. This study supports further development work to better integrate RS techniques, land use information ground-truthing and tools, such as the ET-calculating model ETWatch. Such an advance could provide a basis for the better understanding of the inescapable real-world trade-offs among competing economic, social and environmental goals.

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Author Contributions

Yanmin Yang, Yonghui Yang, Deli Liu and Tom Nordblom are principal authors contributing at irrigation estimation. The order of the authors reflects their level of contribution. Yanmin Yang and Yonghui Yang have written the majority of the manuscript and Deli Liu and Tom Nordblom are responsible for the paper revision in English. Bingfang Wu and Nana Yan contributed in the ET inversion by ET calculation model (ETWatch).

Conflicts of Interest

The authors declare no conflict of interest.

Reference


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