

Article

Evaluation of Different CH₄-CO₂ Replacement Processes in Hydrate-Bearing Sediments by Measuring P-Wave Velocity

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Abstract: The replacement of methane with carbon dioxide in natural gas hydrate-bearing sediments is considered a promising technology for simultaneously recovering natural gas and entrapping CO₂. During the CH₄-CO₂ replacement process, the variations of geophysical property of the hydrate reservoir need to be adequately known. Since the acoustic wave velocity is an important geophysical property, in this work, the variations of P-wave velocity of hydrate-bearing sediments were measured during different CH₄-CO₂ replacement processes using pure gaseous CO₂ and CO₂/N₂ gas mixtures. Our experimental results show that P-wave velocity continually decreased during all replacement processes. Compared with injecting pure gaseous CO₂, injection of CO₂/N₂ mixture can promote the replacement process, however, it is found that the sediment experiences a loss of stiffness during the replacement process, especially when using CO₂/N₂ gas mixtures.

Keywords: gas hydrate; CH₄-CO₂ replacement; P-wave velocity

1. Introduction

Natural gas hydrates have been found to occur naturally in large quantities on the deep ocean floor and in permafrost regions [1,2]. These natural gas hydrates are seen as a potentially vast energy resource, but an economical exploitation method has so far proven elusive [3]. Nowadays, replacement of CH_4 with CO_2 in natural gas hydrate-bearing sediments is recognized as a promising technology for the simultaneous recovery of natural gas and geological entrapment of CO_2 in hydrate form [4–6].

Based on experimental measurements and theoretical calculations, many researchers have proved the feasibility of methane exploitation by CO₂ injection [7–11]. For example, Ohgaki *et al.* [7] found the change in Gibbs free energy (ΔG) for the replacement process of CH₄ in hydrates with CO₂ was negative, which is the first theoretical evidence for the feasibility of CH₄-CO₂ replacement process. Based on the *P*-*T* phase diagrams of CH₄ and CO₂ hydrate [4,6,8–10], it can be seen that the equilibrium pressure of CO₂ hydrate is lower than that of CH₄ hydrate when the temperature is lower than 283 K. This indicates that under the same conditions CO₂ hydrate is more stable than CH₄ hydrate. Seo *et al.* [11] found for CH₄/CO₂ gas mixtures, when the molar fraction of CO₂ is higher than 40%, the molar fraction of CO₂ in hydrate phase will be more than 90%. With the decrease of pressure, the concentration of molar fraction of CO₂ and the selectivity of CO₂ to CH₄ will increase. In addition to pure CO₂, flue gas was also used for exploitation of methane. For example, Park *et al.* [12] found the final methane replacement efficiency was ranged from 71.3% to 83.4% with different CO₂/N₂ molar compositions. Recently, Koh *et al.* [13] investigated the replacement process in clay sediments. They found CO₂/N₂ gas mixtures worked better than pure CO₂.

During the gas injection, CH₄ replacement, and CH₄ production processes, we need to know the variations of the geophysical properties of the hydrate reservoir to avoid production problems like geological disasters. It is known that the acoustic wave velocity is an important geophysical property. It can provide the information of the stiffness evolution of hydrate-bearing sediments, which is related to the stability of gas hydrate reservoirs [14]. For example, compressional wave velocity, *i.e.*, P-wave velocity (V_p), could be related to elastic modulus (M) of the sediment by the following equation [15]:

$$M = \rho V_P^2 \tag{1}$$

where ρ is the density of the sediment. In our previous work [16], we have described an apparatus for measuring V_p of the hydrate bearing sediment and a series of V_p datasets were obtained for different hydrate saturations. In this work, the stability of hydrate reservoirs during the replacement process of CH₄ with CO₂ is examined by measuring and analyzing the variation of V_p .

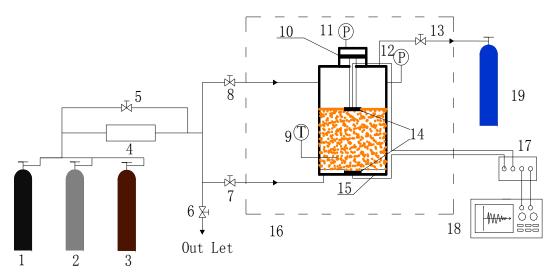
2. Experimental Section

2.1. Experimental Apparatus

A schematic outline of the experimental apparatus for measuring P-wave velocity of hydrate-bearing sediment is shown in Figure 1 [16]. The main parts of the apparatus are a high-pressure reactor, an air bath, a gas filling system, and a data collection system. The reactor is made of stainless steel and the effective volume of it is 2.0 L (φ 130 × 150). Its designed maximum working pressure is 32 MPa. Two piezoelectric transducer detectors were placed at the bottom and in the middle of the reactor, respectively. The top one was connected with a hand lever, which can move the transducer upward or downward in the vertical direction to change the distance between the two acoustic detectors within the maximum distance of 60 mm. The two transducers were connected with a pulse sender/receiver machine (5077PR, Olympus NDT, Waltham, MA, USA). The ultrasonic P-wave signals were generated by one transducer with a frequency from 500 kHz to 1.0 MHz, passed through the sandy

sediment in central axis, and received by another transducer. Then the received signal was sent back to the pulse sender/receiver machine, where it was amplified, digitized, and displayed on a digital oscilloscope (TBS2012B, Tektronix, Inc., Portland, OR, USA). A thermocouple, with a precision of ± 0.1 K, was inserted into the sediment through the wall of the reactor to detect temperature variations. Pressure was monitored by an absolute pressure transducer with an accuracy of 0.5%, which was mounted on the top of the reactor. Generally, the uncertainty of the measured P-wave velocity is always within 0.5%.

Figure 1. Schematic diagram of experimental apparatus for measuring P-wave velocity during CH₄-CO₂ replacement process. 1: N₂ cylinder; 2: CO₂ cylinder; 3: CH₄ cylinder; 4: Gas mass meter; 5–8 and 13: Valve; 9: Temperature transmitter; 10: Piston; 11 and 12: Pressure transmitter; 14: Piezoelectric transducers; 15: High-pressure reactor; 16: Air bath; 17: Ultrasonic pulse sender/receiver; 18: Digital oscilloscope; and 19: Gas recovery tank.



Unconsolidated quartz sand was used to form hydrate-bearing samples, *i.e.*, sandy sediment (shown as an orange color in Figure 1). Before being used, the raw sand was washed several times with water to eliminate any clay. Afterwards, it was dried in an oven at 393.2 K and then sieved into a series of sand samples with different particles size ranges.

2.2. Experimental Procedures

Firstly, the dry sand and the high-pressure reactor were pre-cooled to 268.2 K and the brine was cooled to 273.2 K. Then the cooled sand was homogeneously mixed with the cooled brine to have an initial saturation of brine solution of 30%. Promptly, the partially brine-saturated sandy sediment was loaded into the pre-cooled high-pressure reactor to rapidly freeze the sediment in the air bath at 268.2 K. During the mixing process, part of brine converted into ice because of low temperature of the sand. Additionally, most of water could be converted into ice in a short time during the following freezing process. By these ways, the gravity effect on the homogeneity of water could be eliminated. After 24 h, methane gas was injected from the CH₄ cylinder to form hydrate-bearing sediment in the reactor and the temperature of air bath was adjusted to the hydrate formation temperature. Starting from the injection of methane, temperature, pressure, and P-wave signals were recorded every 10 min by the

computer. When the system temperature, pressure, and P-wave velocity all became constant for a long time, the formation of methane hydrate was assumed to be finished.

After methane hydrate-bearing sample was prepared, a certain amount of methane gas was slowly released to reduce the pressure of free gas in the reactor; however, the pressure was still higher than the equilibrium pressure of methane hydrate at present temperature. When temperature and pressure did not change anymore, we closed the vent at the top of the reactor and maintained the system for another 24 h to make the formation of methane hydrate reach equilibrium again.

After 24 h, CO_2 gas or CO_2/N_2 gas mixture was injected through pipeline at the bottom of the reactor to replace free methane gas in the pore space of the sediment. To keep the pressure of gas in the reactor being slightly higher than methane hydrate equilibrium pressure at present sediment temperature, mixed gas was released with the same injection rate of the flooding gas through the vent on the top of reactor. The released gas was analyzed by a HP7890 gas chromatograph (Agilent Technologies, Santa Clara, CA, USA) and stored in the recovery tank. When the molar fraction of CH₄ in the released gas achieved lower than 2 mol %, gas injection process was stopped and the gas pressure in the reactor was rapidly adjusted to the desired replacement pressure by releasing gas from the top vent. Afterwards, all the valves connected to the reactor were shut off.

During the replacement process, variations of temperature, pressure, and P-wave signals were automatically recorded every 10 min by the computer until the end of the experimental run. The composition of gas phase was analyzed by the HP7890 gas chromatograph every 72 h. When the gas phase composition remained unchanged, we stopped the replacement process and increased the temperature of the high-pressure reactor. With the increase of temperature, hydrate started to decompose. The decomposed gas was collected and its composition was then analyzed.

2.3. Data Processing

Based on the law of the conservation of mass, by combining the molar composition of gases in the reactor and the cylinder and the experimental conditions, we can obtain the mole number of methane retained in hydrate phase $(n_{CH_a,H})$ and that replaced $(n_{CH_a,re})$ at time *t*. They were calculated as follows:

$$n_{CH_4,H} = n_{CH_4,total} - n_{CH_4,\tan k} - n_{CH_4,reactor}$$
⁽²⁾

$$n_{CH_4,re} = n_{CH_4,reactor} - n_{CH_4,reactor}^0$$
(3)

$$n_{CH_4,\tan k} = x_{CH_4,\tan k} \cdot n_{G,\tan k} \tag{4}$$

$$n_{CH_4,reactor} = x_{CH_4,reactor} \cdot n_{G,reactor}$$
(5)

$$n_{CH_4,reactor}^0 = 0.02 n_{G,reactor}^0 \tag{6}$$

where $n_{CH_4,total}$ is the total mole number of methane injected into the reactor; $n_{CH_4,reactor}$ is the mole number of methane in free gas of the reactor during the replacement process; $n_{CH_4,reactor}^0$ is the mole number of methane in free gas of the reactor before the replacement process; $n_{CH_4,tank}^0$ is the mole number of methane in the recovery tank; $x_{CH_4,tank}$ is the molar fraction of methane in recovery tank, and $x_{CH_4,reactor}$ is the molar fraction of methane in free gas of the reactor during the replacement process. The total mole number of free gas, $n_{G,reactor}$, in the reactor during the replacement process were calculated by the following formula:

$$n_{G,reactor} = \frac{P_{reactor}V_{G,reactor}}{ZRT_{reactor}}$$
(7)

where $P_{reactor}$ is the pressure of the reactor; $V_{G,reactor}$ is the volume of the free gas in the reactor; $T_{reactor}$ is temperature of the reactor; R is the universal gas constant; and Z is the compressibility factor corresponding to $T_{reactor}$, $P_{reactor}$, and gas phase composition. Z was calculated with BWR equation of state.

Like to $n_{G,reactor}$, the mole number of free gas in the reactor, $n_{G,reactor}^0$, at the beginning of the replacement process and that in the recovery tank, $n_{G,tan k}$, could also be calculated. At the beginning of the replacement process, it was assumed that the molar fraction of methane in gas mixture is 0.02.

To understand the kinetics of the replacement process better, in this work, the replacement rate was defined as:

$$r_{\rm Re} = \frac{1}{n_{CH_4,H}^i} \frac{dn_{CH_4,H}}{dt}$$
(8)

The efficiency of the replacement process, η , was defined as:

$$\eta(\%) = \frac{n_{CH_4, re}}{n_{CH_4, H}^i} \times 100\%$$
(9)

where $n_{CH_4,H}^i$ is the initial mole number of methane in hydrate phase. The data is summarized in Table 1.

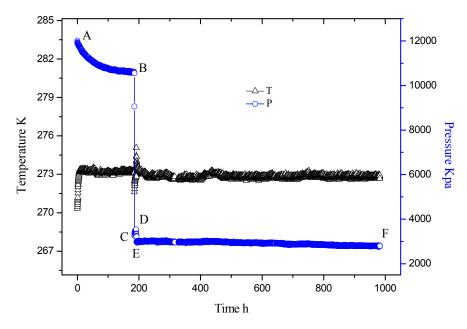
Run	Experimental conditions for methane hydrate formation			Experimental conditions for CO ₂ -CH ₄ replacement process			
	S_{w} (%)	<i>T</i> ₁ (K)	P_1 (MPa)	$T_2(\mathbf{K})$	P_2 (MPa)	$n^i_{CH_4,H}$ (mol)	S_h (%)
1	30	273.2	12.01	273.2	3.0	0.97	29.1
2	30	273.2	8.10	277.2	3.5	0.91	28.1
3	30	273.2	8.60	273.2	5.0	0.94	28.5
4	30	273.2	8.84	273.2	5.1	0.84	27.1

Table 1. Experimental conditions for methane hydrate formation and CO_2 -CH₄ replacement in different runs.

3. Results and Discussion

To investigate the influence of the type of injected gas on the efficiency of CH_4 - CO_2 replacement, we prepared four hydrate bearing sediment samples using the same sand with a sand grain diameter range of 0.30 to 0.45 mm and a porosity of 37.8 vol %. Their properties such as the hydrate saturation are close to each other. The experimental conditions for methane hydrate formation and CH_4 - CO_2 replacement processes as well as the properties of hydrate-bearing sediment samples were listed in Table 1. T_1 and P_1 are the initial temperature and pressure at which hydrate-bearing sediment was formed. T_2 and P_2 are temperature and the initial pressure at which the replacement experiment was carried out. S_w is the initial saturation of water in sediment before hydrate formation, which was defined as the fraction of pore volume occupied by water. S_h is the saturation of hydrate in the prepared hydrate bearing sediment, which was defined as the fraction of pore volume occupied by hydrate. $n_{CH_4,H}^i$ is the initial mole number of methane in the hydrate phase before the replacement. As shown in Table 1, S_h is close to S_w in each experiment run, which indicates that most of water was converted into hydrate. As an example, the evolution of temperature and pressure in the formation of hydrate bearing sediment in experiment run 1 was shown in Figure 2. From it, one can see that gas pressure in reactor decreased with as hydrate formation proceeded. At the beginning stage of hydrate formation, the sediment temperature increased rapidly because the air bath temperature was higher than the initial temperature of the sediment. However, when it reaches the air bath temperature, the sediment temperature remains nearly constant during the long period of hydrate formation. This can be attributed to the reason that hydrate was formed from ice in this work and the heat released in the hydrate formation from liquid water.

Figure 2. The evolvement of system pressure and temperature with the elapsed time during experiment Run 1. A–B, period of hydrate formation; B–C, depressurization above dissociation pressure of methane hydrate; C–D, flooding process of free CH₄ with CO₂; D–E, beginning stage of the replacement of CH₄ with CO₂; and E–F, long stable period of replacement.



Gaseous CO_2 was used as injected gas in the first two experimental runs, while $CO_2 + N_2$ gas mixtures were used in runs 3 and 4. The concentration of N_2 in gas mixtures in runs 3 and 4 were 78 mol % and 96 mol %, respectively.

The replacment processes using pure gaseous CO_2 were carried out at two different temperatures, 273.2 K and 277.2 K, respectively, in experimental runs 1 and 2. The corresponding experimental results were given in Figures 2–5. Figure 2 shows the evolution of the system pressure and temperature with the elapsed time during experiment run 1. One can see there was a sudden decrease of temperature when the free gas pressure decreased (from point B to point C in Figure 2) because of the

Joule-Thomson effect. However, it rebounded in a short time. On the other hand, there was a sudden increase of temperature during the flooding process of free CH_4 with CO_2 . This could be attributed to the fact that the flooding CO_2 temperature is higher than that of sediment. Correspondingly, the pressure also increased a little. At the beginning stage of replacement, there was a rapid decrease of pressure, which corresponds to a rapid increase of P-wave velocity as shown in the following Figure 3. This phenomenon indicates that part of water produced from the dissociation of methane hydrate in flooding process formed hydrate again with CO_2 . During the long period of the replacement process, there was no obvious change in temperature. This phenomenon demonstrates the heat released in the formation of CO_2 hydrate is balanced by the heat absorbed in the dissociation of methane hydrate during the replacement process. However, the pressure in the reactor decreased very slowly during the replacement process. This phenomenon might be attributed to the fact that there was interstitial water in the hydrate and it might form hydrate because the pressure is higher than the equilibrium hydrate formation pressure of CO_2 .

Figure 3. Variation of P-wave velocity with elapsed time during methane hydrate formation and CH₄-CO₂ replacement processes. (**a**) Run1 at 273.2 K and (**b**) run 2 at 277.2 K.

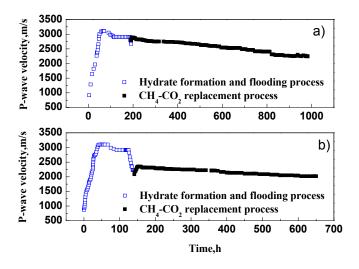


Figure 3 shows the variations of P-wave velocity with elapsed time during methane hydrate formation and CH₄-CO₂ replacement processes in experimental runs 1 and 2. As can be seen from Figure 3, the variations of P-wave velocity have similar trends in the two experimental runs. In the methane hydrate formation stage, P-wave velocity firstly increased to a maximum value, then decreased slightly, and finally tended to be constant in both runs. During the depressurization and flooding process of free methane gas, *i.e.*, injecting CO₂ from the bottom and releasing gas mixture from the top of the reactor, the P-wave velocity decreased in both runs. This phenomenon demonstrates that the flooding process can lead to the dissociation of methane hydrate. As there was a heating process in run 2 in order to increase temperature of sediment from 273.2 to 277.2 K, more hydrate dissociated and a bigger drop of the P-wave velocity rebounded slightly. This phenomenon suggests that part of water produced from the dissociation of methane hydrate again with CO₂. For experiment run 1, almost all the loss of the P-wave velocity caused by flooding was recovered in this rebounding process. Although the rebounding strength is larger in run 2, only a small part of the

P-wave velocity loss was recovered in run 2 because this loss is too large at higher temperature. In the following long standing CO_2 -CH₄ replacement process, the P-wave velocity decreased slowly with elapsed time. This implies a reduction of elastic modulus of hydrate-bearing sediment during CH₄-CO₂ replacement because P-wave velocity (V_p) could be related to elastic modulus (M) of the sediment by Equation (1). The replacement rate and efficiency at different times for the experimental runs 1 and 2 were calculated using Equations (8) and (9) and are plotted in Figures 4 and 5, respectively.

Figure 4. Variation of the replacement rate with elapsed time at different replacement temperatures.

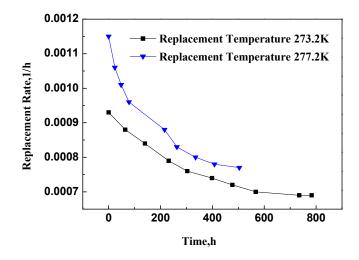
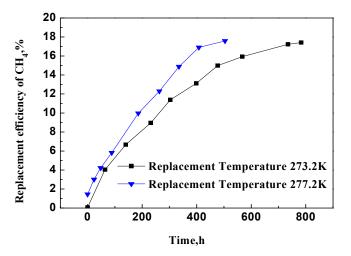


Figure 5. Replacement efficiency of CH₄ with elapsed time during the replacement process at different replacement temperatures.

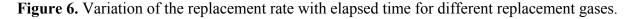


From Figure 4 we can see that replacement rate of methane decreased from $9.3 \times 10^{-4} \text{ h}^{-1}$ to $6.9 \times 10^{-4} \text{ h}^{-1}$ at 273.2 K, while it decreased from $1.15 \times 10^{-3} \text{ h}^{-1}$ to $7.7 \times 10^{-4} \text{ h}^{-1}$ at 277.2 K, indicating that the higher the temperature, the higher the rate of the methane replacement. This can be attributed to the fact that the dissociation rate of methane hydrate is higher at higher temperature. Bai *et al.* [17] demonstrated that the CH₄-CO₂ replacement process is composed of two sequential processes, *i.e.*, the dissociation of methane hydrate and the formation of a CH₄-CO₂ double hydrate. Thus the higher dissociation rate results in a total higher replacement rate. This characteristic of the

 CH_4 - CO_2 replacement process requires that the formed CH_4 - CO_2 double hydrate be of enough porosity for dissociated CH_4 molecules to escape and for CO_2 molecules to diffuse into the interface of methane hydrate and CH_4 - CO_2 double hydrate. Thus, as the replacement proceeds, the porosity of the hydrate-bearing sediment increased and led to the decrease of P-wave velocity as shown in Figure 3.

As shown in Figure 5, although the replacement rate is higher at higher temperature, the influence of temperature on the final replacement efficiency is not obvious. At 273.2 K, the highest efficiency is 17.4% and it is 17.6% at 277.2 K. This conclusion is consistent with that obtained by Qi *et al.* [18] Using a pure water system, they investigated the influence of temperature and pressure on the CO_2 and CH_4 replacement process. Their experimental results show that within a certain range, increasing the replacement temperature can promote the replacement process, but the influence on replacement efficiency is not obvious.

Figures 6–8 show the comparisons between experimental results of run 1 and those of run 3. As denoted before, pure gaseous CO_2 was used as injected gas in experimental run 1 and $CO_2 + N_2$ gas mixture with 78 mol % of N_2 was used in run 3.



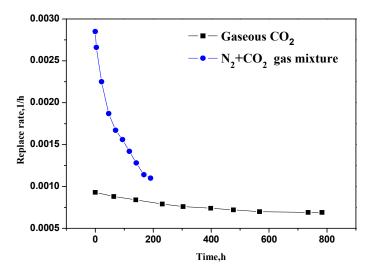


Figure 7. Variation of the replacement efficiency of CH_4 with elapsed time for different replacement gases.

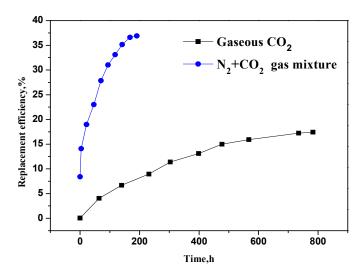
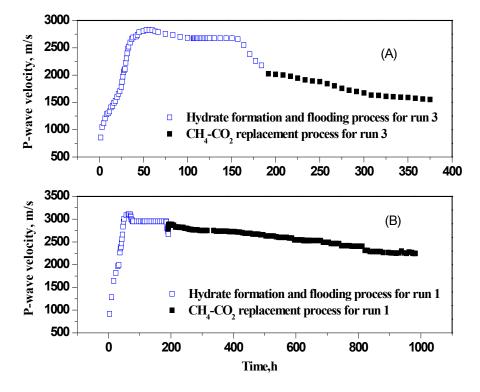


Figure 8. Comparison of variation of P-wave velocity during methane hydrate formation and replacement process using (**A**) 22 mol % CO₂ + 78 mol % N₂ and (**B**) pure gaseous CO₂.

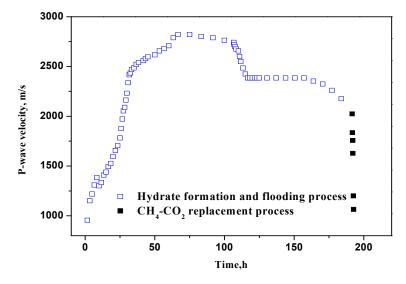


We can see from Figures 6 and 7 that when we used a $CO_2 + N_2$ mixture for the replacement of methane, the replacement rate is much higher than that using pure gaseous CO_2 . The highest replacement efficiency reached 36.9% when we used a $CO_2 + N_2$ mixture, which is much higher than that using pure gaseous CO_2 (17.4%). The reason for this might be that, compared with pure CO_2 , the hydrate formation pressure of $CO_2 + N_2$ gas mixture is much higher at the same temperature. Thus, in the beginning stage, the water produced from the dissociation of methane hydrate could not form hydrate remains at a high level until the molar fraction of methane in the gas mixture reaches a value at which gas mixture can form hydrate with water. This can be demonstrated by Figure 8.

Figure 8 shows a comparison of the variation laws of P-wave velocity in experimental runs 1 and 3. One can see that the decrease of the P-wave velocity is much bigger during the flooding period when using $CO_2 + N_2$ gas mixture, indicating more methane hydrate is dissociated in this case. Additionally, unlike the case using pure CO_2 , no rebound of P-wave velocity was observed when the flooding ended, indicating that $CO_2 + N_2$ gas mixture cannot form hydrate with the dissociated water during the beginning stage of the replacement process. The final P-wave velocity in run 3 is 1554 m/s, which is lower than the values obtained in experimental runs 1 and 2. All these results indicate that when we use CO_2/N_2 gas mixtures for replacing methane, although a higher replacement rate and efficiency could be achieved, the sediment stability decreases to some extent.

In order to further demonstrate the mechanism described above, we performed another experimental run (run 4) for the replacement of methane using a $CO_2 + N_2$ gas mixture containing only 6 mol % of CO_2 . The P-wave velocity profiles obtained in this run are given in Figure 9.

Figure 9. Variations of P-wave velocity as a function of elapsed time during methane hydrate formation and replacement process for run 4.



From Figure 9 we can see a big drop of the P-wave velocity, indicating a remarkable dissociation of methane hydrate occurred during the flooding period. More significantly, Figure 9 shows that the P-wave velocity decreased rapidly during the replacement process. About 2.2 h later, the acoustic signal disappeared completely, indicating that most methane hydrate was dissociated and no hydrate reformed by the gas mixture of CO_2 , N_2 and CH_4 because of the high concentration of N_2 in it. This is because the partial pressure of CO_2 and CH_4 in the gas phase is low. The dissociation of hydrate in sediment results in the rapid decrease of the sediment stiffness. The effect of the sediment stiffness on the attenuation of acoustic energy is obvious, so the P-wave velocity decreased quickly. Ultimately we are unable to collect the acoustic signal. The final replacement efficiency in this group of experiment is 67%, which is significantly higher than that in other experimental runs; however, from Figure 9 we can see that when the mole fraction of N_2 in gas mixture is very high, so the stability of the sediment could not be guaranteed.

Based on the Chen-Guo hydrate model [19], the CO₂ hydrate equilibrium pressures were calculated to be 1.21 MPa and 1.96 MPa at 273.2 K and 277.2 K, respectively. The initial formation pressures of $CO_2 + N_2$ double hydrate at 273.2 K were calculated to be 4.91 MPa and 14 MPa for the gas mixtures containing 78 mol % N₂ and 96 mol % N₂, respectively. Thus, it is possible for injected gases to form hydrate in experiment runs 1 to 3 because the replacement pressures, as listed in Table 1, are higher than their initial hydrate formation pressure. However, it is impossible for injected gas in experiment run 4 to form hydrate as its initial hydrate formation pressure is much higher than the specified replacement pressure of 5.1 MPa. In this case, the replacement process could not occur. What occurred was only the dissociation of methane hydrate because of the low partial pressure in the free gas phase caused by the injection of $CO_2 + N_2$ gas mixture. For experiment run 3, also part of the water from the dissociation of methane hydrate could also not form hydrate as the concentration of CO_2 in the free gas might reach the replacement pressure. Thus, part of the replacement efficiency was contributed by the pure dissociation of methane hydrate in this run. In this case, the saturation of hydrate in the sediment would decrease. That is why the final P-wave velocity was smaller than in those cases (runs 1 and 2)

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using pure CO₂ in. In order to increase the contribution of pure replacement and maintain the sediment stability, a higher replacement pressure should be adopted.

4. Conclusions

In this work, a series of experiments were carried out to measure the variation of P-wave velocity of methane hydrate-bearing sediments during the CH_4 - CO_2 replacement process. The effects of different experimental conditions on rate and efficiency of the replacement process were systematically studied. Based on our results, we found both P-wave velocity and the elastic modulus of hydrate-bearing sediments decreased gradually during the CH_4 - CO_2 replacement process. When the replacement process is finished, the sediment system can still maintain a certain stability. In the replacement process, the higher the temperature, the higher the rate of the replacement of methane; however, the influence of temperature on replacement efficiency is not obvious. In addition, we found both rate and efficiency of the replacement process can be improved significantly by injecting CO_2 and N_2 mixtures instead of injecting pure gaseous CO_2 . When the molar fraction ratio of N_2 and CO_2 is around 8:2, the methane replacement efficiency can be as high as 36.9% in 190 h. Though the elastic modulus of the sediment system decreased during the replacement process, the system can still maintain a certain stability. The knowledge obtained in this work could be useful for understanding and applying CH_4 - CO_2 replacement technology for simultaneously recovering natural gas and entrapping CO_2 .

Acknowledgments

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Conflicts of Interest

The authors declare no conflict of interest.

References

- 1. Klauda, J.B.; Sandler, S.I. Global distribution of methane hydrate in ocean sediment. *Energy Fuels* **2005**, *19*, 459–470.
- 2. Reagan, M.T.; Moridis, G.J. Large-scale simulation of methane hydrate dissociation along the West Spitsbergen Margin. *Geophys. Res. Lett.* **2009**, *36*, L23612.
- 3. Liu, B.; Yuan, Q.; Su, K.H.; Yang, X.; Wu, B.C.; Sun, C.Y.; Chen, G.J. Experimental simulation of the exploitation of natural gas hydrate. *Energies* **2012**, *5*, 466–493.
- 4. Ota, M.; Saito, T.; Aida, T.; Watanabe, M.; Sato, Y.; Smith, R.L.; Inomata, H. Macro and microscopic CH₄-CO₂ replacement in CH₄ hydrate under pressurized CO₂. *AIChE J.* **2007**, *53*, 2715–2721.
- 5. Zhou, X.T.; Fan, S.S.; Liang, D.Q.; Du, J.W. Replacement of methane from quartz sand-bearing hydrate with carbon dioxide-in-water emulsion. *Energy Fuels* **2008**, *22*, 1759–1764.

- 6. Jung, J.W.; Espinoza, D.N.; Santamarina, J.C. Properties and phenomena relevant to CH₄-CO₂ replacement in hydrate-bearing sediments. *J. Geophys. Res.* **2010**, *115*, B10102.
- Ohgaki, K.; Takano, K.; Sangawa, H.; Matsubara, T.; Nakano, S. Methane exploitation by carbon dioxide from gas hydrates—Phase equilibria for CO₂-CH₄ mixed hydrate system. *J. Chem. Eng. Jpn.* 1996, *29*, 478–483.
- 8. Ota, M.; Abe, Y.; Watanabe, M.; Smith, R.L.; Inomata, H. Methane recovery from methane hydrate using pressurized CO₂. *Fluid Phase Equilib*. **2005**, *228*, 553–559.
- 9. Ota, M.; Morohashi, K.; Abe, Y.; Watanabe, M.; Smith, R.L.; Inomata, H. Replacement of CH₄ in the hydrate by use of liquid CO₂. *Energy Convers. Manag.* **2005**, *46*, 1680–1691.
- Uchida, T.; Ikeda, I.Y.; Takeya, S.; Kamata, Y.; Ohmura, R.; Nagao, J.; Zatsepina, O.Y.; Buffett, B.A. Kinetics and stability of CH₄-CO₂ mixed gas hydrates during formation and long-term storage. *ChemPhysChem* 2005, *6*, 646–654.
- 11. Seo, Y.T.; Lee, H.; Yoon, J.H. Hydrate phase equilibria of the carbon dioxide, methane, and water system. *J. Chem. Eng. Data* **2001**, *46*, 381–384.
- Park, Y.; Kim, D.Y.; Lee, J.W.; Huh, D.G.; Park, K.P.; Lee, J.; Lee, H. Sequestering carbon dioxide into complex structures of naturally occurring gas hydrates. *Proc. Natl. Acad. Sci. USA* 2006, *103*, 12690–12694.
- 13. Koh, D.Y.; Kang, H.; Kim, D.O.; Park, J.; Cha, M.; Lee, H. Recovery of methane from gas hydrates intercalated within natural sediments using CO₂ and a CO₂/N₂ gas mixture. *ChemSusChem* **2012**, *5*, 1443–1448.
- 14. Li, D.L.; Wang, D.; Liang, D.Q. P-wave of hydrate-bearing sand under temperature cycling. *Geophysics* **2011**, *76*, E1–E7.
- 15. Dvorkin, J.; Berryman, J.; Nur, A. Elastic moduli of cemented sphere packs. *Mech. Mater.* **1999**, *31*, 461–469.
- Zhang, Q.; Li, F.G.; Sun, C.Y.; Li, Q.P.; Wu, X.Y.; Liu, B.; Chen, G.J. Compressional wave velocity measurements through sandy sediments containing methane hydrate. *Am. Mineral.* 2011, *96*, 1425–1432.
- 17. Bai, D.S.; Zhang, X.R.; Chen, G.J.; Wang, W.C. Replacement mechanism of methane hydrate with carbon dioxide from microsecond molecular dynamics simulations. *Energy Environ. Sci.* **2012**, *5*, 7033–7041.
- 18. Qi, Y.X.; Yu, Z.G.; Yan, H.; Zhou, S.P.; Liu, Y.F. Experimental studies of the thermodynamic effects on the replacement of methane hydrate by CO₂. *Chem. Eng. Oil Gas* **2012**, *41*, 285–288, (in Chinese).
- 19. Chen, G.J.; Guo, T.M. A new approach to gas hydrate modeling. Chem. Eng. J. 1998, 71, 145–151.

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